



UNIVERSITY OF LEEDS

This is a repository copy of *Early Triassic super-greenhouse climate driven by vegetation collapse*.

White Rose Research Online URL for this paper:

<https://eprints.whiterose.ac.uk/id/eprint/227242/>

Version: Accepted Version

Article:

Xu, J., Yu, J., Yin, H. et al. (14 more authors) (Accepted: 2025) Early Triassic super-greenhouse climate driven by vegetation collapse. Nature Communications. ISSN 2041-1723 (In Press)

This is an author produced version of an article accepted for publication in Nature Communications made available under the terms of the Creative Commons Attribution License (CC-BY), which permits unrestricted use, distribution and reproduction in any medium, provided the original work is properly cited.

Reuse

This article is distributed under the terms of the Creative Commons Attribution (CC BY) licence. This licence allows you to distribute, remix, tweak, and build upon the work, even commercially, as long as you credit the authors for the original work. More information and the full terms of the licence here:
<https://creativecommons.org/licenses/>

Takedown

If you consider content in White Rose Research Online to be in breach of UK law, please notify us by emailing eprints@whiterose.ac.uk including the URL of the record and the reason for the withdrawal request.



eprints@whiterose.ac.uk
<https://eprints.whiterose.ac.uk/>

Supplementary Materials for

Early Triassic super-greenhouse climate driven by vegetation collapse

Zhen Xu *et al.*

*Corresponding author: Zhen Xu, Z.xu@leeds.ac.uk; Jianxin Yu, yujianxin@cug.edu.cn; Benjamin Mills,
B.Mills@leeds.ac.uk.

The PDF file includes:

Supplementary text 1
Figures S1 to S6

Other Supplementary Materials for this manuscript include the following:

Table S1 to S7

disappearance of coal, the spore *Dulhuntyispora parvithola* and the appearance of the *Playfordiaspora crenulate*. This zircon age suggests the land crisis was 410–700 Kyrs earlier than the PTB^{89,138–141}.

In Antarctica, the terrestrial PTB is marked by the disappearance of a Gondwana *Glossopteris* flora, appearance of the *Lystrosaurus*, and negative carbon isotope shifts^{142,143}. Argentina is also found in high southern latitudes of Pangea. The PTB is placed in the top of the Puesto Tscherig Formation or the bottom of the Puesto Vera Formation with an imprecise age of 253±2 Ma in the boundary between these two formations^{73,144}. The *Glossopteris* flora extinction happened around the middle Changhsingian in the South Polar area, which is much earlier than the PTB^{73,145}.

North China sat in mid northern latitudes of eastern Tethys. The extinction of the Voltziales flora in the region occurs within the Sunjiagou Formation, at a level dated as 252.21±0.15 Ma^{84,146}. The Early Triassic spinicaudatan assemblage of *Euestheria gutta* - *Palaeolimnadiopsis vilujensis* occurs above this flora. In North China, the disappearance of coal and floral losses are 150–270 Kyrs earlier than the marine extinction, and palaeomagnetic records show that the Smithian/Spathian boundary is possibly in the upper Liujiagou Formation, whilst there is a hiatus at the Olenekian/Anisian boundary^{84,146}.

The PTB terrestrial records of the Xinjiang region show the replacement of late Permian tetrapods (*Striodon magnus*, *Jimusaria (Dicynodon) sinkianensis*, *Dicynodon tienshanensis*) by a *Lystrosaurus* fauna. There is also a complete turnover of the ostracod fauna; these changes predate the marine losses^{91,92,147,148}.

The European area we study was in the continental interior of Pangea and at a similar latitude to North China. Only the western Germanic Basin contains continuous sequences. The terrestrial PTB is placed in the bottom of the Buntsandstein Formation between the boundary of the spinicaudata *Falsisca eotriassica* assemblage and the *F. verchojanica* assemblage, the palynological *Lundbladispora obsoleta* - *Lunatisporites noviaulensis* assemblage and the *Lundbladispora willmotti* - *Lunatisporites hexagona* assemblage, which coincide with a carbon isotope negative excursion, and palaeomagnetic evidence^{36,83,107,149}.

