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1 **Constitutive modelling of Hydro-Mechanical coupling in double porosity**

2 **media based on Mixture Coupling Theory**

3 Yue Ma¹, Jun Feng², Shangqi Ge^{3,*}, Kai Wang⁴, Xiaohui Chen⁵, Aizhong Ding⁶

4 1 School of Civil Engineering, University of Leeds, Leeds, LS2 9JT, UK. Email: cnym@leeds.ac.uk

5 2 School of airport engineering, Civil Aviation Flight University of China, Guanghan, 618307,
6 CHINA. Email: sckid1987@163.com

7 3 Research Center of Coastal and Urban Geotechnical Engineering, Zhejiang University, Hangzhou,
8 310058, CHINA; Geomodelling and Artificial Intelligence Centre, School of Civil Engineering,
9 University of Leeds, Leeds, LS2 9JT, UK. Email: geshangqi@zju.edu.cn

10 4 College of Water Sciences, Beijing Normal University, Beijing, 100875, CHINA. Email:
11 wangkaik@mail.bnu.edu.cn

12 5 School of Civil Engineering, University of Leeds, Leeds, LS2 9JT, UK. Email: x.chen@leeds.ac.uk

13 6 College of Water Sciences, Beijing Normal University, Beijing, 100875, CHINA. Email:
14 ading@bnu.edu.cn

15 * Corresponding author: Shangqi Ge (geshangqi@zju.edu.cn)

16 **Abstract:** Modelling of fluids in deformable geof ormation media has gained great attention in the
17 past decades due to significant applications such as groundwater prediction, shale gas and carbon
18 capture and storage. However, considerable research has been focused on the porous media concept,
19 and dual network (fracture and pores) multiphysics coupled modelling has remained a challenge due
20 to the lack of a systemic theory to bridge the physical deformation of the media (e.g., rocks) and the
21 interaction of water flow in pores and fractures. This paper adopts the non-equilibrium
22 thermodynamics-based approach, the Mixture Coupling Theory, to develop a thermodynamics
23 consistency constitutive model for the fully coupled Hydro-Mechanical behavior in double porosity
24 formation. The energy dissipation due to fluid flow in matrix pore and fracture is given through non-
25 equilibrium thermodynamics, and the relationship between the solid and fluid is linked through
26 Helmholtz free energy. The dynamic evolution of stress, porosity change of the matrix pores and
27 fracture, are derived with respect to mechanical strain, pore pressure, and fracture pressure to account
28 for the flow-deformation interaction. The developed constitutive equations are then solved
29 numerically to show the hydraulic and mechanical behavior of double porosity formation, as well as
30 their sensitivity to parameters.

31 **Keywords:** double porosity; nonequilibrium thermodynamics; hydraulic-mechanical

32 **Introduction**

33 The ubiquity of double porosity media and its completely different characteristics compared to porous
34 media give rise to the importance of studies on water flow in fractured soil or rock in groundwater
35 sources evaluation, underground construction, groundwater contamination, petroleum and shale gas
36 exploitation, underground gas storage, geothermal reservoir (Berkowitz 2002, Rutqvist and
37 Stephansson 2003, Gupta and Yadav 2020). Water flow in the subsurface is driven by both hydraulic
38 gradient and rock mechanical field. The interaction between water flow and the deformation of the
39 solid results in a more complex process in groundwater flow (Segura and Carol 2008, Tsang et al.
40 2015), making it difficult for mathematical modelling.

41 The early modelling work toward the Hydro-Mechanical coupling are the ones by Terzaghi (1943)
42 and Biot (1962), Biot (1972), then followed extensively by many other researchers (Lewis and
43 Schrefler 1987, Vardoulakis et al. 1996, Rutqvist and Tsang 2002, Laloui et al. 2003, Rajagopal and
44 Tao 2005, Tarantino and Tomblato 2005, Wong and Mašin 2014, Zhou and Sheng 2015). In these
45 studies, the porous media is assumed to be homogeneous with single porosity. However, many
46 geomaterials have two scales of void space: the matrix pores and the fracture (Borja and Koliji 2009),
47 as illustrated in Fig. 1.

48 The distinctive fluid transport and pressure distribution in the fracture and matrix pores are quite
49 different from those in the single porosity situation, so that the classic Biot equations fail to capture
50 the feature of the double porosity situation.

51 To describe the coupled hydro-mechanical behavior of the double porosity material, the material is
52 often viewed to be composed of two distinct but overlapping media: one consisting of the porous
53 matrix, in which there are the solid matrix and matrix pores, and the other is the fracture (Barenblatt et
54 al. 1960, Warren and Root 1963), see Fig. 1. The two media can exchange water mass as the porous
55 matrix holds a large storage capacity and low permeability, while the fracture has high permeability
56 and low storability (Song et al. 2019). Based on the above concept, various mathematical formulations
57 representing the fluid flow or hydro-mechanical coupling have been developed by different
58 approaches with different degrees of sophistication.

59 The early double porosity model (Barenblatt et al. 1960, Warren and Root 1963, Aifantis 1980)
60 explored the fluid transport behavior but failed to explore the mechanical deformation and its
61 coupling with fluid. In these research, the coupling between the fluids in the matrix pores and the
62 fracture is achieved by the fluid exchange between the two regions; the flow is assumed to be
63 independent of deformation. Later models (Wilson and Aifantis 1982, Khaled et al. 1984, Beskos and
64 Aifantis 1986, Zhang et al. 2003, Zhang and Roegiers 2005) incorporated the fluid pressure into the
65 strain equation and the strain influence on the fluid transport to achieve the coupling between flow
66 and deformation. However, these models made no progress in the coupling between the fluids in the
67 matrix pores and the fracture, as they only considered the mass exchange. The fact is that the fracture
68 fluid acting on the porous matrix must lead to the change of fracture volume and matrix pore volume,
69 and further influence the fluid transport in the matrix pore, and vice versa. Such a phenomenon is then
70 incorporated in the new fully coupled Hydro-Mechanical models proposed by Khalili (2003), Khalili
71 (2008).

