

1 **Title: PIV analysis of flow-type landslides under suboptimal image conditions**

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10

11 **Abstract**

12 The dynamics of flow-type landslides, including earthflows and debris flows, is still not fully
13 understood, primarily due to the complexity of the physical processes that govern the flow and the
14 challenges in acquiring direct field measurements. In modern monitoring stations, cameras represent
15 cost-effective data sources, providing essential information for characterising documented
16 reactivation events. *Particle Image Velocimetry (PIV)* algorithms have been extensively employed in
17 the literature to reconstruct velocity fields and rheological behaviour of laboratory physical models
18 under ideal conditions. However, the resolution of camera footage in the field typically falls short of
19 being optimal due to lighting and weather conditions, as well as non-zenithal recording geometry,
20 hindering a straightforward application of *PIV*. This study presents two primary sets of laboratory
21 flume tests conducted to explore a broad range of recording conditions, bridging the gap between
22 ideal laboratory settings and actual field acquisitions. The experiments enabled the evaluation of *PIV*
23 performance for each image quality scenario, detailing and quantifying the main uncertainties as well
24 as their impact on the resulting velocity fields while discussing underlying reasons and mitigation
25 measures. The experimental results reveal that, with due adjustments, suboptimal-quality footage can
26 be used to estimate the actual flow velocity field and infer the rheological behaviour of the flow.
27 Furthermore, distortions related to non-zenithal perspectives can be reliably minimised through
28 suitable orthorectification algorithms. These findings support the potential for broader application of
29 the tested *PIV*-based methodological approach in field scenarios to investigate the dynamics of flow-
30 type landslides.

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32 **Keywords:** Flume experiments, Particle Image Velocimetry (PIV), flow velocity field, flow-type
33 landslides

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36 ***Declaration of Competing Interest***

37 The authors declare no conflict of interest.

38

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43

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64 1. Introduction

65 The dynamics and mechanisms underlying the behaviour of flow-type landslides kinematics are still
66 subject to debate and not fully understood in the literature, encompassing both faster debris flows
67 (Kaitna et al. 2016; Nagl et al. 2020) and slower earthflows (Picarelli et al. 2005; Carrière et al. 2018;
68 Berti et al. 2022). This knowledge gap primarily stems from the complexities of the processes and
69 from the difficulties in acquiring direct field measurements (Hürlimann et al. 2003; McArdell et al.
70 2007), as well as in replicating these phenomena through representative laboratory-scaled physical
71 models (Iverson 2015; Turnbull et al. 2015). Modern monitoring stations frequently employ a variety
72 of sensors to investigate flow dynamics, including geophones (Mainsant 2012; Walter 2017; Coviello
73 et al. 2019), infrasound sensors (Leng et al. 2017), pore water pressure transducers and load cells
74 (Hürlimann et al. 2019), typically coupled with video cameras to document the events. Field cameras,
75 in particular, represent a cost-effective and highly valuable tool for data collection, providing essential
76 information to characterise captured flow-type landslides. When flow depth is known, surface
77 velocity observations enable the reconstruction of the transverse distribution of the depth-averaged
78 shear rate. This, in turn, allows at least a qualitative inference of the rheological properties of the
79 flow.

80 Image-based analysis techniques, such as *Particle Image Velocimetry (PIV)* and *Particle Tracking*
81 *Velocimetry (PTV)*, have been extensively reported in the literature to accurately reconstruct internal
82 and surface deformation and velocity fields, and to investigate the rheological behaviour of physical
83 models under ideal and controlled laboratory conditions (e.g. GDR MiDi 2004; Faug et al. 2015).
84 Over the past 30 years, *PIV* algorithms have also been frequently applied to larger scales, referred to
85 as *Large-Scale Particle Image Velocimetry (LSPIV)*, particularly in hydraulic applications to analyse
86 the surface velocity distribution of rivers and evaluate their discharge (e.g. Fujita et al. 1998; Le Coz
87 et al. 2010; Gunawan et al. 2012; Muste et al. 2014). More recently, *LSPIV* has also been employed
88 to estimate the surface velocity of debris flows (e.g. Theule et al. 2018). In particular, Schöffl et al.
89 (2023) combined *LSPV*-derived surface velocity measurements with contextual Pulse-Doppler high-
90 frequency radar observations, while Aaron et al. (2023) and Spielmann & Aaron (2024) proposed a
91 novel approach utilising high-frequency 3D *LiDAR* (Light Detection And Ranging) point clouds as
92 an alternative to conventional camera recordings for *LSPIV* analysis.

93 The determination of the surface velocity distribution through the implementation of these methods
94 with adequate spatial resolution would allow the assessment of flow boundary conditions. Integrating
95 information from qualitative observations on the recorded footage and independent morphological
96 assessments enables inference of the rheological behaviour of the landslide material. However, since
97 *PIV* outcomes are highly dependent on input footage quality (Prasad et al. 1992), its transposition to

98 larger-scale field scenarios and processes is not straightforward. Factors such as adverse weather
99 conditions or inadequate and inconsistent lighting can significantly impact the quality of the acquired
100 images. Physical barriers, including fog, dust or raindrops, may distort captured footage by scattering
101 light or causing uneven focus, resulting in a blurred effect. Furthermore, due to frequently limited
102 accessibility and difficult setup conditions in the field, cameras are often placed at a non-zenithal
103 perspective on the channel and non-optimal recording distances. This potentially introduces severe
104 distortions in the acquired footage, leading to the reconstruction of velocity values and distributions
105 that are not representative of the flowing landslide mass. Additionally, image quality may be further
106 impacted by motion blur, which occurs when the frame rate of the installed cameras, especially in the
107 context of low-cost monitoring stations, is not sufficient to fully capture the movement of the flow.
108 In this study, two sets of flume experiments were conducted to test the applicability of *PIV* algorithms
109 under different input image quality conditions, from ideal laboratory acquisitions to suboptimal
110 recordings that mimic typical field footage. The main objective is to assess *PIV*'s ability to produce
111 reliable flow field data using blurred images and non-zenithal recording angles and to identify
112 uncertainties arising from these suboptimal conditions.

113 **Section 2** outlines the experimental setup and the analysis workflow adopted. **Section 3** presents the
114 main experimental results, while **Section 4** explores the main sources of uncertainties in the
115 techniques employed, considering their application in the field.

116

117 **2. Methods**

118 **2.1 Experimental setup**

119 **Figure 1a** schematically illustrates the experimental setup employed in the two sets of landslide flume
120 experiments, referred to as *F1* and *F2*, conducted at the Department of Civil and Structural
121 Engineering at the University of Sheffield (**Figure 1b**) and the British Geological Survey site in
122 Keyworth (**Figure 1c**), respectively. Both series of tests focused on observing dry granular flows from
123 different perspectives as they were gradually released from the flume hopper. The selection of dry
124 granular materials was driven by the need to design straightforward experimental setups, prioritising
125 data acquisition and processing over the precise replication of in situ phenomena in terms of material
126 properties and scaling. Despite the similarity of the processes observed, the recording conditions in
127 the two sets of experiments, summarised in **Table 1**, were systematically different, as detailed in the
128 following sections.

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130

131

132 2.1.1 FI tests

133 The *FI* tests featured recordings from two optimal perspectives utilising a high-performance high-
134 speed camera, sharply capturing dry granular flows of approximately 3.85 mm Denstone® ceramic
135 beads (Zhao et al. 2023) at 1000 frames per second (fps). The first recording geometry (R_1) positioned
136 the camera above, facing the chute surface perpendicularly to capture the flow from an ideal *zenithal*
137 perspective. The second viewing angle involved positioning the camera perpendicularly to the side
138 of the flume, recording the side of the flow through the smooth Perspex wall and aligning with the
139 fixed chute's slope (θ) at 30° from the horizontal. This ideal perspective, while not accessible in the
140 field, is commonly utilised in flume experiments to investigate internal flow dynamics (e.g.
141 Wiederseiner et al. 2011; Li et al. 2022).

142 These acquisition geometries were established to frame fixed regions on both the surface and side of
143 the chute, as indicated by the green and blue squares in Figure 1a, respectively. This setup enabled
144 consistent recording and characterisation of surface and internal velocity distributions within the same
145 portion of the flow. Cameras were positioned to focus on regions around the flume's midpoint, aiming
146 to capture as uniform flow as possible and minimise potential effects from the material release at the
147 upper slope or the transition to free-fall motion near the chute's end.

148 To facilitate the generation of a regular flow, potentially reaching a steady-state condition over time,
149 upon which to focus the analyses, the granular material was progressively released from the flume
150 hopper through a sluice gate, avoiding *dam-break* release mechanisms, typically employed in debris
151 flow physical models (e.g. Iverson et al. 2010; Eu et al. 2017). The initial thickness of the flow (h_0)
152 and the material's release rate were controlled by the gate opening (H) and the material's head in the
153 hopper, which gradually decreased during the experiment. Care was taken to maximise the initial
154 head while maintaining an even distribution of the material within the hopper, ensuring a uniform
155 release rate throughout a significant portion of the test.

156 With a single high-speed camera available, each flow was repeatedly captured from different
157 recording geometries, while maintaining consistent boundary and initial conditions. Subsequently,
158 the boundary conditions at the chute's base were systematically altered to increase surface roughness,
159 documenting eventual variations in the flow dynamics. The transparent Perspex sidewalls were
160 consistently kept unchanged to ensure continuous observation. As detailed in Table 1, three main
161 tests have been conducted: one employing the standard smooth aluminium base, another with an
162 intermediate roughness base composed of glued ceramic beads of about 1 mm diameter, and a third
163 using a coarse sand base with particle sizes ranging from 1 to 2 mm.

164 Unlike similar experiments described in the literature, which often include a seeding phase, where
165 coloured, fluorescent or reflective tracer particles are added to the material to enhance flow patterns

166 (Savage 1979; Parsons et al. 2001; Lindken 2009), the tests performed omitted this procedure. The
167 texture of the granular materials employed was deemed adequate to enable the reconstruction of flow
168 dynamics, given proper lighting. This was achieved using 50W DC LED floodlights (Figure 1b).

169

170 2.1.2 F2 tests

171 In contrast to the *F1* experiments, which involved the application of the *PIV* algorithm to sharp
172 footage captured from optimal recording angles, the *F2* experiments explored different acquisition
173 geometries and image quality scenarios. In addition to the ideal perspectives (R_1 and R_3) used in the
174 *F1* experiment set, the *F2* tests implemented an additional recording geometry (R_2). This other
175 observation angle involved positioning the camera slightly to the side of the flume, facing downslope
176 with an inclination of 20° relative to the flume slope (θ), capturing the surface of the flow from a
177 more oblique, non-zenithal perspective (Figures 1c and 3). This suboptimal camera orientation
178 mirrors the recording geometry typically employed in the field, where achieving a perfectly zenithal
179 perspective can be challenging (Patalano et al. 2017). In the field, cameras are commonly placed to
180 the side of the channel, slightly inclined downward to minimise interference from rain droplets and
181 at a generally low angle relative to the horizontal.

182 The *F2* experiments featured simultaneous recordings of the same region of the flume from the three
183 described perspectives of dry granular flows composed of fine sand ($d_{mean} \approx 0.28$ mm), utilising
184 regular cameras with maximum frame rates of 240 fps. Multiple tests were conducted at progressively
185 higher flume slope angles (θ) while maintaining the other boundary conditions constant. This
186 approach resulted in increasingly faster flows and, given the fixed camera frame rate, gradually
187 blurrier and lower-quality footage, enabling the consideration of a wide spectrum of image quality
188 conditions.

189 Despite these methodological differences, the *F2* tests were conducted with the same care and
190 precautions as detailed for the *F1* experiments, specifically regarding material preparation and
191 release, as well as camera positioning around the mid-portion of the flume length, ensuring the
192 recording of flows as regular and uniform as possible.

193 As with the *F1* experiments, the *F2* tests were performed under the assumption that no seeding phase
194 was needed, contributing to the generation of challenging conditions for the application of the *PIV*
195 algorithm to suboptimal quality footage (see Section 3.2). Nevertheless, this setup is representative
196 of field conditions, where seeding is not possible. Similar to *F1* tests, lighting was provided using
197 multiple 50W DC LED floodlights (Figure 1c).

198

199

200 **2.2 Analysis workflow**

201 Despite the slight differences in experimental procedures between the two described sets of tests, the
202 acquired footage was processed following the same analysis workflow schematised in [Figure 2](#).

