

1    **Soil–climate interactions drive above-ground biomass in the Caatinga, the largest Neotropical seasonally  
2    dry tropical forest**

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78

79 **Abstract**

80 *Background and Aims:* Soil properties are key drivers of vegetation structure, yet their influence on above-  
81 ground woody biomass (AGB<sub>W</sub>) in seasonally dry tropical forests (SDTFs) remains underexplored, particularly  
82 at larger scales. This gap is evident in the Caatinga, Latin America's largest SDTF, known for its biodiversity  
83 and carbon storage potential. We investigated relationships among soil, climate, and vegetation properties to  
84 understand accumulation patterns of AGB<sub>W</sub> in SDTFs.

85 *Methods:* We used standardised soil and vegetation data from 29 research plots spanning diverse geological  
86 and floristic conditions. Linear mixed models and multi-model inference were applied to analyse relationships  
87 between AGB<sub>W</sub> and environmental variables, including soil texture, fertility, plant-available soil water, mean  
88 annual precipitation (MAP), temperature, and climatic water deficit (CWD). Structural equation modelling  
89 (SEM) was utilised to assess how environmental variables influenced community-weighted maximum stem  
90 diameter, wood density, functional richness, and their combined effects on AGB<sub>W</sub>.

91 *Results:* AGB<sub>W</sub> was influenced by MAP, soil fertility, maximum plant-available soil water, and CWD. SEM  
92 indicated that soil nutrient availability shaped community functional traits, reflecting trade-offs between growth  
93 and water-use strategies. In turn, species' maximum stem diameter and, to a lesser extent, functional richness  
94 positively influenced AGB<sub>W</sub>, underscoring the role of soil-mediated functional traits in determining biomass.

95 *Conclusion:* AGB<sub>w</sub> in the Caatinga is shaped by soil, climate, and their interactions, with soil properties exerting  
96 strong effects on community functional diversity. Our findings highlight patterns of functional trait variability  
97 and biomass storage, offering insights for biodiversity conservation and carbon sequestration in SDTFs under  
98 global environmental change.

99 **Keywords:** Brazilian semi-arid, carbon stocks, drylands, functional traits, global change, dryland soils

100

101 **Introduction**

102 Above-ground woody biomass (AGB<sub>w</sub>) is a key component of the carbon cycle in forest systems, as  
103 it integrates productivity, recruitment, and mortality dynamics (Lloyd et al. 2009). Although seasonally dry  
104 tropical forests (SDTFs) typically have lower AGB<sub>w</sub> stocks than their wetter counterparts, they are important  
105 carbon reservoirs due to their widespread distribution in the tropics (Glenday 2008; Roa-Fuentes et al. 2012;  
106 Corona-Núñez et al. 2018). While this work focuses on AGB<sub>w</sub>, which typically accounts for approximately 60–  
107 70% of the total biomass per unit area in SDTFs, the remaining portion is represented by below-ground biomass  
108 (BGB) (Murphy and Lugo 1986; Menezes et al. 2021), underscoring the importance of both pools for carbon  
109 storage in these ecosystems. Once estimated to comprise 42% of all subtropical and tropical forests (Murphy  
110 and Lugo 1986), SDTFs are now experiencing significant declines, with an 11.4% global loss in cover from  
111 2001 to 2020 (Ocón et al. 2021). These ecosystems have been recognised as highly diverse yet threatened (Miles  
112 et al. 2006; DRYFLOR 2016). The Caatinga region in Brazil, which hosts the largest continuous expanse of  
113 SDTF in Latin America, harbours substantial biodiversity (Queiroz et al. 2017; Fernandes et al. 2020; Londe et  
114 al. 2023) and has significant carbon storage potential (Castanho et al. 2020), highlighting its ecological  
115 importance. However, like other neotropical SDTFs, the Caatinga is under threat from various pressures and  
116 lacks conservation efforts (Oliveira et al. 2012; DRYFLOR 2016). The Caatinga SDTFs have long faced  
117 anthropogenic pressures, including firewood and charcoal extraction, cattle raising and overgrazing, and slash-  
118 and-burn agriculture (Andrade 1977; Araujo et al. 2023). These activities have pushed many previously forested  
119 areas towards ecological thresholds, with only 11% of the original forest coverage remaining, while some areas  
120 are desertified or at risk of desertification (Araujo et al. 2023). Despite the overwhelming influence of human  
121 activities on AGB<sub>w</sub> distribution in the region, it remains crucial to investigate how environmental factors shape

122 AGB<sub>w</sub> in structurally mature stands, as these can offer insights into potential AGB<sub>w</sub> accumulation under semi-  
123 natural conditions.

124 Despite the likely influence of both climate and soils on AGB<sub>w</sub>, research has primarily focused on  
125 climatic factors, with soil properties often underrepresented. Among climatic variables, mean annual  
126 precipitation is widely recognised as a primary driver of AGB<sub>w</sub>, and many studies have associated biomass  
127 accumulation with rainfall gradients (e.g., Brown and Lugo 1982; Becknell et al. 2012; Castanho et al. 2020).  
128 However, plant water availability is also influenced by evapotranspiration, precipitation seasonality, and soil  
129 properties. While the positive correlation between mean annual precipitation and AGB<sub>w</sub> is well-reported, few  
130 studies have taken soil attributes, such as soil texture, clay mineralogy, and nutrient levels, into account, mainly  
131 due to incompatible sampling and analysis protocols or simply due to the absence of soil data (Becknell et al.  
132 2012; Santos et al. 2023). This gap in knowledge limits our understanding of how environmental drivers interact  
133 to shape AGB<sub>w</sub> in SDTFs, particularly in the spatially complex Caatinga region.

134 In this region, distinct geological substrates have given rise to a variety of Reference Soil Groups  
135 (RSGs, the highest categorical level in the WRB–FAO soil classification), ranging from nutrient-rich shallow  
136 soils overlying carbonate rocks to fertile, fine-textured shallow soils over crystalline basements, to less fertile,  
137 deeper soils developed from sedimentary deposits (Sampaio 1995; Oliveira 2011). Earlier studies have  
138 highlighted the significant role of soil properties in shaping vegetation structure and floristic composition in  
139 Brazilian dry forests (e.g., Ratter et al. 1973; 1978; Furley and Ratter 1988). In the Caatinga, soil properties  
140 have been linked to local variations in structural and floristic diversity (e.g., Souza et al. 2019; Maia et al. 2020).  
141 At the regional scale, the combination of soil properties and climate attributes has been shown to more  
142 effectively predict differences in vegetation physiognomies than either soil or climate separately (Oliveira et al.  
143 2019).

144 The relationship between vegetation and environmental conditions can also be explored through the  
145 lens of functional attributes, and associated community functional properties, such as community-weighted  
146 mean maximum stem diameter (CWM<sub>DMAX</sub>), wood density (CWM<sub>WD</sub>), and their derived functional richness  
147 (F<sub>RIC</sub>), i.e., the range and diversity of single or combined traits within each community. These traits have been  
148 shown to influence stand-level biomass and productivity in dry ecosystems (Prado-Junior et al. 2016). In their

149 study, Prado-Junior et al. (2016) tested contrasting ecological hypotheses to explain biomass accumulation  
150 patterns in SDTFs, including the ‘biomass ratio hypothesis’ (Grime 1998), which suggests that the dominant  
151 traits in a community exert the greatest influence on stand-level ecosystem properties; the ‘niche  
152 complementarity hypothesis’ (Tilman et al., 1999), which proposes that species can coexist by using resources  
153 differently, thereby reducing competition; and the ‘soil fertility hypothesis’ (Pastor et al. 1984),  
154 comprehensively tested in this study.

155 Here, we use a *Space-for-Time* approach (Pickett, 1989), which involves examining spatial variation  
156 across environmental gradients as a proxy for temporal ecological changes. This approach is particularly useful  
157 in the Caatinga, where long-term monitoring studies are limited, but pronounced environmental heterogeneity  
158 may reflect ecosystem development over time or responses to long-term drivers. By comparing plots distributed  
159 across climatic and edaphic gradients, we aim to infer how these factors shape current patterns of biomass  
160 accumulation and functional diversity.

161 Specifically, we address the following research questions: (1) How do soil, climate, and their potential  
162 interactions modulate patterns of biomass accumulation in the Caatinga region? (2) How does soil influence  
163 stand-level functional properties, such as wood density, maximum stem diameter, and the associated functional  
164 richness? (3) How do these functional properties affect stand-level above-ground woody biomass? By  
165 integrating standardised soil and vegetation data from 29 research plots across the Brazilian Caatinga, this study  
166 seeks to deepen our understanding of the environmental and functional drivers of biomass in SDTFs, offering  
167 valuable insights for the conservation and management of these ecosystems in the context of global change.

168

## 169 **Material and Methods**

### 170 **Study sites**

171 Data used in this study were compiled from 29 study plots established by the *Nordeste Project* (see  
172 Acknowledgements and Funding for further details). These plots are distributed across the Brazilian Caatinga  
173 region (Table 1, Fig. 1), encompassing seven reference soil groups (RSGs), three geological substrate types,  
174 and distinct mineral assemblages. These mineral assemblages are represented by high-activity clay soils (HAC),  
175 low-activity clay soils (LAC), and highly sandy soils (Arenic), as described below. The sampled soils vary from

176 shallow, slightly weathered soils mostly developed from crystalline rocks ( $S_{CRY}$ ) to much deeper, highly  
177 weathered soils overlying sedimentary substrates ( $S_{SED}$ ), including two study plots located in the Quaternary  
178 dunes of the middle *São Francisco River*. Additionally, the sampling included three vegetation stands on soils  
179 derived from carbonate rocks ( $S_{CAR}$ ), characterised by distinctive properties such as elevated calcium (Ca) and  
180 phosphorus (P) levels, and neutral to basic soil pH. Examples of sampled soil and vegetation are displayed in  
181 Fig. 1 and Figs. 1, 2, and 3 of Online Resource 1. The average mean annual precipitation (MAP) for the study  
182 plots was 802 mm yr<sup>-1</sup>, ranging from 510 mm yr<sup>-1</sup> to 1363 mm yr<sup>-1</sup>, whereas the average mean annual  
183 temperature (MAT) was 23.8 °C, ranging from 20.5 °C to 26.8 °C. The study plots had an average elevation of  
184 535 m asl, varying from 99 to 944 m asl (Table 1). Vegetation structure ranged from open canopies 4–7 m in  
185 height to closed canopies 25–30 m tall. Study sites consist of well-conserved structurally mature stands, though  
186 sporadic grazing and occasional timber logging cannot be fully disregarded in a few study plots. Most plots  
187 were established on flat terrain, with some on slightly sloping reliefs. A more detailed version of Table 1 with  
188 environmental and vegetation data is available in Table 1 of Online Resource 1. Vegetation inventory and soil  
189 sampling were conducted during three fieldwork campaigns in 2017, 2018, and 2019, as part of the inventory  
190 carried out by the *Nordeste Project fieldwork team*. In all years, sampling was consistently carried out during  
191 the late wet season to capture vegetation at its maximum vegetation development stage (Moonlight et al. 2021).