72 The mathematical models are developed by different approaches. There is no certain classification of
73 the approaches for modelling the double porosity problem. Different categories of approaches can be
74 found in (Chen and Teufel 2000, Gelet et al. 2012, Boutin and Royer 2015). Among all the
75 approaches, two noticeable ones are the conventional mechanics approach and the mixture theory
76 approach. Some remarkable mathematic models have been developed by the mechanics approach
77 (Elsworth and Bai 1992, Khalili and Valliappan 1996, Pao and Lewis 2002, Khalili 2003, Khalili
78 2008) and by the mixture theory approach (Aifantis 1977, Aifantis 1979, Aifantis 1980, Wilson and
79 Aifantis 1982, Beskos and Aifantis 1986, Bai et al. 1993, Bai et al. 1993, Borja and Koliqi 2009) and
80 following further work (Wilson and Aifantis 1982, Khaled et al. 1984, Beskos and Aifantis 1986,
81 Berryman and Wang 1995).

82 The mechanics approach is straightforward and simple, but it often requires ad hoc assumptions, and
83 it lacks the ability of systemic self-development (Laloui et al. 2003). For the mixture theory approach,
84 as pointed out by Heidug and Wong (1996), since it maintains the individuality of the solid and fluid
85 phase, it highly relies on the phase interaction information that is very difficult to obtain. Physical

86 intuition and specific assumptions must be required to form the coupling between phases. This may
87 bring difficulties for this approach and restrict the application.

88 In this paper, a non-equilibrium thermodynamics-based approach, the mixture coupling theory, is
89 adopted to develop the fully coupled governing equations for the hydro-mechanical behavior of
90 double porosity media saturated with single-phase flow. This theory origins from Heidug's research
91 for single porosity media with swelling effects (Heidug and Wong 1996). It is modified from the
92 mixture theory by viewing the solid-fluid mixture as a single continuum without explicitly
93 discriminating between the solid and fluid phases, therefore, this theory is more like a hybrid of the
94 Biot poroelasticity view and the mixture theory. Unlike the mixture theory adopting the momentum
95 conservation equation, mixture coupling theory directly works on the free energy conservation,
96 making it easier. This theory provides a rigorous framework to study the coupling effects between
97 multi-physics and multi-phases and has been applied to different couplings in porous media (Chen
98 2013, Chen et al. 2013, Chen et al. 2016, Chen et al. 2018, Ma et al. 2020), and it is the first attempt
99 to apply the theory to the double porosity media to develop the fully coupled Hydro-Mechanical
100 model.

101 By using mixture coupling theory, the very general evolution equation of stress, the porosity of the
102 matrix pores and fracture are obtained with respect to the pore water pressure and fracture water
103 pressure as well as coupling with mechanical strain. The final governing equations are restricted
104 within the small strain and elastic conditions, and are the same as the model proposed by Khalili
105 (2003) through the mechanics approach. The developed mathematical models are solved by the finite
106 element method to illustrate the coupling phenomenon in double porosity media and the sensitivity of
107 parameters.

108 **Balance equation**

109 *Basic definitions and relationships*

110 In a double porosity model, water can pass through the boundary via the porous matrix and the
111 fracture, so that two water flux, namely, porous matrix flux \mathbf{I}^{Mw} and fracture flux \mathbf{I}^{Fw} , are defined

112
$$\mathbf{I}^{Mw} = \rho^{Mw} (\mathbf{v}^{Mw} - \mathbf{v}^s), \mathbf{I}^{Fw} = \rho^{Fw} (\mathbf{v}^{Fw} - \mathbf{v}^s) \quad (1)$$

113 where the subscripts M_w , F_w , s represent the water in the matrix pore, water in fracture and the
 114 solid phase. ρ^{Mw} and ρ^{Fw} are the density of porous water and fracture water, which are relative to
 115 the volume of the whole mixture system. \mathbf{v}^{Mw} , \mathbf{v}^{Fw} , \mathbf{v}^s are the velocity of porous water, fracture
 116 water and the solid.

117 ρ^{Mw} and ρ^{Fw} are related to the true mass density (relative to the volume of porous water and fracture
 118 water) ρ_t^{Mw} and ρ_t^{Fw} through

119
$$\rho^{Mw} = \phi^{Mw} \rho_t^{Mw}, \rho^{Fw} = \phi^{Fw} \rho_t^{Fw} \quad (2)$$

120 where ϕ^{Mw} , ϕ^{Fw} are the porosity of porous matrix and fracture, and they are the volume of the matrix
 121 pore and fracture against the volume of the whole mixture.

122 The Darcy velocity for porous water and fracture water are

123
$$\mathbf{u}^{Mw} = \phi^{Mw} (\mathbf{v}^{Mw} - \mathbf{v}^s), \mathbf{u}^{Fw} = \phi^{Fw} (\mathbf{v}^{Fw} - \mathbf{v}^s) \quad (3)$$

124 ***Balance equation***

125 An arbitrary domain V with a boundary S attached to its surface is selected. The domain includes all
 126 phases and the double porosity feature. Water flow can pass through the boundary while the solid
 127 cannot.

