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Chen, X., Cao, J., Sinsabaugh, R.L. et al. (6 more authors) (2025) Soil extracellular enzymes as drivers of soil carbon storage under nitrogen addition. Biological Reviews. ISSN 1464-7931

https://doi.org/10.1111/brv.70021

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1	Soil extracellular enzymes as drivers of soil carbon storage under
2	nitrogen addition
3	
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27

28 ABSTRACT

29 Enhanced anthropogenic nitrogen (N) inputs to ecosystems may have substantial impacts on 30 microbially mediated soil organic carbon (SOC) cycling. One way to link species-rich soil 31 microbial communities with SOC cycling processes is via soil extracellular enzyme activities 32 (EEAs). However, the effects of N addition on EEAs and the associated driving factors 33 remain poorly understood. By conducting a meta-data synthesis, we find that N addition 34 increases hydrolytic C-degrading EEAs that target simple polysaccharides by 12.8%, but 35 decreases oxidative C-degrading EEAs that degrade complex phenolic macromolecules by 36 11.9%. The net effect of N addition on SOC storage is determined by the shifts between these 37 two types of C-degrading EEAs, and the impacts varied across different ecosystem types. 38 These insights highlight the crucial but understudied roles of hydrolytic and oxidative C-39 degrading EEAs on SOC dynamics with ongoing enhanced anthropogenic N loading. 40 Understanding the mechanisms behind these C-degrading EEAs could help optimize SOC 41 sequestration and inform climate mitigation strategies across different ecosystems. 42 Key words: nitrogen addition, hydrolytic C-degrading enzyme activities, oxidative C-43 44 degrading enzyme activities, soil microorganisms, soil carbon storage, meta-data synthesis. 45 46 **CONTENTS** 47 I. Introduction

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66 I. INTRODUCTION

67 Reactive nitrogen (N) emissions to the atmosphere have increased substantially over the past

decades in most regions, primarily due to human activities (Fenn et al., 2018; Liu et al.,

69 2013). This escalation exhibits significant spatial variation among regions (Liu *et al.*, 2016;

70 Schwede et al., 2018). Increasing N deposition to the land surface has significantly increased

- 71 the productivity of terrestrial ecosystems, notably due to the alleviation of N limitation
- 72 (Mason et al., 2022). Furthermore, N availability is key to ecosystem functioning, the cycling
- of nutrients and the flow of energy through the biosphere, with profound impacts on soil
- 74 organic carbon (SOC) decomposition in terrestrial ecosystems (Ye *et al.*, 2018).

75 Despite numerous studies in recent decades, the net effects of N addition on SOC storage are 76 controversial. Recent meta-analyses of N addition experiments at global or regional scales 77 have shown that N addition can increase SOC storage by 4-11% (Chen et al., 2018a; Hu et 78 al., 2024a,b; Janssens et al., 2010; Liu & Greaver, 2010; Xu et al., 2021; Yue et al., 2017). 79 This increase in SOC storage may occur because N inputs can alleviate N limitation for 80 microbial decomposers and decrease microbial degradation of pre-existing SOC for N 81 acquisition (i.e. negative priming effect and reduced N mining from organic matter) (Craine, 82 Morrow & Fierer, 2007; Hicks, Lajtha & Rousk, 2021; Kuzyakov, 2010). Conversely, N 83 addition could increase soil N availability and the activity of previously N-limited microbial 84 decomposers, thereby enhancing SOC decomposition and reducing SOC storage (Zhang et 85 al., 2014). N addition may lead to soil acidification. It can indirectly impact SOC storage by 86 inhibiting microbial growth (Ontman et al., 2023), reducing enzyme efficiency (Chen et al., 87 2017), decreasing necromass accumulation (Ye et al., 2018), and affecting soil structure 88 (Cotrufo et al., 2015; Liang, Schimel & Jastrow, 2017). These combined effects can lead to a 89 long-term decrease in SOC storage. Several mechanisms may help explain changes in SOC 90 storage following N addition, including altered C inputs due to a change in net primary 91 productivity (Čapek et al., 2018), changes in microbial activity (Crowther et al., 2019) and/or 92 changes in soil aggregation (Lu et al., 2021a), but a comprehensive understanding of the 93 overall effects of N addition on SOC decomposition remains elusive. Specifically, the lack of 94 a mechanistic understanding hinders our ability to predict long-term effects of N addition on 95 SOC dynamics. 96 Changes in SOC storage resulting from N addition are associated with a range of soil

97 microbial properties, including community composition, physiology, and metabolic activities

98 (Jing et al., 2021; Ni et al., 2021). However, establishing a direct link between diverse soil

99 microbial properties and changes in SOC storage remains a formidable challenge due to the

100 intricate feedbacks between plants, soil, and microbes (Chen, Sinsabaugh & van Groenigen, 101 2023). Extracellular enzymes are produced by plants and soil microorganisms, providing 102 another perspective for understanding the effects of N addition on SOC storage. Soil 103 extracellular enzyme activities (EEAs) mediate the transformation of SOC by catalysing rate-104 limiting steps in SOC decomposition (Fanin et al., 2022; Sinsabaugh, 2010), which may 105 directly or indirectly affect SOC storage. Hydrolytic and oxidative C-degrading EEAs are 106 among the most important C-degrading enzymes targeting the degradation of litter and SOC 107 with different qualities. Hydrolytic C-degrading enzymes catalyse the degradation of 108 cellulose, whereas oxidative C-degrading enzymes facilitate the biodegradation of lignin and 109 other phenolic compounds in litter and soil (Ljungdahl & Eriksson, 1985). Hydrolytic and 110 oxidative C-degrading EEAs can provide another mechanistic insight to track SOC dynamics 111 (Chen et al., 2017) and the ratio of oxidative to hydrolytic C-degrading EEAs provides an 112 index that reflects microbial preference in utilizing various pools of litter and SOC (Yang et 113 al., 2019). It is important to note that geography significantly influences the soil's physical 114 and chemical properties (Islam et al., 2020; Lira-Martins et al., 2022). These soil 115 characteristics, in turn, can impact the abundance, diversity, and activity of soil microbes 116 (Choe, Kim & Lee, 2021; Islam et al., 2020; Carson et al., 2007). For instance, soils rich in 117 specific minerals may foster the growth of certain microbial communities, thereby 118 influencing the overall pattern of EEAs (Choe et al., 2021; Carson et al., 2007). Therefore, 119 understanding the responses of soil C-degrading EEAs to N addition is central to predicting 120 changes in SOC cycling and its feedback to climate change, which may offer fresh insights to 121 advance the understanding of the intricate plant-soil-microbial feedback on SOC storage under N addition. 122