Table 1: Study plots, selected environmental and vegetation data: plot code; latitude; longitude; APS = average plot slope; F = flat; SS = slightly sloping; geology; SED = sedimentary; CRY = crystalline; CAR = carbonate; RSG = reference soil group; MAP = mean annual precipitation; MAT = mean annual temperature; elevation; AGB<sub>W</sub> = above-ground woody biomass; clay (fraction); [N]<sub>T</sub> = soil total nitrogen; [P]<sub>T</sub> = soil total phosphorus; pH<sub>H2O</sub> = water-measured soil pH; [Ca]<sub>ex</sub> = soil exchangeable calcium. Study plots are ordered according to increasing MAP. Soil data refers to the upper 0.3 m from the soil surface. Original vegetation and soil data are integrated into the *ForestPlots Network* ([www.ForestPlots.net](http://www.ForestPlots.net)).

Plot Code	Latitude	Longitude	APS	Geology	RSG	MAP (mm)	MAT (C°)	Elevation (m)	AGB <sub>W</sub> (Mg ha <sup>-1</sup> )	Clay (fraction)	[N] <sub>T</sub> (mg g <sup>-1</sup> )	[P] <sub>T</sub> (mg kg <sup>-1</sup> )	pH <sub>H2O</sub>	[Ca] <sub>ex</sub> (mmol <sub>c</sub> kg <sup>-1</sup> )
CND-01	-9.97	-39.01	F	SED	Arenosol	512	22.7	535	20.9	0.04	2.5	69	4.66	1.5
GBR-02	-11.02	-41.41	F	CAR	Leptosol	515	22.9	637	47.3	0.15	16.8	1194	7.48	55.8
GBR-01	-11.01	-41.44	F	CAR	Cambisol	519	23.3	564	74.8	0.22	12	469	7.77	55.7
LGE-01	-9.05	-40.32	F	CRY	Luvisol	591	25.1	390	10.2	0.12	11.9	148	5.47	17.4
MOR-02	-11.50	-41.35	SS	SED	Leptosol	591	20.8	907	15.8	0.16	19	223	4.17	3.4
CGR-01	-7.28	-35.98	F	CRY	Luvisol	599	22.8	487	14.1	0.18	8.8	162	4.71	8.6
MOR-01	-11.49	-41.33	F	SED	Arenosol	602	20.5	944	18.8	0.06	16.8	51	4.49	7.7
IBD-02	-10.79	-42.78	F	SED	Arenosol	684	25.6	411	19.7	0.02	3.3	92	5.68	3.3
IBD-01	-10.79	-42.82	SS	SED	Arenosol	696	25.5	421	18.5	0.01	6.1	61	5.74	5.9
BVT-01	-12.73	-40.71	SS	CRY	Acrisol	724	22.2	495	22.5	0.24	12.1	52	4.32	4.6
SET-01	-7.97	-38.38	F	CRY	Luvisol	752	23.7	472	27.8	0.08	6.3	144	6.13	30.8
SCP-02	-8.86	-42.68	F	SED	Acrisol	768	25.6	487	20.6	0.15	20.5	89	4.3	0.9
MCS-02	-13.06	-42.52	F	CRY	Arenosol	782	24.4	545	33.9	0.06	5.1	43	5.67	12.3
SCP-01	-8.86	-42.70	F	SED	Cambisol	786	25.4	529	11.7	0.21	7.1	302	4.27	2.5
MCS-01	-13.00	-42.71	SS	CRY	Leptosol	789	23.1	770	46.9	0.12	13.4	103	4.54	2.9
PAT-01	-7.01	-37.40	F	CRY	Luvisol	792	26.2	282	18.3	0.18	10.1	204	5.86	25.9
SJO-01	-8.81	-36.41	F	CRY	Arenosol	792	21.4	670	12.7	0.06	6.6	164	5.15	5.3
PAT-02	-7.02	-37.40	F	CRY	Luvisol	808	26.2	283	21.2	0.15	5.9	163	6.27	25.1
SDA-03	-5.12	-40.87	F	CRY	Luvisol	815	25.5	309	55.6	0.11	9.5	163	5.97	30
CJU-01	-14.97	-43.92	F	SED	Arenosol	825	24.2	470	4.9	0.07	6.7	15	4.77	2.5
PFF-01	-5.04	-37.52	F	CAR	Calcisol	864	26.8	99	33.6	0.42	20	360	7.89	65.6
JUV-01	-14.43	-44.16	SS	SED	Luvisol	900	24.2	518	85.7	0.28	11.4	176	6.14	27.4
CTI-01	-14.22	-42.53	SS	CRY	Regosol	939	20.9	938	60.9	0.15	10.7	28	4.38	3.6
SDA-02	-5.14	-40.91	F	SED	Leptosol	969	23.3	640	48.7	0.22	13.2	181	4.21	0.4

SDA-01	-5.15	-40.93	F	SED	Arenosol	973	23	682	22.2	0.07	12.1	129	4.51	0.8
ARI-04	-7.36	-39.48	F	SED	Acrisol	1011	21.4	899	10.6	0.29	15	101	4.47	1.8
ARI-03	-7.27	-39.45	SS	SED	Alisol	1081	22	796	83.2	0.12	9.5	80	4.22	4.3
BTI-01	-3.36	-41.74	SS	SED	Leptosol	1222	26.6	102	46	0.17	28.9	282	4.65	17.2
PSC-02	-4.13	-41.68	F	SED	Regosol	1363	26	238	37	0.09	7.8	156	4.58	0.9

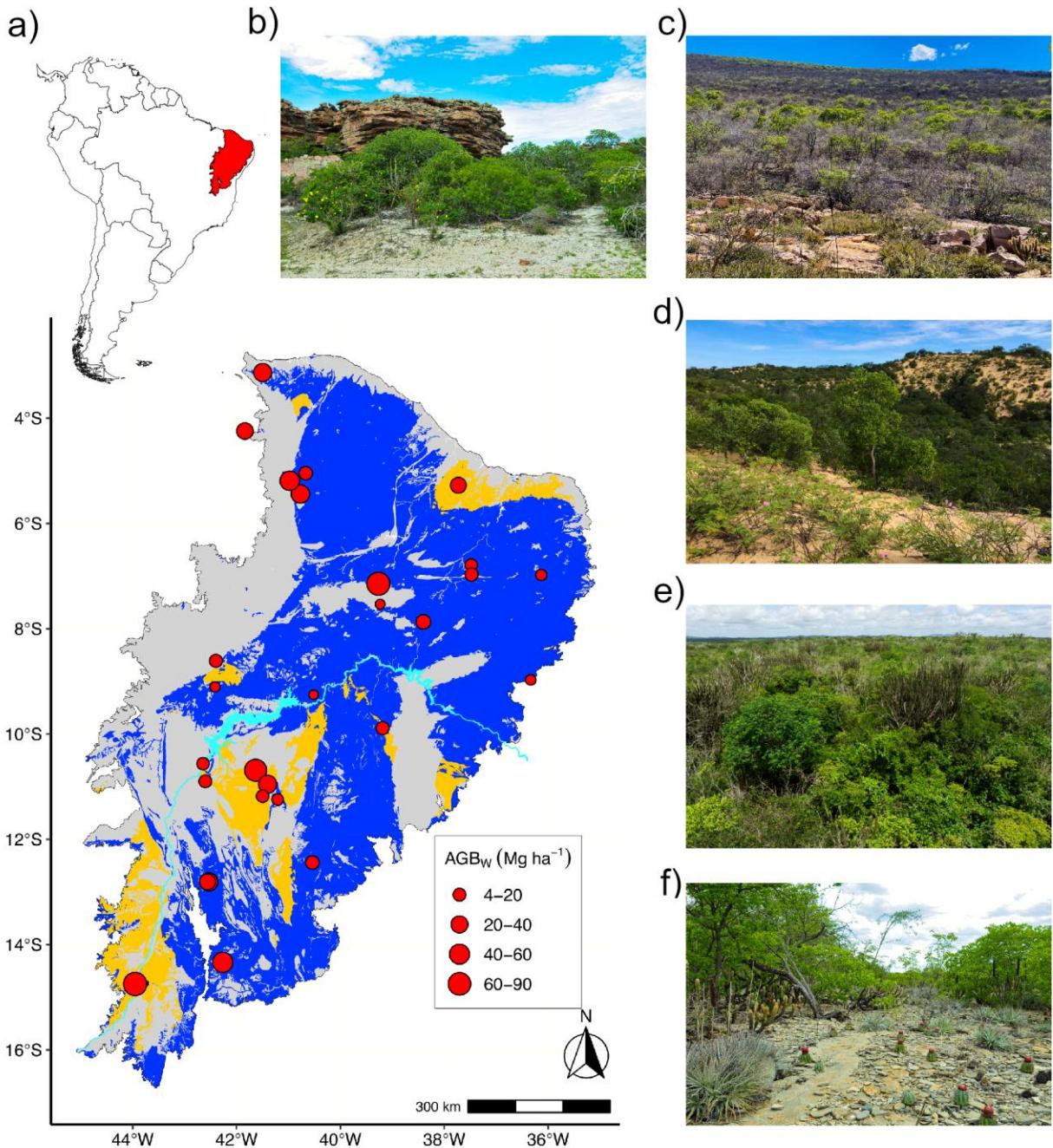


Fig. 1: Study plots and sampled vegetation. a) Geographic location of the Brazilian Caatinga region in South America and the distribution of dry above-ground woody biomass (AGBw) across the *Nordeste Project* study plots. b) Caatinga *sensu stricto* interspersed with exposed sand in the sedimentary Morro do Chapéu Formation (MOR-01); c) Caatinga *sensu stricto* interspersed with arenites outcrops in the sedimentary Morro do Chapéu Formation (MOR-02); d) Caatinga *sensu stricto* in the Quaternary dunes of the middle São Francisco River (IBD-01); e) Caatinga *sensu stricto* in the crystalline São Caetano Formation (CGR-01); f) Caatinga *sensu stricto* overlying carbonate rock outcrops in Gruta dos Brejões (GBR-02). Blue, grey, and yellow areas indicate crystalline, sedimentary, and carbonate substrates, respectively. The light blue course depicts the São Francisco River. The outline of the Caatinga region follows the Brazilian Institute of Geography and Statistics (IBGE, 2019), crystalline and sedimentary substrates according to the geoscience system of the Geological Survey of Brazil (CPRM), and areas with carbonate substrates according to Queiroz et al. (2017). Overlapping study plots have been displaced to allow their visualisation. Datum: SIRGAS2000. Photos by Domingos Cardoso (b, c, d, and f) and Peter Moonlight (e).