128 *1. Balance equation for Helmholtz free energy*

129 In the double porosity model, the Helmholtz free energy change involves four parts: the mechanical
 130 energy, the energy change by porous water flow, the energy change by fracture water flow and the
 131 entropy part, the balance equation writes as

132
$$\frac{D}{Dt} \int_V \psi dV = \int_S \boldsymbol{\sigma} \cdot \mathbf{v}^s dS - \int_S \mu^{Mw} \mathbf{I}^{Mw} \cdot \mathbf{n} dS - \int_S \mu^{Fw} \mathbf{I}^{Fw} \cdot \mathbf{n} dS - T \int_V \gamma dV \quad (4)$$

133 where: ψ is Helmholtz free energy density, $\boldsymbol{\sigma}$ is the Cauchy stress tensor, μ^{Mw} , μ^{Fw} are the
 134 chemical potential of porous water and fracture water, respectively. T is the temperature, which is
 135 regarded as constant in this paper; γ is the entropy produced per unit volume of the mixture.

136 The material time derivative is $\frac{D}{Dt} = \partial_t + \mathbf{v}^s \cdot \nabla$, where ∂_t is the time derivative and ∇ the gradient,

137 then the derivative version of the balance equation (4) for the Helmholtz free energy is

$$138 \quad \dot{\psi} + \psi \nabla \cdot \mathbf{v}^s - \nabla \cdot (\boldsymbol{\sigma} \mathbf{v}^s) + \nabla \cdot (\mu^{Mw} \mathbf{I}^{Mw}) + \nabla \cdot (\mu^{Fw} \mathbf{I}^{Fw}) = -T\gamma \leq 0 \quad (5)$$

139 2. Balance equation for water mass

140 Water in the matrix pore (fracture) changes in two ways: 1. Water flow \mathbf{I}^{Mw} (\mathbf{I}^{Fw}) passes through the
 141 boundary and exchanges with the surroundings; 2. Water exchanges between the matrix pore and the
 142 fracture. The balance equation for water in the matrix pore and the fracture are

$$143 \quad \frac{D}{Dt} \left(\int_V \rho^{Mw} dV \right) = - \int_S \mathbf{I}^{Mw} \cdot \mathbf{n} dS - \int_V r_{ex} dV \quad (6)$$

$$144 \quad \frac{D}{Dt} \left(\int_V \rho^{Fw} dV \right) = - \int_S \mathbf{I}^{Fw} \cdot \mathbf{n} dS + \int_V r_{ex} dV \quad (7)$$

145 where r_{ex} is the exchange rate of fluid mass between the fracture and matrix pore.

146 The time derivative versions of the water balance equations are

$$147 \quad \dot{\rho}^{Mw} + \rho^{Mw} \nabla \cdot \mathbf{v}^s + \nabla \cdot \mathbf{I}^{Mw} + r_{ex} = 0 \quad (8)$$

$$148 \quad \dot{\rho}^{Fw} + \rho^{Fw} \nabla \cdot \mathbf{v}^s + \nabla \cdot \mathbf{I}^{Fw} - r_{ex} = 0 \quad (9)$$

149

150 **Entropy production and transport law**

151 *Entropy production*

152 During irreversible processes, such as heat and mass transfer, entropy will be produced. The entropy
 153 production can be expressed in terms of thermodynamic flows and thermodynamic forces (Kondepudi
 154 and Prigogine 2014). The quantification of entropy production through non-equilibrium
 155 thermodynamics is the core of the Mixture Coupling Theory.

156 During water transport in double porosity media, the entropy is generated from one mechanism, i.e.,
 157 the friction between the solid and water boundary, which consists of three parts: 1. the friction of
 158 porous water \mathcal{G}_{Mw} ; 2. the friction of fracture water \mathcal{G}_{Fw} ; 3. the friction generated when water
 159 exchange between the matrix pore and the fracture networks \mathcal{G}_{ex} . From non-equilibrium
 160 thermodynamics (Katchalsky and Curran 1965), there is

$$161 \quad \mathcal{G}_{Mw} = -\mathbf{I}^{Mw} \cdot \nabla \mu^{Mw}, \quad \mathcal{G}_{Fw} = -\mathbf{I}^{Fw} \cdot \nabla \mu^{Fw} \quad (10)$$

162 According to Gelet et al. (2012) and Coussy (2004), \mathcal{G}_{ex} can be expressed as

$$163 \quad \mathcal{G}_{ex} = r_{ex} (\mu^{Mw} - \mu^{Fw}) \quad (11)$$

164 Then, the overall entropy production of the mixture system can be written as

$$165 \quad 0 \leq T\gamma = -\mathbf{I}^{Mw} \cdot \nabla \mu^{Mw} - \mathbf{I}^{Fw} \cdot \nabla \mu^{Fw} + r_{ex} (\mu^{Mw} - \mu^{Fw}) \quad (12)$$

166 In equation (12), \mathbf{I}^{Mw} , \mathbf{I}^{Fw} , r_{ex} are the thermodynamic flows, and $-\nabla \mu^{Mw}$, $-\nabla \mu^{Fw}$, $(\mu^{Mw} - \mu^{Fw})$
 167 are the corresponding thermodynamic forces that drive the transport processes.

168 The quantification of the entropy production (12) and the following substitution of $T\gamma$ term in the
 169 free energy equation (5) (like what has been done leading to equation (14)) are the key features of the
 170 Mixture Coupling Theory, distinguishing it from other thermodynamics approaches, such as the ones
 171 by Coussy (2004), Gelet et al. (2012), Nakshatrala et al. (2018). The advantages of the two features
 172 have been presented in Ma et al. (2022) when dealing with the dissolution process.