In Siberia, the PTB is placed within the lower Maltsev Formation of the Kuznetsk Basin according to detrital zircon ages and magnetostratigraphy. As in Xinjiang, the ostracod fauna shows a major turnover in the lower Maltsev Formation although standing diversity remains fairly constant⁴⁹. Angara *Cordaites* species dominate the plant assemblages below the Maltsev Formation, but they do not range up into this Formation where ferns, sphenopsids and lycopsids occur instead⁴⁹.

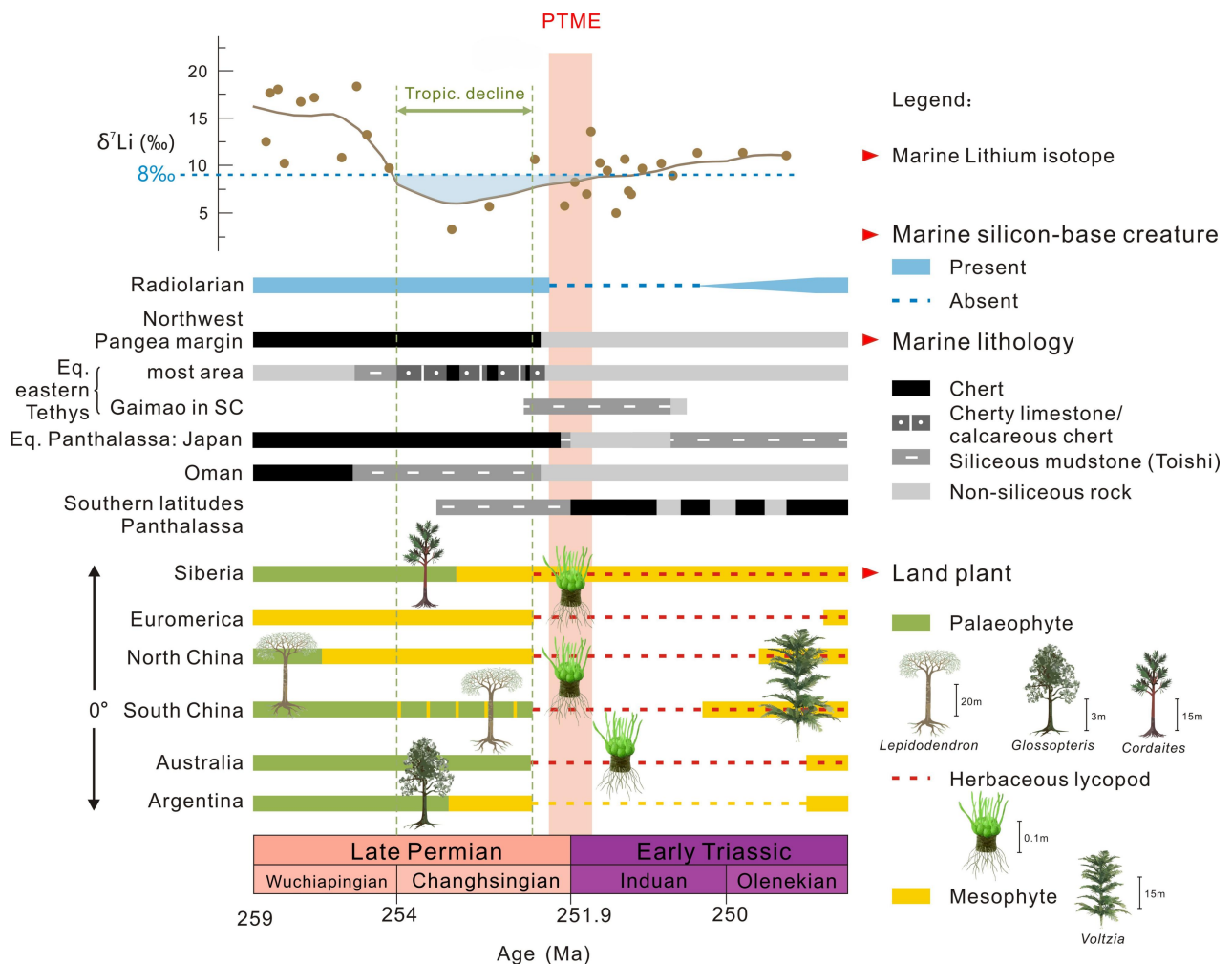
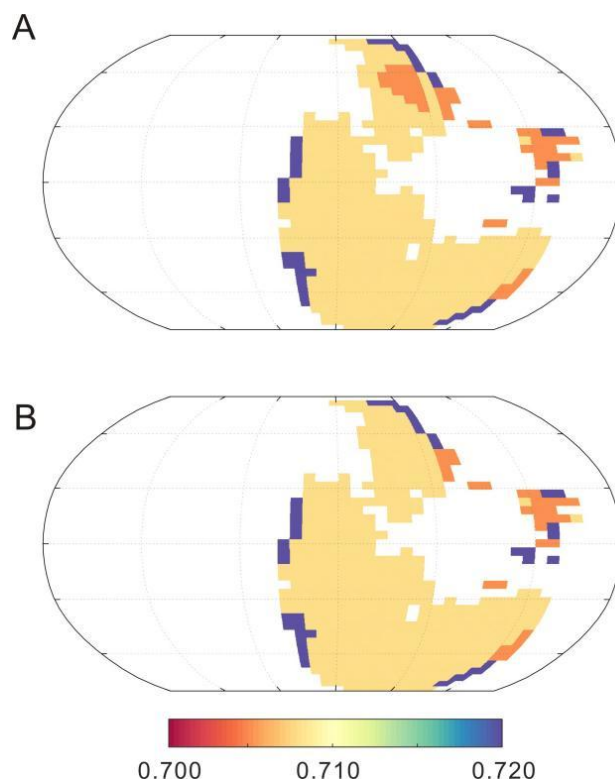


