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Marine redox and nutrient dynamics linked to the Cambrian radiation of animals

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18 **ABSTRACT**

19 The early Cambrian witnessed an increase in metazoan ecosystem complexity, likely
20 linked to enhanced oxygen and nutrient availability. However, while improved stratigraphic
21 and geochemical records suggest that the Cambrian explosion occurred under highly
22 dynamic redox conditions, mechanistic links to nutrient availability and early Cambrian
23 evolutionary innovations are poorly constrained. Here, we report paleoredox and nutrient
24 data for two drill cores documenting late Cambrian Stage 2 to Stage 3 (~522 to 514 Ma)
25 strata from the Yangtze Block, South China. The development of water column euxinia
26 during Cambrian Stage 2 led to extensive recycling of P, fueling elevated primary production
27 and hence an increase in atmospheric and shallow marine oxygen concentrations. The
28 resulting expansion of oxygenated shelf area promoted sedimentary P retention, and in
29 combination with a diminished supply of P from upwelling, drove a transition to oligotrophic
30 shallow marine conditions during Cambrian Stage 3. Reduced primary production and
31 limited water column oxygen consumption allowed for the stabilization of oxygenated
32 continental shelf habitats that supported a burgeoning biotic complexity.

33

34 **INTRODUCTION**

35 The early Cambrian (Fortunian–end Stage 3; ca. 538.8–514 Ma) witnessed the abrupt
36 appearance and rapid diversification of most modern animal phyla during the ‘Cambrian
37 Explosion’ (Knoll and Carroll, 1999; Erwin et al., 2011; Wood et al., 2019). A wide range of
38 potential triggers for the Cambrian Explosion have been proposed, including intrinsic
39 developmental and ecological innovations (Erwin et al., 2011), and extrinsic changes
40 (particularly oxygen levels) to environmental habitability (Smith and Harper, 2013; Knoll and
41 Carroll, 1999; Lenton et al., 2014). The first animals likely had low oxygen demands (Mills
42 et al., 2014), however the larger and more metabolically active skeletonizing animals that
43 appeared during the early Cambrian would have required considerably higher oxygen levels
44 (Sperling et al., 2013). Indeed, potential links between pulses of shallow-ocean oxygenation
45 and biodiversity have been recorded on the early Cambrian Siberian platform (He et al.,

46 2019; Zhuravlev et al., 2022) and in South China (Chen et al., 2015; Li et al., 2017).

47 The availability of the bio-limiting nutrients phosphorus (P) and nitrogen (N) likely also
48 exerted an important control on environmental habitability, with P generally considered to be
49 the ultimate limiting nutrient for primary production on geological timescales (Tyrell, 1999).
50 These nutrients control the degree of regional oxygen production and consumption, as well
51 fueling more complex life (Brocks et al., 2017). Depositional redox conditions also control P
52 recycling, resulting in feedbacks that dictate the evolution of regional marine redox state
53 (Schobben et al., 2020; Bowyer et al., 2023). However, despite their fundamental importance
54 for ecosystem habitability and animal evolution, integrated records of redox and P cycling in
55 the early Cambrian are highly limited (see Creveling et al., 2014).

56 Here, we report a geochemical study of shelf and slope successions from the Yangtze
57 Block, South China. We utilize Fe speciation, redox sensitive trace element systematics and
58 P phase partitioning to evaluate contemporaneous marine redox conditions and controls on
59 primary productivity. These data allow a detailed assessment of redox and nutrient controls
60 on the major biological events that occurred during the early Cambrian.

61

62 **GEOLOGIC SETTING AND MATERIALS**

63 We investigate two drill cores from the Yangtze Block, deposited in an open-marine
64 continental shelf (ZK 402) and slope (CY 1) setting (see Supplemental Material¹ for full
65 details of the geologic setting and methods). Both cores preserve a relatively continuous
66 record spanning the latter part of Cambrian Stage 2 through Cambrian Stage 3, offering a
67 prime opportunity to evaluate biogeochemical changes across an interval that captures a
68 major evolutionary event.

69

70 **RESULTS AND DISCUSSION**

71 **Oceanic redox dynamics during the early Cambrian**

72 We employ a combination of Fe speciation and redox-sensitive trace metal (RSTM; Mo,
73 U, V) systematics (see Text S3) to provide a robust framework for interpreting depositional

74 redox conditions (e.g., Tribovillard et al., 2006; Poulton, 2021). In modern and ancient
75 marine sediments, a highly reactive (Fe_{HR}) to total Fe (Fe_T) ratio of >0.38 provides an
76 indication of anoxia, while $Fe_{HR}/Fe_T < 0.22$ suggests oxic conditions, with values between
77 considered equivocal. When anoxia is indicated, the extent of pyritization (Fe_{Py}) of the Fe_{HR}
78 pool is used to differentiate between ferruginous ($Fe_{Py}/Fe_{HR} < 0.6$) or euxinic ($Fe_{Py}/Fe_{HR} > 0.8$)
79 conditions (Poulton, 2021). Generally, RSTMs tend to be less soluble under reducing
80 conditions, leading to authigenic enrichments in the sediment, with Mo in particular requiring
81 the presence of dissolved sulfide (Tribovillard et al., 2006).