123 In this study, we performed a comprehensive global meta-data synthesis of the responses of 124 soil C-degrading EEAs to N addition. This study aims to provide novel insights into the role

125 of soil EEAs in mediating the response of SOC storage to N addition. Several previous meta-126 analyses have synthesized the impact of N addition on soil C-acquisition EEAs (Chen et al., 127 2017; Jia et al., 2020; Jian et al., 2016; Xiao et al., 2018), but the links between soil EEAs 128 and SOC storage have not been explicitly examined, let alone the underlying mechanisms. Additionally, there has been a surge in field studies on soil C-acquisition EEAs related to N 129 130 addition over recent years, providing a unique opportunity to examine the links between soil 131 EEAs and SOC storage. To this end, we conducted a comprehensive synthesis of data from 132 multiple field experiments that concurrently measured C-degrading EEAs and SOC in 133 response to N addition (Allison et al., 2010; Du et al., 2014; Luo et al., 2019; Rappe-George 134 et al., 2017; Zhu et al., 2020). A meta-data synthesis approach was combined with a state-of-135 the-art model selection analysis to evaluate simultaneously the importance of a range of 136 biotic and abiotic factors on the effects of N addition on SOC and C-degrading EEAs. In 137 particular, we explored the role of C-degrading EEAs related to changes in SOC storage by 138 analysing data from N addition studies that included both SOC storage and C-degrading 139 EEAs. Based on previous experimental and synthesis studies, we hypothesized that: (1) N 140 addition suppresses C-degrading EEAs due to altered N availability; and (2) N addition 141 increases SOC storage, which is associated with N-induced changes in C-degrading EEAs. 142

143 **II. METHODS**

144 (1) Data collection

We systematically searched all peer-reviewed journal articles that investigated the effects of
N addition on SOC storage as well as the underlying enzymatic mechanisms, using *Web of Science* (http://apps.webofknowledge.com/), *Google Scholar* (http://scholar.google.com/),
and *China National Knowledge Infrastructure* (CNKI, www.cnki.net/) up to October 1st,
2022. The literature search was performed following guidelines from Preferred Reporting

150 Items for Systematic Reviews and Meta-Analyses (see online Supporting Information, 151 Appendix S1 for PRISMA flow chart). Various key word combinations were used for the 152 search: (nitrogen addition OR nitrogen amendment OR nitrogen enrichment OR nitrogen 153 fertilizer OR nitrogen elevated OR nitrogen deposition) AND (glucosidase OR cellobiosidase 154 OR xylosidase OR peroxidase OR phenol oxidase OR polyphenol oxidase OR lignin 155 modifying enzymes OR cellulase) AND (soil carbon) AND (terrestrial OR soil OR land). 156 Additional studies that were not covered by these search terms but were cited in Chen et al. 157 (2018b) were also included. We reviewed each article to determine whether the studies met 158 the following criteria: (1) articles must have an N addition treatment with a paired ambient 159 treatment (control). For multifactorial studies, only ambient and N addition treatments 160 (including multiple N addition levels) were included. Data from N addition treatments 161 combined with other factors (e.g. P addition) were excluded. (2) Ambient and N addition 162 treatments must be conducted at the same experimental site, thus the microclimate, ecosystem 163 type, and soil types are similar between ambient and N addition treatments. (3) Details of N 164 addition methods (rate, frequency, form, and duration) must be provided. (4) Activity of at 165 least one kind of C-degrading EEA must be measured. (5) The experiment must be conducted 166 in the field. Studies involving model simulations and incubations in the laboratory (e.g. 167 greenhouse or pot experiments) were not included. We used Plot Digitizer version 2.0 to 168 digitally extract data from figures when the results were graphically reported. When data 169 from different years were available in the same study, we used the latest data in accordance 170 with the commonly used approach in meta-data synthesis studies (Lu et al., 2021a,b). When 171 data for the same variable from multiple soil layers were provided, we only used data from 172 the uppermost soil horizon to maintain independence across the different studies. Following the criteria mentioned above, we obtained a meta-data set of 379 paired 173 174 observations from 102 publications (see Appendix S2 for full list of included studies). In

175 brief, the variables examined in this study included C-degrading EEAs (listed in Table 1), 176 plant properties (expressed as plant productivity, e.g. aboveground biomass, diameter at 177 breast height, root biomass, litter biomass, net primary productivity), soil properties (soil pH, 178 SOC, recalcitrant SOC, soil total N and the ratio of SOC to N) and microbial properties 179 [microbial biomass carbon (MBC), microbial biomass nitrogen (MBN), MBC:MBN, 180 bacterial and fungal biomass, and the ratio of fungal to bacterial biomass]. The following 181 indicators of SOC recalcitrance were used: (i) soil lignin content or the size of the slow-182 decomposing C pool; and (*ii*) the heavy fraction of C or non-hydrolyzable organo-mineral 183 complexed C. To estimate the effect of N addition on microbial biomass, we prioritized the 184 following proxies based on their availability: (1) for MBC and MBN, we tabulated microbial 185 biomass measured by chloroform fumigation (Brookes et al., 1985; Vance, Brookes & 186 Jenkinson, 1987); (2) for bacterial biomass, we tabulated microscopic counts in soil extracts, 187 bacterial phospholipid fatty acid (PLFA) concentrations (Bossio & Scow, 1998; Frostegård, 188 Tunlid & Bååth, 2011) or quantitative PCR analysis; and (3) for fungal biomass, we used 189 microscopic counts in soil extracts, fungal PLFA concentrations, or soil concentrations of 190 ergosterol. Results for specific groups of fungi or bacteria (e.g. arbuscular mycorrhizal fungi, 191 saprotrophic fungi, ammonia-oxidizing bacteria/archaea, and actinomycetes) were excluded 192 to avoid biasing results towards subsets of the fungal or bacterial communities. In our 193 analysis, hydrolytic C-degrading EEAs were represented by the activities of β -1,4-194 glucosidase, β -1,4-xylosidase, and β -1,4-D-cellobiohydrolase, or a combination of these. 195 These hydrolytic C-degrading enzymes accelerate the breakdown of cellulose or 196 hemicellulose. Oxidative C-degrading EEAs were represented by the activities of peroxidase, 197 phenol oxidase, polyphenol oxidase, or a combination of these (Table 1). These oxidative C-198 degrading enzymes mediate the breakdown of relatively recalcitrant molecules such as lignin, 199 phenols, and other aromatics. The above enzymes are among the most important for the