Figure S1. Summary of recent evidence showing the persistence of cherts during the PTME and the early recovery of radiolarians. The land plant extinction is the only event that continues over the whole Early Triassic hothouse. Eq.: Equatorial; SC: South China; Tropic. Decline: Tropical forest decline; PTME: Permian Triassic Mass Extinction. Lithium isotope data comes from ref.¹⁵. Radiolarian data comes from ref.¹⁵⁰. Lithology data of global major marine sections comes from ref.¹⁹. All the fossil plant reconstructions are inspired after ref.²².

1



2

3

4

5

6

7

8

9

10

11

12

Figure S4. Model strontium isotopic $^{87}\text{Sr}/^{86}\text{Sr}$ distribution for continental gridcells. The *SCION* Earth surface is divided into 48x40 gridcells. Continental arcs have more radiogenic felsic values, whereas LIPs and suture zones have more mafic unradiogenic values. All other continental gridcells have an average sediment value. A. Early to Middle Triassic post Siberian Traps emplacement. B. Late Permian prior to Siberian Traps emplacement. See methods and text for literature sources. The palaeogeographic reconstructions are inspired by the PaleoMAP Project (<http://www.scotese.com/Default.htm>).

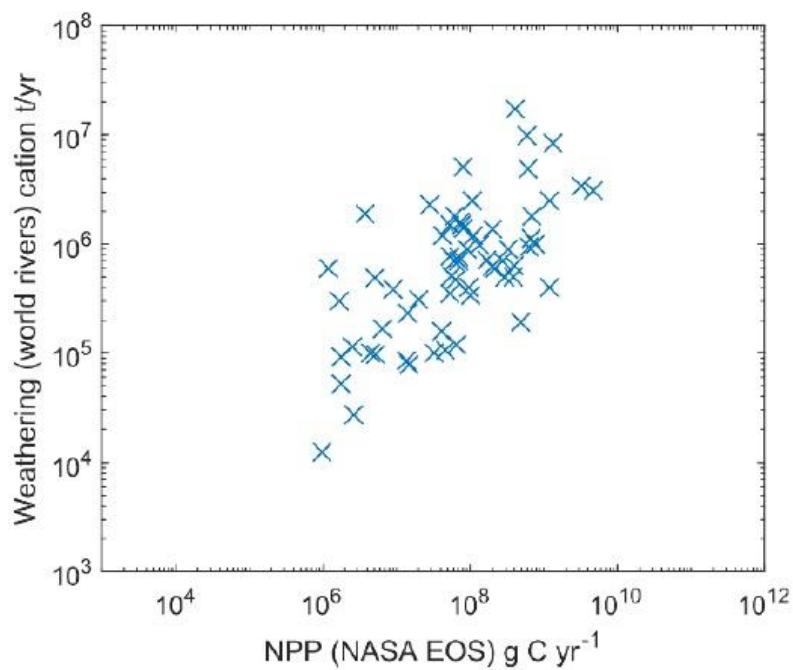
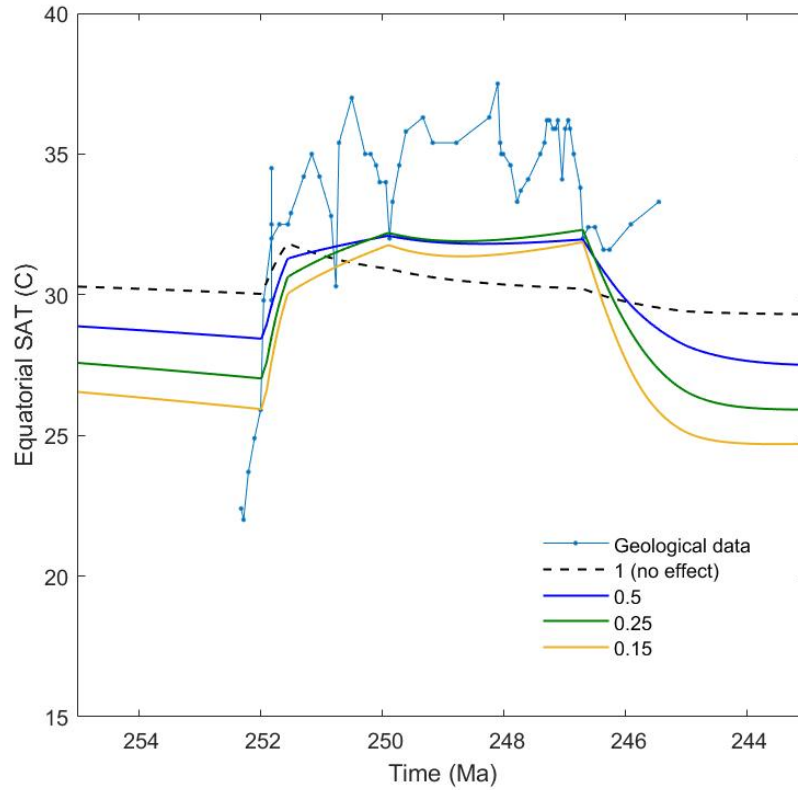


Figure S5. Recent vegetation Net Primary Productivity (NPP) and catchment scale silicate weathering relationship. From Rogger et al.⁵¹ and Gurung et al.^{52,117}.