167 Vegetation structure

168 Standardised floristic and structural inventories were carried out following '*The DryFlor Field Manual*  
169 *for Plot Establishment and Remeasurement*' (Moonlight et al. 2021a; 2021b). In brief, a 100 × 50 m (0.5 ha)  
170 plot was established and sub-partitioned into 50 subplots (10 × 10 m or 0.01 ha). All trees with stems with ≥ 5  
171 cm in diameter were inventoried, measured at both breast height (DBH, 1.3 m from the ground) and 30 cm from  
172 the ground level (DGL), as recommended by the DRYFLOR protocol to ensure comparability across dry forest  
173 networks. This criterion represents a compromise that balances practical field constraints with the need to  
174 capture most of the forest's structure and dynamics. Multi-stemmed individuals were carefully measured stem  
175 by stem. Trees were identified in the field, herbarium, and by taxonomic specialists, to species where possible,  
176 and voucher specimens were deposited into the Herbarium of Feira de Santana State University (HUEFS; Feira  
177 de Santana, Bahia, Brazil). Tree-by-tree data from each '*Nordeste Project*' study plot contribute to the  
178 *ForestPlots network* (ForestPlots.net, 2021) and are curated at [www.ForestPlots.net](http://www.ForestPlots.net). High-resolution images of  
179 the voucher specimens are also publicly accessible through the *speciesLink* network of biodiversity collections  
180 (<http://www.splink.org.br/>).

181

182 Soil sampling

183 We used a standard protocol ([https://rainfor.org/wp-  
184 content/uploads/sites/129/2022/07/soilandfoliarsampling.pdf](https://rainfor.org/wp-content/uploads/sites/129/2022/07/soilandfoliarsampling.pdf)) with some adjustments to accommodate specific  
185 characteristics of Caatinga soils, such as restricted depth range, and a marked presence of stones and rocks in  
186 some cases. This protocol has been widely used in past research conducted in tropical regions, such as the  
187 RAINFOR and TROBIT networks (e.g., Quesada et al., 2010; Lloyd et al., 2015). Specifically, four auger cores  
188 were taken as baseline samples in each plot, with one to three additional cores collected in cases of pronounced  
189 spatial variability within the plot (e.g., irregular topography, rock outcrops, vegetation changes, or markedly  
190 shallow soils). Samples were obtained at standard depth intervals (0–5, 5–10, 10–20, 20–30, 30–50, 50–100,  
191 100–150, and 150–200) or according to the maximum depth achievable at each location. In addition, a soil pit  
192 was excavated just beside each plot to describe soil profiles, also serving as an additional sampling point for  
193 chemical and physical analyses. Subsequently, all samples were air-dried and sent to the Soil and Plant

194 Thematic Laboratory at the National Institute for Research in the Amazon (LTSP, INPA, Manaus, Amazonas,  
195 Brazil).

196 Laboratory analysis

197 In the LTSP, samples were loosened, sieved with a no. 10-mesh sieve (particle size of 2 mm), and any  
198 non-fine earth residues removed (e.g., gravels and vegetation or faunal debris). We determined water-measured  
199 soil pH ( $\text{pH}_{\text{H}_2\text{O}}$ ) using a 1: 2.5 soil-to-deionised water ratio with a glass electrode. Soil exchangeable cations  
200 were determined by the Silver-Tioureia (AgTU) method (Pleysier and Juo 1980). The concentration of each  
201 cation extracted was determined using an Atomic Absorption Spectrophotometer (AAS, Model 100b, Perkin  
202 Elmer, Norwalk, CT, USA). Soil sum of bases ( $\Sigma_B$ ) and effective cation exchange capacity ( $I_E$ ) were calculated  
203 according to Eqn. (1) and (2), respectively:

$$204 \Sigma_B = [\text{Ca}]_{\text{ex}} + [\text{Mg}]_{\text{ex}} + [\text{K}]_{\text{ex}} + [\text{Na}]_{\text{ex}} \quad \text{Eqn. (1)}$$

$$205 I_E = \Sigma_B + [\text{Al}]_{\text{ex}} \quad \text{Eqn. (2)}$$

206 Where Ca, Mg, K, Na and Al refer to calcium, magnesium, potassium, sodium and aluminium, respectively,  
207 while “ex” refers to exchangeable contents.

208 Soil total carbon (C) and nitrogen (N) were determined using dry combustion with an automated  
209 analyser (Vario Max CN, Elementar, Germany). Soil samples were combusted at high temperatures, and the  
210 resulting gases were measured to quantify C and N concentrations. Soil total phosphorus concentrations,  $[\text{P}]_T$ ,  
211 were obtained using composite samples from the 0–5, 5–10, 10–20, and 20–30 cm soil depths. Samples were  
212 digested with concentrated sulfuric acid, followed by the addition of hydrogen peroxide (Tiessen and Moir  
213 1993). Afterwards,  $[\text{P}]_T$  were determined by colourimetry using the molybdenum blue colour development  
214 method (Olsen and Sommers 1982), using a spectrophotometer (Model 1240, Shimadzu, Kyoto, Japan). Soil  
215 texture was determined using the sieve-pipette method (Gee and Bauder 1986). Soil dry bulk density (BD) was  
216 determined using the volumetric ring method (Eijkelkamp Agrisearch Equipment BV, Giesbeek, Netherlands).  
217 Calibration procedures and standard samples were routinely used to ensure the accuracy of the results.

218

219 Soil classification, clay mineralogy, and geological surveying  
220 The soils were classified according to the World Reference Base for Soil Resources (IUSS Working  
221 Group WRB, 2014/2015), with the aid of the WRB Tool 1.1.2.0 (Downloaded in March 2021; OrlovDO, 2017;  
222 <https://apps.microsoft.com/store/detail/wrb-tool/>). This tool streamlines the soil classification process by  
223 guiding users systematically through the key steps, potentially reducing classification errors through its step-  
224 by-step interface. Complete soil classifications are provided in Table 2 of Online Resource 1. Following  
225 Quesada et al. (2020), we categorised the sampled soils as HAC, LAC, and 'Arenic'. HAC soils are those with  
226 CEC clay<sup>-1</sup> > 24 mmol<sub>c</sub> kg<sup>-1</sup> (typically less weathered soil classes such as Luvisols in this study), while LAC  
227 soils are those with CEC clay<sup>-1</sup> < 24 mmol<sub>c</sub> kg<sup>-1</sup> (typically more weathered soils such as Acrisols, Alisols, and  
228 Regosols in this study) (IUSS Working Group WRB, 2014/2015). The third category ('Arenic') was used for  
229 considerably sandy soils, i.e., those soils with loamy sand texture or coarser (Arenosols in this study).  
230 Additionally, we used the geoscience system (GeoSGB; <https://geosgb.sgb.gov.br/geosgb/>) of the Geological  
231 Survey of Brazil (CPRM), and the delineation of areas with carbonate rocks as in Queiroz et al. (2017) to  
232 characterise the geology of each study plot.

233

234 Maximum plant-available soil water ( $\theta_p$ )

235 The maximum plant-available soil water ( $\theta_p$ ) is defined as the difference in volumetric soil water  
236 content ( $\theta_v$ ) values between field capacity ( $\theta_v$  at a matric potential of -10 kPa:  $\theta_{v,FC}$ ) and the permanent wilting  
237 point ( $\theta_v$  at a matric potential of -1500 kPa:  $\theta_{v,WP}$ ):  $\theta_{v,FC} - \theta_{v,WP}$ . These  $\theta_v$  values can be obtained from the water  
238 retention curve (WRC), which for our study was described with the widely used van Genuchten (VG) equation  
239 (van Genuchten, 1980). The VG parameters required for the calculation of the WRC were obtained from Table  
240 6 in Hodnett and Tomasella (2002). These parameters had been obtained using a (soil) class pedotransfer  
241 function approach for tropical soils of the IGBP soils dataset. For each *Nordeste* plot, and each soil layer, the  
242 soil textural class, based on measured sand/silt/clay contents, was determined using the USDA soil texture  
243 triangle, after which the look-up table provided by Hodnett and Tomasella (2002) was used to obtain the VG  
244 parameters. After estimating  $\theta_{v,FC}$  and  $\theta_{v,WP}$  from the constructed WRCs, their difference was integrated over  
245 the maximum measured soil depth (mm), thus providing  $\theta_p$  in mm<sup>3</sup> mm<sup>-2</sup>, or simply mm.

246 Climatic data

247 Climatic data were extracted from the WorldClim database version 2.1. The BioClim variables  
248 represent the historical averages for the 1970 – 2000 period with 30 arc-seconds ( $\sim 1 \text{ km}^2$ ) resolution (Fick and  
249 Hijmans 2017). We selected a few key variables based on *a priori* hypotheses, i.e., mean annual precipitation  
250 (BIO12 in the WorldClim system, MAP in this study), mean annual temperature (BIO1 in the WorldClim  
251 system, MAT in this study), the maximum temperature of the warmest month (BIO5 in the WorldClim system,  
252  $T_{\text{MAX}}$  in this study), and precipitation seasonality (BIO15 in the WorldClim system,  $\Psi$  in this study).  $T_{\text{MAX}}$   
253 reflects high-temperature events throughout the year and can be used to examine whether vegetation properties  
254 are affected by extreme temperature events, while BIO15 is a measure of variation in monthly precipitation  
255 totals over the year (O'Donnell and Ignizio 2012). Potential evapotranspiration (PET) was obtained from the  
256 CGIAR Consortium for Spatial Information – CGIAR-CSI (Zomer et al., 2022). Estimates of climatic water  
257 deficit (CWD) for each study plot were obtained from raster layers (currently available at  
258 <https://zenodo.org/records/14932971>) with 2.5 arc-minute resolution. The CWD variable was found to be  
259 important in determining allometric relationships (Chave et al. 2014) and represents the net balance between  
260 precipitation and potential evapotranspiration (PET) in the dry months (i.e., months where PET exceeds rainfall,  
261 given in mm per year). Note that although CWD values are originally negative, we present them as positive to  
262 indicate ‘millimetres of deficit’.