200 degradation of litter and SOC (Chen et al., 2017; Jian et al., 2016; Margida, Lashermes & 201 Moorhead, 2020). In instances where researchers reported the presence of multiple types of 202 hydrolytic and oxidative C-degrading EEAs, the sum of their values was used to represent the 203 cumulative response of hydrolytic or oxidative C-degrading EEAs (Chen et al., 2018b). 204 In our data set, we also recorded a wide range of environmental variables, including latitude (30.4°S to 69.0°N), longitude (145.7°W to 137.9°E), mean annual temperature (MAT, -5.4 to 205 24.8 °C), mean annual precipitation (MAP, 69-3537 mm), initial soil pH (3.2-8.7), initial 206 SOC $(3.3-179.1 \text{ g} \cdot \text{kg}^{-1})$ and initial total nitrogen (TN, $0.43-12.7 \text{ g} \cdot \text{kg}^{-1})$). The global 207 208 distribution of the experimental sites is presented in Fig. 1. If these data were not reported, 209 we contacted the corresponding author for more information. Otherwise, we obtained MAT 210 and MAP from the WorldClim database (www.worldclim.org/), and background N deposition 211 from the Global N deposition database (http://webmap.ornl.gov/). We classified ecosystem 212 types according to the Whittaker Biome Diagram (Whittaker, 1962), and soil types according 213 to the Food and Agriculture Organization taxonomy (www.fao.org/soils-portal/soil-214 survey/soil-classification/usda-soil-taxonomy/en). 215 216 (2) Data analysis

We quantified the effects of N addition on the studied variables using response ratios (RRs), which were calculated for each observation by taking the natural logarithm of the average EEAs of the ambient and N addition treatments as described in Equation (1), in which $\overline{X_t}$ and $\overline{X_c}$ are the respective means of a given variable in the N addition treatment group and the ambient (control) group, respectively.

222
$$RR = In \frac{\bar{x}_t}{\bar{x}_c}$$
(1)

The variance (v) of RR was calculated following Equation (2), where SD_t and n_t are the standard deviation and sample size of a given variable observed in the treatment group, and \overline{SD}_c and n_c are the respective mean, standard deviation and sample size of a given variable in the control group.

227
$$v = \frac{SD_c^2}{n_c \times \bar{x}_c^2} + \frac{SD_t^2}{n_t \times \bar{x}_t^2}$$
(2)

228 The effects of N addition on plant productivity, soil characteristics, C-degrading EEAs and 229 microorganisms were evaluated by mixed-effects models using the *rma.mv()* function from 230 the R package metafor (Viechtbauer, 2010). The random factor 'experiment' was included to 231 ensure independence among multiple RRs within an experiment (an experiment may have 232 included several N addition rates). Previous meta-analyses suggest that the response of C-233 related processes to N enrichment can vary with ecosystem types and fertilization regimes 234 (Chen et al., 2015; Chen et al., 2018b; Deng et al., 2018). Thus, all observations were 235 subdivided into four categories based on ecosystem type (forest, grassland, farmland and shrubland) and various N addition properties, including N addition rate (high, >10 g N m^{-2} 236 year⁻¹; medium, 5–10 g N m⁻² year⁻¹; low, <5 g N m⁻² year⁻¹), experimental duration (<5 237 years, 5–10 years, and >10 years), N addition frequency (< 4 times year⁻¹, 4–12 times year⁻¹ 238 and > 12 times year⁻¹), and N fertilizer form [mix; organic, i.e. urea; inorganic, i.e. NH₄NO₃, 239 240 NaNO₃, NH₄Cl, (NH₄)₂SO₄]. We chose these thresholds for breakpoints based on results 241 from previous studies (Chen et al., 2018b; Lu et al., 2021a; Song et al., 2019) and general breakpoints used in manipulative N-addition experiments in our data set. The statistical 242 243 results reported include total heterogeneity among all observations, the heterogeneity in the RRs associated with each moderator variable (Q_M) , and the residual error. A significant Q_M 244 245 indicates a significant effect of the moderator variable on the RR (Hedges, Gurevitch & 246 Curtis, 1999). The effects of N addition were considered significant if the 95% confidence 247 interval did not overlap with zero. The results are reported as percentage change with N addition [i.e. $100\% \times (e^{RR}-1)$] to aid interpretation. 248

249 If standard deviations were not reported, we calculated them from the standard error and the 250 number of replicates. Alternatively, we estimated them using the coefficient of variation from 251 all complete cases, following the method described by Bracken (1992). Estimation of 252 standard deviations was performed using the R package metagear (Lajeunesse, 2016; Terrer 253 et al., 2021). We also calculated the ratio of oxidative to hydrolytic C-degrading EEAs, the 254 ratio of soil microbial C to N, and the fungal to bacterial biomass ratio. We employed the 255 error propagation method (Lorber, 1986) to calculate the standard deviation (SD_{ν}) of these 256 ratios using Equation (3), where \overline{r} and \overline{s} are the mean values of r and s, respectively and SD_r 257 and SD_s are the standard deviations, respectively.