203

204 2.2.1 The PIV algorithm

205 *PIV* algorithms are non-intrusive optical measurement techniques used to determine the velocity field
206 of recorded fluid flows. This is achieved by identifying the displacement of similar groups of pixels
207 (or particles) in successive images through cross-correlation methods ([Adrian 1991](#)). The captured
208 frames are initially subdivided into a grid of smaller sub-regions, framed by a moving interrogation
209 window, or *patch*, whose size and shift define the final measurement resolution. For each consecutive
210 image pair, the cross-correlation matrix is evaluated within the moving interrogation patch. The
211 location of resulting cross-correlation peaks provides the most probable displacement vectors,
212 connecting the similar groups of pixels identified between the two consecutive frames along a straight
213 line ([Raffel et al. 2007](#); [Thielicke and Stamhuis 2014](#)). This information, combined with the timestep
214 between images, ultimately yields the frame-by-frame flow velocity field. Unlike Lagrangian *PTV*
215 techniques, which reconstruct the trajectory of individual tracer particles ([Kreizer et al. 2010](#)), *PIV*
216 algorithms are Eulerian methods that enable the reconstruction of the flow vectors regardless of single
217 particle characteristics ([Patalano et al. 2017](#)).

218 For both the *F1* and *F2* experiments, a *PIV* analysis was conducted on the frames extracted from the
219 footage recorded from the different perspectives based on the camera's frame rate, employing the
220 open-source algorithm *PIVlab* ([Thielicke and Stamhuis 2014](#)). This particular algorithm utilises a
221 multi-pass approach, where each flow vector results from successive computational steps during
222 which the original moving interrogation patch is progressively deformed and refined. The first step
223 involves computing the frame-by-frame displacement vectors between similar groups of pixels by
224 evaluating two-dimensional cross-correlation peaks within a moving interrogation patch ([Thielicke
225 and Sonntag 2021](#)), following the described basic *PIV* principles. In the subsequent step, the
226 displacement information obtained within the first interrogation patch is interpolated for every
227 enclosed pixel and used to deform the second-pass patch. Within this finer patch, the cross-correlation
228 matrix peaks are evaluated by repeating the procedure followed in the previous pass, thereby refining
229 the input displacement information. Through this iterative process, the algorithm yields progressively
230 more accurate and refined displacement vectors, enhancing correlation robustness and decreasing
231 random error contributions ([Thielicke 2014](#)).

232

233

234 2.2.2 Optimal images

235 In the flume experiments performed in this study, a two-pass *PIV* approach was adopted in the *PIVlab*
236 environment defining interrogation patch sequences of $64 \times 64 - 32 \times 32$ pixels and $24 \times 24 - 12 \times 12$
237 pixels, for *F1* and *F2* tests, respectively. The interrogation patch size, especially for *F1* experiments
238 where larger particles were employed, was defined to enclose around 5 – 15 particles on average to
239 improve correlation significance (Adrian and Westerweel 2011; Gollin et al. 2017). Additionally, the
240 interrogation patch shift in the moving window algorithm was consistently set to 50% of its size.
241 Consequently, as depicted in Figure 2a, for each pair of extracted frames, the two-dimensional flow
242 displacement and corresponding velocity fields were derived, with a flow vector density determined
243 by the size and shift of the moving patch. The obtained velocity fields, initially expressed in
244 pixels/frame, were converted to real units (m/s) by defining the time step between frames and the
245 pixel size. The latter was determined by calibrating the images against known distances: the chute
246 width for the surface velocity field and a reference distance measured on a ruler taped to the flume
247 sidewall within the camera's field of view for R_3 recordings.

248 After retrieving the frame-by-frame velocity field, the analysis focused on a specific subset where the
249 flow exhibited steady-state behaviour. This approach facilitated more meaningful comparisons
250 between recordings acquired from different geometries, whether obtained in separate tests (*F1*) or
251 simultaneously (*F2*). The extraction of the steady-state frames subset was performed as follows. First,
252 two section traces were defined perpendicular to the flow direction, bounding the edges of an ideal
253 plane transversally crossing both the surface and the side of the flume at the same distance from the
254 hopper, applicable to the R_1 and R_3 recording geometries, respectively. Along these traces, for R_1
255 acquisitions, a short section, indicated as *Section (I)* and represented by a red solid line in Figure 2a,
256 was traced around the mid portion of the chute to capture the central part of the flow surface.
257 Similarly, another short section, limited to the upper portion of the flow, was considered for R_3
258 footage. Subsequently, for each recording geometry, the frame-by-frame mean velocity value along
259 the defined sections was computed and observed over time, as illustrated in the chart in Figure 2b.
260 The subset of frames exhibiting an overall constant mean velocity, indicated by a null slope of the
261 regression line and suggesting steady-state flow conditions, was isolated and extracted, while the
262 preceding frames were discarded. The steady-state surface and side velocity profiles were then
263 extracted for the selected frame subset along the same sections, but this time considering the entire
264 extent of the flow, as indicated by the red dashed line in Figure 2a for R_1 recordings, labelled as
265 *Section (II)*. The resulting velocity profiles, along with their mean distribution resampled at regular
266 intervals, and the corresponding standard deviation values were ultimately plotted as shown in Figure
267 2c to summarise the steady-state surface and internal velocity distributions of the recorded flow. For

268 validation, the reconstructed velocity values were compared with manually derived measurements
269 obtained by visually tracking frame-by-frame individual particle trajectories at multiple flow depths
270 on the same steady-state subset of R_3 recordings. These independent velocity measurements were
271 taken using a ruler affixed to the side of the flume and visible within the camera's field of view.

272

273 2.2.3 Suboptimal images

274 The methodological approach detailed in the previous section for analysing sharp footage from ideal
275 perspectives in both experiment sets was also applied to suboptimal images, with minor
276 modifications. Specifically, the analysis workflow remained unchanged for the blurred footage
277 captured in the $F2$ experiments from the ideal perspectives, whereas it was properly adjusted for R_2
278 acquisitions regardless of image quality. While pixel size is uniform across the image for R_1 and R_3
279 acquisitions, this is not the case for R_2 . In such instances, the non-zenithal recording angle introduces
280 a significant disproportion between pixel sizes in the foreground and background. Background pixels
281 are generally more poorly resolved, with differences and gradients depending on the geometric
282 relationship between the captured portion of the channel and the camera, both in terms of recording
283 angle and distance (Jolley et al. 2021).

284 Consequently, converting and transposing the pixel reference system of the images to real coordinates
285 and units is not as straightforward as with R_1 and R_3 recordings. Applying the same methodology
286 could result in significant deviations of the reconstructed flow velocities from the actual values.
287 Therefore, to assess the contribution and remove such optical distortions, the obtained R_2 surface
288 velocity field (Figure 3a) was appropriately orthorectified using the open-source algorithm *RIVeR*
289 (Patalano et al. 2017). This software enables orthorectification operations directly on *PIV*-derived
290 flow velocity fields within a user-defined two-dimensional region of interest. This is accomplished
291 by computing the homography matrix (Corke 2011), which uniquely maps real-world coordinates to
292 their planar projection on the image within the selected area. This matrix is solved using the *Camera*
293 *Calibration Toolbox* (Vision Caltech 2009) for MATLAB by specifying the real-world and image
294 coordinates of a series of known control points (Figure 3b). While four control points are sufficient
295 to perform 2D orthorectification, solving the homography matrix in three dimensions requires the
296 spatial coordinates of at least six points. Given the controlled conditions of the present experiments,
297 which focused on steady-state flows over flat and regularly shaped chutes, the simpler 2D
298 orthorectification approach using four control points was deemed adequate. However, as noted in
299 Section 4.3, more complex field scenarios may necessitate 3D orthorectification to account for
300 irregular geometries and unsteady flow conditions.

301 The calculated homography matrix is then utilised to extend the image-to-real-world coordinate
302 system relationship to the entire region of interest, accurately determining the actual pixel size (Figure
303 3c) and reprojecting the flow velocity vectors as if derived from a zenithal observation (Figure 3d).
304 The resulting R_2 orthorectified velocity profiles were ultimately compared and plotted against the
305 corresponding uncorrected ones and those derived from the zenithal perspective (R_I), which ideally
306 represent the actual surface velocity distribution of the flow.

307 Despite the existence of advanced techniques for mitigating motion blur and restoring image
308 sharpness, including deconvolution algorithms (Lucy 1974; Hosseini & Plataniotis 2020; Satish et al.
309 2020) and more sophisticated machine learning approaches (Zhang et al. 2017; Lian & Wang 2023;
310 Chen 2024), the described methodology was applied directly to the unprocessed camera footage,
311 prioritising operational simplicity and straightforward implementation for future field applications.

312

313 **3. Results**

314 This section presents the results obtained from the $F1$ (Section 3.1) and $F2$ (Section 3.2) laboratory
315 flume tests, conducted employing the experimental setup and processing procedures detailed in
316 Section 2.

317

318 **3.1 $F1$ tests**

319 Figure 4 summarises the results of the three main flume experiments in the $F1$ set, comparing the
320 mean velocity profiles and corresponding standard deviation bands of the steady-state flows recorded
321 from R_I (Figure 4a) and R_3 (Figure 4c) perspectives, under the three varying boundary conditions at
322 the base of the flow indicated in Table 1. The velocity profiles reconstructed for the flows with the
323 smooth aluminium, intermediate roughness (1 mm ceramic beads), and coarse angular sand beds are
324 displayed in blue, green, and red, respectively. The blue curves exhibit the highest velocities, with
325 values of approximately 1.7 m/s around the central region and 1.5 m/s towards the shallower portions
326 of the flow in the surface and side velocity distributions, respectively. Conversely, the green and
327 particularly the red curves consistently display lower velocities, with mean values of 1.6 m/s and 1.3
328 m/s, and 1.2 m/s and 0.7 m/s, respectively, evaluated in the same regions of the flow. These curves
329 not only highlight a significant overall velocity reduction with increasing chute surface roughness but
330 also display a consistent marked continuity between surface and side velocity distributions, revealing
331 comparable values at the intersections between the analysed section pairs.

332 Examining the side velocity profiles (Figure 4c), the employed sluice gate opening (H) of 85 mm,
333 corresponding to a flow thickness normalised by particle size ratio (h/d) of approximately 22, resulted
334 in varying flow thicknesses across the different bed types. The angular sand bed produced the thickest

335 flows, measuring around 38 mm, while values of 27 and 24 mm were documented for the smooth and
336 intermediate roughness bases, respectively. In terms of normalised h/d ratios, these values correspond
337 to approximately 10, 7 and 6.

338 Additionally, the side velocity profiles reveal progressively different flow conditions as bed
339 roughness increases and overall flow velocity decreases. The faster flows over smooth and
340 intermediate roughness beds share similar velocity distributions, with progressively lower values with
341 depth that remain visibly above zero at the bottom, indicating a *slip-flow* rheological behaviour (Nagl
342 et al. 2020). In contrast, the high-roughness bed produces distinct flow conditions, characterised by a
343 steeply decreasing velocity profile, reaching zero at the base. This particular rheological behaviour,
344 alternatively referred to as *heap flow* (GDR MiDi 2004; Jop et al. 2005, 2006), where the flow is
345 concentrated in the upper portions, marks the transition from inclined plane flow to flow on an
346 erodible bed. Within this flow, the grains effectively move over a lower, thinner *quasi-static* layer
347 (Jop 2015) of interlocked particles due to the base roughness.

348 The distinct rheological behaviours observed are also quantitatively highlighted by the velocity
349 coefficient α , computed as the ratio between the depth-averaged velocity along the vertical
350 investigated in the proximity of the transparent sidewall and the corresponding surface velocity.
351 Flows over smooth and intermediate roughness bases, exhibiting similar rheological behaviour,
352 reveal comparable α values of 0.73 and 0.68, respectively, aligning with the 0.6 – 0.8 range reported
353 in the literature for flows in dense granular media (Cui et al. 2018; Nagl et al. 2020; Aaron et al. 2023;
354 Spielmann & Aaron 2024). In contrast, flow conditions over the rough sand base yield a remarkably
355 lower α value of 0.34.

356 Although this rheological behaviour transition is remarkable in the side profiles, the corresponding
357 surface ones (Figure 4a), net of the overall varying velocities, do not reflect any indication of the
358 documented changes in internal flow conditions. This is further highlighted in the normalised profiles
359 displayed in Figure 4b, with velocity distributions consistently collapsing onto a common trend,
360 making it virtually impossible to infer the actual rheological behaviour and internal flow structure
361 from surface observations alone. This complexity, common in flume experiments, arises from
362 systematically different boundary conditions at the bottom and the side of the flow, framed by a
363 progressively rougher base and smooth Perspex walls, resulting in undifferentiated slip-flow
364 conditions along the sides.

365 Overall, *FI* tests indicate the effectiveness of the *PIV* technique in characterising flows under optimal
366 conditions. This is evidenced by the consistent agreement between *PIV*-derived velocity and
367 independent manual measurements performed at four distinct flow depths, which systematically fall
368 within the standard deviation bands of the mean steady-state profiles displayed in Figure 4c. The low

369 standard deviation values relative to the mean enable the proper identification of the flow velocity
370 profiles. Notably, the quality of the reconstructed side velocity distributions allows for the observation
371 of varying flow regimes, which align with the expected behaviour of the flow under the defined
372 boundary conditions.

