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He, X. orcid.org/0000-0002-2214-9338, Jiang, X. orcid.org/0000-0003-4141-1538, Spracklen, D.V. orcid.org/0000-0002-7551-4597 et al. (8 more authors) (2023) Global distribution and climatic controls of natural mountain treelines. *Global Change Biology*. ISSN 1354-1013

<https://doi.org/10.1111/gcb.16885>

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1 **Global distribution and climatic controls of natural mountain treelines**

2 **Running title: Climate treeline patterns and drivers**

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21

22 **Abstract**

23 Mountain treelines are thought to be sensitive to climate change. However, how climate impacts
24 mountain treelines is not yet fully understood as treelines may also be affected by other human
25 activities. Here we focus on “closed-loop” mountain treelines (CLMT) that completely encircle a
26 mountain and are less likely to have been influenced by human land-use change. We detect a total
27 length of ~916,425 km of CLMT across 243 mountain ranges globally and reveal a bimodal
28 latitudinal distribution of treeline elevations with higher treeline elevations occurring at greater
29 distances from the coast. Spatially, we find that temperature is the main climatic driver of treeline
30 elevation in boreal and tropical regions, whereas precipitation drives CLMT position in temperate
31 zones. Temporally, we show that 70% of CLMT have moved upwards, with a mean shift rate of
32 1.2 m/year over the first decade of the 21st century. CLMT are shifting fastest in the tropics (mean
33 of 3.1 m/year), but with greater variability. Our work provides a new mountain treeline database
34 that isolates climate impacts from other anthropogenic pressures, and has important implications
35 for biodiversity, natural resources, and ecosystem adaptation in a changing climate.

36

37 **Keywords:** treeline, forest boundary, tree cover, climate, mountain ecosystems, alpine area

38

39 **1. Introduction**

40 The mountain treeline is the upper altitudinal limit of tree growth toward the top of mountains, a
41 transitional zone from forests to treeless alpine vegetation (Körner & Paulsen, 2004). Treeline
42 ecotones play important environmental roles, including as habitats for endemic species and by
43 contributing to water supply (Grace, 1989). Mountain treelines are important indicators of the
44 impact of climate change on upland ecosystems (Verrall & Pickering, 2020; Lu et al., 2021) as they
45 are strongly associated with growing season lengths and minimum daily temperatures (Paulsen &
46 Körner 2014). Consequently, as a response to global warming, mountain treelines are expected to
47 shift upward as high elevations become more favourable for tree establishment under a changing
48 climate (Holtmeier & Broll, 2005; Du et al., 2018). Furthermore, treeline shifts give rise to novel
49 high-elevation vegetation patterns and could redefine habitable area for forest-dependent species
50 in a warmer future world (Bolton et al., 2018; Mohapatra et al., 2019). However, the treelines in
51 many mountain regions have been heavily altered by land-use change and land-use management
52 (Gehrig-Fasel et al., 2007; Ameztegui et al., 2016). Such land-use driven treelines are generally
53 lower than the elevation of the local theoretical climatic treelines, making it difficult to isolate
54 potential influences of climate on treeline position and obscuring the impact of climate change on
55 treeline shifts. Therefore, accurate and reproducible detection of natural mountain treelines and
56 their shifts are of great importance to understanding global climate change and the associated
57 response of vegetation dynamics in alpine areas in natural systems.

58
59 Previous studies reporting local treeline sites have mainly relied on field investigation (Wardle &
60 Coleman, 1992; Liang et al., 2014; Elliott & Cowell, 2015; Sigdel et al., 2018). While such studies
61 have enhanced our understanding of treeline patterns, a key limitation of field-based studies is
62 sparse geographic coverage. Remote sensing can overcome such a limitation by providing globally
63 consistent coverage, but the determination of treeline positions only through visually interpreting
64 satellite imagery (Paulsen & Körner, 2014; Irl et al., 2016; Karger et al., 2019) is time-consuming

65 and labour-intensive at large spatial scales. Recently, regional attempts to combine remote sensing
66 data with automated image processing techniques have emerged (Wei et al., 2020; Xu et al., 2020;
67 Wang et al., 2022; Birre et al., 2023), but inconsistent analytical approaches and treeline definitions
68 complicate regional comparisons and make it difficult to generalize global patterns. Early
69 assessment at the global scale suggested that low temperatures limited tree growth at treelines
70 (Körner & Paulsen, 2004), but there is also regional evidence that tree growth at the treeline does
71 not increase under global warming due to moisture limitations (Liang et al., 2014; Lyu et al., 2019;
72 Camarero et al., 2021). A generalizable pattern of the climatic limiting factors of global treelines
73 is still lacking.

74
75 The aforementioned challenges and limitations associated with delineating treelines and
76 determining climatic influences on treeline positions have hindered our understanding of the global
77 impact of climate on treelines in natural systems. To address this issue, we focused on “closed-
78 loop” mountain treelines (CLMT)—treelines with a continuous band of tree cover around a
79 mountain. Such systems are less likely to have been influenced by land-use change. By focusing
80 on this subset of treelines, we are better able to exclude treelines that may be impacted by
81 topographic constraints or anthropogenic land use in order to isolate the effects of climate on
82 mountain treelines in natural systems. An advance over previous studies that only provide a handful
83 of data points for each treeline is a complete depiction of treeline at 30 m resolution. Our approach
84 allows us to calculate the treeline elevation around the entire treeline, providing unprecedented
85 detail on the variability of treeline elevation at the local scale. More importantly, using CLMT as a
86 proxy for natural treelines with little influence from land-use change allows us to make a new and
87 more robust assessment of how natural treelines are responding to changes in climate.

88
89 Here, we map closed-loop treelines in mountain regions globally in 2000 based on remote sensing,
90 via integrating a high-resolution tree cover map (Hansen et al., 2013) with a digital elevation model

91 at the same spatial resolution (Tachikawa et al., 2011). For this purpose, we develop a novel
92 automatic detection algorithm that can produce consistent characterizations of CLMT across space.
93 Our detection of mountain treeline is based on tree cover data that consider trees as any vegetation
94 taller than 5 m (Hansen et al., 2013), using a 5% tree cover threshold to delineate forested and non-
95 forested areas. The algorithm starts from the highest elevation point for each mountain range and
96 generates a forest boundary map from which we extract the closed-loop treelines. To further ensure
97 that our CLMT are natural treelines that are not impacted by anthropogenic disturbances, we
98 conduct a manual inspection of high-resolution imagery to remove treelines with any indication of
99 anthropogenic land use and restrict our analysis to regions where the human footprint is low (Mu
100 et al., 2022). To understand which bioclimatic factors control the position of natural mountain
101 treelines from global to local scales, we use the gradient boosting decision trees (GBDT) model
102 (Friedman, 2001) to calculate the feature importance of each temperature or precipitation variable.
103 Further, we map the new natural treeline positions in 2010 using the same algorithm above and the
104 amount of tree cover in 2010 (Hansen et al., 2013) to explore the shifting of mountain treelines in
105 natural systems.

106

107 **2. Methods**

108 **2.1. Tree canopy cover data**

109 We used a high-resolution remote sensing global map of tree canopy cover for the year 2000
110 (Hansen et al., 2013) to delineate forested and non-forested areas. The dataset was produced at a
111 30 m resolution based on multiple types of forest sample data and spectral curves of Landsat time
112 series using a decision tree method (Hansen et al., 2013). To test which tree cover threshold is
113 suitable for treeline mapping, we undertook a sensitivity analysis with different thresholds in
114 mountains, finding there is little difference among different thresholds from 0 to 10% (examples
115 refer to Figs. S1–S3). Thus, we took the mean value of 0 to 10%, namely 5%, as the tree cover
116 threshold, and define the treeline to be the transition zone above which tree cover is $\leq 5\%$ and below

117 which tree cover is >5%. We then binary-classified the tree canopy cover data using the threshold,
118 assigning a value of 1 for the alpine land zone (the area above treeline) with tree cover $\leq 5\%$ (non-
119 forested area), and 0 for pixels with greater than 5% tree cover (forested area).