258
$$SD_y = \frac{\bar{r}}{\bar{s}} \times \sqrt{\left(\frac{SD_r}{r}\right)^2 + \left(\frac{SD_s}{s}\right)^2}$$
(3)

259 The oxidative:hydrolytic C-degrading EEA ratio is an effective indicator for microbial 260 substrate preference (Romero-Olivares, Allison & Treseder, 2017; Sinsabaugh, 2010; Yang et 261 al., 2019), with higher ratios indicating relatively greater investment in the decomposition of 262 chemically recalcitrant C pools (Ren et al., 2017; Romero-Olivares et al., 2017). 263 Publication bias was examined using Begg's test and Egger's test (Begg & Mazumdar, 1994, 264 1994; Egger et al., 1997; Leimu & Koricheva, 2004). Begg & Mazumdar (1994) proposed 265 testing the interdependence of variance and effect size using Kendall's method, and Egger et 266 al. (1997) proposed a test for asymmetry based on a funnel plot. Our results showed that most 267 variables did not have publication bias, and where there was potential publication bias for 268 some variables, this would not affect the results because Rosenthal's fail-safe number was 269 much greater than 5n+10 (where *n* is number of observations; Table S1). 270 We examined the relative influence of multiple controlling factors on the responses of SOC 271 storage, hydrolytic C-degrading EEAs, and oxidative C-degrading EEAs to N addition using

- 272 model-selection analysis. We used the '*rma.mv* ()' function from the R package *metafor*
- 273 (Viechtbauer, 2010) and the 'glmulti()' function from the glmulti R package (Calcagno & de

274 Mazancourt, 2010) to automate the fitting of all possible models containing the predictors. 275 We utilized the Akaike information criterion to select the most parsimonious models with the lowest Akaike value. The relative importance of a given predictor was ascertained by 276 277 summing the Akaike weights, representing the probability that a particular model is the most plausible, for all models in which the predictor was present. A cut-off of 0.8 was set to 278 279 differentiate between essential and non-essential predictor variables (Calcagno & de Mazancourt, 2010). We used 'lmer' function in the lme4 package to conduct linear mixed-280 281 effects models based on a restricted maximum likelihood approach (Bates et al., 2015) to test 282 the overall effects of significant predictors in Model Selection Analysis, as well as their 283 interactions. To explore the possible mechanisms of changes in SOC storage under the 284 condition of N addition, we also conducted meta-regressions between the RRs of SOC 285 storage and the response of C-degrading EEAs, initial SOC content, N addition rate, and the 286 ratio of N addition rate to background N deposition rate. We used Spearman's rank 287 correlation analysis to evaluate the relationships of C-degrading EEAs with environment 288 properties, soil properties, and N addition properties. 289 Based on known relationships, we fitted the data to a structural equation model to evaluate 290 the direct and indirect effects of N addition on SOC storage. The model was fitted by 291 maximum likelihood estimation in R using the lavaan package (Rosseel, 2012). The fit of the 292 final model was evaluated using a chi-squared test. Models are considered to have a good fit 293 when $0 \le \text{chi-squared/df} \le 2$, as well as when $0.05 < P \le 1.00$ (Schermelleh-Engel, 294 Moosbrugger & Müller, 2003).

295

296 III. RESULTS

297 (1) Responses of hydrolytic and oxidative C-degrading EEAs to N addition

298 Averaged across all studies, N addition had contrasting effects on hydrolytic and oxidative C-

299 degrading EEAs (Fig. 2). Nitrogen addition significantly increased hydrolytic C-degrading

EEAs by 12.8% with a range of 7.5% to 18.4%. This increase was primarily due to an

301 increase in activities of β -1,4-glucosidase and β -1,4-D-cellobiohydrolase by 16.7% and

302 10.9%, respectively. By contrast, N addition significantly decreased oxidative C-degrading

EEAs by 11.9% (ranging from -17.6% to -5.9%), due to a decrease in peroxidase, phenol

304 oxidase and polyphenol oxidase activity by 7.8%, 14.6%, and 14.0%, respectively (Fig. 2C).

305 Model selection analysis suggested that the responses of hydrolytic C-degrading EEAs to N

addition were mostly explained by the initial SOC, N addition rate, and the climatic

307 parameters MAT and MAP (Fig. 3B). In particular, the response of hydrolytic C-degrading

308 EEAs to N addition was greater when initial SOC levels were high (Fig. S1B). By contrast,

309 oxidative C-degrading EEAs were mostly explained by ecosystem type, soil type, and MAT

310 (Fig. 3C). Specifically, N addition decreased oxidative C-degrading EEAs by 12.1% for both

311 forest and grassland ecosystems (Fig. 4C). The responses of oxidative:hydrolytic C-degrading

EEA ratio were mostly explained by ecosystem type, initial SOC, and N addition rate (Fig.

313 3D). N addition decreased the oxidative:hydrolytic C-degrading EEA ratio by 19.4% in

forests and by 29.1% in farmland (Fig. 4D). In addition, N addition had more pronounced

315 negative effects on oxidative:hydrolytic C-degrading EEA ratio when soils were C-rich (Fig.

316 S1D) or when N addition rates were high (Fig. S2D). There were no significant relationships

between the ratio of N addition rate to background N deposition rate (BND) and any of the
RRs considered (Fig. S3).

319

312

320 (2) Linking shifts in C-degrading EEAs to changes in SOC storage with N addition 321 For studies that reported changes in SOC storage, N addition enhanced SOC storage by an 322 average of 7.2% and increased soil recalcitrant SOC by 16.2% (Fig. 2B). The N addition-323 induced increases in SOC storage were predicted by ecosystem type, N-induced changes in oxidative C-degrading EEAs, and N addition rate (Fig.3A). Furthermore, we found that the 324 325 response of SOC storage to N addition was positively related to hydrolytic and negatively 326 correlated with oxidative C-degrading EEAs (Fig. 5A, B). The relationships between N-327 induced changes in C-degrading EEAs and SOC storage were significant over a variety of 328 ecosystems and soil types (Figs 5 and S4). As a result, we found a strong negative 329 relationship between the overall response of SOC storage and the responses of 330 oxidative:hydrolytic C-degrading EEA ratio to N addition (Fig. 5C). According to the 331 structural equation model, N-induced changes in SOC storage and oxidative C-degrading 332 EEAs were strongly mediated by MBC rather than by changes in the ratio of fungal to 333 bacterial biomass (Fig. 6). Specifically, reductions in MBC following N addition negatively 334 affected oxidative C-degrading EEAs ultimately resulting in enhanced SOC storage. In 335 addition, soil pH had a direct and significant effect on MBC, suggesting that N-induced soil 336 acidification suppressed the decomposition of SOC by limiting the growth of soil 337 microorganisms (Fig. 6).