373

374 **3.2 F2 tests**

375 **Figure 5** summarises the results from *F2* experiments testing and assessing the applicability of the
376 described *PIV* methodology under varying input image quality conditions. Among the various
377 experiments performed, two reference cases are presented, depicting ideal and sub-optimal
378 acquisition conditions within the spectrum of recording scenarios explored. These configurations are
379 represented by the blue and red curves in the figure, corresponding to slower flows with a slope angle
380 (θ) of 28° and faster ones with $\theta = 30^\circ$, respectively. As illustrated by the frames in **Figure 5a**, for a
381 fixed camera frame rate, increasingly faster flows due to higher slope angles (θ) translate into
382 progressively poorer image quality, transitioning from ideal, sharp images to blurrier frames. Similar
383 to **Figure 4** for the *F1* experiments, the curves displayed represent the mean velocity distributions
384 along with the corresponding standard deviation values of the steady-state flows recorded from the
385 R_1 (**Figure 5a**) and R_3 (**Figure 5b**) camera geometries. Apart from a consistently greater dispersion
386 around the mean values for the blurrier footage, the curves reveal that the applied methodology can
387 coherently reproduce the flow, estimating velocity distributions in both scenarios. This is further
388 corroborated by the positive correspondence with independent, visually determined velocities at three
389 different flow depths, which consistently align with and fall within the standard deviation values of
390 the R_3 profiles shown in **Figure 5b**. Despite the different flow velocities, the two sets of curves exhibit
391 similar distributions with *slip-flow* behaviour at the interfaces with the smooth aluminium base and
392 Perspex walls, consistent with the rheological behaviour expected for a dry granular flow governed
393 by the described boundary conditions. The observed similarity in flow conditions is further supported
394 by the numerical values of the velocity coefficient (α), calculated consistently with the *F1* experiment
395 set. The tests conducted at $\theta = 28^\circ$ and $\theta = 30^\circ$ yield analogous α values of 0.83 and 0.81, respectively,
396 both comparable with the previously reported literature range of 0.6 – 0.8 for such phenomena.

397 The side velocity profiles in **Figure 5b** reveal slight differences in flow thickness. With an initial gate
398 opening (H) of 40 mm ($h/d \approx 143$), thicker flows of about 10.5 mm were observed at $\theta = 30^\circ$, while
399 slightly lower values around 9.5 mm were observed at $\theta = 28^\circ$. These thicknesses correspond to
400 normalised h/d ratios of approximately 34 and 38, respectively.

401 Additionally, as highlighted in the *F1* tests, the velocity profiles in **Figure 5** display a systematic
402 continuity between the surface and side velocity fields, with values of approximately 0.1 m/s and 0.3

403 m/s documented for $\theta = 28^\circ$ and $\theta = 30^\circ$ flows, respectively, near the intersection of the analysis
404 sections.

405 **Figure 6** condenses the results of the *F2* tests focused on evaluating the effects of a *non-zenithal*
406 recording geometry (R_2) on reconstructing the surface velocity distribution of the flow compared to
407 its actual velocity field, ideally determined from a zenithal perspective (R_1). The R_1 surface velocity
408 profiles reconstructed for the flows observed at $\theta = 28^\circ$ and $\theta = 30^\circ$, described in **Figure 5a**, are
409 reported as green curves in **Figures 6a** and **6b**, respectively. In both figures, the blue profiles represent
410 the surface velocity distributions obtained for the same flows from recording geometry R_2 , where
411 pixel size was considered homogeneous across the entire image and defined uniquely based on a
412 known distance coinciding with the chute width. These curves, for which input image distortion was
413 not corrected, significantly and systematically underestimate the real flow velocity distribution, with
414 an average offset of approximately 55% for $\theta = 28^\circ$ and 65% for $\theta = 30^\circ$. Conversely, a consistent
415 alignment with the real surface velocity distribution is documented in both scenarios for the
416 orthorectified R_2 profiles, depicted in red in **Figures 6a** and **6b**, exhibiting comparable offsets of about
417 2% on average. This highlights that sub-optimal quality footage from non-zenithal recording
418 geometries can be effectively employed for estimating the surface velocity field of a flow when
419 appropriate orthorectification algorithms are applied.

420

421 **4. Discussion**

422 **4.1 Effects of suboptimal image quality**

423 The results from the *F2* flume tests revealed that the methodological approach employed enables a
424 reasonable estimate of flow velocity fields even under suboptimal image quality conditions (**Section**
425 **3.2**). However, to consistently assess and quantify the effects of image blur on the *PIV*-derived
426 velocity distributions, the *F2* experimental footage captured at $\theta = 28^\circ$ from the zenithal perspective
427 (R_1) was considered. The quality of the original sharp footage was progressively artificially decreased
428 by applying a *Gaussian smoothing* algorithm. *Gaussian smoothing* is a two-dimensional convolution
429 operation that utilises a moving kernel represented by a Gaussian function to transform the enclosed
430 original image pixels, reducing their noise and detail, and resulting in blurred images (**Marr and**
431 **Hildreth 1980; Shapiro and Stockman 2001**). The blur level is controlled by the standard deviation
432 (σ) of the Gaussian distribution employed. A wide spectrum of image blur scenarios was considered,
433 using values of σ ranging from 1 to 5. For each σ , the mean surface velocity profile was reconstructed
434 along the same section and compared to the original.

435 The mean velocity profiles obtained reveal a consistent increase in uncertainty with σ , with standard
436 deviation values 20% higher than the original for $\sigma = 1$, and up to 2 or 5 times the original for $\sigma = 3$

437 and $\sigma = 5$, respectively. Additionally, as displayed in [Figure 7a](#), the shape of the original velocity
438 distribution reported in green is gradually less recognisable transitioning to higher blur levels. Besides
439 the greater uncertainty, these velocity profiles consistently and sensibly underestimate the original
440 values, up to an average difference of approximately 45 % for $\sigma = 5$. This bias is attributable to the
441 poorer quality of the input images, which results in broader and less distinct peaks in the cross-
442 correlation matrix evaluated in the *PIV* algorithm, leading to a decrease in the signal-to-noise ratio,
443 and, consequently, reduced correlation accuracy ([Elsinga et al. 2005](#)). This causes the generation of
444 spurious velocity vectors that do not accurately represent the recorded flow. These vectors exhibit
445 orientations that can significantly deviate from the actual flow direction, as displayed qualitatively in
446 the plan view ([Figures 7b, c](#)) and quantitatively along a reference section ([Figures 7d, e](#)) for the same
447 frames pair under optimal and blurred ($\sigma = 5$) image quality conditions. While in the former case, the
448 flow vectors are uniformly oriented at about 90° relative to the analysis section, they become more
449 chaotic for $\sigma = 5$, revealing diverging orientations characterised by peak differences of approximately
450 $\pm 60^\circ$, indicating fictitious transversal velocity components. These transversal components may lead
451 to a substantial reduction in the velocity magnitude values extracted along the reference section,
452 thereby motivating the consistent velocity underestimation observed.

453 As synthetically shown in [Figure 7](#) and further detailed in [Figures 8](#) and [9](#) in the supplementary
454 materials, the effect of image blur may be partially mitigated by defining a sequence of larger
455 interrogation patches in the *PIV* analysis. For instance, [Figure 7](#) highlights how the adoption of an
456 interrogation window sequence of 96 – 48 pixels, four times larger than the 24 – 12 pixels patches
457 used in the initial analysis, produces a clear reduction in uncertainty and underestimation of the flow
458 velocity as σ increases, albeit with a reduced spatial resolution. In this case, for $\sigma = 5$, the deviation
459 of the velocity distribution from the original values, indicated by the red and green dashed curves in
460 [Figure 7a](#), respectively, decreases from approximately 45% to 20%. Additionally, the distribution of
461 the velocity vector orientations along the analysis section, depicted by the black dashed curves in
462 [Figures 7d](#) and [7e](#), more closely aligns with the original, with deviations consistently below $\pm 10^\circ$
463 from the true flow direction.

464 While the usage of larger interrogation windows visibly enhances *PIV* analysis results for suboptimal-
465 quality footage, with the improvement proportional to the level of image blur, the effect is less
466 pronounced for sharp images. Maximum differences of approximately 2% were observed between
467 the velocity profiles reconstructed from sharp frames employing interrogation patch sequences of 24
468 – 12 pixels and 96 – 48 pixels, depicted by the two green curves in [Figure 7a](#).

469

470 The fact that the accurate reconstruction of flow dynamics is not hampered when individual particles
471 within the flow are not clearly recognisable, provided the discussed adjustments are applied, supports
472 the reasonable extension of the adopted methodology to the analysis of more complex multiphase
473 flows in both laboratory and field settings.

474

475 **4.2 Effects of inadequate frame rate**

476 Another crucial parameter in field acquisitions, particularly in the context of low-cost monitoring
477 systems, is the recording camera's frame rate. To investigate the impact of inadequate sampling
478 frequency on flow velocity field reconstruction, the *F2* experimental footage captured at $\theta = 28^\circ$ from
479 the zenithal perspective (R_I) was again used as a reference. The original 240 fps footage was
480 progressively undersampled down to 30 fps, with *PIV* analysis performed at each intermediate step
481 to compare the reconstructed surface velocity distributions.

482 The resulting mean velocity profiles reveal a marked increase in uncertainty towards lower frame
483 rates. Standard deviation values were 20% higher than the original at 120 fps, and up to 2 and 4 times
484 higher at 60 fps and 30 fps, respectively. Similar to the motion blur effect described in the previous
485 section, [Figure 10](#) illustrates how the shape of the original velocity distribution (reported in green)
486 becomes less recognisable towards lower frame rates, consistently underestimating velocity values.
487 This discrepancy stems from the reduced temporal resolution in the undersampled footage, where
488 actual particle displacements exceed the size and shift of the *PIV* moving interrogation window. This
489 aliasing effect leads to a systematic loss of correlation between frame pairs ([Raffel et al. 2007](#)),
490 generating biased flow velocity vectors that inaccurately reflect the observed flow. These spurious
491 vectors display orientations that diverge remarkably from the true flow direction, as shown
492 qualitatively from a zenithal perspective in [Figures 10b](#) and [10c](#), and numerically along a reference
493 section in [Figures 10d](#) and [10e](#), for the same image pairs captured at the original and reduced (30 fps)
494 frame rates. While properly sampled footage, produces a homogeneously oriented velocity field at
495 about 90° relative to the analysis section, undersampled recordings exhibit a more chaotic distribution
496 of vectors, with some even directed upstream and orientation deviations of up to $\pm 120^\circ$ from the
497 actual direction.

498 [Figure 10](#) also briefly illustrates that undersampling effects can be effectively mitigated by adopting
499 larger interrogation patch sequences during *PIV* analysis. Specifically, employing a window sequence
500 of 96 – 48 pixels, four times larger than the original, leads to a sharp reduction in uncertainty and
501 underestimation of flow velocity at lower frame rates, albeit at the cost of a coarser spatial resolution.
502 For example, at 30 fps, the velocity deviation from the original values drops approximately from 80%
503 to 1%. Similarly, the velocity vector orientations along the analysis section, highlighted by the black

504 dashed curves in [Figures 10d](#) and [10e](#), closely align with the original, with deviations consistently
505 under $\pm 2^\circ$ from the true flow direction.

506 The effects of using larger interrogation windows are further illustrated in [Figures 11](#) and [12](#) in the
507 supplementary materials, concerning patch sequences of 48 – 24 pixels and 94 – 48 pixels,
508 respectively.

509 **4.3 Orthorectification uncertainties**

510 The methodological approach adopted in the *F2* experiments highlighted that the use of appropriate
511 orthorectification algorithms on the *PIV*-derived surface velocity field can mitigate distortions
512 introduced by non-zenithal recording geometries (R_2). The two-dimensional orthorectification
513 algorithm employed requires specifying the coordinates of four known control points, framing an
514 ideal plane where the transposition between the pixel reference system and the real-world coordinates
515 is performed based on the actual size of the pixels enclosed. While defining these control points as
516 well as their respective distances in a controlled, small-scale laboratory environment is
517 straightforward, in the field, these operations may present a more significant challenge depending on
518 the accessibility of the landslide channel. This could lead to inaccuracies in determining these
519 distances and, consequently, in estimating the actual flow velocity field.

520 To assess and quantify the potential effects of measurement errors on the final surface velocity
521 distribution, the orthorectification procedure was repeated on the same footage captured at $\theta = 28^\circ$
522 multiple times, artificially modifying the control points' distance values. The real distances were
523 progressively altered considering errors of $\pm 5\%$, 10% and 15% . The resulting velocity profiles are
524 compared in [Figure 13](#). Overestimations of the actual distances result in profiles that increasingly
525 overestimate both the channel width and flow velocity, whereas underestimations yield the opposite
526 effect. These deviations are represented in the figure by curves grading from green (actual values) to
527 red (overestimation) or blue (underestimation). The displayed distributions reveal that, in our
528 experiments, the input percentage error in the distances approximately corresponds to the deviation
529 observed with respect to the real velocity profile, both positively and negatively.