120

121 **2.2. Topography data**

122 We combined global mountain polygons with a high-resolution digital elevation model to restrict
123 the search area of mountain treelines. Mountain boundaries were collected from the Global
124 Mountain Biodiversity Assessment (GMBA) inventory (version 1.2; Körner et al., 2017). The
125 GMBA inventory delineated global mountains into discrete regions (polygons) based on
126 topographic ruggedness metrics and expert delineation (Körner et al., 2017). The elevation
127 information in mountains was provided by the Advanced Spaceborne Thermal Emission and
128 Reflection Radiometer Global Digital Elevation Model (version 3; Tachikawa et al., 2011) at a
129 spatial resolution of 30 m.

130

131 **2.3. Iterative mountain treeline extraction algorithm**

132 We developed an algorithm to automatically detect CLMT (Fig. S4). We first determined the
133 coordinates of the highest peak within each mountain region. The algorithm starts at this peak point
134 if it is within the alpine area that is non-forested, then expands outward (i.e., downslope), and
135 determines all other pixels of the image that are connected to the point and equivalent (marked as
136 “1”). The eight neighbourhood region of the pixel $I(x, y)$ is expressed as:

$$137 \quad R8 = \{(x + i, y + j); i, j \in (-1, 1)\} \quad (1)$$

138 where i, j are integers. In the collection of the eight neighbourhood pixels, if $I(x, y) = I(x + i, y +$
139 $j)$, there are connected relationships. The connected domain generated by this method is the
140 connected alpine area. Because the algorithm determines the starting search point, we marked only
141 one connected domain (namely the treeline zone) after one iteration.

142

143 To accelerate the efficiency of the algorithm, we set search blocks to determine the full altitudinal
144 range of treelines within mountain ranges (Fig. S4). Specifically, the first round of the search takes
145 the highest point of the mountain as the centre and the buffer zone with a side length of R as the
146 search area for the treeline. After testing, the square area with 8,000 rows/ranks (side length R
147 about 240 km) covered most alpine areas of mountains. For some of the mountaintops larger than
148 this range, we expanded the side length to ~720 km to ensure that all close-loop mountain treelines
149 of the world's mountaintops were covered.

150
151 There may be multiple treelines within a mountain range because a mountain may have multiple
152 peaks. To account for this, we next searched for the second highest starting point (i.e., the highest
153 point of the unsearched part) and repeated the process until the selected highest point was covered
154 by forests (tree cover >5%).

155
156 After each iteration, the loops that were determined to be “open” were removed. Focusing only on
157 closed treeline loops generated from the algorithm, we then visually inspected all loops using
158 Google Earth (with spatial resolution ranging from 15 m to ~15 cm) to further exclude treelines
159 with apparent signs of anthropogenic disturbances, such as roads, buildings, or croplands and
160 removed the part of water bodies (i.e., pixels that were determined to be water). Last, we filled all
161 the holes in the closed-loop polygons using the “imfill” function and extracted the edges of the
162 binary images using the “bwperim” function in Matlab R2019a to obtain the CLMT positions.

163
164 To validate the robustness of the elevational distribution of CLMT derived from satellite images,
165 at the pixel level, we used an independent validation dataset by manual interpretation using Google
166 Earth's high-resolution images. We randomly produced 100 validation samples at a spatial
167 resolution of 30 m. On a larger scale, we validated our CLMT database by comparison with *in situ*

168 measures ($n = 62$; Table S1). For each treeline site, we corresponded it to the closest treeline loop
169 detected in this study and compared its elevation with the range of the corresponding treeline loop.
170

171 **2.4. Climate data**

172 Considering the effect of climatic lag effects on treelines (Harsch et al., 2009), we used the climate
173 data from WorldClim (version 2.1; Fick and Hijmans, 2017), which provided the average for the
174 years 1970–2000 at a resolution of 30 seconds ($\sim 1 \text{ km}^2$), to understand which climate variables are
175 important in controlling treeline elevations. We used bioclimatic variables, which were derived
176 from monthly temperature and precipitation. A total of eight temperature variables and eight
177 precipitation variables were included, representing annual trends, seasonality, and extreme or
178 limiting environmental factors. They are annual mean temperature (annual T), temperature
179 seasonality (T seasonality; calculated as the standard deviation of the monthly mean temperatures,
180 then multiply by 100), the maximum temperature of the warmest month (maximum T), the
181 minimum temperature of the coldest month (minimum T), mean temperature of the wettest quarter
182 (wet season T), mean temperature of the driest quarter (dry season T), mean temperature of the
183 warmest quarter (warm season T), mean temperature of the coldest quarter (cold season T), annual
184 precipitation (annual P), precipitation of the wettest month (maximum P), precipitation of the driest
185 month (minimum P), precipitation seasonality (P seasonality; calculated as the coefficient of
186 variation, which is the ratio of the standard deviation to the mean), precipitation of the wettest
187 quarter (wet season P), precipitation of the driest quarter (dry season P), precipitation of the
188 warmest quarter (warm season P), and precipitation of the coldest quarter (cold season P). A
189 ‘quarter’ here refers to any consecutive three months. For example, the coldest quarter consists of
190 the three months that are colder than any other set of three consecutive months. For each pixel
191 determined to be on a CLMT, we extracted the values of all 16 climate variables.

192

193 **2.5. Gradient boosting decision trees (GBDT) model**

194 We applied a GBDT method to model the treeline elevation as a function of climate factors. The
195 GBDT model is a type of tree model with good interpretability for feature values (Friedman, 2001),
196 which assembles and iterates over multiple regression trees, with the values of the negative gradient
197 of the loss function in the model as an approximation of the residuals of the lifting tree algorithm
198 in the regression problem (Ke et al., 2017). It is flexible in handling large amounts of data and often
199 performs well in dealing with complex relationships in data (Ke et al., 2017). The GBDT initializes
200 a weak learner, estimating a constant value of the loss of function minimization, and then creates
201 decision trees according to the datasets and performs iterative training on them. Next, it calculates
202 the negative gradient for loss of function (residuals) corresponding to each tree, fits a regression
203 tree to the residuals to obtain the leaf node region of the m-th tree, and minimizes loss of function
204 by estimating the values of all leaf node regions using a linear search. Last, GBDT repeats the
205 above steps until the target evaluation indicator is optimal. Using this model, we calculated the
206 feature importance of each variable and determined the dependent correlations for each factor after
207 the model was built. The GBDT analysis was undertaken in Python 3.7 with the “sklearn.ensemble”
208 module.

209
210 We carried out the GBDT analyses at global and local scales, as well as separately for different
211 climatic belts (i.e., boreal, temperate, and tropical regions). At the global and regional scales, we
212 considered each treeline loop as a sample, namely, using the mean of treeline elevation in each loop
213 for the analysis. A total of 1,690 samples (treeline loops) were used for the global model. At the
214 local scale, we regarded one treeline pixel as a sample. Hence, in each treeline loop, the repeated
215 GBDT model represents the local effect of climate factors on treeline positions.

