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Article:

Roberts, Neil and Walsh, Kevin orcid.org/0000-0003-1621-2625 (2025) Palaeo-perspectives on anthropogenic soil loss and landscape resilience: An introduction. Journal of Quaternary Science. ISSN: 0267-8179

<https://doi.org/10.1002/jqs.70021>

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Palaeo-perspectives on anthropogenic soil loss and landscape resilience: an introduction

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Abstract

This paper provides an introduction and context for a series of empirical-historical case studies in a joint Virtual Special Issue (VSI) of *Geoarchaeology* and *Journal of Quaternary Science* that offer a palaeo-perspective on anthropogenic soil loss and landscape resilience. The articles cover a range of different environments and timescales, and seek to address the overarching question of how much erosion has been accelerated due to human action. They potentially allow the identification of a pre-disturbance baseline on which a subsequent increase of anthropogenic causation was superimposed. In several cases, this onset – along with associated actions to conserve soil (e.g. via agrarian terrace construction) - can be traced back to around 4000 years ago.

How fast is Planet Earth being eroded? How much of this erosion is 'natural', and how much has it accelerated due to human action? More than a century ago, Sherlock (1922) claimed that the amount of sediment moved by humankind exceeded that of any other geomorphic agent. But was he right? The question has been much debated in the years since then (Hooke 2000, Syvitski et al. 2005, Wilkinson 2005, Giusan et al. 2017). Alongside numerical modelling, two principal empirical approaches have been adopted in attempts to answer this question. The first involves the measurement of contemporary rates of sediment flux (i.e., sediment transport) via different agencies. Flux of sediment reflects landscape stability and instability, whether of climatic, tectonic or human origin, or some combination. Transporting agencies include glacier ice and the wind (mainly aeolian dust), but globally, the dominant means of sediment movement from land to sea is via rivers, as suspended or dissolved load (Gibbs 1970; Douglas 1990). Studies have ranged in spatial scale from measurement of experimental soil erosion plots under different land cover types (e.g. Rapp 1975), through small catchment monitoring (Boardman et al. 1990), to analysis of sediment discharge from major river basins like the Yangtze and Ganges (Walling and Webb 1983; Summerfield and Hulton 1994). The alternative approach is historical, commonly involving the examination of sedimentary 'archives' in lakes, river valleys and estuaries to establish

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46 longer-term rates of change. The papers in this joint Virtual Special Issue (VSI) of
47 Geoarchaeology and Journal of Quaternary Science adopt this second strategy via a
48 series of empirical-historical case studies that cover a range of different environments
49 and timescales (Table 1). It is an approach based on the assumption that at some point
50 in the past, human impact on erosion rates was minimal relative to natural agencies,
51 and that this provides a baseline for assessing any subsequent increase of
52 anthropogenic causation. In regions such as Europe and South Asia, this baseline is
53 believed to date to mid-Holocene times, whereas it continued into the late Holocene in
54 areas such as temperate North America (Van Vliet-Lanoë et al. 1992; Dearing and Jones
55 2003; Hoffmann et al. 2010).

56
57 While a shared theme underpins this joint special issue, that of soil erosion, and the
58 study of the complex interplay between climate and anthropogenic mechanisms, the
59 allocation or “badging” of articles as either Geoarchaeology or Journal of Quaternary
60 Science was based on the extent to which colleagues explicitly engaged with
61 archaeological phenomena in their studies. For example, von Zon et al.’s (2025)
62 assessment of the role of Roman road networks in altering erosion dynamics in the Gete
63 catchment, Belgium, or the two papers that review recent geoarchaeological
64 approaches in the study of terraces, Brown et al. (2025) and Kinnaird et al. (2025).

65
66 The sustainability of anthromes (anthropogenic biomes) for human use depends on
67 maintaining soil and ecosystem functions, and these have often been threatened by
68 mismanagement and landscape degradation. Soil erosion, for example, affects not only
69 upstream source areas but also downstream zones of sediment deposition in valley
70 bottoms and via coastal progradation. Storage of sediment in “sinks” such as basal
71 hillslope colluvium represents an important methodological impediment to scaling up
72 micro-scale studies to larger river basin flux rates, as Trimble (2009) showed for small
73 tributaries of the Mississippi. Alluvial stratigraphies have revealed how overbank
74 sedimentation in the Mid-West USA increased historically linked to agricultural
75 intensification, including several major increments of deposition from large floods
76 (Knox 2006). Lowland regions such as this typically had relatively low natural
77 denudation sites, and land-use conversion to cropland often increased erosion by at
78 least an order of magnitude, but much of this eroded soil failed to enter downstream
79 river systems. Observations about denudation rates are thus scale-dependent, and
80 sediment yields derived from small erosion plots cannot be applied to sub-continental
81 drainage basins (and vice-versa). Soil erosion can also permanently alter the potential
82 natural and agro-ecosystems that emerge during subsequent periods of landscape
83 recovery. The uplands of southern Anatolia, for example, saw a switch in forest
84 composition during the late Holocene, with pine trees becoming dominant during early
85 Medieval reforestation. This followed a period of cultural land use in antiquity, the

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86 Beyşehir Occupation phase, during which erosion permanently altered soil conditions,
87 favouring pine relative to other tree taxa (Roberts 1990; van Loo et al. 2017).