530 In addition to evaluating uniformly distributed errors across the region of interest, the potential effect
531 of measurement errors affecting a single control point was also considered. Artificially altering the
532 inter-distances associated with one control point, simulating its misidentification in the field,
533 produced a pronounced distortion of the orthorectified velocity field towards the modified point. This
534 resulted in surface velocity profiles that not only exhibited different flow widths but also displayed
535 markedly asymmetric trends, deviating from the original distribution with an intensity proportional
536 to the error magnitude. The abnormal shape of these velocity distributions, characterised by values
537 progressively increasing towards one side of the channel, could serve as an indicator in field

538 applications, suggesting the presence of errors concentrated on the identification of a single control
539 point.

540 Given the potential significance of such errors, it is crucial in the field to measure the channel
541 geometry within the field of view of the monitoring camera as accurately as possible.

542 In natural environments, challenges in orthorectification extend beyond the discussed complex
543 channel geometries, which may hinder the uniform distribution of control points and compromise
544 their precise positioning. Further complexities arise from the transient nature of the observed
545 phenomena. While the flume experiments described in this study focused on steady-state flows with
546 constant thickness and velocity, such stable conditions are rarely encountered in real-world scenarios.
547 Under fluctuating flow levels, the recording distance between the camera and the flowing mass
548 becomes variable, and pronounced changes can significantly alter the area encompassed within the
549 region of interest identified for *PIV* analysis and orthorectification procedures. These fluctuations
550 complicate the accurate mapping of real-world coordinates into the 2D image plane where the flow
551 velocity field is reconstructed (Li et al. 2019). For instance, flow level underestimation leads to
552 exaggerated horizontal distances and, in turn, velocity overestimation, while flow level
553 overestimation produces opposite effects (Dramais et al. 2011). Consequently, flow level variability
554 represents a primary source of uncertainty in *LSPIV* analyses (Le Boursicaud et al. 2016).

555 Recent literature shows that *LSPIV* has predominantly been applied in riverine environments (Muste
556 et al. 2008; Zhu & Lipeme Kouyi 2019). Studies indicate that 2D orthorectification algorithms,
557 similar to those employed in the flume tests described, are suitable for relatively narrow channels
558 (approximately 10 – 20 m wide) with regular geometries, where the four specified control points and
559 flowing water can be reasonably assumed to lie within the same plane (Patalano et al. 2017; Bodart
560 et al. 2024). Conversely, in cases involving more complex channel geometries, irregular riverbank
561 topographies, or fluctuating flow levels, more advanced 3D orthorectification algorithms are
562 generally required (Le Coz et al. 2010; Detert 2021). These methods typically necessitate identifying
563 at least six control points, preferably up to ten for redundancy and better reliability (Fujita and Kunita
564 2011; Jolley et al. 2021), distributed uniformly along the observed channel segment. Notably, simply
565 increasing the number of control points does not inherently reduce orthorectification uncertainty.
566 Rather, the precision in locating these points is critical, making fewer but highly accurate control
567 points generally preferable to a larger number with less precise measurements (Le Coz et al. 2021;
568 Bodart et al. 2024).

569 Episodic flow-like landslides, such as debris flows, typically exhibit transitions between a front, a
570 central body, and a tail, potentially grading into hyperconcentrated water flows (Turnbull et al. 2015).

571 These events may feature successive surges, resulting in significant flow depth variations up to

572 several meters (Zanuttigh & Lamberti 2007; Meyrat et al. 2022). Therefore, adopting 3D
573 orthorectification algorithms is essential for field applications. Additionally, segmenting the acquired
574 footage into subsets based on homogeneous flow levels can help mitigate orthorectification
575 uncertainty caused by fluctuating flow depths (Theule et al. 2018).

576

577 Further laboratory flume tests would be required to thoroughly investigate and quantify errors and
578 limitations of the orthorectification process, particularly as a function of the non-zenithal (R_2) camera
579 placement. This would involve systematically varying the viewing angle, recording distance, and
580 potentially the flow level, enabling a comprehensive understanding of how these factors impact the
581 accuracy of reconstructed velocity fields.

582

583 **5. Conclusions**

584 In this study, two sets of laboratory flume experiments were performed between the Department of
585 Civil and Structural Engineering at the University of Sheffield (*F1*) and the British Geological Survey
586 site in Keyworth (*F2*). The primary goal was to assess *PIV* algorithm capabilities across a range of
587 recording scenarios, from ideal laboratory settings to suboptimal conditions typical of field
588 observations, and validate a methodology that could be employed to reconstruct the surface velocity
589 distribution of flow-type landslides and acquire insights into their rheological behaviour.

590 Based on the research findings, the following conclusions can be drawn:

591 1) The experiments performed under optimal conditions demonstrate the efficacy of the *PIV*
592 algorithm in reconstructing the velocity distributions of observed flows. Additionally, the processing
593 of tests monitored using zenithal and lateral cameras emphasises the critical role of understanding the
594 boundary conditions at the base and sides of the flow to reasonably estimate its rheological behaviour.
595 This is particularly crucial in field applications, especially if attempts are made to infer the internal
596 dynamics of the flow based on the exclusively available surface velocity distribution.

597 2) Non-zenithal footage, commonly available from field acquisitions, can be effectively utilised to
598 retrieve the real surface velocity distribution of the flow by adopting appropriate orthorectification
599 techniques. The analyses underscore the importance of accurately defining suitable control points in
600 the field for the orthorectification algorithm, noting that percentage errors in measuring their inter-
601 distances result in comparable deviations in the reconstructed velocity profiles.

602 3) Footage from widely deployed low-cost field camera monitoring systems, often characterised by
603 blurriness, suboptimal quality, or insufficient sampling rates, can still be used to estimate the surface
604 velocity distribution of flows, albeit with non-negligible uncertainty. This uncertainty may include

605 significant underestimations of the actual flow velocity, which can be detected and reasonably
606 mitigated by adopting sequences of larger interrogation patches in the *PIV* analysis, at the cost of
607 slightly lower spatial resolution.

608

609 The present work establishes the foundation for applying the detailed and tested methodological
610 framework to the in-depth characterisation of field-recorded debris flow events, addressing both
611 rheological behaviour and hydrodynamic parameters. A comprehensive analysis is currently
612 underway, including rigorous comparisons with field observations and numerical modelling analyses,
613 which will be presented in a forthcoming study.

614

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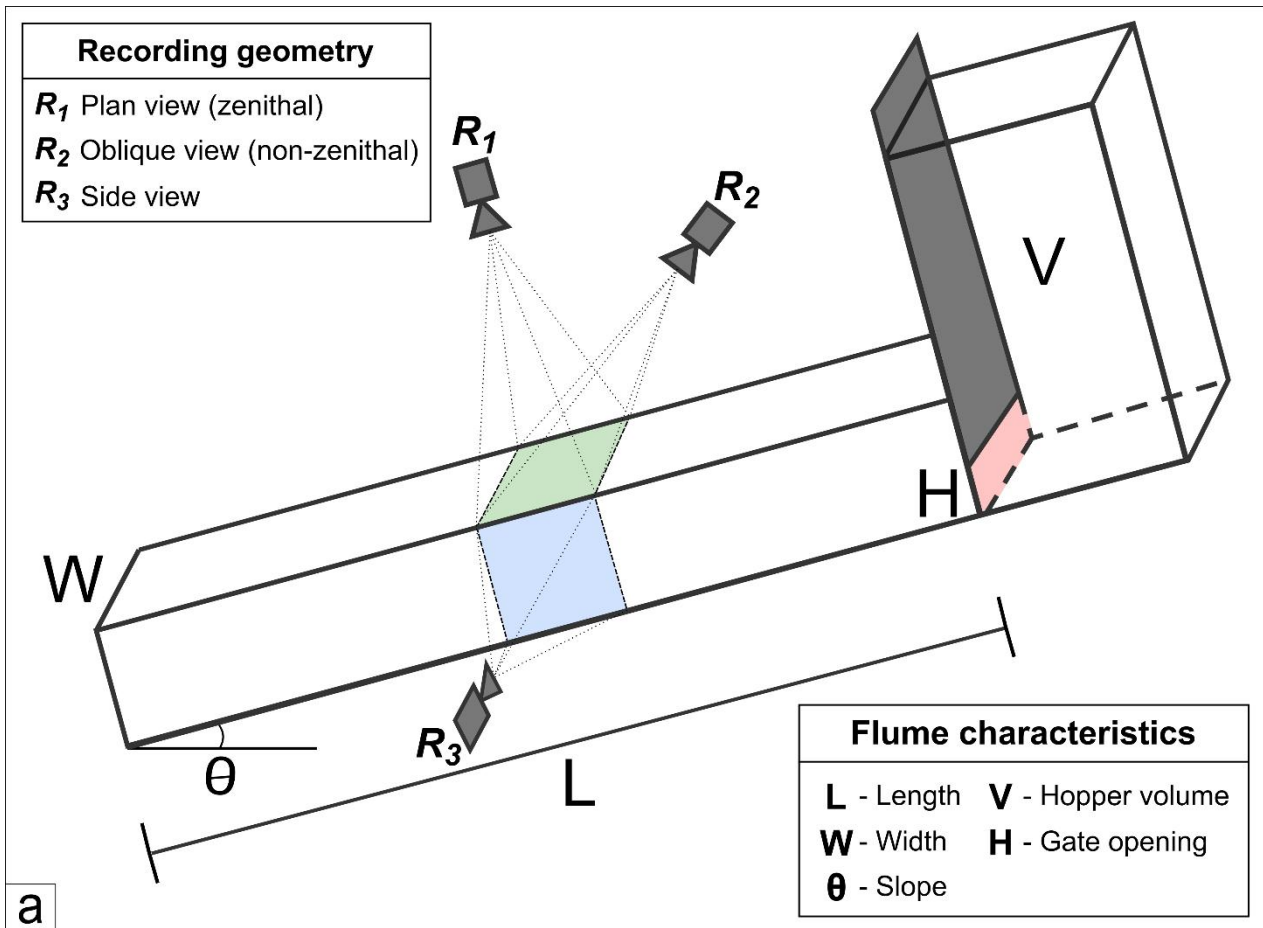
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797 Fig. 1 – Experimental setup scheme (a) and pictures of the different laboratory equipment employed for $F1$ (b) and $F2$
 798 (c) flume test sets.