338

339 (3) Effect of ecosystem type on N addition-induced changes in SOC storage and EEAs

Ecosystem type was identified as an important predictor for N addition-induced changes in SOC storage and C-degrading EEAs (Fig. 4A). Overall, N addition increased SOC storage by 7.0% in forests and by 15.1% in farmlands while having no discernible effect on SOC storage in grasslands and shrublands. Regarding hydrolytic C-degrading EEAs, N addition resulted in an increase of 13.3% in forests and 27.2% in farmlands (Fig. 4B). Conversely, we observed a 345 reduction of 12.1% in oxidative C-degrading EEAs in both forests and grasslands following 346 N addition (Fig. 4C). Moreover, changes in SOC storage induced by N addition were 347 significantly related to the responses of oxidative C-degrading EEAs in all ecosystems (Fig. 348 5E) while a significant relationship between the response of SOC storage and the responses 349 of hydrolytic C-degrading EEAs was only observed for forest ecosystems (Fig. 5D). 350 In farmland ecosystems, our analysis revealed that the responses of hydrolytic C-degrading 351 EEAs were negatively correlated with soil clay content and N addition duration (Fig. 4E). 352 The responses of oxidative C-degrading EEAs were positively correlated with background 353 soil total N content and clay content, and negatively correlated with background soil C:N 354 ratio and N addition frequency. In forest ecosystems, significant correlations were observed 355 between C-degrading EEAs and soil properties. Specifically, hydrolytic C-degrading EEAs 356 showed a positive correlation with initial SOC and total nitrogen (N) content. In contrast, 357 oxidative C-degrading EEAs exhibited negative correlations with these properties (Fig. 4E).

358

359 IV. DISCUSSION

360 (1) New insights into SOC dynamics under N addition from EEAs

361 Our investigation provides compelling evidence that N addition can significantly influence 362 soil microbial anabolism, as manifested by alterations in C-degrading EEAs. Our analysis 363 demonstrated that N addition augmented hydrolytic C-degrading EEAs while suppressing 364 oxidative C-degrading EEAs (Fig. 2). These results are consistent with several recent studies 365 that also reported these same patterns (Chen et al., 2017; Jian et al., 2016; Xiao et al., 2018). 366 N limitation can stimulate the production of oxidase enzymes, as N-containing molecules are 367 often chemically bound within complex, recalcitrant substrates (Kuzyakov & Xu, 2013). By reducing microbial N limitation and increasing relative microbial C limitation with N 368 369 addition (Sinsabaugh, 2010), hydrolytic C-degrading EEAs may be stimulated while

370 oxidative C-degrading EEAs are suppressed. While a previous study identified experimental 371 duration to be a crucial factor influencing the outcomes of C dynamics in plant and soil 372 responses to N addition (Xu et al., 2021), in our analyses these relationships between changes 373 in C-degrading EEAs and SOC storage were unaffected by experimental duration (Fig. S5). This suggests that the strategies for using carbon employed by soil microbes remain 374 375 consistent across different durations of N addition (Leff et al., 2015). These robust 376 relationships offer fresh perspectives into the mechanisms that regulate SOC storage in 377 response to N addition.

378 N addition can alter soil C-degrading EEAs, thereby influencing both SOC decomposition 379 and sequestration (Cao et al., 2021; Riggs et al., 2015). Microorganisms produce oxidative 380 enzymes in part to mineralize structurally complex C sources to obtain N, which is often 381 protected or shielded by recalcitrant substrates such as lignin (Manzoni et al., 2012). Thus, 382 the reduction in oxidative C-degrading EEAs and an increase in hydrolytic C-degrading 383 EEAs may increase mineral-associated organic matter derived from the decay of cellulose 384 (Margida et al., 2020). This shift could increase the accumulation of microbial products and the formation of stable organic matter (Cotrufo et al., 2013). This could also explain the 385 386 significant increase in recalcitrant SOC under N addition (Fig. 2) and aligns with recent 387 comprehensive meta-analyses of N-induced changes in recalcitrant soil carbon (Qi et al., 388 2023; Xu et al., 2024). Indeed, Chen et al. (2018b) showed that suppression of oxidative C-389 degrading EEAs by N addition exerted greater control over SOC storage than climatic and 390 edaphic factors. A novel perspective acknowledges the substantial and direct impact of 391 microbial biomass in the accumulation of organic matter. This process involves the 392 transformation of microbial necromass into stable fractions of SOC (Kallenbach, Frey & 393 Grandy, 2016; Ye et al., 2018). For example, Lu et al. (2021a) reported an increase in soil 394 aggregate formation and SOC sequestration under N addition. Finally, increased SOC storage under N addition may result from decreased microbial N mining from recalcitrant SOC,
known as a negative priming effect (Fanin, Alavoine & Bertrand, 2020; Hicks, Lajtha &
Rousk, 2021). In the presence of excess N, soil microbes switch to using exogenous N
sources rather than degrading SOC for N (Allison & Vitousek, 2005). The reduced 'mining'
for organic nutrients can conserve organic matter and result in greater SOC storage
(Blagodatskaya *et al.*, 2007; Nottingham *et al.*, 2015).

In summary, N addition can have multifaceted effects on SOC storage through its impact on
microbial growth and EEAs. The equilibrium between oxidative and hydrolytic C-degrading
EEAs plays a pivotal role in determining the net effect of N addition on SOC storage. Further
research is imperative to advance understanding of the mechanisms underlying these effects
and to devise strategies for managing N addition to enhance SOC storage.

406

407 (2) Factors affecting the enzyme control of SOC storage under N addition

408 We provide a framework (Fig. 6) proposing several factors modulating the enzymatic 409 regulation of SOC storage under N addition. First, our data demonstrate that N addition 410 significantly stimulated plant productivity (Fig. 2), but this increased plant productivity did 411 not directly influence SOC storage (Fig. 6). Although previous studies have demonstrated 412 that N addition significantly increased the aboveground productivity of various terrestrial 413 ecosystems (Čapek et al., 2018; Schulte-Uebbing & de Vries, 2018), our results suggest that 414 these new C inputs do not necessarily increase SOC storage (Cotrufo et al., 2015; Niu et al., 415 2016; Lu et al., 2021a). This is particularly true because greater C inputs coupled with more 416 favourable soil C:N ratios, may accelerate microbial metabolic activity and SOC 417 decomposition (Liang et al., 2017). Consequently, an increase in plant biomass does not 418 result systematically in an increase in SOC storage, and N addition can also have neutral or 419 even negative effects on SOC storage depending on the specific driving mechanisms

420 (Crowther *et al.*, 2019; Liu & Greaver, 2010; Lu *et al.*, 2022). These negative responses can
421 be ascribed to either N saturation over time or the accumulated deleterious effects of long422 term N addition on plant and microbial growth (Treseder, 2008). Therefore, stimulated plant
423 productivity alone cannot be considered as the sole determinant of alterations in SOC storage
424 in terrestrial ecosystems (Xu *et al.*, 2021), with other mechanisms related to microbial
425 functions remaining largely unexplored.