263

264 Above-ground woody biomass (AGB<sub>w</sub>) calculations

265 Estimates of AGB<sub>w</sub> for individual tree stems were calculated using an allometric equation specifically  
266 developed for Caatinga trees (Sampaio and Silva 2005):

$$267 \quad \text{AGB}_T = 0.0644 \times \text{DGL}^{2.3948} \quad \text{Eqn. (3)}$$

268 Where AGB<sub>T</sub> is the dry tree above-ground biomass (kg) and DGL is the diameter at the ground level.  
269 Biomass of individual cacti (1,098 stems) was estimated using a separate specific equation for cacti from  
270 Sampaio and Silva (2005), and palm tree biomass (35 stems) was calculated using a formula from Saldarriaga  
271 et al. (1988). At the plot level, total AGB<sub>w</sub> was the sum of the above-ground dry biomass of all stems measured,

272 with individual biomass for multi-stemmed individuals calculated and then summed. It is worth noting that the  
273 equation determines the total biomass of trees (including leaves), with “w” referring to woody species.

274

275 Community functional composition

276 Following Prado-Junior et al. (2016), we calculated two community-weighted trait means of strong  
277 ecological significance: community-weighted maximum stem diameter ( $CWM_{D\text{MAX}}$ ) and community-weighted  
278 mean wood density ( $CWM_{WD}$ ). Both traits represent fundamental life history strategies and are closely linked  
279 to resource storage, structural resistance, hydraulic safety, and the ability to adapt to environmental stressors  
280 (Larjavaara and Muller-Landau 2010; Reich 2014). Species maximum stem diameter reflects adult sizes and  
281 was calculated as the upper 0.95 percentile of those trees with a stem diameter  $\geq 0.1 \times$  the diameter (cm) of the  
282 thickest tree observed in each population. We adopted this approach since it was considered the least sensitive  
283 to varying sample sizes while providing robust estimates for both larger and smaller species (King et al. 2006;  
284 Prado-Junior et al. 2016). Species’ wood density values were extracted from the global wood density database  
285 (Chave et al. 2009; Zanne et al. 2009). When unavailable at the species level, we used wood density values at  
286 the genus or family levels. Botanical names were checked and adjusted according to the *Flora do Brasil* 2020  
287 with the *flora* package version 0.3.5 (Carvalho 2020). Each trait was weighted according to the basal area of  
288 individual species, as this is expected to reflect plant performance better than abundance (Prado-Junior et al.  
289 2016). The distribution of a given trait across a niche space can be summarised along orthogonal axes, from  
290 which functional diversity indices can be calculated (Mason et al. 2005). Specifically, we derived the functional  
291 richness index ( $F_{RIC}$ ) from community-weighted maximum stem diameter and wood density, measuring the  
292 amount of functional trait space occupied by a community. The functional richness index reflects the diversity  
293 of ecological strategies present and is calculated as the convex hull volume in a multidimensional trait space.  
294 Community-weighted mean traits and  $F_{RIC}$  were computed using the FD package in R (Laliberté et al. 2014).

295

296 Data analysis

297 Initially,  $AGB_w$  values were plotted against three categorical predictors: geological substrates, clay  
298 activity, and RSGs. To assess whether these categories explained variations in  $AGB_w$ , a robust non-parametric

299 Kruskal-Wallis test ( $\chi^2$ ) was performed. Comparisons were limited to categories with at least 5 observations,  
300 including Arenosols, Leptosols, and Luvisols for RSGs; HAC, LAC and Arenic for clay activity; and crystalline  
301 (SCRY) *versus* sedimentary (SED) for geological substrates.

302 To test the hypothesis that AGB<sub>w</sub> is influenced by multiple environmental factors and their  
303 interactions, a linear mixed-effects model was employed, combined with multi-model/model averaging  
304 inferences. For this purpose, two 0.25 ha subplots were considered within each 0.5 ha plot (29 total), resulting  
305 in 58 observations. This procedure was adopted to increase the number of observations and to provide greater  
306 flexibility for including multiple predictors in the same models, while reducing the risk of overfitting (Harrison  
307 et al. 2018). The global model included mean annual precipitation and climatic water deficit to represent water  
308 input and water balance, and maximum temperature of the warmest month as a key thermal variable.  
309 Correlations between climatological variables are provided in Table 3 of Online Resource 1.

310 Since soil predictors were strongly correlated (Table 4, Online Resource 1), they were carefully  
311 selected by systematically replacing them one at a time in the global model (Eqn. 4). We evaluated relative  
312 importance values (RIV), variance inflation factors (VIF), marginal  $r^2$  (fixed effects), and Akaike Information  
313 Criterion corrected (AICc). The VIF values of all predictors in Eqn. (4) were checked to prevent overfitting.  
314 Given the high correlation between potential evapotranspiration and temperature variables ( $\rho = 0.84, p < 0.001$   
315 for MAT; and  $\rho = 0.83, p < 0.001$  for  $T_{MAX}$ , Table 3 of Online Resource 1), and considering that PET was  
316 incorporated into CWD, we did not include it in the global model of Eqn (4). In all analyses, we used the upper  
317 0-30 cm soil layer, commonly used in vegetation ecology studies (e.g., Quesada et al. 2012; Lloyd et al. 2015).  
318 To facilitate interpretation, all predictors were standardised (subtracting the mean and dividing by the standard  
319 deviation) with the aid of the *caret* package in R (Max et al. 2020), providing comparative effect sizes among  
320 predictors. Since the analysis involved non-independent observations, sites were treated as random structures  
321 within the model (Harrison et al. 2018). The final global model is expressed by Eqn. (4):

$$322 \log(\text{AGB}_w) = \beta_0 + \beta_1 \theta_p + \beta_2 \text{MAP} + \beta_3 [\text{Ca}]_{\text{ex}} + \beta_4 \text{CWD} + \beta_5 \theta_p \times \text{CWD} + \beta_6 [\text{Ca}]_{\text{ex}} \times \text{CWD} + \beta_7 T_{MAX} \\ 323 + \beta_8 \log[\text{N}]_T + \beta_9 \log[\text{P}]_T + (1|site) + \varepsilon \quad \text{Eqn. (4)}$$

324 Where AGB<sub>w</sub> is above-ground woody biomass;  $\theta_p$  is maximum plant-available soil water; MAP is  
325 mean annual precipitation;  $[\text{Ca}]_{\text{ex}}$  is soil exchangeable calcium; CWD is climatic water deficit;  $T_{MAX}$  is the

326 maximum temperature of the warmest month;  $[N]_T$  is soil total nitrogen;  $[P]_T$  is soil total phosphorus;  $(1|site)$   
327 represents the random intercept for site; and  $\epsilon$  is the residual error. We evaluated the distribution of residuals  
328 both statistically and graphically. The response variable was log-transformed to meet the normality assumption  
329 and reduce heteroscedasticity. The potential presence of spatial structures was checked using the methods  
330 outlined by Bauman et al. (2018a; 2018b). Specifically, the *listw.candidates* function from the *adespatial*  
331 package in R (Dray et al. 2021) was used to test a few distance and graph-based spatial weighting matrices.  
332 Spatial autocorrelations were assessed through Moran's *I* coefficient with a significance level of  $p \leq 0.05$ .

333 We tested all possible predictor combinations using the *MuMin* package in R (Bartón 2020).  
334 Collinearity issues were further controlled by preventing predictors with Pearson's correlation  $|r| \geq 0.6$  in the  
335 same models. This process generated 137 unique models, with the maximum number of predictors in each  
336 model being constrained to 6, thus ensuring nearly 10 observations per model term. Model selection followed  
337 an information-theoretic (I-T) approach, retaining models with  $\Delta AICc < 4$  (Burnham et al. 2011; Harrison et  
338 al. 2018). From the 19 retained models ( $\Delta AICc < 4$ ), coefficients were averaged using the *model.avg* function  
339 of the *MuMin* R package. Full averaging was used for model predictions, providing more reliable  $\beta$  estimates  
340 when multiple models have support (Mazerolle 2023). The 'full' averaging approach dictates that each variable  
341 is included in every model (setting the coefficients to zero in the models where the term is absent), whereas the  
342 'conditional' average approach considers only those models where the parameter appears (Bartón 2020). In  
343 both cases, average coefficients were weighed according to Akaike weights. Model marginal  $r^2$  ( $r^2m$ ) was  
344 reported to represent fixed effects. Additionally, we performed a series of 95th-percentile linear mixed-effects  
345 model relationships to explore the predictive ability of individual soil and climate predictors on AGBw.

346 A Spearman's rank correlation matrix ( $\rho$ ) was computed to explore relationships between soil  
347 properties and community functional properties, with significant relationships graphically represented (Fig. 7  
348 of Online Resource 1). Subsequently, the most supported variables identified through the multi-model inference  
349 approach were used to investigate the relationships between environmental factors, vegetation properties, and  
350 their combined effects on AGBw. For this task, we employed a Structural Equation Modelling (SEM)  
351 framework using the *lavaan* R package (Rosseel 2012). The model assessed direct and indirect effects of  
352 edaphic and climatic variables, i.e., mean annual precipitation, climatic water deficit, soil exchangeable

353 calcium, and maximum plant-available soil water on community-weighted maximum stem diameter, wood  
354 density, and associated functional richness, as well as their potential effects on above-ground biomass (AGB<sub>w</sub>).

355 To overcome the non-normality of some variables included in the SEM, we utilised the Maximum  
356 Likelihood Estimator with Robust Standard Errors (MLR), which adjusts for non-normal distributions and  
357 potential heteroscedasticity. Robustness was further enhanced with the *Yuan-Bentler* scaling correction, which  
358 is appropriate for handling non-normality and small sample sizes. The SEM was evaluated using multiple fit  
359 indices, including the Comparative Fit Index (CFI), Tucker-Lewis Index (TLI), Root Mean Square Error of  
360 Approximation (RMSEA), and Standardised Root Mean Square Residual (SRMR), with robust versions  
361 addressing non-normality. Overall model performance was assessed using the chi-square test statistic ( $\chi^2$ ) and  
362 its associated *p*-values to determine significance.

363 Finally, we used the *alphahull* R package (Pateiro-López and Rodríguez-Casal 2019) to make the  
364 heatmaps presented in Figs. 3 and 4. We used heatmaps to represent interaction terms, as they provide an  
365 intuitive visualisation of how AGB<sub>w</sub> responds across the gradients of two predictors simultaneously. Heatmap  
366 simulations were constrained to the actual environmental conditions found in the dataset. All graphs were  
367 created using the *ggplot2* R package (Wickham et al. 2016), and all analyses were conducted in the R  
368 environment, Version 4.1.1 (R Core Team 2021).