216 217 **2.6. Mountain treeline shift rate**

218 We mapped the new treeline positions in 2010 based on the global 2010 tree cover data (Hansen et
219 al., 2013; Potapov et al., 2015), which is an update of the 2000 tree cover product. Using this dataset,

220 we re-ran the algorithm around treelines to detect the new closed-loop treelines in 2010. Starting
221 from the highest elevation point we detected before, we expanded the rectangular area of the
222 original treeline around by 10 km as the search area. Then we manually checked the results from
223 the 1,690 treeline loops to (i) exclude treelines without closed loops; (ii) isolate examples of
224 “broken treeline loops” and restrict them to corresponding areas in 2000 and 2010 (Fig. S5); and
225 (iii) remove outliers (>95th percentile of both increasing and decreasing rates) to avoid the
226 inclusion of any special cases with extremely steep changes. This filtering resulted in 1,110 treeline
227 loops in 2010 (65.7% of all treelines initially assessed) being available for analysis of the treeline
228 change. The main reason for the reduction in number of treeline loops between 2000 and 2010 is
229 that some of the closed-loop treelines detected in 2000 did not form closed loops in 2010. We then
230 calculated the mean elevation of closed-loop treelines in 2010 and the corresponding treelines in
231 2000 and used the difference to represent the treeline change over the 10-year period. The treeline
232 shift rate (m/year) at each treeline loop was calculated as follows:

$$233 \quad \text{shift rate} = \frac{\text{mean elevation 2010} - \text{mean elevation 2000}}{10 \text{ yrs}} \quad (2)$$

234

235 **3. Results**

236 **3.1. A map of global closed-loop mountain treelines**

237 We detected 27,468,662 closed-loop mountain treeline positions (pixels at 30 m resolution) across
238 243 mountain ranges globally. The total length of CLMT we detected is ~916,425 km. Those
239 treeline pixels form 1,690 treeline loops covering all continents except Antarctica, ranging from
240 64°N (Khrebet Polyarnyy, Russia) to 46°S (Princess Mountains, New Zealand), with mean
241 elevations spanning from 489 ±283 m on Khrebet Chayatyn (Russia) to 4,528 ±104 m on
242 Ruwenzori (Uganda, Kenya). The average length of these treeline loops is 542 km, and the average
243 alpine land area above them is 142 km². To visualize global patterns of the elevation of CLMT, we
244 calculated the mean elevation for each treeline loop and plotted their locations using the mean
245 latitude and longitude of treeline pixels at 30 m resolution in each loop (Fig. 1a). The CLMT

246 derived from satellite tree cover data are consistent with fine resolution remote sensing images
247 available on Google Earth (Fig. 1b-g). At the pixel level, the CLMT showed good agreement with
248 manually interpreted data at 30 m resolution ($R^2 = 0.96$; Fig. S6). On a larger scale, the validity of
249 our CLMT database was further supported by corroboration against *in situ* measures from previous
250 studies ($n = 62$ measurements; Table S1), which fall within the elevation range of treeline loops
251 ($R^2 = 0.98$; Fig. S7).

252
253 We found a bimodal pattern for the closed-loop mountain treeline elevation along latitude, with
254 peaks at the equator and $\sim 25^\circ\text{N}$ (Fig. 2a). Between 0° and 10° , the elevation of CLMT is
255 symmetrical in the northern and southern hemispheres, but beyond this range, treeline elevations
256 in the northern hemisphere are higher than those in the southern hemisphere at equivalent latitudes
257 (Fig. 2a), which is attributed to the oceanic influence on a smaller southern landmass (Cieraad et
258 al., 2014). Our global CLMT distribution is consistent with previous global assessments, though
259 there are some differences. In the tropics, the elevation of CLMT reaches up to 3,500 m (Fig. 2), a
260 lower elevation than in a recent global assessment by Testolin et al. (2020) that reported tropical
261 treelines higher than 4000 m. This discrepancy may be due to our strict definition of trees, >5 m
262 height, as well as the exclusion of some unilateral and non-closed treelines in high mountains. At
263 low latitude (especially at 0 - 20°N), there is large variation in the range of CLMT elevation (Fig.
264 2a). Among different continents, South America has a large CLMT elevation range variation. At
265 50°N - 60°N and 20°N - 30°N , many mountains in Asia and North America have similar treeline
266 elevations, whereas there is a rather different behaviour at 30°N - 50°N where treelines in North
267 America are higher than those in Europe and Asia (Fig. 2a). To help understand what causes this
268 behaviour, we calculated the distance to the coast for each treeline. We found lower treelines in
269 coastal mountains at the same latitude (Fig. 2a) as has been suggested in the literature (Irl et al.,
270 2016), which can be largely attributed to the thermo-dynamic effect of large high-elevation
271 landmasses (Karger et al., 2019). At 30°N - 60°N , mountains close to the coast have lower treelines

272 than their latitude might suggest (i.e., fall below the fitted curve; Fig. 2a). Similarly, along with
273 longitude decreasing from 150°W to 100°W, treeline elevations in North America increase due to
274 an increase in the distance to the coast (Fig. 2b).

275

276 **3.2. Climatic determinants of closed-loop mountain treelines**

277 We found that T seasonality, cold season P, and warm season T predict nearly 60% of the spatial
278 distribution of CLMT globally (Fig. 3a). We then assessed how the three leading factors modulated
279 the elevation of CLMT spatially. The results showed the abrupt transition of CLMT elevation
280 occurring at the T seasonality threshold of ~9°C, but attenuated transitions in areas where T
281 seasonality exceeded 10°C (Fig. S8a). Similarly, there is a CLMT elevation gradient that is spatially
282 driven by cold season P, with abrupt transitions occurring at the thresholds of 320 mm and 450 mm
283 along the gradient of cold season P (Fig. S8b). By contrast, we did not find such a dramatic
284 transition of CLMT elevation along the warm season T gradient (Fig. S8c).

285

286 Collectively, temperature-related factors (64%) are more important than precipitation-related
287 factors for limiting CLMT elevations on a global scale (Fig. 3a). In different latitudinal belts,
288 temperature-related factors are most important in boreal and tropical regions, especially the
289 temperature of the warmest and the wettest quarters, respectively, while precipitation dominates
290 the CLMT elevation in temperate regions (Fig. 3b,c,d). We found that T seasonality is the most
291 important individual factor (30%) at global scale, whereas its importance is lower than 10% for
292 boreal and tropical regions (Fig. 3). These patterns may be because thermal limitation to growth at
293 treelines during the summer is most critical in the cold boreal regions, while in the tropics where
294 temperature is high throughout the year, temperature of the wettest season plays a key role in
295 limiting tree growth at treelines. Our results confirm the importance of temperature during the
296 warm part of the year in the boreal zone (Jobbágy & Jackson, 2000), but suggest that precipitation
297 is more important than temperature in temperate regions. It agrees with climatic sensitivity of tree

298 growth in the Northern Hemisphere (Gao et al., 2022). Especially under dry environmental
299 conditions, moisture availability is crucial to limiting tree growth in the treeline ecotone (Liang et
300 al., 2014; Ren et al., 2018).

301
302 Our study provides vastly more data points for each treeline compared to previous global
303 assessments (Jobbágy & Jackson, 2000; Körner & Paulsen, 2004), allowing us to explore for the
304 first time what controls treeline position at a local scale. We found that temperature remains the
305 dominant explanation for the altitudinal variation of 76% of the treeline within a single treeline
306 loop with similar climatic conditions (Fig. S9).

307

308 **3.3. Shifts in closed-loop mountain treelines**

309 Between 2000 and 2010, mountain treelines have shifted upwards at 777 out of the 1,110 treeline
310 loops (70%) and downward at 333 treeline loops (Fig. 4a). The mean global treeline shift rate was
311 an upward shift of 1.2 m/year, which is consistent with case studies of treeline change, with rates >1
312 m/year reported in the literature (Table S2). A synthesis of treeline shift rates reported in the
313 literature suggests the rate was 0.67 m/year before 1970 compared to 4.36 m/year after 1970 and
314 6.16 m/year after 2000 (Fig. S10; Table S2). This provides evidence that the rate of change in
315 treeline elevation is accelerating, possibly due to recent rapid climate change (Bolton et al., 2018).
316 Treeline shift rates in the tropics (mean of 3.1 m/year) were higher than those in boreal and
317 temperate regions (Fig. 4b). The faster changes in the tropics could be related to hydrothermal
318 conditions: in the tropics, higher temperature and more abundant precipitation bring a longer
319 growing season, which naturally favours the growth of seedlings and young trees. By contrast, there
320 is a slight downward shift in temperate regions (an average of -0.5 m/year), where the position of
321 the treeline is dominated by precipitation (Fig. 3c). This could be due to decreasing precipitation
322 in some mountain areas of the temperate zone, for example in northern China (Piao et al., 2010).