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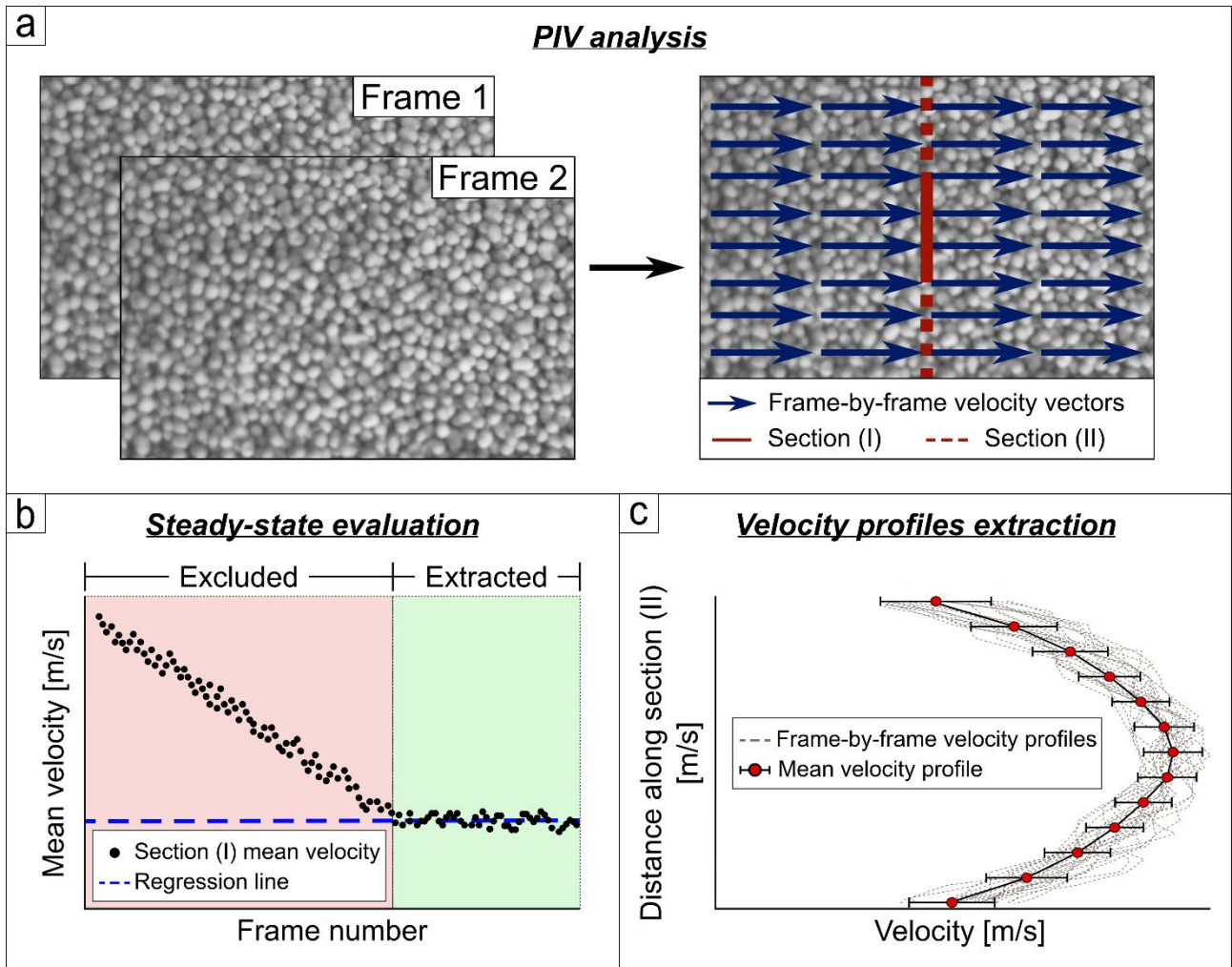
| Experimental setup | F1 | F2 |
|---------------------------------|---|---|
| <u>Flume length</u> (L) | 1.2 m | 4 m |
| <u>Flume width</u> (W) | 0.1 m | 0.1 m |
| <u>Flume slope</u> (θ) | 30° | 0 – 33° |
| <u>Hopper volume</u> (V) | 0.021 m ³ | 0.05 m ³ |
| <u>Gate opening</u> (H) | 0.085 m | 0.04 m |
| <u>Chute surface</u> | <ul style="list-style-type: none"> - Smooth (aluminium) - Ceramic beads ($d \approx 1$ mm) - Angular sand ($d \approx 1 - 2$ mm) | Smooth (aluminium) |
| <u>Material employed</u> | Denstone [®] ceramic beads ($d_{mean} \approx 3.85$ mm) | Fine sand ($d_{mean} \approx 0.28$ mm) |
| <u>Material volume</u> | 0.008 m ³ | 0.03 m ³ |
| <u>Cameras</u> | <i>Phantom Miro M310</i> (Frame rate = 1000 fps) | <i>GoPro Hero 10 [R₁, R₃] – 11 [R₂]</i> (Frame rate = 240 fps) |
| <u>Recording geometries</u> | R_1, R_3 | R_1, R_2, R_3 |
| <u>Recording Distances</u> | <ul style="list-style-type: none"> - $R_1 = 0.45$ m - $R_3 = 0.40$ m | <ul style="list-style-type: none"> - $R_1 = 0.25$ m - $R_2 = 0.30$ m - $R_3 = 0.20$ m |

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Table 1 – Summary of the primary characteristics of the flume apparatus, granular material, cameras and recording geometries utilised in the *F1* (blue) and *F2* (orange) experiment sets.



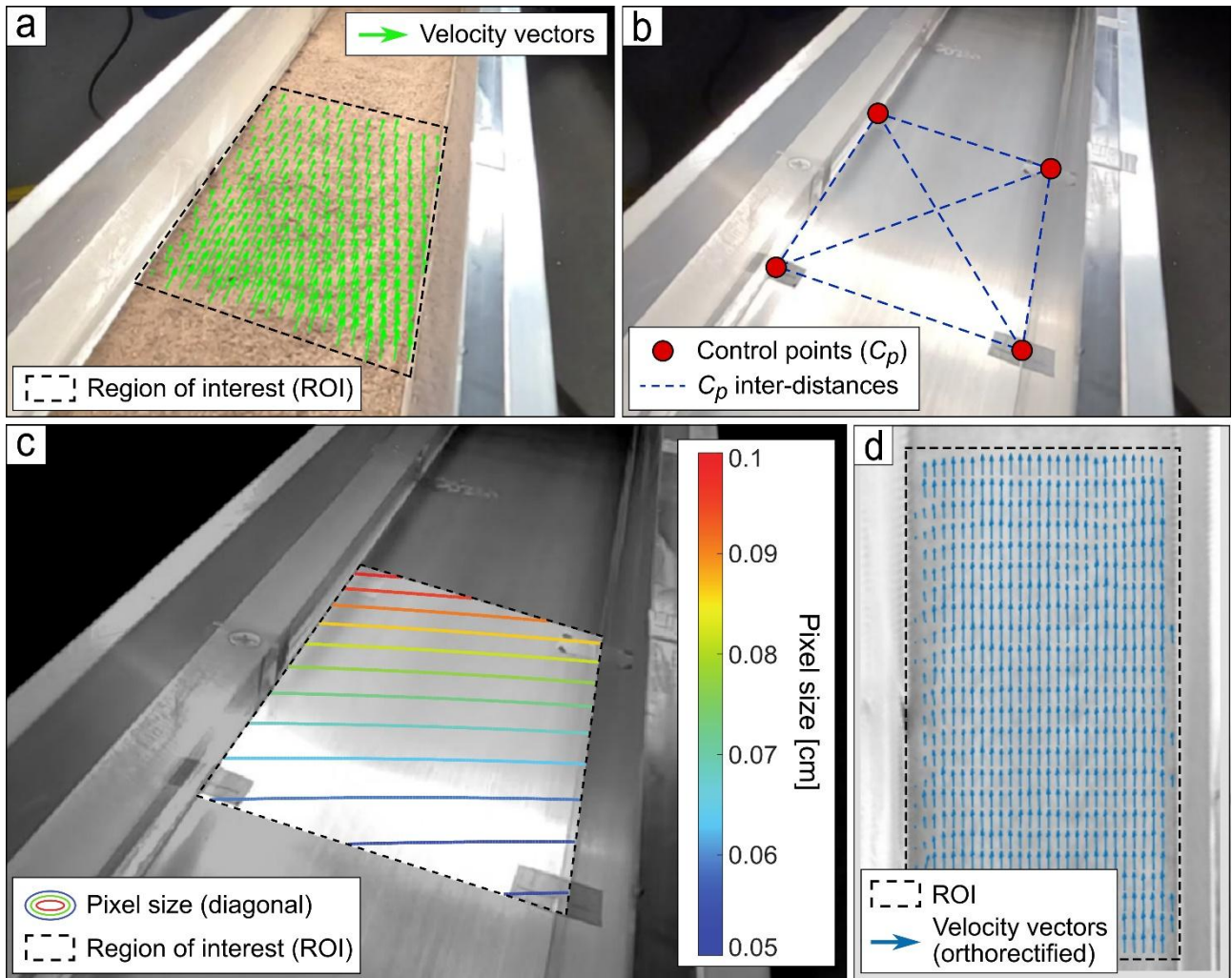
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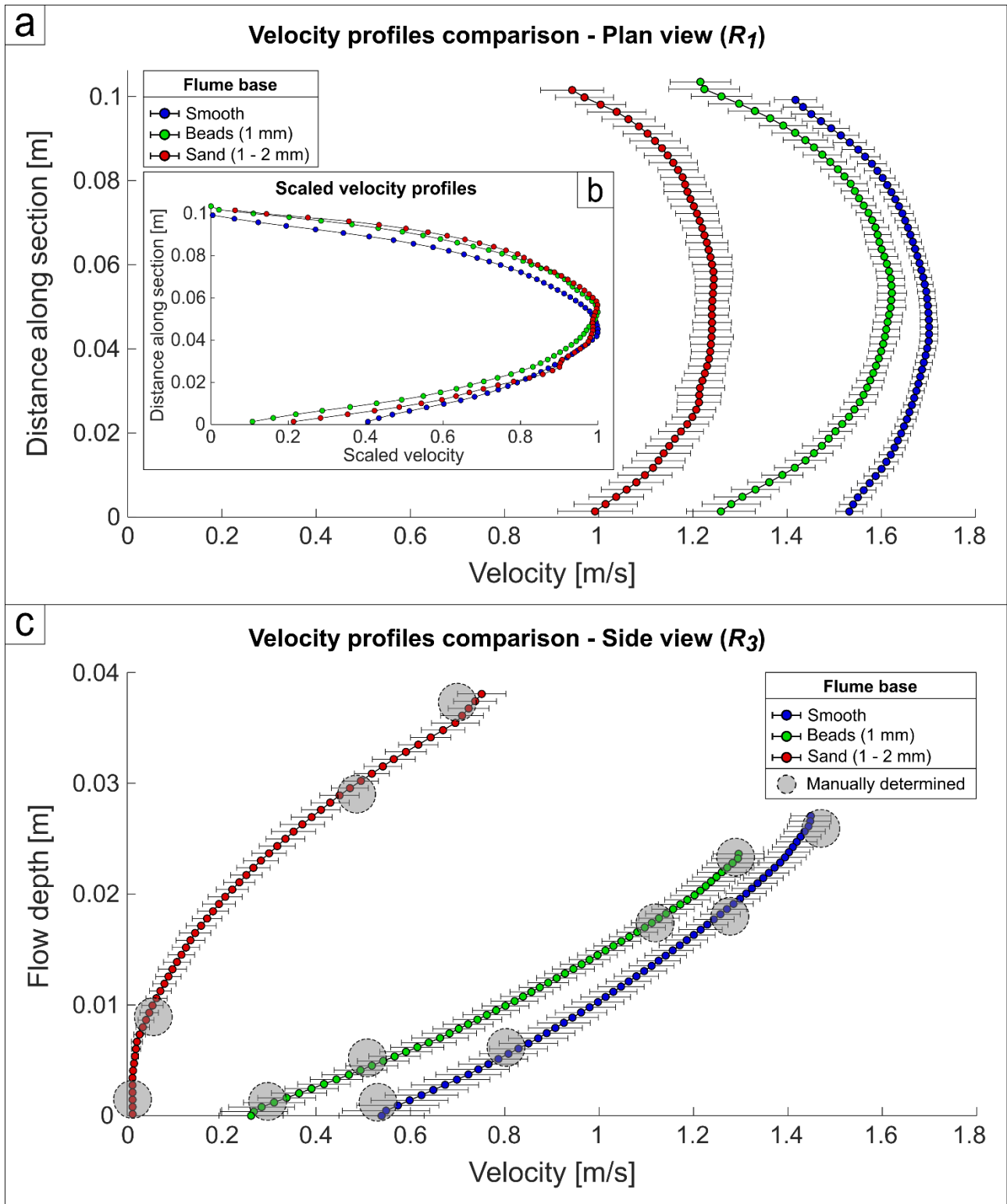
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Fig. 2 – Summary of the processing chain employed to reconstruct the velocity profiles of the recorded flows, including determination of the flow velocity vectors through *PIV* (a), evaluation of the quasi-steady-state of the flow (b), and extraction of the velocity profiles for the identified steady-state frames subset along reference sections (c).