Changes in microbial community size and composition may contribute to changes in 426 427 hydrolytic and oxidative C-degrading EEAs, resulting in changes to SOC storage. This is 428 because individual microbial taxa produce only a subset of the enzymes required to degrade 429 complex organic matter (Condron et al., 2010). While many kinds of soil microorganisms can 430 secrete hydrolytic C-degrading enzymes, only a small number of microorganisms produce 431 oxidative C-degrading enzymes, such as white-rot basidiomycetes and xylacarious 432 ascomycetes (Carreiro et al., 2000). N addition often results in a decrease in fungal biomass 433 and an increase in bacterial biomass (He, Ruan & Jia, 2024; Huang et al., 2023). With a 434 lower fungi-to-bacteria ratio, the soil's efficiency in forming stable SOC is reduced, 435 potentially diminishing long-term SOC storage (Ali et al., 2021). Fungi and bacteria perform 436 distinct roles in SOC storage. Bacteria typically decompose organic matter more rapidly, 437 leading to faster SOC turnover and less stable SOC storage (Don et al., 2017). Conversely, 438 fungi are more efficient at decomposing complex organic materials and forming stable soil 439 aggregates, which enhances SOC storage (Fan et al., 2022). Mycorrhizal fungi, in particular, 440 produce a variety of EEAs that aid in organic matter decomposition (Parihar et al., 2020). 441 They initially stimulate the decomposition of fresh residues and later tend to suppress the 442 decomposition of older or more decomposed SOC (Wei et al., 2019). The extensive hyphal 443 networks of these fungi contribute to SOC by transporting carbon from plants to soil and 444 forming stable soil aggregates (Wang et al., 2023; Kohler et al., 2017). Saprophytic fungi are

445 the primary cellulose degraders, while mycorrhizae (including many basidiomycetes and 446 ascomycetes) likely function as recalcitrant nitrogen miners. Augmented by their plant hosts' energy budgets, these mycorrhizae are also the main producers of phenol oxidase (Burke & 447 448 Cairney, 2002). Additionally, N addition promotes SOC accumulation by enhancing the 449 relative contribution of roots compared to hyphae in SOC accrual (Yuan et al., 2024). 450 Moreover, previous studies found that N addition can reduce microbial biomass by 5–20% 451 (Lu et al., 2011; Xiao et al., 2018; Zhang, Chen & Ruan, 2018), possibly contributing to 452 shifts in taxonomic composition. Our analysis revealed positive associations between 453 oxidative C-degrading EEAs (but not hydrolytic EEAs) and MBC under N addition (Table 454 S3), demonstrating that increasing total microbial biomass drives the increase in soil 455 oxidative C-degrading EEAs (Fig. 6). These results suggest that shifts in C-degrading EEAs 456 with N addition were associated with changes in microbial community composition. 457 However, the underlying mechanism of this response remains unclear. High-resolution 458 analyses that provide more detailed taxonomic information are necessary to assess the 459 structural response of soil microbial communities to N addition. 460 Finally, N-induced alterations in soil pH could contribute to enzymatic control of SOC 461 storage. Soil pH affects the binding of substrates to enzymes and the formation of enzyme 462 proteins (Sinsabaugh, 2010). Our shifts in C-degrading EEAs with the reduction in soil pH 463 are consistent with previous explanations for changes in SOC accumulation under N addition 464 due to changes in microbial biomass and activities (Funakawa et al., 2014). We found that N 465 addition-induced soil acidification (Fig. 2) inhibited microbial growth and oxidative Cdegrading EEAs, which were accompanied by decreased MBC (Fig. 6), microbial respiration 466 467 and SOC decomposition (Averill & Waring, 2018; Janssens et al., 2010; Niu et al., 2016; Wang et al., 2018). Moreover, the consistently negative effect of high soil pH on SOC 468

storage suggests that low soil pH may increase the capacity for SOC storage and nutrient
supply in specific biome types (Chen *et al.*, 2018*c*).

471 In summary, the responses of soil C-degrading EEAs to N addition are complex. The net 472 effect depends on how both enzyme production and turnover are affected by changes in both 473 soil environment and resource availability. Enzyme synthesis and secretion in natural 474 environments are influenced by a combination of abiotic factors such as soil pH and 475 temperature, as well as biotic processes including the composition of the soil microbial 476 community and root-microbe interactions. These abiotic and biotic factors, however, can 477 vary across different ecosystem types. Consequently, EEAs are likely to be responsive to N 478 addition across various ecosystem types, as well as to other climate changes (Chen et al., 479 2018b, 2020a). These changes will have important consequences for ecosystem functions 480 such as decomposition and nutrient cycling, which will ultimately affect SOC storage.

481

482 (3) Uncertainties and implications

483 Our findings provide novel insights into the relationships between N addition, soil 484 extracellular EEAs, microbial communities and SOC dynamics, however several 485 uncertainties warrant further investigation. First, atmospheric N deposition involves low-dose 486 and continuous N inputs into ecosystems, whereas many N addition experiments used only 487 single pulses. High-dose inputs of N may elicit stronger short-term effects on SOC storage 488 than N applied in several doses across the year (Cao et al., 2020). Therefore, it is necessary to 489 assess N addition at lower levels over extended timeframes. In addition, the N addition rate in 490 most experiments exceeded background N deposition rates, and our results may therefore 491 overestimate the responses of SOC storage to N addition (Fornara & Tilman, 2012). In our 492 analyses, we computed the ratio of the N addition rate to the background N deposition rate 493 (Rate:BND). This ratio serves as a measure of the additional N introduced into an