323

324 Although the tropical CLMT have the fastest shift rates, their variability is the largest, ranging from
325 -10.2 to 16.9 m/year (Fig. 4b). In the tropics, treeline shift rates greater than 10 m/year in the
326 mountains of Malawi, Papua New Guinea, and Indonesia may reflect a more extreme trend in these
327 tropical systems. In other regions, there are also some treelines that have shifted much more than
328 expected (>10 m/year; Fig. 4b): in boreal regions, these expectations are mainly in Russia and
329 Canada; in temperate regions, they are geographically concentrated in East Asia (North Korea,
330 Japan, and China). On the contrary, there are also cases of treelines receding at a high rate, possibly
331 driven by fire in some areas, either through the physical destruction of trees that acts to lower the
332 existing treelines, or through the destruction of seedlings established upslope that acts to prevent
333 treeline advances (Kim & Lee, 2015). For example, treelines have significantly receded in the
334 western USA where climate and vegetation are favourable for fire (Seven Devils Mountains, Swan
335 Range, etc.; Fig. 4a).

336
337 In addition, independent analysis for the changes in annual maximum Normalized Difference
338 Vegetation Index (NDVI) at CLMT that we identified for the year 2000 shows the NDVI has
339 significantly increased by 3.3% by 2020, at a rate of 0.0012 per year ($P < 0.01$; Fig. S11a). There
340 are significant positive trends in NDVI at treeline zones in boreal, temperate, and tropical regions
341 during 2000-2020 ($P < 0.01$), and tropical areas have the highest rate, approaching 0.0016 per year
342 (Fig. S11b). The increase in NDVI occurred at most treeline zones (~90%; Fig. S11c). This
343 greening at the treeline may also be conducive to upward movement of the treeline in the future.

344

345 **4. Discussion**

346 **4.1. Comparison of treeline datasets before and after considering human footprint**

347 Although we have examined CLMT by manual interpretation to remove anthropogenic treelines,
348 we further conduct a stricter assessment of human pressures to check whether our results would
349 still be impacted by human activity. We used a global Human Footprint dataset (Mu et al., 2022)

350 and found 83% of our CLMT in wilderness (Human Footprint < 1) or in highly intact areas (Human
351 Footprint <4). We then removed those treelines with human footprint values ≥ 4 , re-ran the analysis
352 with the higher human footprint values excluded, and updated all the results above (Figs. S12-S14).
353 By comparing these new results with those using the whole dataset, we found a similar pattern
354 along latitude and longitude gradients (Figs. 2 and S12). The results regarding climate dominants
355 (Figs. 3 and S13) and treeline shift rates (Figs. 4b and S14) were also consistent using either
356 approach. Thus, the additional criterion to further focus our analysis on treelines with no human
357 disturbance does not alter our overall results or conclusions, and further confirms that our CLMT
358 product can well represent the change and pattern of climatic treelines.

359

360 **4.2. Implications of treeline shifts for carbon, biodiversity, and hydrology**

361 Changing treeline position can affect the carbon cycle, biodiversity, and hydrological processes in
362 mountain environments. Mountain treelines moving upward to higher elevations increase woody
363 biomass at and above the treeline, accumulating carbon and increasing their ability to act as carbon
364 sinks (Lopatin et al., 2006; Tarnocai et al., 2009). However, such increases may be offset by
365 increases in soil respiration, leading to a net loss of ecosystem carbon (Wilmking et al., 2006;
366 Hartley et al., 2012). The ascent of mountain treelines also substantially influences biodiversity
367 patterns at high elevations, with enhanced habitat loss of endemic alpine species within a narrow
368 range of mountains (Wang et al., 2022) and potential expansion of habitat for forest-dependent
369 species whose upper range limits coincide with the treeline ecotones (Elsen et al., 2017). For alpine
370 species isolated at the top of mountains, upward treeline shifts could increase the risk of extinction,
371 where there is not enough room for the alpine zone to move upward under future climate change
372 (Dirnböck et al., 2011). In Siberia, for example, we show many treelines have shifted upwards (Fig.
373 4b), inevitably reducing the area of the tundra, which is rich in floristic and species diversity and
374 supports indigenous land use types. The expansion of Siberian forests has been predicted to
375 continue, thus causing huge losses of tundra in the future (Kruse & Herzsuh, 2022). While we

376 focused here on treeline shifts in areas with minimal human impacts, treeline ascent in areas with
377 pronounced human disturbance will further hinder species' ability to track vegetation changes and
378 likely lead to more pronounced population declines (Feeley & Silman, 2010; Elsen et al., 2020).
379 There are many instances with high pressure in high-elevation areas, especially from burning,
380 grazing, and wood harvesting (Bader et al., 2008; Jiménez-García et al., 2021). The combined
381 impact of shifting treelines and human disturbances may also affect local livelihoods and act as a
382 double-blow for sensitive alpine species. In addition, tree expansions into the formerly treeless area
383 may alter downstream water supply. Recent advances of the treeline have decreased the area of
384 alpine tundra, thereby affecting its critical role as a reservoir of freshwater resources and in water
385 release (Barredo et al., 2020).

386

387 **4.3. Uncertainties and caveats**

388 To isolate the impacts of climate on treelines, our analysis identifies CLMT that completely
389 encircle a mountain. However, focusing on this kind of treelines could omit some climate-related
390 treelines as climatic treelines may not be in a closed loop shape in some cases. We acknowledge
391 that our CLMT database does not include all climatic treelines, but is a subset of climatic treelines
392 that specifically form a closed loop, because these enable us to analyse climatic determinants with
393 greater confidence. We also note that tree cover can increase in various ways, either through new
394 or existing trees growing above the 5 m height threshold, or existing trees having increased canopy
395 cover. However, our analysis is based on the definition of treeline according to remotely sensed
396 tree cover, and we used this definition to assess treeline position at two time periods and assess
397 change. While our analysis period is short and errors will exist at a pixel scale, our global detection
398 of a shifting treeline provides an early indication of climate-induced changes that need to be
399 carefully monitored in the future. To reduce uncertainties and further advance our understanding
400 of treeline dynamics, future studies require more high-resolution remote sensing products for a
401 longer period and more field data in alpine treeline zones for cross-validation.

402

403 **5. Conclusion**

404 Our study develops a novel remote sensing-based algorithm to map closed-loop treelines across
405 global mountain regions, isolating the effects of climate on treeline position. Our approach provides
406 a globally consistent way of detecting and monitoring closed-loop treelines around mountains,
407 which are more likely to reflect natural systems with minimal impact of land-use change. Focusing
408 on these closed-loop treelines as a proxy for natural treelines allows us to isolate the impacts of
409 climate and climate change on the elevation distribution and change of treelines. We found
410 temperature was the dominant control on natural treelines both at a global and local scale. Our
411 results indicated an upward migration of treelines over the period 2000 to 2010 in boreal and
412 tropical regions but a slight downward shift in temperate zones. Our new findings and the global
413 closed-loop mountain treeline database produced in this study also provides a useful tool for
414 biodiversity and carbon assessments, ecological modelling, and analyses of adaptation of species
415 to future climate change.