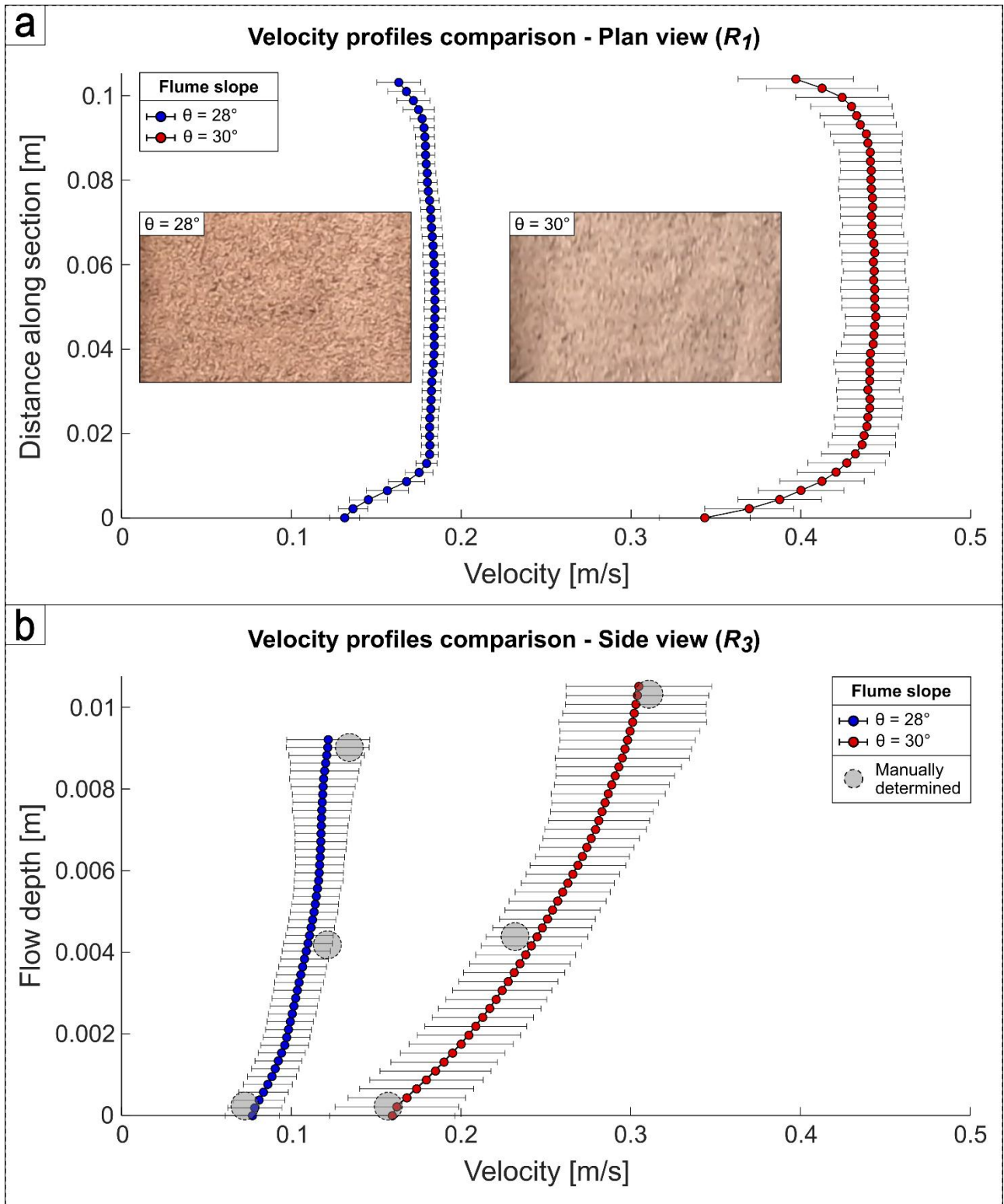
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810 Fig. 3 – Camera perspective and example of the distorted *PIV*-derived flow velocity field from R_2 recording geometry in
 811 *F2* flume tests (a). Locations of the four control points on the flume chute (b) required for applying the orthorectification
 812 algorithm in *RIVeR* (Patalano et al. 2017) to mitigate distortions from non-zenithal acquisitions by calculating the
 813 distribution of the actual pixel sizes across the image (c) and reprojecting the flow velocity vectors as if observed from a
 814 zenithal perspective (d).



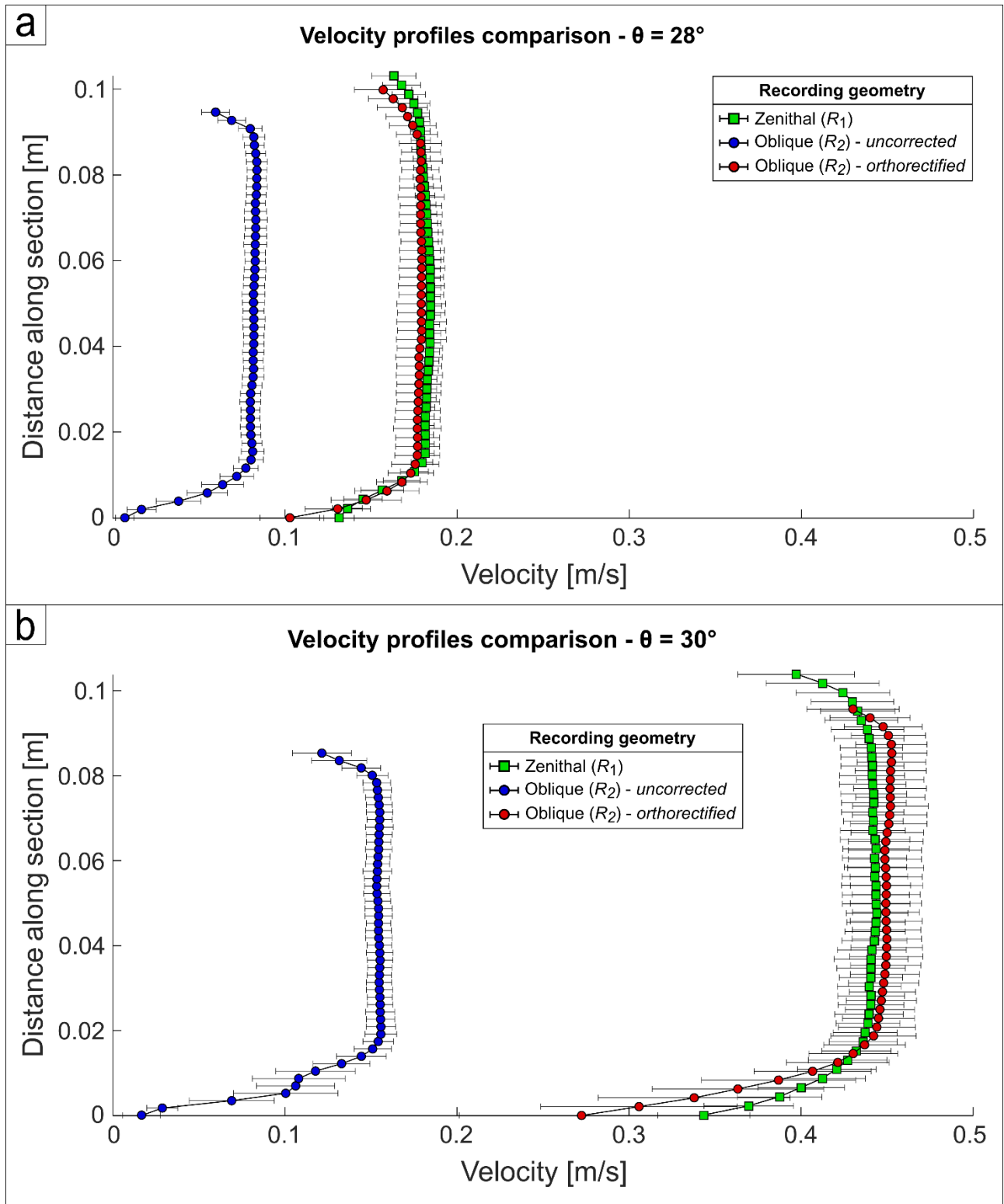
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816 Fig. 4 – Surface velocity distributions, both original (a) and normalised (b), along with side velocity profiles (c) relative
 817 to the flows recorded in the *FI* flume experiments set. The velocity distributions are colour-coded based on the boundary
 818 conditions imposed at the base of the flow, achieved by varying its roughness: smooth (blue), intermediate (green,
 819 constituted of 1mm ceramic beads), and coarse sand (red). The grey circles in (b) indicate manually determined velocity
 820 values, with circle size corresponding to the diameter of visually tracked beads.



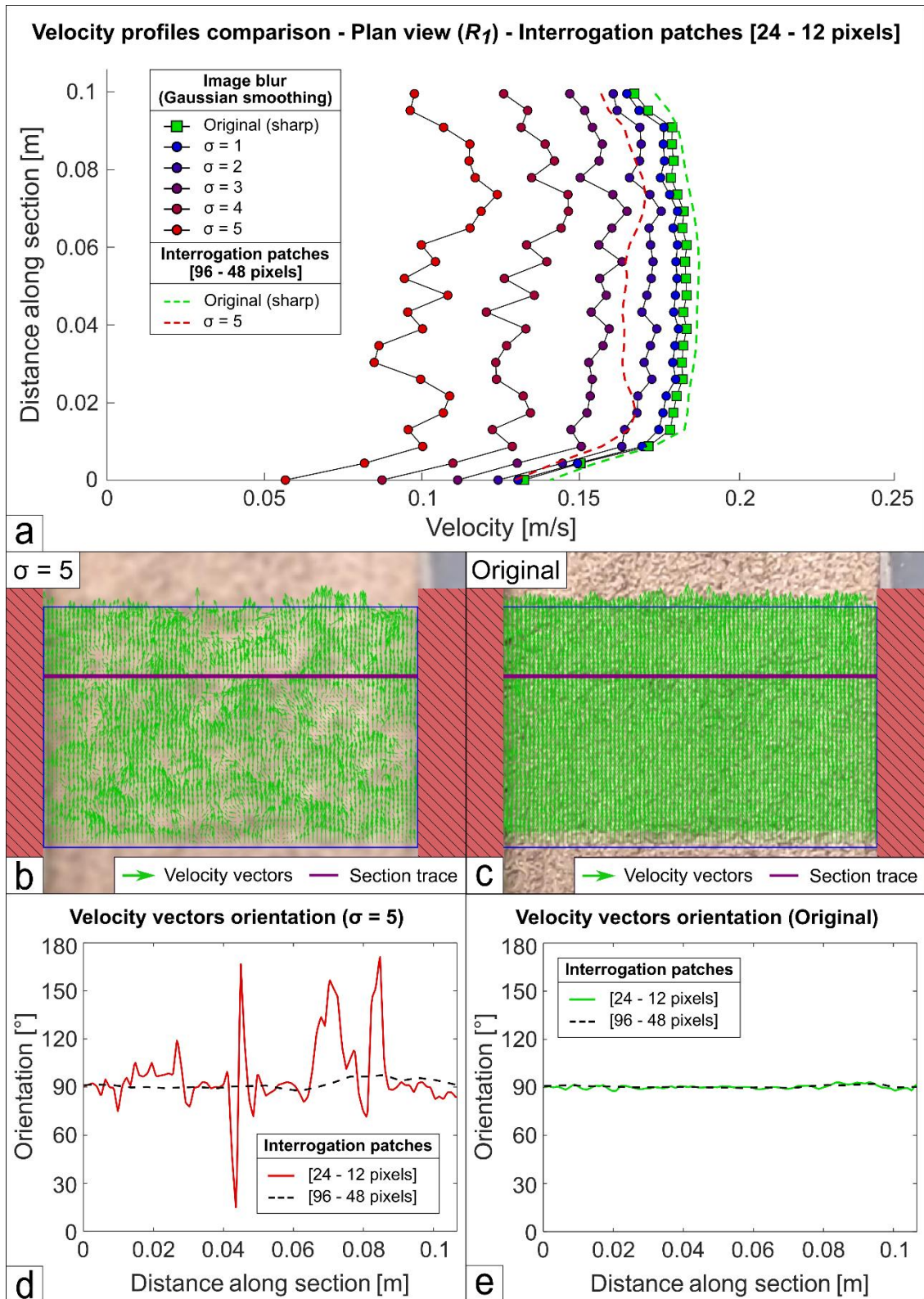
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822 Fig. 5 – Surface (a) and side (b) velocity profiles for the flows captured in the $F2$ flume experiments set. The velocity
 823 distributions are depicted in blue and red for the flows recorded at $\theta = 28^\circ$ and $\theta = 30^\circ$, respectively. Examples of plan
 824 view frames captured for the two image quality scenarios are also provided (a). The mild asymmetry in the surface velocity
 825 distributions in (a) may be attributed to a slight tilt in the flume chute. The grey circles in (b) indicate manually determined
 826 velocity values, with circle size corresponding to the diameter of visually tracked particles.