369

## 370 **Results**

371 Stand structure and categorical predictors

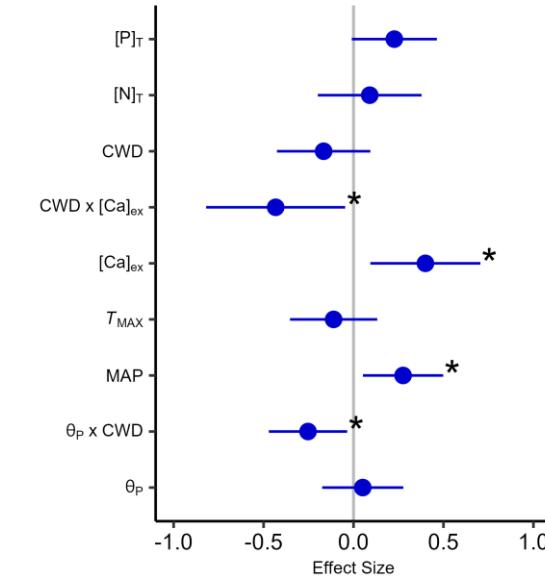
372 Altogether, 18,201 individual stems with a diameter at the ground level  $\geq 5$  cm were recorded across  
373 the 29 study plots, including 1,098 cacti and 35 palm trees. These individuals encompass 331 unique species,  
374 176 genera, and 50 identified tree families. The mean  $\pm$  standard deviation of AGB<sub>w</sub> was  $32.55 \pm 22.35$  Mg ha<sup>-1</sup>  
375 (min = 4.87 Mg ha<sup>-1</sup>; max of 85.65 Mg ha<sup>-1</sup>), mean stem density (stems ha<sup>-1</sup>) was  $1255 \pm 489$  (min = 492; max  
376 = 2,534), and mean basal area ( $B_A$ ) was  $12.89 \pm 7.10$  m<sup>2</sup> ha<sup>-1</sup> (min = 2.44; max = 28.79 m<sup>2</sup> ha<sup>-1</sup>). Above-ground  
377 woody biomass did not differ among geological substrates ( $\chi^2 = 0.08$ ;  $p = 0.775$ ; Fig. 4-a of Online Resource  
378 1), types of clay activity ( $\chi^2 = 3.70$ ;  $p = 0.157$ ; Fig. 4-b of Online Resource 1), and RSGs ( $\chi^2 = 3.39$ ;  $p = 0.183$ ;  
379 Fig. 4-c of Online Resource 1).

380 Above-ground woody biomass modelling

381 Above-ground woody biomass was influenced by both edaphic and climatic factors, and their  
382 interactions, as indicated by the multi-model Inference and Information-Theoretic approaches. Of the 19 models  
383 selected within the  $\Delta\text{AICc} < 4$  range, six included climate, soil chemistry, and soil physics terms; 10 included  
384 only climate and soil chemistry; and three included only soil chemistry (Table 5 of Online Resource 1). The  
385 most strongly supported terms in the conditional average model were the interaction between exchangeable  
386 calcium and climatic water deficit ( $\beta = -0.43$ ), exchangeable calcium ( $\beta = 0.40$ ), mean annual precipitation ( $\beta$   
387 = 0.28), and the interaction between maximum plant-available soil water  $\times$  climatic water deficit ( $\beta = -0.25$ )  
388 (Fig. 2-A). Relative importance values for these terms were: exchangeable calcium (0.95), mean annual  
389 precipitation (0.92), climatic water deficit (0.53), the interaction between exchangeable calcium and climatic  
390 water deficit (0.40), maximum plant-available soil water (0.33), max temperature of the warmest month (0.18),  
391 the interaction between maximum plant-available soil water  $\times$  climatic water deficit (0.18), total nitrogen (0.15),  
392 and total phosphorus (0.02) (Fig. 2-b).

393 Replacing collinear soil predictors in the global model (Eqn. 4) showed that  $r^2m$  decreased in the  
394 following order: exchangeable calcium (0.49) > sum of bases (0.48) > effective cation exchange capacity (0.46)  
395 >  $\text{pH}_{\text{H}_2\text{O}}$  (0.45) > exchangeable magnesium (0.41) > exchangeable aluminium (0.40) > exchangeable sodium  
396 (0.39) > exchangeable potassium (0.35). Correspondingly, AICc values increased, with exchangeable calcium  
397 showing the lowest (92.95) and exchangeable potassium the highest (100.78). RIV decreased in the order:  
398 effective cation exchange capacity (0.96) > exchangeable calcium (0.95) > sum of bases (0.87) >  $\text{pH}_{\text{H}_2\text{O}}$  (0.64)  
399 > exchangeable magnesium (0.46) > exchangeable sodium (0.39) > exchangeable potassium (0.31) >  
400 exchangeable aluminium (0.15). Based on these metrics, and noting the dominance of exchangeable calcium in  
401 the soil cation exchange complex across most sites (Fig. 5 of Online Resource 1), exchangeable calcium was  
402 selected over its alternatives. Model-simulated responses suggest that exchangeable calcium levels have a  
403 marked influence on  $\text{AGB}_w$ . This effect, however, was also influenced by the intensity of the climatic water  
404 deficit (Fig. 3-a).

a)



b)

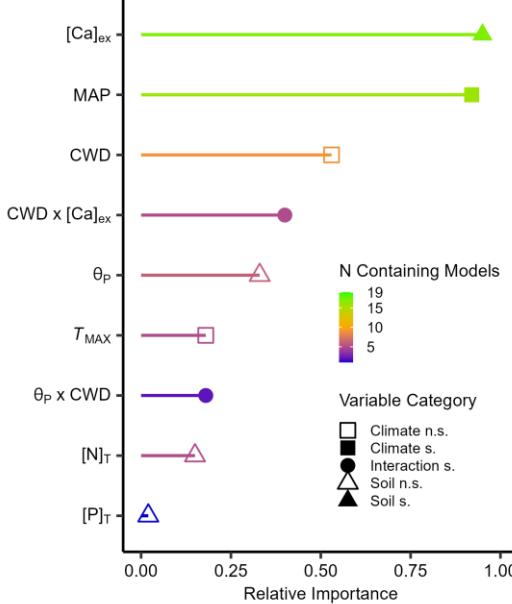


Fig. 2: Multi-model inference statistics. a) edaphic and climatic effects on above-ground woody biomass (AGB<sub>w</sub>) of 29 Caatinga's SDTFs study plots. Points represent conditional average model coefficients. Coefficients were standardised, thus representing changes in log(AGB<sub>w</sub>) for a one standard deviation change in the predictor variable (effect size). [Ca]<sub>ex</sub> = exchangeable calcium; MAP = mean annual precipitation; CWD = climatic water deficit;  $\theta_P$  = maximum plant-available soil water;  $T_{MAX}$  = temperature of the warmest month; [N]<sub>T</sub> = soil total nitrogen; [P]<sub>T</sub> = soil total phosphorus. Error bars show 95% confidence intervals. Asterisks denote statistically significant coefficients. b) Relative importance values (RIV) of each variable included in the final model. Variable category and the frequency of each term across the 19 models selected via AICc ("N Containing Models") are shown. Filled symbols represent significant relationships, while empty symbols represent non-significant relationships.

405 Using coefficients from the best AICc-ranked model (Model 1; Table 5 of Online Resource 1),  
 406 predicted above-ground woody biomass ( $\hat{AGB}_w$ ) was virtually constant across climatic water deficit when  
 407 exchangeable calcium was low ( $\sim 5 \text{ mmol}_c \text{ kg}^{-1}$ ). However, at high exchangeable calcium ( $\sim 40.37 \text{ mmol}_c \text{ kg}^{-1}$ ),  
 408  $\hat{AGB}_w$  ranged from 19.19 to 86.12 Mg ha<sup>-1</sup> as climatic water deficit varied from 1238 to 909 mm. Similarly,  
 409 model predictions showed that  $\hat{AGB}_w$  increases with higher maximum plant-available soil water and lower

410 climatic water deficit (Fig. 3-b). For instance, at a maximum plant-available soil water = 200 mm,  $\hat{AGB}_w$  was  
411 62.15 Mg ha<sup>-1</sup> when climatic water deficit was low (620 mm), but dropped to 14.30 Mg ha<sup>-1</sup> at high climatic  
412 water deficit (1290 mm). The second-best model (Model 2; Table 5 of Online Resource 1) highlighted the  
413 influence of mean annual precipitation and exchangeable calcium on  $\hat{AGB}_w$ . For example, at a mean annual  
414 precipitation of ~800 mm,  $\hat{AGB}_w$  varied from 20.14 to 60.81 Mg ha<sup>-1</sup> as exchangeable calcium increased from  
415 1.37 to 62.87 mmol<sub>c</sub> kg<sup>-1</sup> (Fig. 3-c).

416 Finally,  $\hat{AGB}_w$  predictions across the mean annual precipitation gradient in the dataset were simulated  
417 under varying soil conditions (Fig. 4): A) optimal – exchangeable calcium = 50.67 mmol<sub>c</sub> kg<sup>-1</sup>, and maximum  
418 plant-available soil water = 270 mm; B) moderately high – exchangeable calcium = 32.65 mmol<sub>c</sub> kg<sup>-1</sup>, and  
419 maximum plant-available soil water = 210 mm; C) intermediate – exchangeable calcium = 14.62 mmol<sub>c</sub> kg<sup>-1</sup>,  
420 and maximum plant-available soil water = 150 mm; and D) poor conditions – exchangeable calcium = 1 mmol<sub>c</sub>  
421 kg<sup>-1</sup>, and maximum plant-available soil water = 80 mm. While  $\hat{AGB}_w$  increased with mean annual precipitation  
422 in all cases, this effect was strongest under favourable soil conditions, i.e., increased nutrient content and higher  
423 water storage capability.