416

417 **Acknowledgments**

418 X.H. was funded by a PhD scholarship from Southern University of Science and Technology hosted
419 jointly with the University of Leeds. Z.Z. was supported by the National Natural Science
420 Foundation of China (no. 42071022) and the start-up fund provided by Southern University of
421 Science and Technology (no. 29/Y01296122). We thank Hansen/UMD/Google/USGS/NASA for
422 providing the high-resolution tree cover data; NASA and Japan's Ministry of Economy, Trade and
423 Industry for providing the elevation data; Körner for providing the GMBA inventory; Fick and
424 Hijmans for providing WorldClim data; and Google Earth for providing very-high-resolution
425 satellite imagery.

426

427 **Author contributions**

428 X.H., Z.Z., D.V.S. and J.H. designed the research; X.H. performed the analysis and wrote the draft;
429 and all the authors contributed to the interpretation of the results and the writing of the paper.

430

431 **Data availability statement**

432 The global tree cover data in 2000 are available at [https://earthenginepartners.appspot.com/science-](https://earthenginepartners.appspot.com/science-2013-global-forest/download_v1.7.html)
433 [2013-global-forest/download_v1.7.html](https://earthenginepartners.appspot.com/science-2013-global-forest/download_v1.7.html) and the tree cover data in 2010 are available at
434 https://glad.umd.edu/Potapov/TCC_2010/. The ASTER elevation data are available at
435 <https://earthdata.nasa.gov/>. The GMBA inventory is available at
436 https://ilias.unibe.ch/goto_ilias3_unibe_cat_1000515.html. The WorldClim climate data are
437 available at <https://www.worldclim.org/data/worldclim21.html>. The MODIS combined 16-day
438 NDVI is available at [https://developers.google.com/earth-](https://developers.google.com/earth-engine/datasets/catalog/MODIS_MCD43A4_006_NDVI)
439 [engine/datasets/catalog/MODIS_MCD43A4_006_NDVI](https://developers.google.com/earth-engine/datasets/catalog/MODIS_MCD43A4_006_NDVI). The scripts used to generate the
440 results/figures are available in GitHub at <https://github.com/hexinyue33/treeline>. The global
441 closed-loop mountain treeline (CLMT) database developed in this study can be accessed through
442 <https://hexinyue33.users.earthengine.app/view/clmt>.

443 **References**

- 444 Ameztegui, A., Coll, L., Brotons, L., & Ninot, J.M. (2016). Land-use legacies rather than climate
445 change are driving the recent upward shift of the mountain tree line in the Pyrenees. *Global*
446 *Ecology and Biogeography*, 25(3), 263-273. <https://doi.org/10.1111/geb.12407>
- 447 Bader, M.Y., Rietkerk, M., & Bregt, A.K. (2008). A simple spatial model exploring positive
448 feedbacks at tropical alpine treelines. *Arctic, Antarctic, and Alpine Research*, 40(2), 269-278.
449 [https://doi.org/10.1657/1523-0430\(07-024\)\[BADER\]2.0.CO;2](https://doi.org/10.1657/1523-0430(07-024)[BADER]2.0.CO;2)
- 450 Barredo, J. I., Mauri, A., & Caudullo, G. (2020). Alpine tundra contraction under future warming
451 scenarios in Europe. *Atmosphere*, 11(7), 698. <https://doi.org/10.3390/atmos11070698>
- 452 Birre, D., Feuillet, T., Lagalis, R., Milian, J., Alexandre, F., Sheeren, D., Serrano-Notivoli, R.,
453 Vignal, M., & Bader, M.Y., (2023). A new method for quantifying treeline-ecotone change
454 based on multiple spatial pattern dimensions. *Landscape Ecology*, 1-18.
455 <https://doi.org/10.1007/s10980-022-01589-4>
- 456 Bolton, D.K., Coops, N.C., Hermosilla, T., Wulder, M.A., & White, J.C. (2018). Evidence of
457 vegetation greening at alpine treeline ecotones: three decades of Landsat spectral trends
458 informed by lidar-derived vertical structure. *Environmental Research Letters*, 13(8), 084022.
459 <https://doi.org/10.1088/1748-9326/aad5d2>
- 460 Camarero, J.J., Gazol, A., Sánchez-Salguero, R., Fajardo, A., McIntire E J B, Gutiérrez E, Batllori
461 E, Boudreau S, Carrer M, Diez J, Dufour-Tremblay G, Gaire NP, Hofgaard A, Jomelli V,
462 Kirilyanov AV, Lévesque E, Liang E, Linares JC, Mathisen IE, Moiseev PA, Sangüesa-
463 Barreda G, Shrestha KB, Toivonen JM, Tutubalina OV, & Wilmking M. (2021). Global
464 fading of the temperature-growth coupling at alpine and polar treelines. *Global Change*
465 *Biology*, 27(9), 1879-1889. <https://doi.org/10.1111/gcb.15530>
- 466 Cieraad, E., McGlone, M. S., & Huntley, B. (2014). Southern Hemisphere temperate tree lines are
467 not climatically depressed. *Journal of Biogeography*, 41(8), 1456-1466.
468 <https://doi.org/10.1111/jbi.12308>

469 Dirnböck, T., Essl, F., & Rabitsch, W. (2011). Disproportional risk for habitat loss of high-altitude
470 endemic species under climate change. *Global Change Biology*, 17(2), 990-996.
471 <https://doi.org/10.1111/j.1365-2486.2010.02266.x>

472 Du, H., Liu, J., Li, M.-H., Büntgen, U., Yang, Y., Wang, L., Wu, Z., & He, H.S. (2018). Warming-
473 induced upward migration of the alpine treeline in the Changbai Mountains, northeast China.
474 *Global Change Biology*, 24(3), 1256-1266. <https://doi.org/10.1111/gcb.13963>

475 Elliott, G.P., & Cowell, C.M. (2015). Slope aspect mediates fine-scale tree establishment patterns
476 at upper treeline during wet and dry periods of the 20th century. *Arctic, Antarctic, and Alpine*
477 *Research*, 47(4), 681-692. <https://doi.org/10.1657/AAAR0014-025>

478 Elsen, P. R., Tingley, M. W., Kalyanaraman, R., Ramesh, K., & Wilcove, D. S. (2017). The role
479 of competition, ecotones, and temperature in the elevational distribution of Himalayan birds.
480 *Ecology*, 98(2), 337-348. <https://doi.org/10.1002/ecy.1669>

481 Elsen, P. R., Monahan, W. B., & Merenlender, A. M. (2020). Topography and human pressure in
482 mountain ranges alter expected species responses to climate change. *Nature Communications*,
483 11(1), 1-10. <https://doi.org/10.1038/s41467-020-15881-x>

484 Feeley, K. J., & Silman, M. R. (2010). Land-use and climate change effects on population size and
485 extinction risk of Andean plants. *Global Change Biology*, 16(12), 3215-3222.
486 <https://doi.org/10.1111/j.1365-2486.2010.02197.x>

487 Fick, S.E., & Hijmans, R.J. (2017). WorldClim 2: new 1-km spatial resolution climate surfaces for
488 global land areas. *International Journal of Climatology*, 37(12), 4302-4315.
489 <https://doi.org/10.1002/joc.5086>

490 Friedman, J.H. (2001). Greedy function approximation: a gradient boosting machine. *Annals of*
491 *Statistics*, 29, 1189-1232.

492 Gao, S., Liang, E., Liu, R., Babst, F., Camarero, J.J., Fu, Y.H., Piao, S., Rossi, S., Shen, M., Wang,
493 T., & Peñuelas, J. (2022). An earlier start of the thermal growing season enhances tree growth

494 in cold humid areas but not in dry areas. *Nature Ecology & Evolution*, 6(4), 397-404.
495 <https://doi.org/10.1038/s41559-022-01668-4>

496 Gehrig-Fasel, J., Guisan, A., & Zimmermann, N.E. (2007). Tree line shifts in the Swiss Alps:
497 Climate change or land abandonment? *Journal of Vegetation Science*, 18, 571-582.
498 <https://doi.org/10.1111/j.1654-1103.2007.tb02571.x>

499 Grace, J. (1989). Tree lines. *Philosophical Transactions of the Royal Society of London. B,*
500 *Biological Sciences*, 324(1223), 233-245.