18

19 **Figure S6. *SCION* model outputs with different assumptions for biotic weathering enhancement.**

20 Line colours show model runs with different values for relative ‘preplant’ weathering rates. The figure in
 21 the main paper uses 0.25, taken from the *GEOCARB* models. A lower preplant value results in a large
 22 temperature change when the vegetation is limited. The effect of this biotic weathering outweighs the
 23 effect of both Siberian Traps degassing and changes to organic carbon burial, which are demonstrated
 24 with the black dashed line in which biotic weathering does not change. See main text for further
 25 discussion. Blue line with dots shows temperature proxy data from ref.¹⁰.

26

<i>Zhutheca (Fascipteris) densata</i>			
<i>Zuberia</i> sp.			
<i>Alisporites</i>			
<i>Arberiella</i>			
<i>Corisaccites</i>			
<i>Crescentipollenites</i>			
<i>Cycadopites</i>			
<i>Ephedripites</i>			
<i>Eretmonia</i>			
<i>Falcisporites</i>			
<i>Faunipollenites</i>			
<i>Guttulapollenites</i>			
<i>Klausipollenites</i>			
<i>Lunatisporites</i>			
<i>Pityosporites</i>			
<i>Platysaccus</i>			
<i>Striatopodocarpites</i>			
<i>Strotersporites</i>			
<i>Weylandites</i>			

									UK	-1.7	52.3
									UK	-1.6	54.8
									UK	-2.1	52.3
									USA	-71.4	42.4
									USA	-110.3	34.9
									USA	-105.1	35.2
									CN-Xinjiang	88.8	44
									CN-Xinjiang	81.8	41.8
									CN-Xinjiang	84.8	45.7
									CN-Xinjiang	88.2	44.2