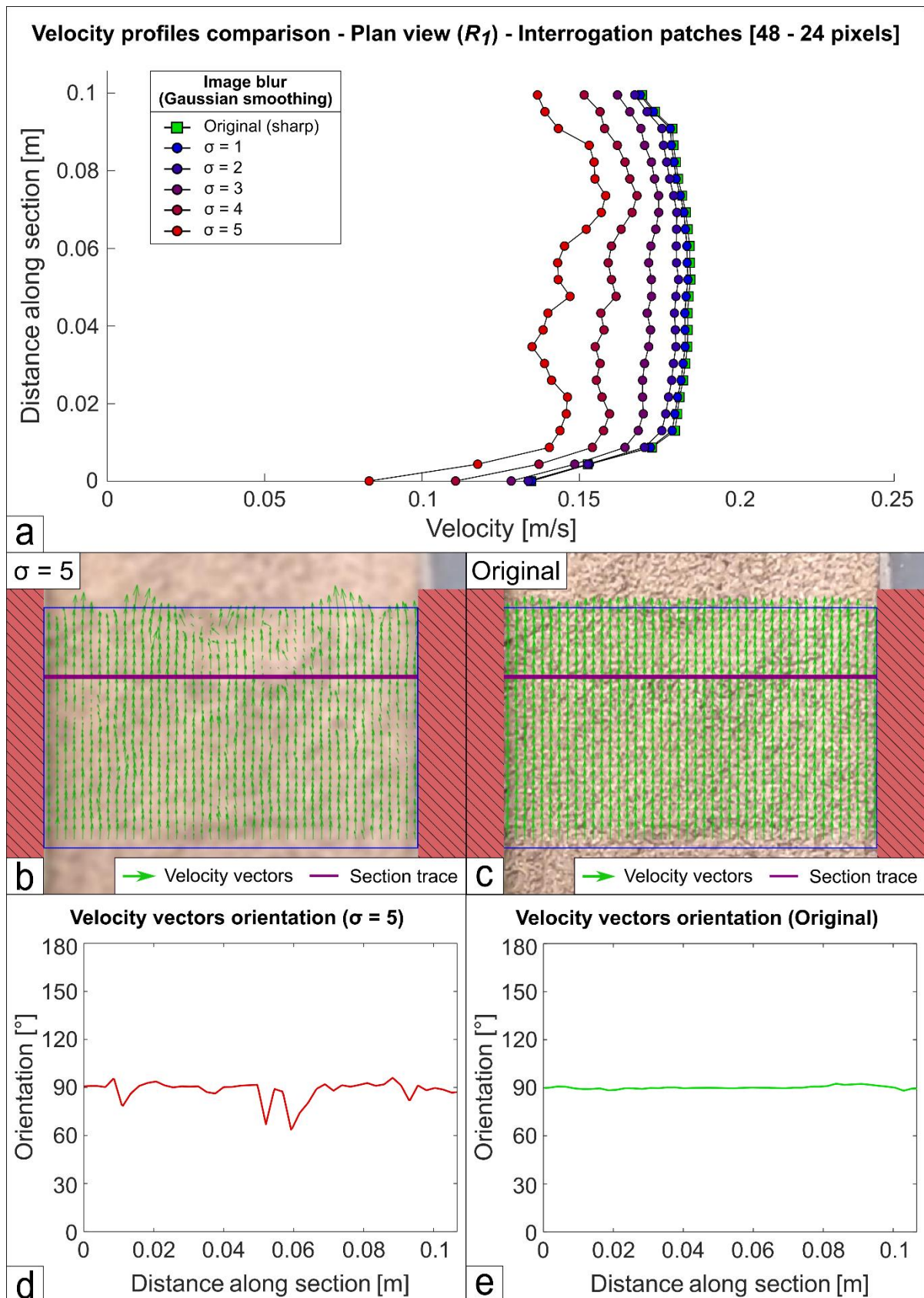
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828 Fig. 6 – Comparison of the surface velocity distributions reconstructed along the same section from the R_1 perspective
 829 (green) and R_2 uncorrected (blue) and orthorectified (red) flow velocity fields, corresponding to two reference image
 830 quality scenarios of $\theta = 28^\circ$ (a) and $\theta = 30^\circ$ (b) considered in the $F2$ experiments. The mild asymmetry in the surface
 831 velocity distributions in (a) may be attributed to a slight tilt in the flume chute.



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833 Fig. 7 – Surface velocity distributions reconstructed along a reference section from the R1 perspective using a sequence
 834 of moving interrogation patches of 24 – 12 pixels in the PIV analysis of the F2 test performed at $\theta = 28^\circ$. Curves are
 835 colour-coded based on the image blur level, expressed by the Gaussian smoothing filter parameter σ , increasing from blue
 836 to red. The green profile refers to the original sharp images. Dashed green and red lines represent the velocity profiles
 837 obtained from the real frames and highly blurred images ($\sigma = 5$), respectively, using an interrogation patch sequence of
 838 96 – 48 pixels (a). Zenithal view of flow vectors reconstructed for the same frame pair analysed at $\sigma = 5$ (b) and for clear
 839 footage (c). Orientation of flow velocity vectors extracted along the purple line shown in (b) and (c) for highly blurred
 840 footage (d) and original images (e). The dashed black lines represent velocity vector orientations determined along the
 841 same section employing an interrogation patch sequence of 96 – 48 pixels in the PIV analysis.



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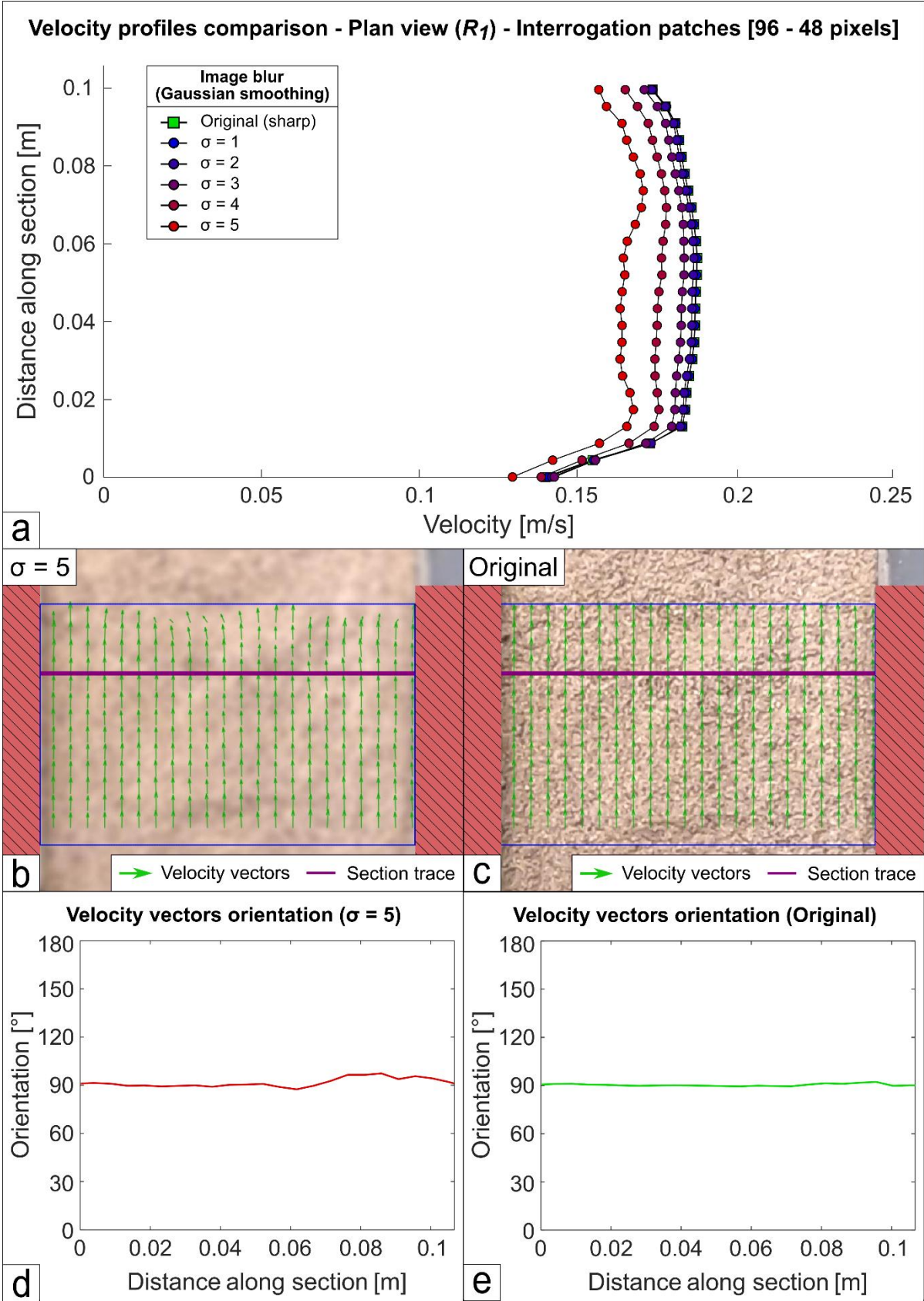
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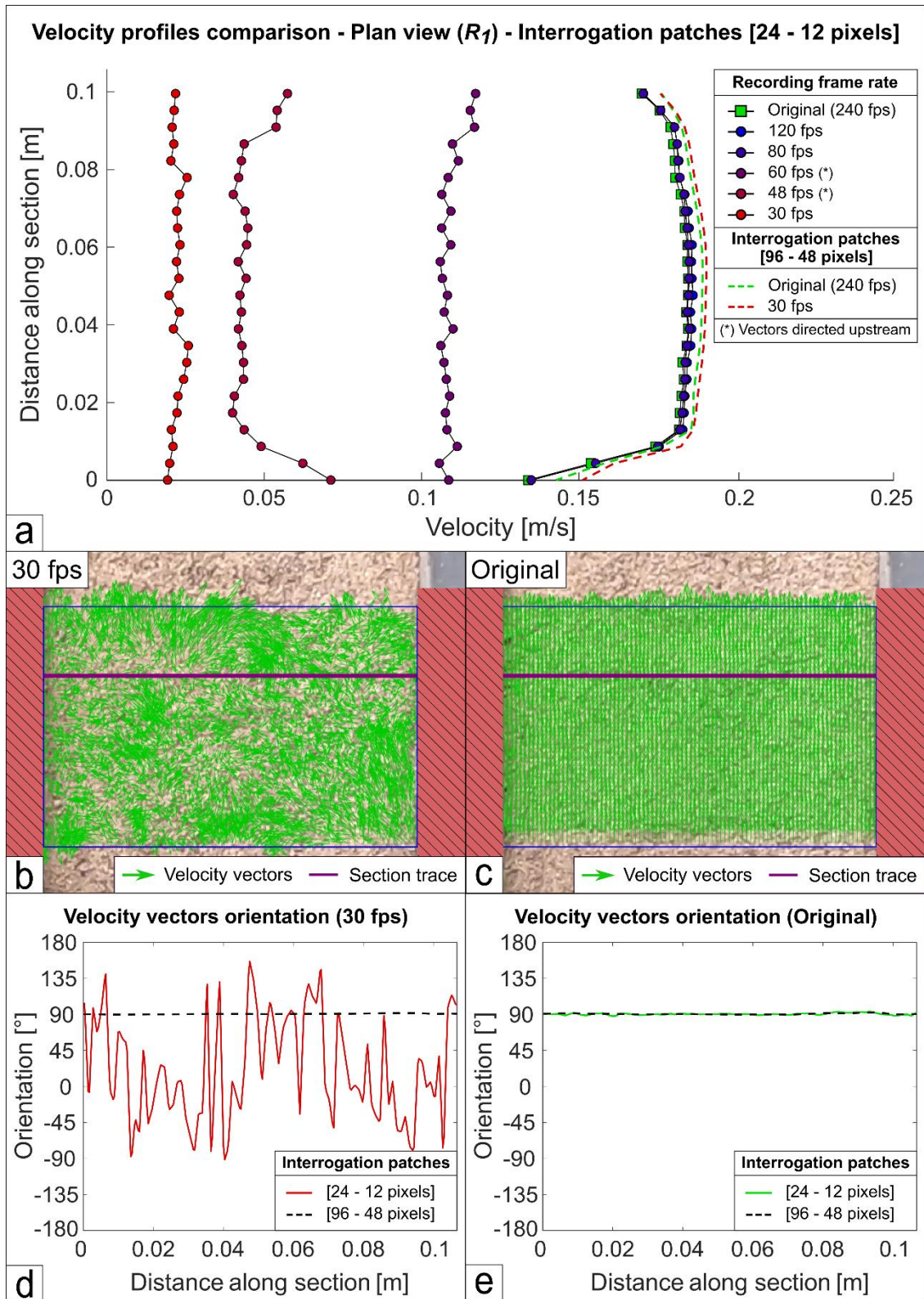
Fig. 8 – Surface velocity distributions reconstructed along a reference section from the R_1 perspective using a sequence of moving interrogation patches of 48 – 24 pixels in the PIV analysis of the F_2 test performed at $\theta = 28^\circ$. Curves are colour-coded based on the image blur level, expressed by the Gaussian smoothing filter parameter σ , increasing from blue to red. The green profile refers to the original sharp images (a). Zenithal view of flow vectors reconstructed for the same frame pair analysed at $\sigma = 5$ (b) and for clear footage (c). Orientation of flow velocity vectors extracted along the purple line shown in (b) and (c) for highly blurred footage (d) and original images (e).