501 Hansen, M.C., Potapov, P.V., Moore, R., Hancher, M., Turubanova, S.A., Tyukavina, A., Thau,
502 D., Stehman, S.V., Goetz, S.J., & Loveland, T.R. (2013). High-resolution global maps of 21st-
503 century forest cover change. *Science*, 342(6160), 850-853.
504 <https://doi.org/10.1126/science.1244693>

505 Harsch, M. A., Hulme, P. E., McGlone, M. S., & Duncan, R. P. (2009). Are treelines advancing?
506 A global meta-analysis of treeline response to climate warming. *Ecology Letters*, 12(10),
507 1040-1049. <https://doi.org/10.1111/j.1461-0248.2009.01355.x>

508 Hartley, I.P., Garnett, M.H., Sommerkorn, M., Hopkins, D.W., Fletcher, B.J., Sloan, V. L.,
509 Phoenix, G.K., & Wookey, P.A. (2012). A potential loss of carbon associated with greater
510 plant growth in the European Arctic. *Nature Climate Change*, 2(12), 875-879.
511 <https://doi.org/10.1038/nclimate1575>

512 Holtmeier, F. & Broll, G. (2005). Sensitivity and response of northern hemisphere altitudinal and
513 polar treelines to environmental change at landscape and local scales. *Global Ecology*
514 *Biogeography*, 14(5), 395-410. <https://doi.org/10.1111/j.1466-822X.2005.00168.x>

515 Irl, S.D., Anthelme, F., Harter, D.E., Jentsch, A., Lotter, E., Steinbauer, M.J., & Beierkuhnlein, C.
516 (2016). Patterns of island treeline elevation—a global perspective. *Ecography*, 39(5), 427-436.
517 <https://doi.org/10.1111/ecog.01266>

518 Jiménez-García, D., Li, X., Lira-Noriega, A., & Peterson, A.T. (2021). Upward shifts in elevational
519 limits of forest and grassland for Mexican volcanoes over three decades. *Biotropica*, 53, 798-
520 807. <https://doi.org/10.1111/btp.12942>

521 Jobbágy, E.G. & Jackson, R.B. (2000). Global controls of forest line elevation in the northern and
522 southern hemispheres. *Global Ecology Biogeography*, 9(3), 253-268.
523 <https://doi.org/10.1046/j.1365-2699.2000.00162.x>

524 Karger, D.N., Kessler, M., Conrad, O., Weigelt, P., Kreft, H., König, C., & Zimmermann, N.E.
525 (2019). Why tree lines are lower on islands—Climatic and biogeographic effects hold the
526 answer. *Global Ecology Biogeography*, 28(6), 839-850. <https://doi.org/10.1111/geb.12897>

527 Ke, G., Meng, Q., Finley, T., Wang, T., Chen, W., Ma, W., Ye, Q., & Liu, T.-Y. (2017). Lightgbm:
528 A highly efficient gradient boosting decision tree. *Advances in Neural Information Processing*
529 *Systems*, 30, 3146-3154.

530 Kim, J.W., & Lee, J.S. (2015). Dynamics of alpine treelines: positive feedbacks and global,
531 regional and local controls. *Journal of Ecology and Environment*, 38(1), 1-14.
532 <https://doi.org/10.5141/ecoenv.2015.001>

533 Körner, C. & Paulsen, J. (2004). A world-wide study of high altitude treeline temperatures. *Journal*
534 *of Biogeography*, 31(5), 713-732. <https://doi.org/10.1111/j.1365-2699.2003.01043.x>

535 Körner, C., Jetz, W., Paulsen, J., Payne, D., Rudmann-Maurer, K., & M. Spehn, E. (2017). A global
536 inventory of mountains for bio-geographical applications. *Alpine Botany*, 127(1), 1-15.
537 <https://doi.org/10.1007/s00035-016-0182-6>

538 Kruse, S., & Herzsuh, U. (2022). Regional opportunities for tundra conservation in the next 1000
539 years. *eLife*, 11, e75163. <https://doi.org/10.7554/eLife.75163>

540 Liang, E., Dawadi, B., Pederson, N., & Eckstein, D. (2014). Is the growth of birch at the upper
541 timberline in the Himalayas limited by moisture or by temperature? *Ecology*, 95(9), 2453-
542 2465. <https://doi.org/10.1890/13-1904.1>

543 Lyu, L., Zhang, Q.B., Pellatt, M.G., Büntgen, U., Li, M.H., & Cherubini, P. (2019). Drought
544 limitation on tree growth at the Northern Hemisphere's highest tree line. *Dendrochronologia*,
545 53, 40-47. <https://doi.org/10.1016/j.dendro.2018.11.006>

546 Lopatin, E., Kolström, T., & Spiecker, H. (2006). Determination of forest growth trends in Komi
547 Republic (northwestern Russia): combination of tree-ring analysis and remote sensing data.
548 *Boreal Environment Research*, 11, 341-353.

549 Lu, X., Liang, E., Wang, Y., Babst, F., & Camarero, J.J. (2021). Mountain treelines climb slowly
550 despite rapid climate warming. *Global Ecology and Biogeography*, 30(1), 305-315.
551 <https://doi.org/10.1111/geb.13214>

552 Mohapatra, J., Singh, C.P., Tripathi, O.P., & Pandya, H.A. (2019). Remote sensing of alpine
553 treeline ecotone dynamics and phenology in Arunachal Pradesh Himalaya. *International*
554 *Journal of Remote Sensing*, 40(20), 7986-8009.
555 <https://doi.org/10.1080/01431161.2019.1608383>

556 Mu, H., Li, X., Wen, Y., Huang, J., Du, P., Su, W., Miao, S., & Geng, M. (2022). A global record
557 of annual terrestrial Human Footprint dataset from 2000 to 2018. *Scientific Data*, 9(1), 176.
558 <https://doi.org/10.1038/s41597-022-01284-8>

559 Paulsen, J. & Körner, C. (2014) A climate-based model to predict potential treeline position around
560 the globe. *Alpine Botany*, 124(1), 1-12. <https://doi.org/10.1007/s00035-014-0124-0>

561 Potapov, P., Turubanova, S., Tyukavina, A., Krylov, A., McCarty, J., Radeloff, V., & Hansen, M.
562 (2015). Eastern Europe's forest cover dynamics from 1985 to 2012 quantified from the full
563 Landsat archive: *Remote Sensing of Environment*, 159, 28-43.
564 <https://doi.org/10.1016/j.rse.2014.11.027>

565 Piao, S., Ciais, P., Huang, Y., Shen, Z., Peng, S., Li, J., Zhou, L., Liu, H., Ma, Y., Ding, Y. and
566 Friedlingstein, P. (2010). The impacts of climate change on water resources and agriculture in
567 China. *Nature*, 467(7311), 43-51. <https://doi.org/10.1038/nature09364>