82 During late Cambrian Stage 2 to early Stage 3 (interval I), the development of persistent
83 euxinia at both sites is suggested by a combination of high Fe_{HR}/Fe_T (>0.7) and high
84 Fe_{Py}/Fe_{HR} ratios (0.69 ± 0.11 for ZK 402 and 0.92 ± 0.04 for CY 1; Fig. 1; Table S1). This
85 interpretation is supported by high enrichment factors (EFs; see Text S2) for the RSTMs and
86 high Mo_{EF}/U_{EF} ratios. During the initial stages of interval II (middle to late Stage 3), a decline
87 in Fe_{HR}/Fe_T ratios, EF values and Mo_{EF}/U_{EF} ratios occurs in both drill cores (Fig. 1).
88 Subsequently, on the shelf (ZK 402), Fe_{HR}/Fe_T ratios and RSTM enrichments remain
89 persistently low. This suggests a transition to dominantly oxic conditions, although
90 occasional minor increases in Fe_{HR}/Fe_T ratios, EF values and Mo_{EF}/U_{EF} ratios may indicate
91 transient, short-lived anoxia. On the slope (CY 1), despite decreasing trends in redox proxy
92 data from intervals I to II, values remain elevated, indicating continued euxinia, perhaps
93 punctuated by episodic ferruginous intervals (lower Mo_{EF}/U_{EF} ratios; Fig. 1).

94 A compilation of global Fe speciation data (Text S4; Table S2) provides further insight into
95 spatio-temporal variability in marine redox state (Fig. 2A). While the majority of data are from
96 South China, a stratified early Cambrian ocean is apparent, with oxic surface waters
97 overlying ferruginous deeper waters (with occasional euxinia), prior to the latter part of Stage
98 2 (Fig. 2). Euxinia then became common in outer shelf to slope settings, before extending
99 into shallow shelf settings at the Stage 2-Stage 3 boundary. A clear transition to expanded
100 shelf oxygenation then occurred, while deeper waters remained largely anoxic during Stage
101 3. The local marine redox state inferred from Fe speciation data is consistent with estimates
102 of global marine redox from U and Mo isotopes (Fig. 2). Carbonates from South China,

103 Siberia and Morocco in the latest Fortunian and Stage 2 document negative $\delta^{238}\text{U}$
104 excursions (Wei et al., 2018; Dahl et al., 2019), suggesting episodic expansions of seafloor
105 anoxia during these intervals. By contrast, consistently higher $\delta^{238}\text{U}$ data recorded in
106 carbonates from Siberia and shales from South China in Stage 3 (Dahl et al., 2019; Wei et
107 al., 2020) provide evidence for more widespread oxygenation. Furthermore, modern-like
108 $\delta^{98}\text{Mo}$ signatures in Cambrian Stage 3 also suggest a major expansion of oxic seafloor
109 (Chen et al., 2015).

110

111 **Marine eutrophication and animal diversification**

112 High total organic carbon (TOC) and total P (P_{tot}) concentrations recorded in the lower
113 part of interval I in both drill cores (Fig. 3) potentially indicate high marine primary productivity
114 and nutrient levels. The greenhouse climate of the early Cambrian (Hearing et al., 2018),
115 likely drove enhanced chemical weathering (Peters and Gaines, 2012; Lipp et al., 2021) and
116 hence a high terrigenous nutrient influx. Alternatively, or in addition, high nutrient levels could
117 have occurred via upwelling of nutrient-replete deep ocean waters (e.g., Behrenfeld et al.,
118 2006; Bowyer et al., 2017). Upwelling can be evaluated using Co (ppm)×Mn (wt%) values,
119 whereby upwelling systems are characterized by low sedimentary Co and Mn
120 concentrations due to their depletion in upwelled deep-ocean waters, in addition to recycling
121 of Mn back to the water column under reducing conditions. However, restricted basins are
122 characterized by relatively high Co and Mn concentrations due to relatively high riverine
123 inputs and limited water exchange with the open ocean (Sweere et al., 2016).