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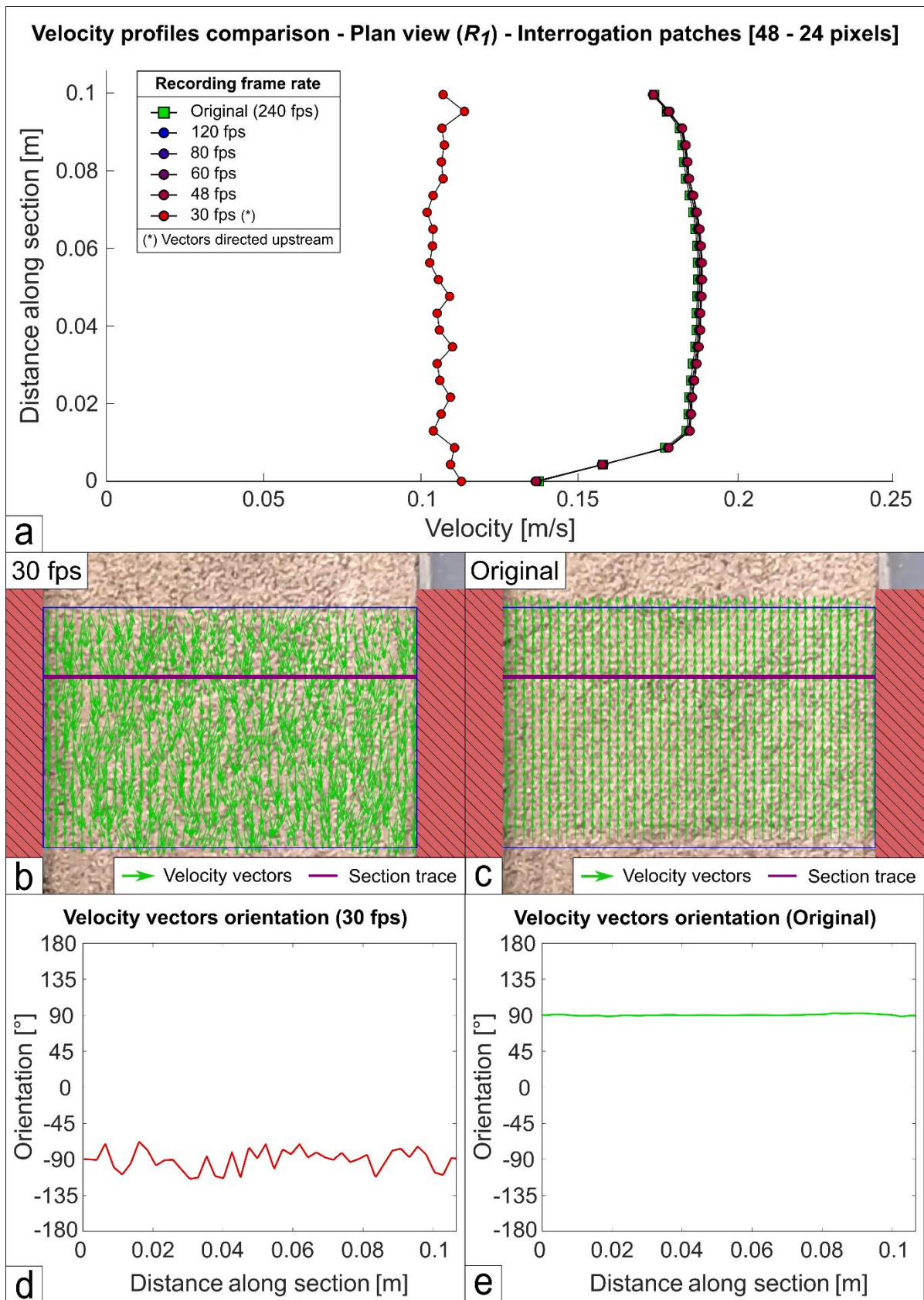
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Fig. 9 – Surface velocity distributions reconstructed along a reference section from the R1 perspective using a sequence of moving interrogation patches of 96 – 48 pixels in the PIV analysis of the F2 test performed at $\theta = 28^\circ$. Curves are colour-coded based on the image blur level, expressed by the Gaussian smoothing filter parameter σ , increasing from blue to red. The green profile refers to the original sharp images (a). Zenithal view of flow vectors reconstructed for the same frame pair analysed at $\sigma = 5$ (b) and for clear footage (c). Orientation of flow velocity vectors extracted along the purple line shown in (b) and (c) for highly blurred footage (d) and original images (e).



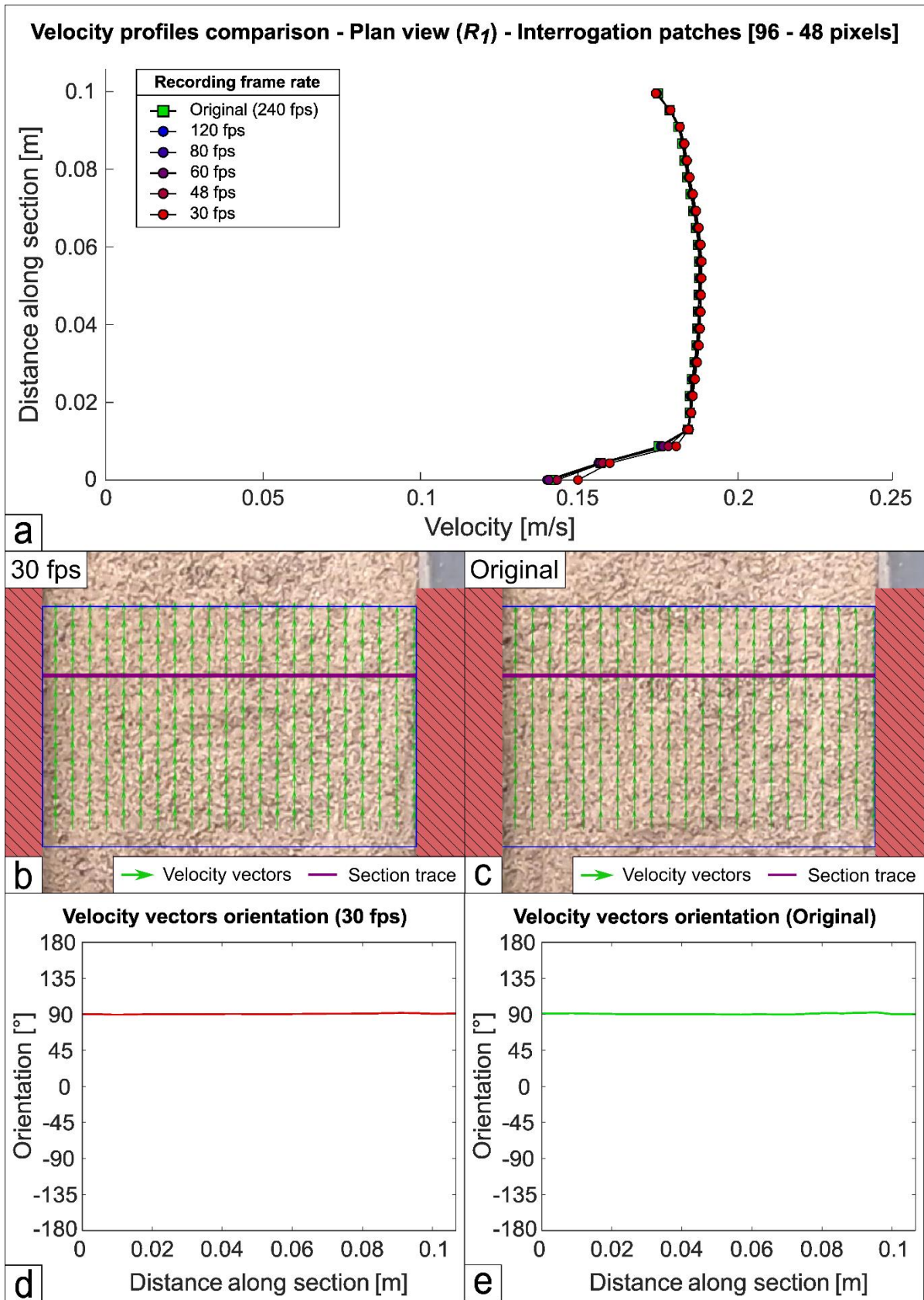
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857 Fig. 10 – Surface velocity distributions reconstructed along a reference section from the R_1 perspective using a sequence
 858 of moving interrogation patches of 24 – 12 pixels in the PIV analysis of the F_2 test performed at $\theta = 28^\circ$. Curves are
 859 colour-coded based on the recording frame rate, decreasing from blue to red. The green profile refers to the original
 860 images captured at 240 fps. Dashed green and red lines represent the velocity profiles obtained from the original images
 861 and footage undersampled at 30 fps, respectively, using an interrogation patch sequence of 96 – 48 pixels (a). Zenithal
 862 view of flow vectors reconstructed for the same frame pair analysed at 30 fps (b) and 240 fps footage (c). Orientation of
 863 flow velocity vectors extracted along the purple line shown in (b) and (c) for undersampled footage (d) and original
 864 images (e). The dashed black lines represent velocity vector orientations determined along the same section employing
 865 an interrogation patch sequence of 96 – 48 pixels in the PIV analysis.



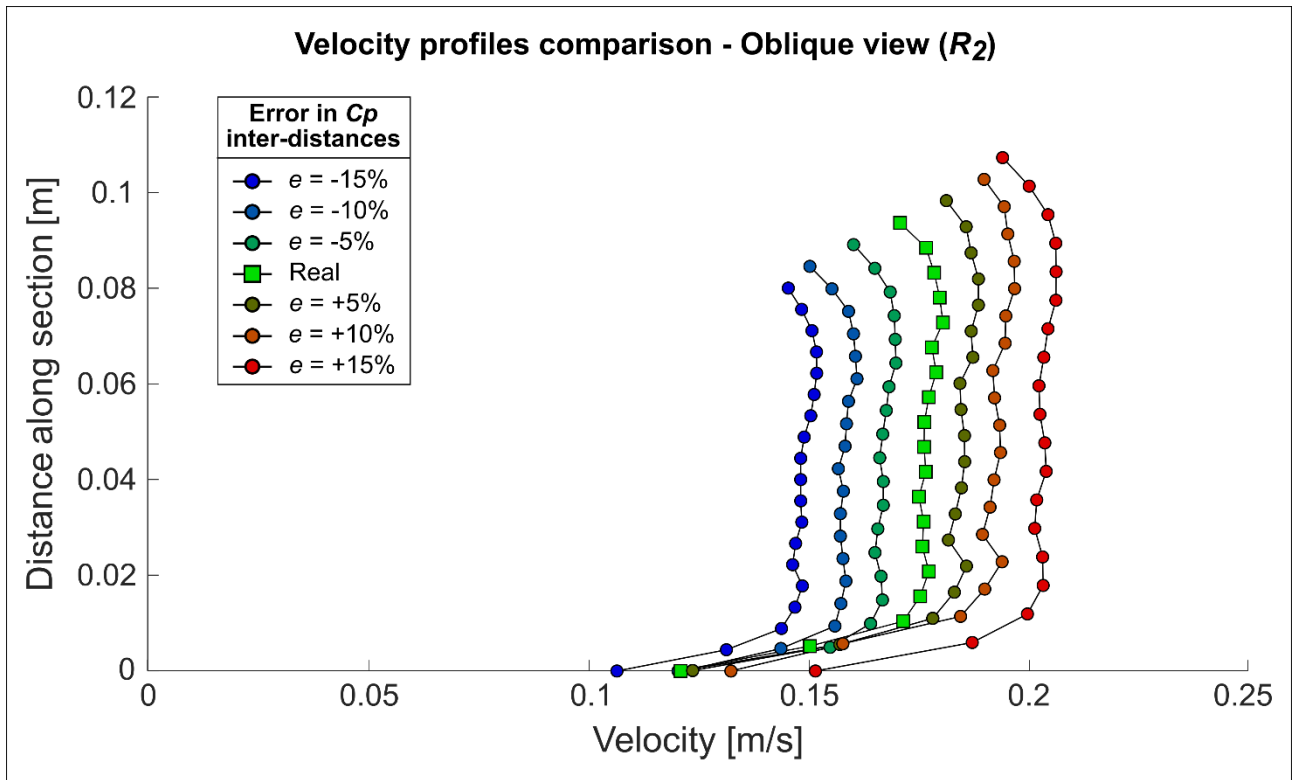
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867 Fig. 11 – Surface velocity distributions reconstructed along a reference section from the R_1 perspective using a sequence
 868 of moving interrogation patches of 48 – 24 pixels in the PIV analysis of the F_2 test performed at $\theta = 28^\circ$. Curves are
 869 colour-coded based on the recording frame rate, decreasing from blue to red. The green profile refers to the original
 870 images captured at 240 fps (a). Zenithal view of flow vectors reconstructed for the same frame pair analysed at 30 fps (b)
 871 and 240 fps footage (c). Orientation of flow velocity vectors extracted along the purple line shown in (b) and (c) for
 872 undersampled footage (d) and original images (e).



873

874 Fig. 12 – Surface velocity distributions reconstructed along a reference section from the R_1 perspective using a sequence
 875 of moving interrogation patches of 96 – 48 pixels in the PIV analysis of the F_2 test performed at $\theta = 28^\circ$. Curves are
 876 colour-coded based on the recording frame rate, decreasing from blue to red. The green profile refers to the original
 877 images captured at 240 fps (a). Zenithal view of flow vectors reconstructed for the same frame pair analysed at 30 fps (b)
 878 and 240 fps footage (c). Orientation of flow velocity vectors extracted along the purple line shown in (b) and (c) for
 879 undersampled footage (d) and original images (e).



880

881 Fig. 13 – Comparison of the surface velocity distributions obtained from the orthorectified velocity fields from R_2 non-
 882 zenithal recordings of the $F2$ test performed at $\theta = 28^\circ$, assuming errors in the estimation of distances between control
 883 points of $\pm 5\%$, 10% , and 15% in the orthorectification procedure. The actual orthorectified velocity profile is shown in
 884 green. The curves grading towards red refer to overestimations, while those grading towards blue indicate
 885 underestimations.