494 environment compared to the naturally deposited amount. Although this ratio did not 495 contribute significantly to explaining the responses of SOC storage or C-degrading EEAs to 496 N addition (Fig. 3) and the relationships were not significant (Fig. S3), it is important to note 497 that specific impacts can vary. These variations depend on ecosystem type and the form of N 498 being added. Therefore, this ratio should be used in conjunction with other measurements and 499 observations to comprehend fully the effects of N deposition. Second, the data set in this 500 study was predominantly from Asia (Fig. 1), but geological differences among sites were not 501 fully considered. Future studies should aim to include a wider range of geographically diverse 502 samples to investigate the effects of geological influences. This would provide a more 503 comprehensive understanding of the interplay between geology, soil microbes, and EEAs, 504 enhancing our ability to predict and manage soil health and function under changing 505 environmental conditions. Third, different responses of hydrolytic and oxidative C-degrading 506 EEAs to N addition revealed divergent microbial nutrient acquisition strategies that regulate 507 SOC storage, but it remains challenging to link the response of C-degrading EEAs to 508 microbial community structure and diversity. Nitrogen addition often induces changes in 509 microbial community structure that are consistent with changes in EEAs (Waldrop, Zak & 510 Sinsabaugh, 2004a). As C mineralization is a complex process that involves the enzyme-511 mediated catalysis of various organic fractions (Wu et al., 2022), the presumption of 512 inflexible microbial functional traits during N addition may lead to inaccurate predictions of 513 SOC storage. Although this meta-data synthesis establishes a framework for plant-soil-514 microbe interactions, future studies using new methods and technologies, such as advanced 515 genome sequencing and probe-based technologies, remain necessary to predict accurately the 516 overall consequences of N addition on SOC storage. In addition, given the strong enzymatic 517 control of SOC storage under N addition, explicitly representing soil enzymatic processes as 518 an important component of land surface models will improve our ability to generate a more

realistic simulation of the responses of SOC storage to environmental changes (Zhang *et al.*,
2022).

521

522 V. CONCLUSIONS

523 (1) Based on a comprehensive global meta-data synthesis, we offer novel insights into the

524 role of soil extracellular enzymes in mediating the response of SOC storage to N addition.

525 (2) Our analysis demonstrates that N addition significantly enhanced SOC storage, with this

526 enhancement strongly correlated with shifts in soil C-degrading EEAs. Specifically, we

527 observed an increase in hydrolytic C-degrading EEAs and a decrease in oxidative C-

528 degrading EEAs following N addition, suggesting that soil microorganisms modulate the

529 cycling of different C pools through distinct mechanisms.

530 (3) The net effect of N addition on SOC storage is determined by the balance between these

two types of C-degrading EEAs, with impacts varying across different ecosystem types.

532 (4) These results highlight the profound effects of N addition on microbially mediated SOC

533 storage in terrestrial ecosystems, emphasizing the importance of understanding these

534 processes to predict SOC storage better under future climate change scenarios.

535

536 VI. ACKNOWLEDGEMENTS

537 We sincerely thank all the scientists whose data and studies were included in this meta-data

538 synthesis. We also appreciate the insightful and constructive comments from the editor and

anonymous reviewers that contributed to improving the manuscript. This study was

540 financially supported by the National Natural Science Foundation of China (Grant number

541 42205167). J.Ch. is funded by the National Natural Science Foundation of China (32471685,

542 42361144886) and Shanxi Province Natural Science Foundation for Distinguished Young

- 543 Scholar (2024JC-JCQN-32). A.T.N. is supported by a UK Natural Environment Research
- 544 Council grant (NE/T012226).
- 545

546 VII. DATA ACCESSIBILITY STATEMENT

- 547 All data needed to evaluate the conclusions in the paper are present in the paper and/or the
- 548 supplementary materials and are also available at Figshare
- 549 (https://figshare.com/s/41394218cf59d54b6608). Requests for additional materials and
- 550 databases should be addressed to Ji Chen (chenji@ieecas.cn) or Xiao Chen
- 551 (chenxiao@mail.iap.ac.cn).

552

553 VIII. AUTHOR CONTRIBUTIONS

- J.Ch. and J.Ca. designed the research; X.C. performed the research, analysed data and wrote
- the paper; and R.L.S., D.L.M., R.D.B., N.F., A.T.N., J.Ch. and X.Z. contributed substantially
- 556 to revisions.
- 557

558 IX. REFERENCES

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- 560 Information.

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893 X. SUPPORTING INFORMATION

- Additional supporting information may be found online in the Supporting Information section
- at the end of the article.
- 896 Appendix S1. PRISMA flowchart for assessment of eligible studies.
- 897 Appendix S2. List of included studies.
- 898 **Table S1.** Results for publication bias using Begg's test and Egger's test.

Table S2. Evaluation of model parameters used to explain SOC storage under N addition.

Table S3. Evaluation of model parameters used to explain C-degrading extracellular enzyme
activities (EEAs) under N addition.

902 **Table S4.** Summary of mixed effect models of significant predictors in Model Selection

Analysis, as well as their interactions on the response of SOC storage and C-degrading EEAs

904 to nitrogen addition.

905 Fig. S1. Relationship between initial soil organic carbon (SOC) content and the response

906 ratios of SOC storage, hydrolytic C-degrading extracellular enzyme activities (EEAs),

907 oxidative C-degrading EEAs, and oxidative:hydrolytic C-degrading EEA ratio to N addition.

908 Fig. S2. Relationship between N addition rate and the response ratios of soil organic carbon

909 (SOC) storage, hydrolytic C-degrading extracellular enzyme activities (EEAs), oxidative C-

910 degrading EEAs, and oxidative:hydrolytic C-degrading EEA ratio to N addition.

911 Fig. S3. Relationship between N addition rate:background N deposition rate (Rate:BND) and

912 the response ratios of soil organic carbon (SOC) storage, hydrolytic C-degrading extracellular

913 enzyme activities (EEAs), oxidative C-degrading EEAs, and oxidative:hydrolytic C-

914 degrading EEA ratio to N addition.

915 Fig. S4. Relationships between the response ratios (RRs) of soil organic carbon (SOC)

916 storage and extracellular enzyme activities (EEAs) to N addition for studies categorized by

soil type.

918 Fig. S5. Effect of experimental duration on the relationship between the response of soil

919 carbon storage and C-degrading extracellular enzyme activities (EEAs) under N addition.