124 The low Co×Mn values in the lower part of interval I are similar to those of modern
125 upwelling systems (<0.4 ppm%; Sweere et al., 2016), suggesting strong upwelling (Text S5).
126 This internal marine nutrient supply was augmented by nutrient recycling from the sediment,
127 which we evaluate via the phase partitioning of P. This approach quantifies four
128 operationally-defined P pools (Thompson et al., 2019), including Fe-bound P (P_{Fe}),
129 authigenic carbonate fluorapatite, biogenic apatite and CaCO_3 -bound P (P_{auth}), detrital
130 apatite (P_{det}), and organic-bound P (P_{org}). Reactive P (P_{reac}) represents potentially mobile P
131 during transport and early diagenesis, and is calculated as the sum of P_{Fe} , P_{auth} and P_{org} .

132 P_{org} is preferentially released from organic matter (C_{org}) via anaerobic remineralization,
133 resulting in C_{org}/P_{org} ratios above the Redfield ratio of 106/1 (Ingall et al., 1993). In addition,
134 P_{Fe} may be released during reductive dissolution of Fe (oxyhydr)oxide minerals (Slomp et
135 al., 1996). A proportion of this P may undergo 'sink switching' to authigenic phases in the
136 sediment (Slomp et al., 1996). However, P may also be recycled back to the water column,
137 particularly under sulfidic conditions, potentially driving a positive productivity feedback (Van
138 Cappellen and Ingall, 1996).

139 In samples from interval I, molar C_{org}/P_{org} and C_{org}/P_{reac} ratios are significantly greater than
140 the Redfield ratio of 106/1 (Fig. 3), demonstrating release of P from organic matter (C_{org}/P_{org})
141 and recycling back to the water column (C_{org}/P_{reac}). This recycling occurs alongside high P_{tot}
142 contents relative to average shale (0.009; Turekian and Wedepohl, 1961), suggesting
143 particularly high levels of bioavailable phosphate in the overlying water column. Expanded
144 euxinia (Fig. 2) would have resulted in a high degree P recycling in continental margin
145 settings, promoting a strong positive productivity feedback following upwelling to the surface
146 ocean (Van Cappellen and Ingall, 1996), and providing a food supply for early mobile
147 animals (Fig. 4). Furthermore, increased TOC burial in shelf and slope environments as a
148 consequence of enhanced primary productivity likely contributed to gradual atmospheric and
149 shallow ocean oxygenation (He et al., 2019). Hence, increased oxygen availability and food
150 supply in the shallow ocean drove the major evolutionary event recorded by the Chengjiang
151 and Qingjiang lagerstätte, which occur coincident with the first appearance of mineralized
152 trilobites in Cambrian Stage 3 (Fig. 2; Sun et al., 2022). The gradual decline in TOC and P
153 in the upper part of interval I may be related to the development of weakly euxinic conditions
154 and weakened upwelling, leading to a decrease in bioavailable P and hence lower TOC
155 burial.

156

157 **Oligotrophic stabilization of shelf oxygenation**

158 Total P (as P_{tot} and P_{tot}/Al ratios) and TOC remain low throughout interval II in both cores,
159 suggesting reduced levels of productivity and nutrient availability (Fig. 3). Variability in TOC
160 and P may also be influenced by changes in redox conditions, lithology and sedimentation

161 rate. However, the decrease in TOC and P in both cores initially occurs in shale in the upper
162 part of interval I, which was deposited under euxinic conditions. Furthermore, despite the
163 lithological differences between the two cores in interval II (and possible corresponding
164 changes in sedimentation rate), a synchronous decline in TOC and P is observed in both
165 cores, suggesting that productivity was the dominant control on TOC and P variability (Text
166 S6).

167 Given the generally enhanced continental weathering flux indicated for the early Cambrian
168 (including during intervals I and II; Peters and Gaines, 2012; Lipp et al., 2021), the decline
169 in nutrient levels likely reflects a reduction in sedimentary P recycling and P upwelling from
170 deeper waters. During interval II, P_{tot}/Al ratios remain below average shale in both drill cores
171 (Fig. 3). On the shelf (ZK 402), ferruginous samples in the lower part of interval II have
172 elevated $C_{\text{org}}/P_{\text{org}}$ and $C_{\text{org}}/P_{\text{reac}}$ ratios relative to the Redfield ratio, but with values that are
173 considerably lower than those of interval I (Fig. 3). This suggests that the extent of P
174 recycling was considerably diminished, with lower P_{tot}/Al ratios implying a major overall
175 decrease in bioavailable P. Furthermore, while the dominantly oxic samples from the rest of
176 interval II also have high $C_{\text{org}}/P_{\text{org}}$ ratios, $C_{\text{org}}/P_{\text{reac}}$ ratios plot below the Redfield ratio. This
177 suggests that the P released during early diagenesis was fixed in the sediment as authigenic
178 phases (supported by an increased proportion of P_{auth} ; Fig. 3). This C/P pattern implies very
179 low bioavailable P, consistent with most modern oligotrophic settings, which are
180 characterized by higher $C_{\text{org}}/P_{\text{org}}$ ratios and lower $C_{\text{org}}/P_{\text{reac}}$ ratios (Slomp et al., 2013).