- 568 Ren, P., Rossi, S., Camarero, J.J., Ellison, A.M., Liang, E., & Peñuelas, J. (2018). Critical
569 temperature and precipitation thresholds for the onset of xylogenesis of *Juniperus przewalskii*
570 in a semi-arid area of the north-eastern Tibetan Plateau. *Annals of Botany*, 121(4), 617-624.
571 <https://doi.org/10.1093/aob/mcx188>
- 572 Sigdel, S.R., Wang, Y., Camarero, J.J., Zhu, H., Liang, E., & Peñuelas, J. (2018). Moisture-
573 mediated responsiveness of treeline shifts to global warming in the Himalayas. *Global Change*
574 *Biology*, 24(11), 5549-5559. <https://doi.org/10.1111/gcb.14428>
- 575 Tachikawa, T., Hato, M., Kaku, M., & Iwasaki, A. (2011). Characteristics of ASTER GDEM
576 version 2. In: 2011 IEEE International Geoscience and Remote Sensing Symposium: IEEE,
577 3657-3660. <https://doi.org/10.1109/IGARSS.2011.6050017>
- 578 Tarnocai, C., Canadell, J.G., Schuur, E.A.G., Kuhry, P., Mazhitova, G., & Zimov, S. (2009). Soil
579 organic carbon pools in the northern circumpolar permafrost region. *Global Biogeochemical*
580 *Cycles*, 23, GB2023. <https://doi.org/10.1029/2008GB003327>
- 581 Testolin, R., Attorre, F., & Jiménez-Alfaro, B. (2020). Global distribution and bioclimatic
582 characterization of alpine biomes. *Ecography*, 43(6), 779-788.
583 <https://doi.org/10.1111/ecog.05012>
- 584 Verrall, B. & Pickering, C.M. (2020). Alpine vegetation in the context of climate change: A global
585 review of past research and future directions. *Science of the Total Environment*, 748, 141344.
586 <https://doi.org/10.1016/j.scitotenv.2020.141344>
- 587 Wang, X., Wang, T., Xu, J., Shen, Z., Yang, Y., Chen, A., Wang, S., Liang, E., & Piao, S. (2022).
588 Enhanced habitat loss of the Himalayan endemic flora driven by warming-forced upslope tree
589 expansion. *Nature Ecology & Evolution*, 6(7), 890-899. [https://doi.org/10.1038/s41559-022-](https://doi.org/10.1038/s41559-022-01774-3)
590 [01774-3](https://doi.org/10.1038/s41559-022-01774-3)
- 591 Wardle, P. & Coleman, M. (1992). Evidence for rising upper limits of four native New Zealand
592 forest trees. *New Zealand Journal of Botany*, 30(3), 303-314.
593 <https://doi.org/10.1080/0028825X.1992.10412909>

594 Wei, C., Karger, D.N., & Wilson, A.M. Spatial detection of alpine treeline ecotones in the Western
595 United States. *Remote Sensing of Environment*, 2020, 240, 111672.
596 <https://doi.org/10.1016/j.rse.2020.111672>

597 Wilmking, M., Harden, J., & Tape, K. (2006). Effect of tree line advance on carbon storage in NW
598 Alaska. *Journal of Geophysical Research*, 111, 1-10. <https://doi.org/10.1029/2005JG000074>

599 Xu, D., Geng, Q., Jin, C., Xu, Z., & Xu, X. (2020). Tree Line Identification and Dynamics under
600 Climate Change in Wuyishan National Park Based on Landsat Images. *Remote Sensing*,
601 12(18), 2890. <https://doi.org/10.3390/rs12182890>

602

603 **Figure Legends**

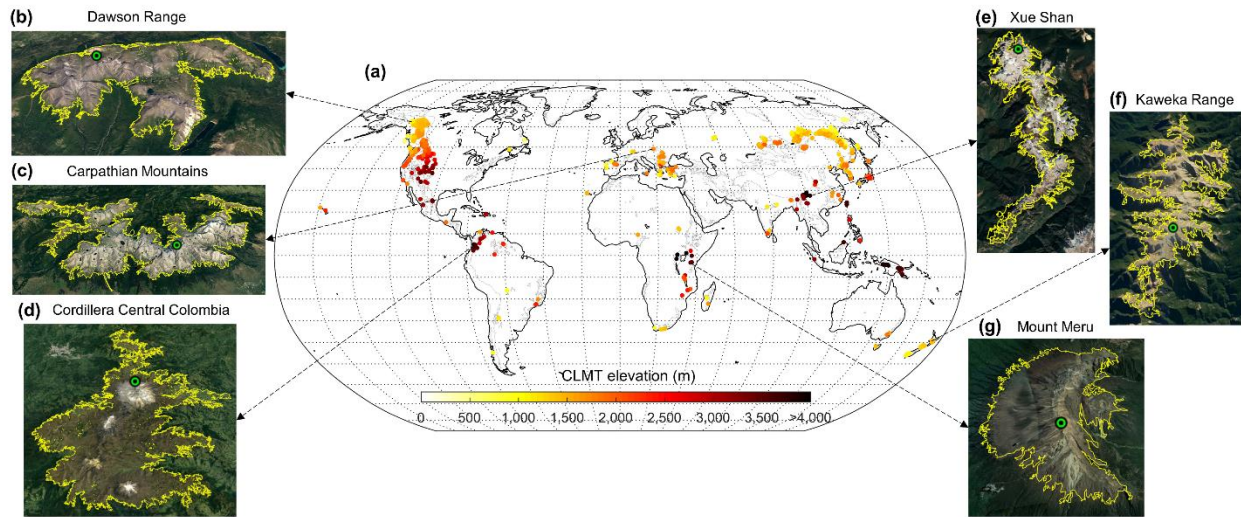
604 **Figure 1. Global distribution of closed-loop mountain treeline (CLMT) elevation.** To improve
605 readability, plot (a) is based on the mean value of each closed-loop mountain treeline (at each 30-
606 m pixel). Grey boundaries indicate mountain regions defined by GMBA inventory data. (b)–(g)
607 show examples of CLMT extraction results superimposed with Google Earth images. The yellow
608 line represents the position of the treeline, and the green circle shows the highest elevation point
609 that formed the starting point of each search by the treeline algorithm.

610
611 **Figure 2. Global latitudinal (a) and longitudinal (b) variation of closed-loop mountain treeline**
612 **(CLMT) elevation.** Different symbols represent different regions and colours represent the
613 distance to the coast. The data points show the mean elevation of all of the pixels in the CLMT.
614 The error bar is the elevation range of the corresponding treeline loop.

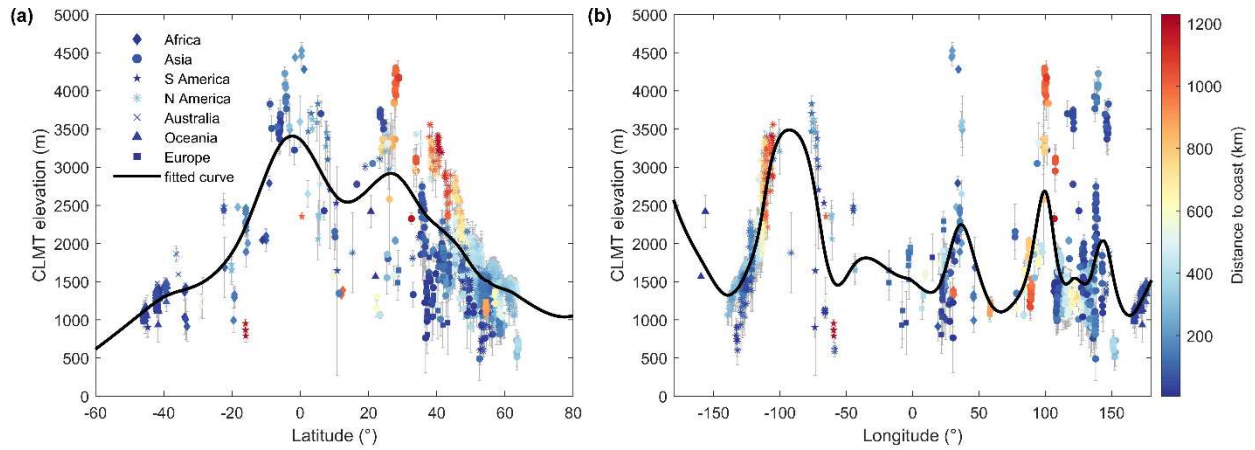
615
616 **Figure 3. Climate drivers controlling the variability in treeline elevation for the globe (a),**
617 **boreal ($\geq 50^\circ\text{N}$, b), temperate ($23.5^\circ - 50^\circ\text{N/S}$, c) and tropical ($23.5^\circ\text{N} - 23.5^\circ\text{S}$, d) regions.**

618
619 **Figure 4. Closed-loop mountain treeline (CLMT) shift rate during 2000-2010.** (a), Spatial
620 pattern of CLMT shift rate. (b), Box-plot showing CLMT shift rate in boreal ($\geq 50^\circ\text{N}$), temperate
621 ($23.5^\circ - 50^\circ\text{N/S}$) and tropical ($23.5^\circ\text{N} - 23.5^\circ\text{S}$) regions (central line: median; red dot: mean; box:
622 25th and 75th percentiles, respectively; error bar: maximum and minimum whisker values; +:
623 maximum and minimum values). The black dashed line is the zero line. Numbers of the studied
624 CLMT are shown above the boxes.