173 ***Transport Law***

174 The entropy production in section 3.1, on the one hand, can be used to develop the transport law, like
 175 what has been done in Chen et al. (2018). This paper focuses on the coupling between the flow and
 176 deformation. To simplify the discussion, the fluid flow is assumed to obey the Darcy's law. The
 177 Darcy's law for porous water and the fracture water can be derived through the entropy production,
 178 Gibbs-Duhem equation and phenomenological equation, as (Chen 2013)

179
$$\mathbf{u}^{Mw} = -\frac{k^{Mw}}{\nu^w} \nabla p^{Mw}, \quad \mathbf{u}^{Fw} = -\frac{k^{Fw}}{\nu^w} \nabla p^{Fw} \quad (13)$$

180 where k^{Mw} , k^{Fw} are the intrinsic permeability for porous matrix and fracture, p^{Mw} and p^{Fw}
 181 correspond to the porous water pressure and fracture water pressure, respectively, ν^w is the water
 182 viscosity.

183

184 **Constitutive equation**

185 *Basic equation for deformation*

186 Assuming the material maintains mechanical equilibrium so that there is $\nabla \cdot \boldsymbol{\sigma} = \mathbf{0}$. Substituting the
 187 entropy production (12) into the Helmholtz free energy balance equation (5), the Helmholtz free
 188 energy change of the mixture system can be written as

189
$$\dot{\psi} + \psi \nabla \cdot \mathbf{v}^s - \nabla \cdot (\boldsymbol{\sigma} \mathbf{v}^s) + \mu^{Mw} \nabla \cdot \mathbf{I}^{Mw} + \mu^{Fw} \nabla \cdot \mathbf{I}^{Fw} + r_{ex} (\mu^{Mw} - \mu^{Fw}) = 0 \quad (14)$$

190 Equation (14) gets rid of the entropy term in equation (5), and enables us to explore the
 191 Helmholtz free energy change through the mass flux \mathbf{I}^{Mw} , \mathbf{I}^{Fw} and r_{ex} . Multiplying μ^{Mw} ,
 192 μ^{Fw} on both sides of equation (8), (9), and substituting the corresponding results into
 193 equation (14), the Helmholtz free energy change becomes

194
$$\dot{\psi} + \psi \nabla \cdot \mathbf{v}^s = \nabla \cdot (\boldsymbol{\sigma} \mathbf{v}^s) + \mu^{Mw} (\dot{\rho}^{Mw} + \rho^{Mw} \nabla \cdot \mathbf{v}^s) + \mu^{Fw} (\dot{\rho}^{Fw} + \rho^{Fw} \nabla \cdot \mathbf{v}^s) \quad (15)$$

195 In equation (15), the term $\dot{\rho}^{Mw} + \rho^{Mw} \nabla \cdot \mathbf{v}^s$ is the pore water mass change in the mixture, and
 196 therefore $\mu^{Mw} (\dot{\rho}^{Mw} + \rho^{Mw} \nabla \cdot \mathbf{v}^s)$ represents the free energy change due to pore water mass
 197 change. Equation (15) indicating that the free energy change of the system is the result of the
 198 mechanical energy and the mass energy.

199 Next, classic continuum mechanics method is adopted to measure the deformation state. Some basic
 200 relationships are required (Wriggers 2008)

201
$$\mathbf{F} = \frac{\partial \mathbf{x}}{\partial \mathbf{X}}(\mathbf{X}, t), \mathbf{E} = \frac{1}{2}(\mathbf{F}^T \mathbf{F} - \mathbf{I}), \mathbf{T} = J \mathbf{F}^{-1} \boldsymbol{\sigma} \mathbf{F}^{-T} \quad (16)$$

202 where \mathbf{X} is an arbitrary reference configuration, \mathbf{x} is the position, \mathbf{E} represents Green strain, \mathbf{F} is the
 203 deformation gradient, \mathbf{T} and $\boldsymbol{\sigma}$ are the second Piola-Kirchhoff stress and Cauchy stress. J is the
 204 Jacobian of \mathbf{F} ($J = \det \mathbf{F}$), and satisfies $J = \frac{dV}{dV_0}$ (V, V_0 are the volume in the current and reference
 205 configuration.)

206 The time derivation of J satisfies the Euler's formula

207
$$\dot{J} = J \text{div} \mathbf{v}_s \quad (17)$$

208 From equation (15), with the relationships (16) and (17), the Helmholtz free energy equation (15) can
 209 be switched into the reference configuration as

210
$$\dot{\Psi} = \text{tr}(\mathbf{T} \dot{\mathbf{E}}) + \mu^{Mw} \dot{m}^{Mw} + \mu^{Fw} \dot{m}^{Fw} \quad (18)$$

211 in which: $\Psi = J \psi$ is the Helmholtz free energy in the reference configuration, $m^{Mw} = J \rho^{Mw}$,
 212 $m^{Fw} = J \rho^{Fw}$ are the mass density of porous water and fracture water in the reference configuration.

213 ***Helmholtz free energy density of porous/fracture water***

214 According to classical thermodynamics, the free energy density of porous matrix water and fracture
 215 water can be written as

216
$$\psi_{porous} = -p^{Mw} + \rho_t^{Mw} \mu^{Mw} \quad (19)$$

217
$$\psi_{fracture} = -p^{Fw} + \rho_t^{Fw} \mu^{Fw} \quad (20)$$

218 Using the Gibbs-Duhem equation for porous water and fracture water, it leads to

219
$$\dot{p}^{Mw} = \rho_t^{Mw} \dot{\mu}^{Mw} \quad (21)$$

220
$$\dot{p}^{Fw} = \rho_t^{Fw} \dot{\mu}^{Fw} \quad (22)$$

221 Invoking equation (21), (22) into the time derivation of equation(19), (20), the following relationships
 222 can be obtained

223
$$\dot{\psi}_{porous} = \dot{\rho}_t^{Mw} \mu^{Mw} \quad (23)$$

224
$$\dot{\psi}_{fracture} = \dot{\rho}_t^{Fw} \mu^{Fw} \quad (24)$$

225 **Free energy density of the solid matrix**

226 The free energy of the solid-porous-fracture mixture system consists of three parts: the free energy of
 227 the porous water, the free energy of the fracture water and the free energy of the solid matrix. By
 228 subtracting the free energy of the porous water and fracture water from the free energy of the mixture
 229 system, the free energy of the solid matrix can be obtained.