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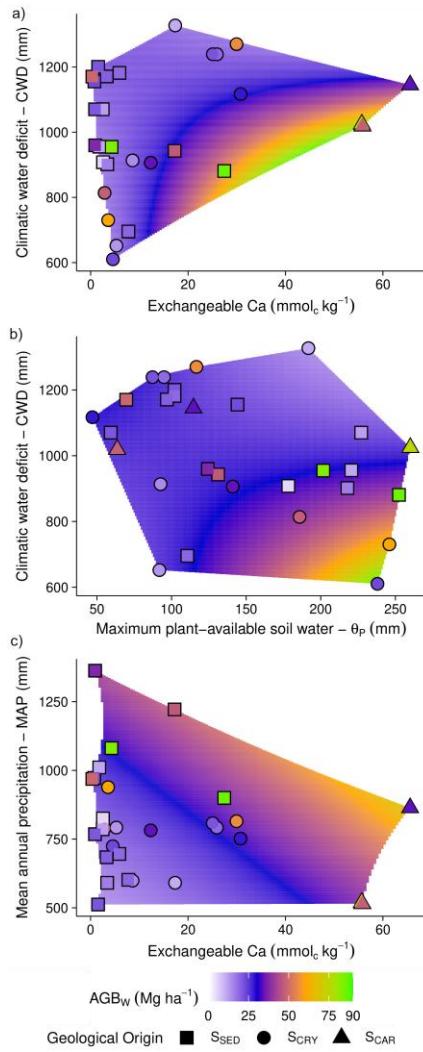
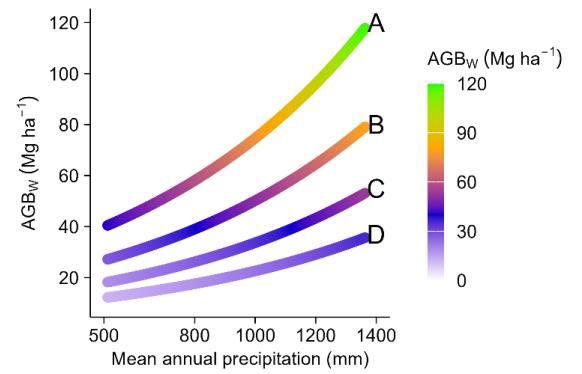


Fig. 3: Modelled responses of above-ground woody biomass ( $\hat{AGB}_W$ ) based on the two best AICc-ranked models. a)  $\hat{AGB}_W$  as a function of the interaction between exchangeable calcium ( $[Ca]_{ex}$ ) and climatic water deficit (CWD). b)  $\hat{AGB}_W$  as a function of the interaction between CWD and maximum plant-available soil water ( $\theta_p$ ). c)  $\hat{AGB}_W$  as an additive function of  $[Ca]_{ex}$  and mean annual precipitation (MAP). Study plots are shown within their respective environmental domains and geological categories. Note: CWD, originally expressed as negative values, is presented here as positive values representing “millimetres of deficit.” Geological categories: SED = sedimentary, CRY = crystalline, CAR = carbonate.

Fig. 4: Modelled responses of above-ground woody biomass as ( $\hat{AGB}_W$ ) a function mean annual precipitation under four edaphic scenarios: A) optimal – maximum plant-available soil water ( $\theta_p$ ) and exchangeable calcium ( $[Ca]_{ex}$ ) (+2 SD above the mean); B) moderately high –  $\theta_p$  and  $[Ca]_{ex}$  (+1 SD above the mean); C) intermediate –  $\theta_p$  and  $[Ca]_{ex}$  (at their means); and D) poor conditions –  $\theta_p$  and  $[Ca]_{ex}$  (-1 SD below their means).



436        Regarding bivariate relationships, we found significant associations between AGB<sub>w</sub> and soil variables,  
437        including exchangeable calcium, sum of bases, effective cation exchange capacity, sand, and silt, across the full  
438        dataset, while climatic variables showed no significant linear relationships with AGB<sub>w</sub> (Table 6 of Online  
439        Resource 1 and Fig. 6 of Online Resource 1).

440

441        Associations between soil properties and community functional composition

442        A Spearman's correlation matrix showed that community-weighted wood density was inversely  
443        correlated with several soil properties, including exchangeable calcium, sum of bases, effective cation exchange  
444        capacity, and silt content, and positively correlated with soil sand. Conversely, community-weighted maximum  
445        stem diameter was positively associated with exchangeable calcium and soil sum of bases. The functional  
446        richness index was positively correlated with multiple soil properties, including exchangeable Ca, Mg, K, sum  
447        of bases, and effective cation exchange capacity. Significant Spearman's coefficients are shown in Fig. 7 of  
448        Online Resource 1, and all tested relationships are summarised in Table 7 of Online Resource 1. No significant  
449        correlations were found between climatic variables and functional properties. Finally, forests in crystalline  
450        environments had higher functional richness values than those in sedimentary substrates ( $\chi^2 = 7.71$ ;  $p = 0.005$ ;  
451        Fig. 8-c of Online Resource 1), while community-weighted mean wood density and maximum stem diameter  
452        did not differ significantly among these categories, though forests in the carbonate category showed a tendency  
453        for lower community-weighted wood density and higher stem diameters values.

454

455        Structural Equation Modelling (SEM)

456        Among the metrics utilised to evaluate our SEM (Fig. 5), the Robust Comparative Fit Index (CFI) was  
457        0.992 (standard) and 0.989 (robust), suggesting an excellent fit. The Tucker-Lewis Index (TLI) values were  
458        0.917 (standard) and 0.884 (robust), indicating a good overall model fit. The Root Mean Square Error of  
459        Approximation (RMSEA) was 0.076 (standard) and 0.092 (robust), both within acceptable thresholds for good  
460        fit (Browne and Cudeck 1992). The Standardised Root Mean Square Residual (SRMR) was 0.063, indicating a  
461        good fit for the data. Although this statistic can be sensitive to sample size, the *Chi-square* was 2.339 (standard)  
462        and 2.552 (scaled), with a *p-value* of 0.282, indicating an acceptable model fit. Noting that, except for AGB<sub>w</sub>,

463 all variables were standardised, significant relationships included the negative effect of exchangeable soil  
 464 calcium on community-weighted mean wood density ( $\beta = -0.59, p < 0.001$ ), its positive impact on community-  
 465 weighted mean maximum stem diameter ( $\beta = 0.42, p = 0.031$ ), and its marginal effect on functional richness ( $\beta$   
 466 = 0.24,  $p = 0.093$ ). Above-ground biomass (AGBw) was significantly influenced by community-weighted mean  
 467 maximum stem diameter ( $\beta = 10.17, p < 0.001$ ), mean annual precipitation ( $\beta = 7.89, p = 0.001$ ), and soil  
 468 available calcium ( $\beta = 9.27, p = 0.012$ ) (Fig. 5).

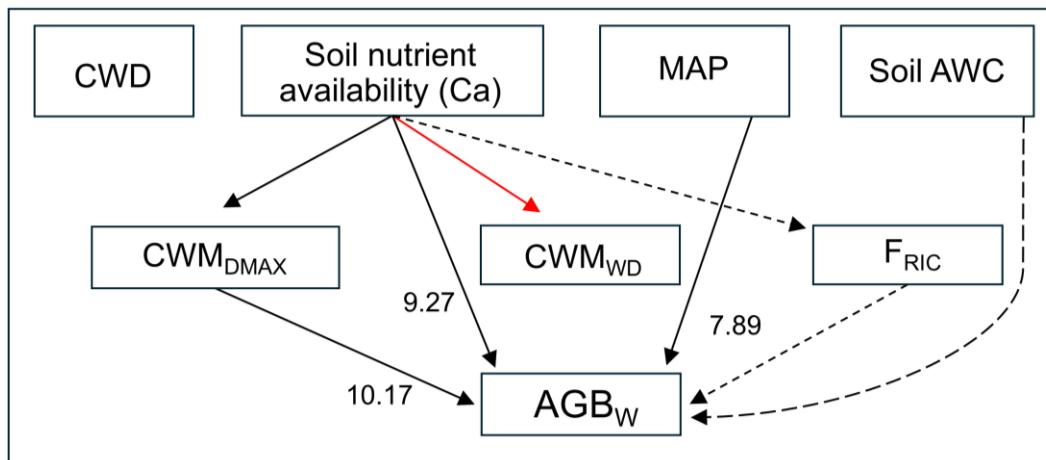


Fig. 5: Structural equation model (SEM) showing the relationships between environmental variables [climatic water deficit (CWD), soil nutrient availability (exchangeable Ca), mean annual precipitation (MAP), and maximum plant-available soil water content (AWC)] and vegetation attributes [community-weighted maximum diameter (CWM<sub>DMAX</sub>), community-weighted mean wood density (CWM<sub>WD</sub>), and functional richness (FRIC)] with above-ground biomass (AGBw). Soil nutrient availability was significantly related to CWM<sub>DMAX</sub>, CWM<sub>WD</sub>, and AGBw, and marginally to FRIC. MAP and CWM<sub>DMAX</sub> also significantly influenced AGBw, while soil AWC and FRIC showed marginal effects (dashed lines). Numbers (path coefficients,  $\beta$ ) represent standardised regression weights. Black solid arrows indicate positive relationships, and the red arrow indicates a negative relationship. Except for AGBw, all variables were standardised; therefore, model estimates are shown only for significant paths towards AGBw to avoid misinterpretation.

469 Moreover, functional richness and maximum plant-available soil water showed marginally significant  
 470 effects on AGBw ( $\beta = 4.27, p = 0.082$  for functional richness; and  $\beta = 5.69, p = 0.066$  for maximum plant-  
 471 available soil water). In the SEM framework, climatic water deficit did not affect vegetation traits or AGBw ( $\beta$   
 472 = -0.33,  $p = 0.915$ ). Finally, a significant negative covariance between community-weighted mean wood  
 473 and maximum stem diameter (estimate = -0.21;  $p = 0.037$ ) was detected in the model, indicating that higher  
 474 wood density is associated with smaller diameters. Significant variances in community-weighted mean wood  
 475 density and maximum stem diameter, functional richness, and AGBw point to considerable variability in the  
 476 data.

477 **Discussion**

478 We aimed to identify the key environmental factors influencing above-ground woody biomass in the  
479 seasonally dry vegetation of the Caatinga region, addressing gaps in the literature by adopting standardised  
480 sampling and analysis protocols. Our research covers a large spatial scale, encompassing the largest and most  
481 continuous SDTF area in Latin America. We confirm our main hypothesis that, alongside climate, soils affect  
482 biomass directly and indirectly by mediating structural traits such as wood density and maximum stem diameter.  
483 This study highlights the critical role of soils in shaping vegetation properties in the dry tropics, providing new  
484 insights into the ecology of these understudied ecosystems.

485

486 Region-wide AGB<sub>w</sub> is driven by complex soil–climate interactions

487 The AGB<sub>w</sub> range in this study (4.87–85.65 Mg ha<sup>-1</sup>) aligns with Santos et al. (2023), who reported  
488 similar values (2.85–80.88 Mg ha<sup>-1</sup>) in seasonally dry vegetation of Bahia State, Brazil. Such variation reflects  
489 the spatial heterogeneity of the region, notably diverse edaphic conditions and distinct vegetation  
490 physiognomies. We highlighted that, while we cannot assert to what extent some of our study sites have  
491 undergone more drastic changes in the past, SDTFs are known for faster recovery after disturbances than  
492 moister forests due to their simpler structure (Josse and Balslev 1994; Pennington et al. 2006; Becknell et al.  
493 2012).