88

89 Geomorphological environments which record soil erosion histories can be divided into
90 those of net erosion and those of net deposition. The former ought to provide the more
91 accurate soil loss data, but on their own, degraded hillslopes are an unreliable guide to
92 erosion history. Their mere existence provides no clue about the age or origin of erosion,
93 and this has given rise to long-standing debates, such as discussions about the causes
94 of gully (arroyo) incision in regions such as the American southwest (Cooke and Reeves
95 1976). A southern African equivalent is provided by the question of the origin of dongas
96 (incised gullies), the subject of critical evaluation by Lyons et al. (2024; this VSI). They
97 used OSL dating to show that incision at three sites preceded European settlement and
98 land-use conversion. Rather, they argue that gully initiation was triggered by alternating
99 droughts and floods linked to climate changes during the millennium prior to the 19th
100 century. The need for accurate dating underpins all attempts at reconstructing
101 historical changes in sediment flux, including those focused on the point of deposition,
102 where sediment yield is used as a surrogate for soil loss. Ideally, this should involve not
103 only accumulation rates in a single section or core, but laterally to allow three-
104 dimensional volumetric calculations. For quantitative reconstructions using
105 sedimentary records, it is best to use relatively small, 'closed' catchments such as lake
106 drainage basins with a high trap efficiency (Dearing 1994). Two of the studies in this VSI
107 involve lake basins, one a study from Mexico (Aguilera Lara and Metcalfe 2025), the
108 other in Malaysia (McGowan et al. 2025). The former involves a comparison of two lake
109 basins within the highlands of Michoacán, one of which was known to have suffered land
110 degradation and erosion, while the other is considered to be relatively undisturbed. In
111 fact, palaeolimnological research shows that, even at the time of the Spanish conquest,
112 neither was pristine, due to pre-Hispanic landscape disturbance. The latter study of a
113 small tropical wetland in peninsular Malaysia takes a shorter time frame, focussing on
114 organic carbon flux, whose sources shifted following forest clearance in the 1940s and
115 subsequent plantation of rubber and palm oil trees. Both studies have been able to
116 make use of historical archives that record landscape changes, and they contribute a
117 long-term perspective to current debates over sustainability and conservation.

118

119 Unlike lakes, river valleys do not represent closed sedimentary systems, with much of
120 the eroded soil and substrate being exported, thus making historical sediment budgets
121 harder to calculate. On the other side, river valleys are widespread in the landscape and
122 have been the locus of human settlement, now and in the past. In a pair of coupled
123 articles in this VSI, Hoevers et al. (2024) and von Zon et al. (2025) report the results of
124 investigations into the Holocene history of land cover and floodplain geoecology in
125 Belgium's loess belt and sandy Campine regions, which can be linked to variations in
126 human impacts, notably via changes in land cover. These studies exploit the potential

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127 of multi-proxy data sources, including pollen, testate amoebae and stable isotopes to
128 reconstruct causal factors and response variables within the same catchments.
129 Another important contribution of palaeo-science has been to show how sediment
130 sources typically shifted during the Holocene from bedrock to topsoil derivation, using
131 sediment tracing methods such as mineral magnetism (Thompson and Oldfield 1986). In
132 many environments, this is an effective way of testing Sherlock's (1922) assertion, by
133 separating material denuded by natural processes from those added by human actions.

134
135
136 Further south, the Mediterranean basin has long been the subject of debate about the
137 causes of the historical erosion. The existence of at least one major cycle of alluvial
138 infilling and incision during the late Holocene was demonstrated by Claudio Vita-Finzi in
139 his seminal "Mediterranean Valleys" book (1969) and is the subject of a re-evaluation in
140 this VSI by Bintliff (2025). Vita-Finzi believed that these hydro-geomorphological changes
141 were caused primarily by post-Roman climatic variations, whereas others (e.g.,
142 Wagstaff 1981) argued for primarily human causation. In a synthesis of records from
143 different Mediterranean regions, Walsh et al. (2019) showed that overall rates of
144 sediment flux increased substantially during the last four millennia, coincident with
145 increases in human population and a decrease in woodland cover. While this is
146 consistent with anthropogenic causation, alluviation did not increase linearly but
147 alternated over time, in line with Vita-Finzi's model. The timing of valley sedimentation
148 and incision phases may therefore have been determined primarily by prevailing
149 climatic conditions, acting as a pacemaker. Bini et al (2025) exploit the potential of the
150 fact that river valleys are often spatially congruent with rich archaeological field
151 evidence to reconstruct the complex history of river channel avulsion between Lucca
152 and Pisa in west central Italy. Comparison with nearby cave speleothem records shows
153 that a major flood phase during the 1st century CE coincided with a period of wetter
154 climate in this region. Berger and Brochier (2025) also report the history of a river
155 debouching into the Mediterranean Sea, in this case the Rhône, which straddles the
156 divide between temperate, alpine and Mediterranean zones. Like Bini et al. (2025), they
157 were able to exploit geoarchaeological field evidence, in this case uncovered during
158 construction work for the high-speed TGV rail line. This has produced detailed and well-
159 dated cross-sections through river alluvium, interbedded with colluvium and buried
160 soils, showing multiple alternating phases of sedimentation and landscape
161 stabilisation during the mid-late Holocene. Thus, it seems that both human and
162 climatic factors played a part in changes in fluvial sediment flux, the former often acting
163 as an ultimate cause, the latter as a proximal one.