230 From equation (18), (23), (24) and using the density relationship (2), the free energy of the solid
 231 matrix is

232
$$\left(\Psi - J\phi^{Mw}\psi_{porous} - J\phi^{Fw}\psi_{fracture} \right)^\square = tr(\mathbf{T}\dot{\mathbf{E}}) + \dot{\nu}^{Mw} p^{Mw} + \dot{\nu}^{Fw} p^{Fw} \quad (25)$$

233 where $\nu^{Mw} = J\phi^{Mw}$, $\nu^{Fw} = J\phi^{Fw}$ are the porosity of porous matrix and fracture in the reference
 234 configuration.

235 Subtracting the contribution of porous water pressure and fracture water pressure, that is

236
$$W = \left(\Psi - \phi^{Mw}\psi_{porous} - J\phi^{Fw}\psi_{fracture} \right) - \nu^{Mw} p^{Mw} - \nu^{Fw} p^{Fw} \quad (26)$$

237 Substituting equation (25) into the time derivation of the equation (26), the evolution of W can be
 238 obtained as below, enabling us to use the pressure \dot{p}^{Mw} and \dot{p}^{Fw} as variables

239
$$\dot{W} = tr(\mathbf{T}\dot{\mathbf{E}}) - \nu^{Mw} \dot{p}^{Mw} - \nu^{Fw} \dot{p}^{Fw} \quad (27)$$

240 where W is a function of \mathbf{E} , p^{Mw} and p^{Fw} .

241 From equation (27), there must be

242
$$T_{ij} = \left(\frac{\partial W}{\partial E_{ij}} \right)_{p^{Mw}, p^{Fw}}, \nu^{Mw} = - \left(\frac{\partial W}{\partial p^{Mw}} \right)_{E_{ij}, p^{Fw}}, \nu^{Fw} = - \left(\frac{\partial W}{\partial p^{Fw}} \right)_{E_{ij}, p^{Mw}} \quad (28)$$

243 So that

244
$$\dot{W}(\mathbf{E}, p^{Mw}, p^{Fw}) = \left(\frac{\partial W}{\partial E_{ij}} \right)_{p^{Mw}, p^{Fw}} \dot{E}_{ij} + \left(\frac{\partial W}{\partial p^{Mw}} \right)_{E_{ij}, p^{Fw}} \dot{p}^{Mw} + \left(\frac{\partial W}{\partial p^{Fw}} \right)_{E_{ij}, p^{Mw}} \dot{p}^{Fw} \quad (29)$$

245 Differentiating equation (28), with the help of equation (29), the evolution of stress, porosity matrix
 246 porosity and fracture porosity can be obtained.

$$247 \quad \dot{T}_{ij} = L_{ijkl} \dot{E}_{kl} - M_{ij} \dot{p}^{Mw} - S_{ij} \dot{p}^{Fw} \quad (30)$$

$$248 \quad \dot{\nu}^{Mw} = M_{ij} \dot{E}_{ij} + Q \dot{p}^{Mw} + B \dot{p}^{Fw} \quad (31)$$

$$249 \quad \dot{\nu}^{Fw} = S_{ij} \dot{E}_{ij} + B \dot{p}^{Mw} + Z \dot{p}^{Fw} \quad (32)$$

250 where the parameters $L_{ijkl}, M_{ij}, S_{ij}, H_{ij}, B, Q, Z$, are as following group equations

$$251 \quad L_{ijkl} = \left(\frac{\partial T_{ij}}{\partial E_{kl}} \right)_{p^{Mw}, p^{Fw}} = \left(\frac{\partial T_{kl}}{\partial E_{ij}} \right)_{p^{Mw}, p^{Fw}}, \quad M_{ij} = - \left(\frac{\partial T_{ij}}{\partial p^{Mw}} \right)_{E_{ij}, p^{Fw}} = \left(\frac{\partial \nu^{Mw}}{\partial E_{ij}} \right)_{E_{ij}, p^{Fw}}$$

$$252 \quad S_{ij} = - \left(\frac{\partial T_{ij}}{\partial p^{Fw}} \right)_{E_{ij}, p^{Mw}} = \left(\frac{\partial \nu^{Fw}}{\partial E_{ij}} \right)_{E_{ij}, p^{Mw}}, \quad Z = \left(\frac{\partial \nu^{Fw}}{\partial p^{Fw}} \right)_{E_{ij}, p^{Mw}} \quad (33)$$

$$253 \quad B = \left(\frac{\partial \nu^{Fw}}{\partial p^{Mw}} \right)_{E_{ij}, p^{Fw}} = \left(\frac{\partial \nu^{Mw}}{\partial p^{Fw}} \right)_{E_{ij}, p^{Mw}}, \quad Q = \left(\frac{\partial \nu^{Mw}}{\partial p^{Mw}} \right)_{E_{ij}, p^{Fw}}$$

254

255 **Coupled hydro-mechanical governing equations**

256 *Assumptions and simplifications*

257 Equations (30), (31), (32) are the general coupled equations for stress, strain, porous/fracture pressure
 258 and porous/fracture porosity, they allow us to explore the Hydro-Mechanical coupling in a broad way
 259 including anisotropy, large deformation, et.al. This paper forms the final governing equations in a
 260 simple elastic-isotropy cases, therefore, some simplifications and assumptions are made below.