625



626
 627 **Figure 1. Global distribution of closed-loop mountain treeline (CLMT) elevation.** To improve
 628 readability, plot (a) is based on the mean value of each closed-loop mountain treeline (at each 30-
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 632 that formed the starting point of each search by the treeline algorithm.
 633



634

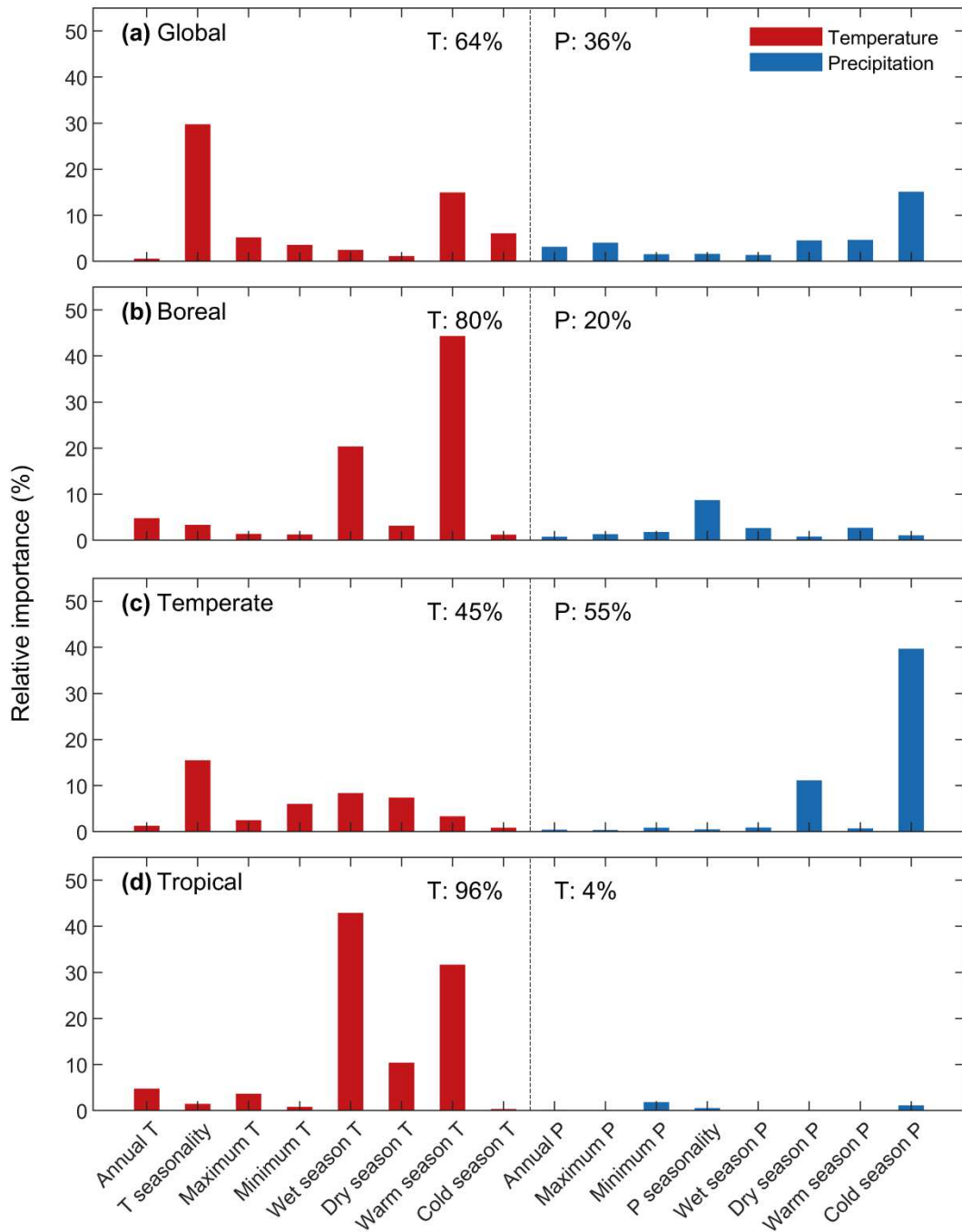
635 **Figure 2. Global latitudinal (a) and longitudinal (b) variation of closed-loop mountain treeline**

636 **(CLMT) elevation.** Different symbols represent different regions and colours represent the

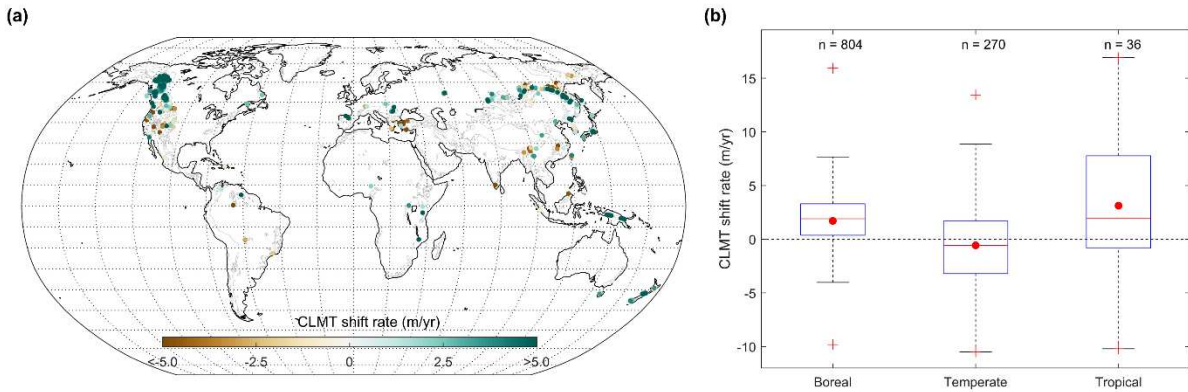
637 distance to the coast. The data points show the mean elevation of all of the pixels in the CLMT.

638 The error bar is the elevation range of the corresponding treeline loop.

639



640
 641 **Figure 3. Climate drivers controlling the variability in treeline elevation for the globe (a),**
 642 **boreal ($\geq 50^\circ\text{N}$, b), temperate ($23.5^\circ - 50^\circ\text{N/S}$, c) and tropical ($23.5^\circ\text{N} - 23.5^\circ\text{S}$, d) regions.**
 643



644
 645 **Figure 4. Closed-loop mountain treeline (CLMT) shift rate during 2000-2010.** (a), Spatial
 646 pattern of CLMT shift rate. (b), Box-plot showing CLMT shift rate in boreal ($\geq 50^{\circ}\text{N}$), temperate
 647 ($23.5^{\circ} - 50^{\circ}\text{N/S}$) and tropical ($23.5^{\circ}\text{N} - 23.5^{\circ}\text{S}$) regions (central line: median; red dot: mean; box:
 648 25th and 75th percentiles, respectively; error bar: maximum and minimum whisker values; +:
 649 maximum and minimum values). The black dashed line is the zero line. Numbers of the studied
 650 CLMT are shown above the boxes.