181 In the slope setting (CY 1), $C_{\text{org}}/P_{\text{org}}$ ratios are elevated and $C_{\text{org}}/P_{\text{reac}}$ ratios scatter close
182 to the Redfield ratio during interval II, with higher values for euxinic samples and lower
183 values for ferruginous samples (Fig. 3 and Fig. S4). These is consistent with a dynamic P
184 response to the prevailing water column redox conditions, with intermittently enhanced P
185 recycling back to the water column under euxinic conditions, although the lower P_{tot} contents
186 relative to interval I (Fig. 3) suggest that water column P concentrations were likely greatly
187 diminished overall. In addition, elevated Co×Mn values suggest that this limited flux of
188 regenerated P was likely not efficiently transported to the photic zone, due to much lower
189 rates of upwelling, thereby resulting in the development of oligotrophic conditions on the

190 shelf (Fig. 4). This would have reduced primary productivity in surface waters, thus
191 decreasing the amount of dissolved O₂ consumed during the oxidation of sinking organic
192 matter, and thereby stabilizing oxic conditions in shelf environments.

193 Our data demonstrate that variability in ocean circulation and sedimentary P recycling
194 played a critical role in marine redox and nutrient dynamics during the early Cambrian.
195 Prevailing euxinia along continental margins and active ocean circulation during late
196 Cambrian Stage 2 to early Stage 3 resulted in intensified sedimentary P recycling in deep
197 waters and enhanced nutrient supply to the surface ocean (Fig. 4). Increased primary
198 production would have led to higher atmospheric and shallow marine oxygen levels, and
199 hence an expansion of oxic waters. The enhanced oxygen availability and food supply likely
200 facilitated the early Cambrian bioradiation that occurred at this time. The expanded oxic
201 seafloor area and diminished ocean circulation that occurred during middle to late Cambrian
202 Stage 3 then promoted the retention of sedimentary P in shelf areas, and along with reduced
203 upwelling, resulted in a decrease in shallow marine nutrient levels. Under the ensuing
204 oligotrophic conditions, stable oxygenated continental shelves developed, thereby creating
205 habitats that supported a burgeoning diversity in animal life.

206

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213

214 **Data Availability**

215 All data that support the findings of this study are provided in Supplementary Tables S1
216 and S2.

217

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346 **FIGURE CAPTIONS**

347 Fig. 1. Stratigraphic plots of Fe speciation, RSTM EFs and Mo_{EF}/U_{EF} ratios for drill cores
348 ZK402 and CY 1. Fm, Formation; Ed, Ediacaran; Fo, Fortunian; DY, Dengying; LCP,
349 Liuchapo.

350

351 Fig. 2. Marine redox conditions and biodiversity patterns during the early Cambrian. A.
352 Compilation of Fe speciation data for platform, shelf, slope and basin environments from
353 541 Ma to 514 Ma. Blue and orange lines represent the LOESS curves. B. Global diversity
354 of phyla and classes (Erwin et al., 2011) and major Cambrian lagerstätten (Sun et al., 2022).
355 C. Schematic evolution of local marine redox conditions based on iron speciation data. D.
356 Schematic evolution of global marine redox conditions based on Mo and U isotope records
357 (Chen et al., 2015; Dahl et al., 2019; Wei et al., 2020).

358

359 Fig. 3. Stratigraphic distribution of TOC, $\delta^{13}C_{org}$, P_{tot} , P_{tot}/Al , C_{org}/P_{org} , C_{org}/P_{reac} , $Co \times Mn$ and
360 P phase partitioning.

361

362 Fig. 4. Summary of marine redox and nutrient cycling during the early Cambrian. A.

363 Persistently euxinic conditions on the shelf and slope promoted P regeneration into the water
364 column. Recycled P was transported to the photic zone by active upwelling, resulting in a
365 positive productivity feedback that led to increased C_{org} burial, and ultimately atmospheric
366 and shallow marine oxygenation. B. P was effectively fixed in shelf sediments. Continued
367 euxinia in slope environments may have resulted in intermittent P recycling to the water
368 column, but P in deep waters was not effectively transported to the photic zone due to
369 decreased upwelling, which limited primary production and O₂ consumption.

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371 ¹Supplemental Material. Detailed geological setting, methods, supplemental notes, Figures
372 S1–S4 and Tables S1–S2. Please visit <https://doi.org/10.1130/XXXX> to access the
373 supplemental material, and contact editing@geosociety.org with any questions.