920 Figure S6 Relationship between the responses ratio of SOC storage to N addition and the

921 responses of hydrolytic C-degrading EEAs, oxidative C-degrading EEAs, and

922 oxidative:hydrolytic C-degrading EEA ratio to N addition in natural ecosystems (forest,

923 grassland and shrubland).

- 924 **Figure S7** Relationships between the response ratios of pH to N addition (RR-pH) and the
- 925 difference in soil pH between N addition and control treatments (ΔpH).
- 926 Figure S8 Meta-data synthesis of the average effects (%) of N addition on SOC storage,
- 927 hydrolytic C-degrading EEAs, oxidative C-degrading EEAs, and oxidative:hydrolytic C-
- 928 degrading EEA ratio for N addition properties.
- 929 **Figure S9** The distribution of relevant studies published over the past several years.

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Table 1. Overview of the enzymes included in our meta-data synthesis. EC is the enzyme

Туре	Extracellular enzyme	EC	Target
Hydrolytic	β-1,4-glucosidase	3.2.1.21	Cellulose degradation products
	β-1,4-xylosidase	3.2.1.37	Hemicellulose degradation products
	β-D-cellobiosidase	3.2.1.91	Cellulose degradation products
Oxidative	Peroxidase	1.11.1.7	Lignin and other complex compounds
	Phenol oxidase	1.10.3.2	Lignin and other complex compounds
	Polyphenol oxidase	1.14.18.1	Lignin and other complex compounds

932 commission number classification for that enzyme or class of enzymes.



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935 Fig. 1. Global distribution of the sites included in this meta-data synthesis. The numbers in

936 parentheses in the key are the number of observations for each ecosystem type.





Fig. 2. Results of meta-data synthesis of the average effects (%) of N addition on (A) plant productivity, (B) soil characteristics, (C) C-degrading extracellular enzyme activities (EEAs), and (D) microorganismal biomass. Error bars represent 95% confidence intervals. Asterisks indicate statistical significance: *P < 0.05; **P < 0.01; ***P < 0.001. The vertical solid grey line is the zero line, and the numbers on the right of each panel indicate the sample size. MBC, microbial biomass carbon; MBN, microbial biomass nitrogen; Soil N, soil total nitrogen; SOC, soil organic carbon.



946 Fig. 3. Model-averaged importance of the predictors of the effects of N addition on (A) soil 947 organic carbon (SOC) storage, (B) hydrolytic C-degrading extracellular enzyme activities 948 (EEAs), (C) oxidative C-degrading EEAs, and (D) oxidative:hydrolytic C-degrading EEA 949 ratio. The importance value is based on the sum of the Akaike weights derived from the model selection using corrected Akaike's information criteria. The cut-off was set at 0.8 to 950 951 explore the most important variables. In this analysis, duration (year) was treated as a numeric variable. BND: background N deposition rate $(g m^{-2} yr^{-1})$; MAP: mean annual 952 precipitation; MAT: mean annual temperature; RR, response ratio; TN, total nitrogen. 953 954 Ecosystem type: farmland, forest, grassland, and shrubland; Soil type: Alfisol, Aridisol, 955 Entisol, Gelisol, Histosol, Inceptisol, Mollisol, Oxisol, Spodosol, Ultisol; Rate: N addition rate (g m⁻² yr⁻¹); Form: NH₄⁺-N, NO₃⁻-N, NH₄NO₃, urea, and mixed fertilizer. 956 957



Fig. 4. Average effects (%) of N addition on (A) soil organic carbon (SOC) storage, (B) 959 960 hydrolytic C-degrading extracellular enzyme activities (EEAs), (C) oxidative C-degrading 961 EEAs, and (D) oxidative:hydrolytic C-degrading EEA ratio for various ecosystem types. The 962 error bars represent the 95% confidence intervals. Asterisks indicate statistical significance: * P < 0.05; ** P < 0.01; *** P < 0.001. The vertical solid grey line is the zero line, and the 963 964 numbers on the right of each panel indicate the sample size. (E) Relationships between different variables and the response ratio (RR) of SOC storage, hydrolytic C-degrading 965 966 EEAs, oxidative C-degrading EEAs, and oxidative:hydrolytic C-degrading EEA ratio. MAP: 967 mean annual precipitation; MAT: mean annual temperature; MBC, microbial biomass carbon; MBN, microbial biomass nitrogen; pH, background soil pH; SOC, background soil 968 969 organic carbon; TN, background soil total nitrogen. Clay, background clay content expressed 970 in %.



973 Fig. 5. Relationship between the response ratio (RR) of soil organic carbon (SOC) storage to 974 N addition and the responses of hydrolytic C-degrading extracellular extracellular enzyme activities (EEAs) (A, D), oxidative C-degrading EEAs (B, E), and the ratio of oxidative to 975 976 hydrolytic C-degrading EEAs (C, F) to N addition across various ecosystem types. The 977 regression lines indicate a significant relationship across all ecosystems (A–C) and within 978 each ecosystem type (D-F), as determined by a mixed-effects meta-regression model with a 979 95% confidence interval. The bubbles represent individual experiments included in the meta-980 data synthesis, with the size of the bubbles being proportional to the model weights.

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983 Fig. 6. Structural equation models investigating multivariate effects on soil organic carbon (SOC) storage under nitrogen addition ($\chi^2 = 7.39$, df = 4, P = 0.12; $R^2 = 0.36$). The orange and 984 985 blue rectangles represent the significant positive and negative effects of N addition, 986 respectively. The grey rectangles represent non-significant effects. Numbers in parentheses 987 indicate the average effects (%) of N addition. Arrows indicate the hypothesized direction of 988 causation, with red and blue arrows representing significantly positive and negative 989 relationships, respectively. Grey dashed arrows indicate non-significant relationships. 990 Numbers next to arrows are standardized path coefficients. BG, β -1,4-glucosidase; BX, β -991 1,4-xylosidase; CBH, β -D-cellobiosidase; EEA, extracellular enzyme activity; MBC, 992 microbial biomass carbon; PER, peroxidase; PO, phenol oxidase; PPO, polyphenol oxidase. Fungi:bacteria is fungal:bacterial biomass ratio. Asterisks indicate statistical significance: * P 993 994 < 0.05; ** *P* < 0.01; *** *P* < 0.001.

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