494 Above-ground woody biomass in the Caatinga was shaped by both soil and climate, as well as their  
495 interactions (Fig. 2). Notably, MAP significantly influenced AGB<sub>w</sub> only in a multivariate context, suggesting  
496 its effect is conditioned by other factors. This contrasts with the notion that, given the water-limited ecology of  
497 SDTFs, a coarse index like MAP is an adequate proxy for biomass content in these environments (Becknell et  
498 al. 2012). The MAP range in our study was on the drier end for global SDTFs, with 17 of 29 plots having MAP  
499  $\leq 800 \text{ mm yr}^{-1}$ . Menezes et al. (2021) noted that Caatinga is even drier than Mexican ‘very dry deciduous  
500 forests’ studied by Lebrija-Trejos et al. (2008), where MAP lies around 900 mm yr<sup>-1</sup>. While Becknell et al.  
501 (2012) noted a clear tendency of higher and lower above-ground biomass values above and below a MAP of  
502 900 mm, respectively, our study did not confirm these differences, possibly due to the underrepresentation of  
503 high-MAP sites. The DryFlor Network (2016) sets an upper limit of 1800 mm MAP for SDTFs, but areas near  
504 this limit resemble semi-deciduous Atlantic forests. Cardoso et al. (2021) defined the ‘core Caatinga’ boundary

505 at 1300 mm MAP, above which distinct functional and floristic traits emerge. We reiterate that, while such  
506 studies relate biomass to MAP, integrated metrics that incorporate potential evapotranspiration and soil water  
507 availability may provide more informative insights in the Caatinga, where water availability is among the lowest  
508 in the tropics.

509 We found no significant effect of temperature variables or elevation on AGB<sub>w</sub>, in contrast to the  
510 findings of Santos et al. (2023), who reported that MAP, MAT, and elevation together explained 46% of AGB<sub>w</sub>  
511 variation across sharp climatic and topographic gradients in Bahia's Chapada Diamantina. At the broader scale  
512 of our study, other environmental drivers likely become more influential. In particular, soils with higher  
513 maximum plant-available water may buffer the intense seasonal drought typical of the Caatinga. Although  
514 shallow impermeable layers can retain moisture beyond the rainy season (Lloyd et al. 2015), rapid  
515 evapotranspiration generally depletes these reserves in the Caatinga's shallow soils (Sampaio 2010). The  
516 importance of water availability metrics observed in this study aligns with Terra et al. (2018), who demonstrated  
517 that water availability is an important determinant of vegetation structure, function, and diversity across  
518 Caatinga–Atlantic rainforest–Cerrado transitions.

519 Our results indicate that soil secondary macronutrients play an important role in shaping stand-level  
520 AGB<sub>w</sub>. Although required in smaller quantities than primary macronutrients (N, P, and K), secondary  
521 macronutrients (Ca, Mg, and S) are essential to plant growth, metabolism, and structure (Marschner 2012).  
522 Specifically, soil calcium was strongly supported in our modelling. Beyond its structural role in cell walls,  
523 calcium enhances antioxidant activities during heat stress (Jiang and Huang 2001) and provides osmoprotection  
524 under water deficit conditions (Jaleel et al. 2007). It also regulates a complex signalling network that helps  
525 plants respond to various stresses (Tong et al. 2021), with its cytosolic concentration being linked to soil calcium  
526 levels (Song et al. 2008; Sharma and Kumar 2021). Calcium is vital for root exocytosis and growth, enabling  
527 roots to exploit soil resources (Wilkins et al. 2016), which may support the survival of Caatinga trees.  
528 Furthermore, calcium plays a crucial role in multiple photosynthetic pathways by stomatal movement and  
529 photosynthetic proteins (Wang et al., 2019). In contrast, our results indicate no significant role for magnesium  
530 or potassium in determining AGB<sub>w</sub> in Caatinga dry forests. Despite magnesium potentially alleviating  
531 aluminium toxicity (Chen et al. 2018), this effect is unlikely to apply here due to the predominantly moderate  
532 acid to alkaline soils found in our dataset. While potassium, in combination with plant-available soil water, has

533 been shown to positively influence tropical woody vegetation (CWAK hypothesis—Lloyd et al. 2015; Ametsitsi  
534 et al. 2020), these studies were conducted in forest-savanna ecotones with markedly different climate conditions  
535 and vegetation characteristics. Soil exchangeable sodium concentrations were minimal, with salinity not being  
536 an issue in most Caatinga soils (Pessoa et al. 2022).

537 Variations in total soil P and N concentrations appeared to have a limited impact on biomass stocks  
538 in our study plots, showing lower relative importance values in our analyses. Although total soil P does not  
539 represent readily available forms, it can reflect overall P availability and serve as a proxy in forest ecosystems  
540 (Quesada et al. 2010). Moreover, while only a small fraction of total soil P is directly available to plants, it may  
541 still indicate long-term P availability across stages of pedogenesis (Cross and Schlesinger 1995). Mechanisms  
542 such as ‘P buffering capacity,’ in which less bioavailable P pools are accessed during periods of scarcity  
543 (Kitayama et al. 2000; Quesada et al. 2010), remain unexplored in the Caatinga, despite its generally P-deficient  
544 soils (Sampaio 2010).

545 Regarding N, while many Caatinga legumes have the potential for biological nitrogen fixation (BNF),  
546 only a small fraction effectively fix nitrogen (Freitas et al. 2010; Silva et al. 2017), at least in part because BNF  
547 is an energy-intensive process (Gutschick 1981). Many legumes, particularly those in the Detarioideae and  
548 Caesalpinoideae subfamilies, cannot even nodulate (Sprent 2009), and no correlation between Fabaceae  
549 biomass and soil  $\delta^{15}\text{N}$  (a potential indicator of the BNF degree) was observed by Brunello et al. (2024) for the  
550 same plots evaluated in this study. Studies on ‘nutrient use efficiency’ mechanisms (Vitousek 1982; 1984)  
551 could deepen our understanding of nutrient resorption from senescing leaves (Aerts 1996) in seasonal SDTFs.  
552 The soil N: P ratios found in this study suggest potential nutrient limitations. As an indicative metric, soil N: P  
553 ratios below 10 generally point to N limitation, whereas ratios above 20 suggest phosphorus limitation  
554 (Güsewell 2004). In our dataset, most sites were consistent with N limitation, although five values exceeded  
555 20 (Table 1 of Online Resource 1), suggesting possible P limitation in some areas. Leaf nitrogen and  
556 phosphorus concentrations strongly influence photosynthetic traits, such as maximum carboxylation rate  
557 ( $V_{\text{cmax}}$ ) (Walker et al., 2014), which in turn affects canopy growth. Therefore, variations in these nutrient levels  
558 may have influenced AGB<sub>w</sub> depending on the plant species composition at each site, even though these patterns  
559 were not explored in detail in our analysis.

560 Finally, while differences in AGB<sub>w</sub> were not statistically significant among geological substrates, a  
561 trend toward higher values in vegetation stands growing on carbonate-derived soils (S<sub>CAR</sub>) was noticeable  
562 ('Karst' in Fig. 4-a of Online Resource 1). The lack of statistical significance might be due to the low number  
563 of observations in this category (n = 3). Recently, Muñoz et al. (2023) found that tropical dry forests growing  
564 on limestone-derived soils exhibit higher structural complexity and diversity (i.e., higher basal area, stand-level  
565 above-ground biomass, tree density, and species richness) compared to forests growing in phyllite-derived soils  
566 in southern Mexico.

567 Interestingly, AGB<sub>w</sub> in GBR-01 was 58% higher than in GBR-02, despite both study plots sharing  
568 similar climatic and soil nutrient conditions. This difference may be related to the markedly shallower soil  
569 observed at GBR-02 (average depth = 28 cm), which could limit root anchorage and water storage. In contrast,  
570 the deeper soil at GBR-01 (average depth = 127 cm) likely offers greater physical support and functions as a  
571 larger reservoir for soil water. Within the S<sub>CAR</sub> plots, PFF-01 exhibited the lowest AGB<sub>w</sub>, even though it  
572 receives approximately 300 mm more annual precipitation than the other S<sub>CAR</sub> sites. This discrepancy could be  
573 attributed to shallow or rocky soils, species composition effects, or potential unaccounted human disturbance,  
574 as the plot is located near small farms.

575

576 Relationships between soil properties and community functional composition

577 In this study, we found negative correlations between community-weighted wood density and several  
578 soil properties, notably exchangeable calcium, magnesium, potassium, zinc, and the silt fraction, while wood  
579 density was positively correlated with sand content (Fig. 7 of Online Resource 1). These correlations are likely  
580 to reflect hydraulic safety and water-use efficiency patterns. Specifically, low-wood-density species may have  
581 greater sapwood water capacitance, as wood density is generally correlated with xylem density. Low-density  
582 trees may store more water in their parenchymatic tissues, which are responsible for the storage of water,  
583 nutrients, and carbohydrates (Sarmiento et al. 2011; Lira-Martins et al. 2019). Osmotically active cations, such  
584 as potassium, improve water-use efficiency by enhancing plant cell capacitance (Quesada et al. 2012), which  
585 can be particularly important under water-limited conditions. An inverse relationship between wood density  
586 and these cations may indicate an evolutionary strategy in low-wood-density species, as a response to

587 anatomical constraints that increase embolism susceptibility, such as larger vessels (Lira-Martins et al. 2019).  
588 It is important to note that deciduousness is closely linked to embolism avoidance. However, evidence is not  
589 entirely consistent: Lima et al. (2018) demonstrated that lignin composition, rather than wood density alone,  
590 was the main factor explaining differences in xylem embolism resistance and leaf lifespan, with some high-  
591 wood-density species shedding their leaves earlier than expected. By contrast, other studies have suggested that  
592 high-wood-density species typically retain their leaves longer during dry periods and are generally considered  
593 the last to avoid embolism by shedding leaves, with their narrow vessels playing a crucial role in this process  
594 (Markesteijn et al. 2011; Lima et al., 2021). Noteworthily, leaf flushing is strongly dependent on soil water  
595 availability in the Caatinga (Paloschi et al. 2021). Lima et al. (2012) also identified distinct functional groups  
596 in the Caatinga, i.e., evergreen, low-wood-density, and high-wood-density species, and showed that  
597 phenological events (leaf flush and flowering) are driven by water availability in high-wood-density species  
598 and by photoperiod in low-wood-density species.