261 1. The mechanical behavior is restricted to small strain condition, therefore the Green Strain tensor
 262 E_{ij} and Piola-Kirchhoff stress T_{ij} can be replaced by strain tensor ε_{ij} and Cauchy stress σ_{ij} .

263 2. Although many double porosity formations show the features of anisotropic and heterogeneous,
 264 following some other research (e.g. Berryman and Wang (1995)), it is roughly assumed that the
 265 material is isotropic. Therefore, material-dependent constants M_{ij}, S_{ij} can be substituted by a form of
 266 scalar multiplied by Kronecker delta.

$$267 \quad M_{ij} = \zeta^{Mw} \delta_{ij}, \quad S_{ij} = \zeta^{Fw} \delta_{ij} \quad (34)$$

268 and considering only the elasticity, the elastic stiffness L_{ijkl} can be a fourth-order isotropic tensor

$$269 \quad L_{ijkl} = G(\delta_{ik}\delta_{jl} + \delta_{il}\delta_{jk}) + \left(K - \frac{2G}{3}\right)\delta_{ij}\delta_{kl} \quad (35)$$

270 in which G and K denote the shear modulus and bulk modulus of the material.

271 According to equations (34), (35) and assumption (1), the stress evolution equation (30) can be
272 simplified to

$$273 \quad \dot{\sigma}_{ij} = \left(K - \frac{2G}{3}\right)\dot{\varepsilon}_{kk}\delta_{ij} + 2G\dot{\varepsilon}_{ij} - \zeta^{Mw}\dot{p}^{Mw}\delta_{ij} - \zeta^{Fw}\dot{p}^{Fw}\delta_{ij} \quad (36)$$

274 With assumption 1, 2, the porosity evolution equations (31) and (32) can be simplified as

$$275 \quad \dot{v}^{Mw} = \zeta^{Mw}\dot{\varepsilon}_{ii} + Q\dot{p}^{Mw} + B\dot{p}^{Fw} \quad (37)$$

$$276 \quad \dot{v}^{Fw} = \zeta^{Fw}\dot{\varepsilon}_{ii} + B\dot{p}^{Mw} + Z\dot{p}^{Fw} \quad (38)$$

277 **Parameter identification**

278 **Identification of ζ^{Mw} and ζ^{Fw}**

279 Consider a situation where the porous matrix block is blocked from the fracture, which means that
280 there is no mass exchange between the matrix pores and the fracture. In this situation, the porous
281 matrix block can be viewed as a non-porous material. The stress/strain change is purely induced by
282 fracture water. This situation is the same as the traditional single porosity research as the porous block
283 can be viewed as solid grain and the fracture become the pores.

284 In this situation, the incremental relationship between stress and pore fluid pressure is

$$285 \quad \dot{\sigma}_{ij} = -\dot{p}^{Fw}\delta_{ij} \quad (39)$$

286 strain rate is related to the fracture water pressure through

$$287 \quad \dot{\varepsilon}_{ij} = -\frac{\dot{p}^{Fw}}{3K_{pb}}\delta_{ij} \quad (40)$$

288 where K_{pb} is the bulk modulus of the porous matrix block.

289 Substituting equation (39) and (40) into equation (36) for the above situation

290
$$-\dot{p}^{Fw} \delta_{ij} = \left(K - \frac{2G}{3} \right) \left(-\frac{\dot{p}^{Fw}}{K_{pb}} \delta_{ij} \right) + 2G \left(-\frac{\dot{p}^{Fw}}{3K_{pb}} \delta_{ij} \right) - \zeta^{Fw} \dot{p}^{Fw} \delta_{kl} \quad (41)$$

291 From equation (41), the expression of ζ^{Fw} can be obtained as

292
$$\zeta^{Fw} = 1 - \frac{K}{K_{pb}} \quad (42)$$

293 Next, to identify the parameter ζ^{Mw} , a second situation is assumed where the porous water pressure
 294 and fracture water pressure change in the same way by \dot{p}^Δ . In this situation, there is no
 295 discrimination between matrix pores and fracture, they are all viewed as the void space in the system.
 296 The incremental relationship between the stress and fluid pressure, and the strain response are

297
$$\dot{\sigma}_{ij} = -\dot{p}^\Delta \delta_{ij}, \quad \dot{\epsilon}_{ij} = -\frac{\dot{p}^\Delta}{3K_s} \delta_{ij} \quad (43)$$

298 Invoking equation (43) into equation (36) for the second situation, it's easy to obtain

299
$$\zeta^{Mw} + \zeta^{Fw} = 1 - \frac{K}{K_s} \quad (44)$$

300 Then, with equation (42), the expression of ζ^{Mw} can be obtained

301
$$\zeta^{Mw} = \frac{K}{K_{pb}} - \frac{K}{K_s} \quad (45)$$

302 *Identification of Q, B and Z*

303 In the first situation considered in section 5.2.1, the fracture volume fraction change is related to
 304 fracture water pressure through

305
$$\dot{v}^{Fw} = -\phi^{Fw} \frac{\dot{p}^{Fw}}{K_{pb}} \quad (46)$$

306 Invoking (46), (40) and the expression for ζ^{Fw} into equation (38) for the first situation leads to

307
$$-\phi^{Fw} \frac{\dot{p}^{Fw}}{K_{pb}} = \left(1 - \frac{K}{K_{pb}} \right) \left(-\frac{\dot{p}^{Fw}}{K_{pb}} \right) + Z \dot{p}^{Fw} \quad (47)$$