164
165 Human land conversion, notably from forest to farmland, has reduced rainfall
166 interception and increased the potential for soil particle entrainment. At the same time,
167 people have a strong incentive to conserve soil, especially in agricultural terrain, and

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168 many landscapes are littered with evidence for past soil conservation, notably via
169 agricultural terraces, as highlighted in this VSI by Brown et al. (2025) and Kinnaird et al.
170 (2025). These culturally created landscapes have a long history and continued potential
171 for future sustainable development. OSL dating shows that terrace construction and re-
172 building did not always occur at times expected from historical narratives, highlighting
173 how rural economies could sometimes function independently of urban centres and
174 wider polities.

175
176 Palaeo-scientific research has highlighted the complex historical interaction between
177 human and natural agencies in altering Hutton's rock cycle. It has also pointed to
178 mechanisms that may not be evident in studies based on monitoring and measurement
179 of contemporary processes alone. The ergodic approach, for instance, in which space
180 is substituted for time, fails to capture the significance of short-lived phases of
181 landscape instability during transitions between one land cover regime and another.
182 This was demonstrated 50 years ago by Margaret Bryan Davis in her classic (1976) study
183 of Frains Lake in the US Midwest, which revealed how the ploughing up of woodland in
184 the mid-19th century led to positive feedback and landscape instability, before the
185 system eventually stabilised at a new, but higher level of sediment flux. Palaeo studies
186 such as this consequently make it possible to establish the time taken for system
187 recovery following disturbance. At Nar Lake in central Anatolia, whose sediments are
188 annually laminated, Roberts et al. (2019) were able to show that erosion rates fell back
189 to pre-disturbance levels within a few decades after land abandonment and rewilding in
190 the late 7th century CE. Records such as this therefore offer hope for the future, by
191 telling us that many human-degraded landscapes are capable of self-healing. This has
192 repercussions for how palaeoenvironmental sciences, including geoarchaeology, can
193 contribute to contemporary debates on landscape management and the mitigation of
194 changing climates and agricultural practices. As von Zon et al. (2025) suggest, the
195 variation in the timing (or lag) in the response of geomorphic systems to different forms
196 of perturbation has implications for when and how societies intervene and what
197 baseline environment they are aiming to recreate and protect. If we accept that
198 agricultural terraces constitute one of the most significant forms of anthropogenic
199 erosion mitigation practices, a key question for contemporary landscape managers is
200 how or even if terrace systems should be preserved in landscapes where agriculture has
201 waned or is no longer practised. Such questions have repercussions for the design of
202 research projects and methodologies; specifically, the need to incorporate catchment-
203 based studies of alluvial systems and erosion with studies of agricultural
204 landscapes/slopes and colluvial systems to interrogate evidence for breaks in the
205 sedimentary record due to positive human mitigation.

206
207 Finally, we can return to a key question posed earlier in this paper, namely, at one point
208 in time do palaeoenvironmental and geoarchaeological data indicate that an upward

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209 shift in the rate of soil loss occurred? In other words, can we identify regional thresholds
210 in landscape stability when anthropogenic activities raised erosion rates outside the
211 envelope of the long-term geological norm? In the empirical studies presented in this
212 VSI, it is clear that the mid-late Holocene transition (i.e., the start of the Meghalayan
213 around 4000 years ago) represents a key tipping point in many regions of both Old and
214 New Worlds. In western Europe, sediment discharge and flooding within the catchment
215 of the river Rhône, began to depart from the previous, relatively stable trajectory around
216 2500 BCE (Berger and Brochier 2025), while in Belgium’s loess belt, a comparable shift
217 took place from ~1900 BCE (Hoevers et al 2024). In the highlands of Mexico, lake basin
218 records similarly show evidence of anthropogenic disturbance of erosion regimes from
219 at least 1500 BCE (Aguilera Lara and Metcalfe 2025). Significantly, the 2nd millennium
220 BCE also marks some of the earliest definitive dates in Europe for the construction of
221 agricultural terraces designed to conserve soil and moisture, in locations ranging from
222 Britain to the Aegean (Brown et al. 2025). Thus, human-induced acceleration of soil loss
223 and measures to counteract this can be detected at a regional scale back to Bronze Age
224 times (or their equivalent) in at least some parts of the world. If this can be considered
225 as a mark of a regional Anthropocene, then its onset can be traced back in time well
226 before the mid-20th century “Great Acceleration” (Steffen et al., 2015).

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