599         Regarding the inverse relationship between soil zinc and wood density, zinc has been shown to  
600 enhance the activity of osmoregulation substances during drought stress (Wu et al. 2015). This suggests that  
601 zinc likely participates in structural and biochemical trade-offs within cells, potentially improving drought  
602 resilience. Soil texture also influenced wood density, with sand content showing a positive association and silt  
603 content a negative association. This relationship may be difficult to interpret due to the strong correlation  
604 between soil texture and cation availability (Table 4 of Online Resource 1), which complicates the separation  
605 of their individual effects. However, soil texture is known to influence plant and soil hydraulic properties, as  
606 well as tree mechanical stability, factors that can affect wood density (Quesada et al. 2012). Moreover, the  
607 observed positive association between sand content and community wood density (Fig. 7 of Online Resource  
608 1) may reflect an adaptive strategy whereby trees tolerate and cope with, rather than avoid, water scarcity. In  
609 coarse-textured soils, water drains more rapidly, and nutrient retention may be lower. Thus, species with denser  
610 wood may be favoured due to their ability to withstand drought stress under such conditions.

611         Maximum stem diameters were positively associated with stand functional richness, suggesting that  
612 stands with larger trunks also occupy more niche space. Soil properties, specifically exchangeable calcium and  
613 the sum of bases, are significantly related to maximum stem diameter, highlighting the importance of soil bases  
614 for secondary growth, as already observed in other Brazilian dry forests (Angélico et al. 2021).

615 The relationships between functional richness and all measured soil cations (excluding exchangeable  
616 aluminium) suggest that variations in soil properties may drive differences in plant physiology and anatomy,  
617 yielding optimal trade-offs between secondary growth and water-use efficiency strategies. Our results suggest  
618 that increased soil nutrient availability across different geological formations in the Caatinga enables a broader  
619 range of conservative and acquisitive strategies, as reflected in the community functional properties studied  
620 here, thereby maximising functional diversity at the regional scale.

621

622 Soil-mediated effects of functional assemblage on above-ground woody biomass

623 Our Structural Equation Model (SEM; Fig. 5) highlights how soil properties, specifically nutrient  
624 availability, indirectly shape biomass by mediating community functional composition. The SEM shows that  
625 soil nutrient availability, tree diameter, and mean annual precipitation are crucial for stand-level biomass  
626 accumulation in the Caatinga region. Specifically, soil calcium not only directly impacts above-ground biomass  
627 through mechanisms already discussed in the previous sections, but also influences wood density, maximum  
628 stem diameter, and functional richness, aligning with previous studies on the role of soil nutrients in vegetation  
629 structure and community assembly in the Caatinga region (Souza et al. 2019; Oliveira et al. 2019; Maia et al.  
630 2020).

631 Considering its effect size, the community-weighted maximum stem diameter was the strongest  
632 predictor of biomass, consistent with the ‘biomass ratio’ hypothesis (Grime 1998), in which predominant traits  
633 are crucial for determining vegetation stand-level attributes in a given community. This suggests that the  
634 abundance of typically larger species reflects biomass patterns at the stand level. Rather than serving solely as  
635 a biomass predictor, community-weighted maximum stem diameter captures ecological filtering, reflecting the  
636 ability of certain species to establish and dominate in the community. In the SEM, mean annual precipitation  
637 influenced biomass positively alongside other variables, underscoring the importance of rainfall totals for  
638 biomass accumulation, as comprehensively discussed in this paper. However, variables related to water  
639 availability, such as climatic water deficit and maximum plant-available soil water, showed weaker or no  
640 significant effects on biomass in the SEM. The multi-model inference framework tested a broader set of  
641 environmental variables, including interaction terms, while the SEM provides a more integrative picture of the

642 relationships among environmental and vegetation variables. These approaches were conceived as  
643 complementary rather than directly comparable.

644 The inverse relationship found between wood density and maximum stem diameter indicates that  
645 thicker trees tend to have lower wood densities, which may reflect different plant life-history strategies. Wood  
646 density is linked to plant hydraulic safety and construction costs, with thicker trunks often resulting in higher  
647 respiration costs, which may not be optimal for Caatinga trees (Bosc et al. 2003; Larjavaara and Muller-Landau  
648 2010). Additionally, wood density has been strongly associated with mortality rates in tropical forests, with  
649 higher survival rates generally associated with denser wood (Kraft et al., 2010). Alongside these findings, it is  
650 worth noting that although our study plots are considered structurally mature, older, thicker trees are relatively  
651 rare in many areas of the Caatinga due to chronic wood extraction by local communities.

652 The functional richness index exhibited only a weak, marginally positive association with AGB<sub>w</sub>,  
653 providing little support for the niche complementarity hypothesis (Tilman 1999). Despite this marginal effect,  
654 our result contrasts with the findings of Prado-Junior et al. (2016), who observed a positive effect of functional  
655 divergence and evenness on biomass, rather than functional richness. Their study suggested that communities  
656 with functionally distinct, yet evenly abundant individuals, are more likely to exhibit higher biomass over time.  
657 Prado-Junior et al. (2016) included specific leaf area (SLA) in their functional diversity index, although this  
658 trait was less significant in explaining biomass in their work.

659

660 Caveats and future directions

661 Our *Space-for-Time* approach (Pickett 1989) supports existing ecological hypotheses while contrasting  
662 others. For example, we found no significant impact of the maximum temperature of the warmest month on  
663 AGB<sub>w</sub> in the Caatinga. High temperatures can induce tree mortality via hydraulic failure and carbon starvation  
664 (McDowell et al. 2018), but the Caatinga flora is adapted to endure extreme heat and drought. High temperatures  
665 also affect photosynthesis via stomatal closing, which depends on the optimal/maximum values for each  
666 species. Adaptive mechanisms include deciduousness, leaf trait modifications, osmoprotectant accumulation  
667 (Medina 1983; Mathur et al. 2014; Jajoo and Allakhverdiev 2017), deeper root systems, and arbuscular

668 mycorrhizal associations (Hodge 2009; Smith and Smith 2011). However, the lack of temporal data limits our  
669 conclusions. Long-term monitoring is crucial for accurately assessing the impact of temperature and other  
670 environmental variables on vegetation structure and functioning in SDTFs.

671 While AGB<sub>W</sub> values were not significantly influenced by the RSGs, clay mineral types, or geological  
672 substrates evaluated here, we do not generalise these findings to the entire region. A broader sampling  
673 incorporating more observations, a wider range of clay mineral proportions, additional RSGs (e.g., Ferralsols),  
674 and geological substrates would be necessary to more comprehensively test this hypothesis.

675 Another limitation of the current approach is that the estimation of maximum plant-available soil  
676 water does not account for stoniness. The presence of stones and rock fragments was recorded only semi-  
677 quantitatively during field sampling, making it unsuitable for volume correction without introducing  
678 considerable uncertainty. Additionally, the pedotransfer function employed was not calibrated to accommodate  
679 significant coarse fragment content. While the work of Saxton and Rawls (2006) incorporates rock fragment  
680 corrections, it is not considered appropriate for the edaphoclimatic conditions of the semi-arid Caatinga. It is  
681 also important to note that most of our study sites are located in sedimentary terrains, where stoniness is  
682 generally negligible or absent. In crystalline landscapes, although rock fragments may be more frequent, the  
683 depth of soil is likely the primary constraint on plant-available water storage. Future work could benefit from  
684 more detailed assessments of coarse fragment contents and their implications for water retention, particularly  
685 in rocky landscapes.

686 Although our SEM demonstrated reasonably good fit indices, it is important to acknowledge the  
687 exploratory nature of the model and its relatively small sample size, which can reduce statistical power and  
688 increase the risk of Type II errors, where true relationships may go undetected. A larger sample size would  
689 strengthen the robustness of the estimates and enhance the generalisability of the findings. Furthermore, while  
690 the model tested various explanatory pathways, incorporating additional environmental variables or alternative  
691 pathways could reveal relationships not captured in the current analysis. Therefore, future studies with larger  
692 datasets are needed to disentangle other intricate relationships among environmental factors, ecosystem  
693 structure, and functioning. Furthermore, future studies should explore the role of clay mineralogy on soil  
694 hydraulic properties (water retention curve and hydraulic conductivity) and related effects on root zone storage

695 and root water uptake, as well as on root viability, which could enhance plant resilience under water-limited  
696 conditions. Finally, given the significant environmental heterogeneity of the Caatinga region and its long  
697 history of human alterations, caution is needed to avoid overgeneralising our results.

698

699 **Conclusions**

700 Our study unravels the complex interplay between climate, soil properties, and vegetation properties  
701 in SDTFs of the Caatinga region. The multi-model inference approach employed proved effective in capturing  
702 these relationships, while the structural equation model provided a comprehensive picture of how environmental  
703 factors and functional attributes collectively influence above-ground woody biomass.

704 Soil nutrient availability, mean annual precipitation, and the interaction between climatic and edaphic  
705 factors emerged as key drivers of above-ground woody biomass in the Caatinga. Beyond their direct influence  
706 on stand-level biomass, soil cations played a significant role in shaping community-weighted traits and  
707 functional richness. In synthesis, more favourable soil conditions (i.e., higher nutrient availability and greater  
708 water storage capacity) and higher mean annual precipitation, altogether, positively influenced above-ground  
709 woody biomass.

710 While our study provides valuable insights into the ecology of SDTFs, limitations such as a relatively  
711 small sample size and the absence of temporal data restrict the generalisability of our findings. Nevertheless,  
712 our research advances understanding of the role of functional attributes in AGBw accumulation patterns within  
713 SDTFs, supporting the forecasting of potential tipping points and ecosystem state shifts, as highlighted by  
714 Muñoz et al. (2023). These findings carry important implications for biodiversity conservation and carbon  
715 sequestration initiatives in dry tropical regions, offering guidance for policymaking in the face of global  
716 environmental change.

717

718

719

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729

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1053 **Statements and declarations**

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1066 **Competing Interests**

1067 The authors declare that they have no known competing financial interests or personal relationships  
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1069 **Author contributions**

1070 All authors contributed to the study writing, review and editing. Conceptualisation: A. T. B. and J. L.;  
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1076 **Data availability**

1077 Original vegetation and soil data are integrated into the *ForestPlots Network* ([www.ForestPlots.net](http://www.ForestPlots.net)).  
1078 High-resolution images of the voucher specimens are also publicly accessible through the *speciesLink* network  
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