308 The expression of Z can be derived as

309
$$Z = \frac{1}{K_{pb}} \left(1 - \frac{K}{K_{pb}} - \phi^{Fw} \right) \quad (48)$$

310 Then, in the second situation in section 5.2.1, the fracture volume change is

311
$$\dot{v}^{Fw} = -\phi^{Fw} \frac{\dot{p}^\Delta}{K_s} \quad (49)$$

312 With equation (49), (43), (48), equation (38) for the second situation can be written as

313
$$-\phi^{Fw} \frac{\dot{p}^\Delta}{K_s} = \left(1 - \frac{K}{K_{pb}} \right) \left(-\frac{\dot{p}^\Delta}{K_s} \right) + B \dot{p}^\Delta + \frac{1}{K_{pb}} \left(1 - \frac{K}{K_{pb}} - \phi^{Fw} \right) \dot{p}^\Delta \quad (50)$$

314 Then, the expression B can be obtained as

315
$$B = \left(1 - \frac{K}{K_{pb}} - \phi^{Fw} \right) \left(\frac{1}{K_s} - \frac{1}{K_{pb}} \right) \quad (51)$$

316 Applying the second situation to equation (37), with the strain and porosity evolution equations

317
$$\dot{\epsilon}_{ij} = -\frac{\dot{p}^\Delta}{3K_s} \delta_{ij}, \quad \dot{v}^{Mw} = -\phi^{Mw} \frac{\dot{p}^\Delta}{K_s} \quad (52)$$

318 It leads to

319
$$-\phi^{Mw} \frac{\dot{p}^\Delta}{K_s} = \zeta^{Mw} \left(-\frac{\dot{p}^\Delta}{K_s} \right) + Q \dot{p}^\Delta + \left(1 - \frac{K}{K_{pb}} - \phi^{Fw} \right) \left(\frac{1}{K_s} - \frac{1}{K_{pb}} \right) \dot{p}^\Delta \quad (53)$$

320 So that the expression for Q can be obtained as

321
$$Q = \left(\frac{1}{K_{pb}} - \frac{1}{K_s} \right) \left(1 - \frac{K}{K_{pb}} + \frac{K}{K_s} - \phi^{Fw} \right) - \frac{\phi^{Mw}}{K_s} \quad (54)$$

322 **Governing equations**

323 **Effective stress and mechanical equation**

324 With the expression of ζ^{Mw} and ζ^{Fw} , equation (36) can be written as

325
$$\dot{\sigma}_{ij} = \left(K - \frac{2G}{3} \right) \dot{\epsilon}_{kk} \delta_{ij} + 2G \dot{\epsilon}_{ij} - \zeta^{Mw} \dot{p}^{Mw} \delta_{ij} - \zeta^{Fw} \dot{p}^{Fw} \delta_{ij} \quad (55)$$

326 From which the effective stress can be written as

327
$$\dot{\sigma}_{ij} = \dot{\sigma}'_{ij} - \zeta^{Mw} \dot{p}^{Mw} \delta_{ij} - \zeta^{Fw} \dot{p}^{Fw} \delta_{ij} \quad (56)$$

328 where σ'_{ij} is the effective stress.

329 The proposed effective stress includes the influence of matrix water pressure p^{Mw} and fracture
 330 water pressure p^{Fw} through the coefficient ζ^{Mw} and ζ^{Fw} . It is the same as the one proposed by
 331 Callari and Federico (2000) and the one by Khalili et al. (2005), Khalili (2008) reduced for saturated
 332 condition. When there is no fracture, the effective stress can be reduced to the effective stress in
 333 porous media by regarding $K = K_{pb}$.

334 Assuming the mechanical equilibrium condition $\partial \sigma_{ij} / \partial x_j = 0$, and using displacement variables

335 $d_i (i=1,2,3)$ through $\varepsilon_{ij} = \frac{1}{2} (d_{i,j} + d_{j,i})$, it leads to

336
$$G \nabla^2 \mathbf{d} + \left(\frac{G}{1-2\theta} \right) \nabla (\nabla \cdot \mathbf{d}) - \zeta^{Mw} \nabla \dot{p}^{Mw} - \zeta^{Fw} \nabla \dot{p}^{Fw} = 0 \quad (57)$$

337 in which θ is Poisson's ratio

338 ***Hydraulic behavior***

339 According to the parameters identified in section 5.2, the porosity change equation (37) and
 340 (38) can now be quantitated as

341
$$\dot{v}^{Mw} = \zeta^{Mw} \dot{\varepsilon}_{ii} + \left[\frac{\zeta^{Mw}}{K} (1 - \zeta^{Mw} - \phi^{Fw}) - \frac{\phi^{Mw}}{K_s} \right] \dot{p}^{Mw} + \left[-\frac{\zeta^{Mw}}{K} (\zeta^{Fw} - \phi^{Fw}) \right] \dot{p}^{Fw} \quad (58)$$

342
$$\dot{v}^{Fw} = \zeta^{Fw} \dot{\varepsilon}_{ii} + (\zeta^{Fw} - \phi^{Fw}) \left(\frac{1}{K_s} - \frac{1}{K_{pb}} \right) \dot{p}^{Mw} + \frac{1}{K_{pb}} (\zeta^{Fw} - \phi^{Fw}) \dot{p}^{Fw} \quad (59)$$

343 The two equations represent the pore and fracture porosity change with respect to the
 344 volumetric strain, the porous water pressure and the fracture water pressure in a fully coupled
 345 way, which are the same as the porosity change equation in Khalili (2003).

346 ***1. Porous matrix water***

347 From water partial mass equation (8), water density relationship (2), Darcy velocity equation (3),
 348 water flux equation (1) and Euler identity, the conservation equation of water can be written as

349
$$\left(\nu^{Mw} \rho_t^{Mw}\right)^{\square} + \nabla \cdot \left(\rho_t^{Mw} \mathbf{u}^{Mw}\right) + r_{ex} = 0 \quad (60)$$

350 Expanding the first term in equation (60) and considering the variation of true density as

351
$$\dot{\rho}_t^{Mw} = \rho_t^{Mw} \left(\frac{1}{\rho_t^{Mw}} \frac{\partial \rho_t^{Mw}}{\partial p^{Mw}} \right) \frac{\partial p^{Mw}}{\partial t} = \rho_t^{Mw} \frac{1}{K_w} \dot{p}^{Mw} \quad (61)$$

352 where K_w is the compressibility of water.

353 With equation (61), equation (60) becomes

354
$$\dot{\nu}^{Mw} \rho_t^{Mw} + \nu^{Mw} \rho_t^{Mw} \frac{1}{K_w} \dot{p}^{Mw} + \nabla \cdot \left(\rho_t^{Mw} \mathbf{u}^{Mw}\right) + r_{ex} = 0 \quad (62)$$

355 Invoking the porous volume fraction evolution equation (37) and Darcy's law (13), the governing
356 equation for porous matrix water transport is

357
$$\zeta^{Mw} \nabla \cdot \dot{\mathbf{d}} + \left(Q + \frac{\nu^{Mw}}{K_w} \right) \dot{p}^{Mw} + B \dot{p}^{Fw} + r_{ex} = \frac{k^{Mw}}{\nu^{Mw}} \nabla^2 p^{Mw} \quad (63)$$

358 with $\zeta^{Mw} = \frac{K}{K_{pb}} - \frac{K}{K_s}$, $Q = \frac{\zeta^{Mw}}{K} (1 - \zeta^{Mw} - \phi^{Fw}) - \frac{\phi^{Mw}}{K_s}$, $B = -\frac{\zeta^{Mw}}{K} \left(1 - \frac{K}{K_{pb}} - \phi^{Fw} \right)$.

359

360 2. Fracture water

361 Similar to the above steps, the governing equation for the fracture water transport can be obtained as

362
$$\zeta^{Fw} \nabla \cdot \dot{\mathbf{d}} + B \dot{p}^{Mw} + \left(Z + \frac{\nu^{Fw}}{K_w} \right) \dot{p}^{Fw} - r_{ex} = \frac{k^{Fw}}{\nu^{Fw}} \nabla^2 p^{Fw} \quad (64)$$

363 with $\zeta^{Fw} = 1 - \frac{K}{K_{pb}}$, $B = \left(\zeta^{Fw} - \phi^{Fw} \right) \left(\frac{1}{K_s} - \frac{1}{K_{pb}} \right)$, $Z = \frac{1}{K_{pb}} \left(\zeta^{Fw} - \phi^{Fw} \right)$.

364 **Verification and discussion**

365 Equations (30)-(32) provide very general coupled formulations for the coupled evolution of PK-2
366 stress, porous matrix porosity and fracture porosity, along with the dynamic change of porous water
367 pressure and fracture water pressure, and Green strain. These formulations are for general cases like
368 large deformation, isotropic and anisotropic. With the assumption of small deformation and isotropic,

369 the governing equations (57), (63), (64) derived in this paper are the same as the ones proposed based
370 on the mechanics approach, such as Khalili et al. (1999), Khalili (2003) or the ones by Pao and Lewis
371 (2002), Khalili (2008) reduced for saturated condition which have been verified by comparing with
372 experimental data (Khalili 2003) and further verified through comparing with different double
373 porosity models (Ashworth and Doster 2019). This indicates that the governing equations developed
374 by mechanics approaches are only specific cases of the general constitutive equations by using
375 mixture coupling theory.

376 Additionally, the mixture coupling theory framework developed in this paper is more rigorous and
377 realistic, with the least assumptions. For example, in the mechanics approach, the assumptions (e.g.
378 isotropic) were made at the very beginning of the derivation process, therefore, restricting its
379 following derivation process, limiting any possibility of extension to other conditions, for example,
380 anisotropic. In other words, the mechanics approach can only obtain the constitutive relations for a
381 specific case. However, through the mixture coupling theory adopted in this paper, no specific
382 assumptions are required at the beginning, so that the general constitutive relations (i.e. equation
383 (30),(31),(32)) can be obtained. Such a relationship applies to many conditions, e.g., large
384 deformation, isotropic/anisotropic. Following the general constitutive relations, elasticity and isotropy
385 conditions are selected to obtain the final governing equations. It can thus be concluded that the
386 mixture coupling theory approach is more flexible and widely applicable.

387 The key research object of Mixture Coupling Theory is the Helmholtz free energy change of the
388 system, i.e., equation (15) and (25), which are achieved through the mass and energy balance, as well
389 as the entropy production and the Gibbs-Duhem equation. Since many dynamics processes can be
390 described through energy dissipation, Mixture Coupling Theory can be used in a lot of fields, such as
391 thermo-hydro-mechanical-chemical coupling (Ma et al. 2022), swelling (Chen et al. 2016), dissolution
392 (Ma et al. 2022), or potentially wave propagation as wave is mainly a movement of energy through a
393 medium (Kumar et al. 2021, Rajak et al. 2022), or other fields like biological tissue. However, all the
394 aforementioned research topics are mainly for porous media. This paper is the first attempt to develop
395 the fully coupled equations for the double porosity media. It is noticed that Aghighi et al. (2021) used
396 a similar approach to study the sorption in double porosity media, but their equations are not

397 presented in a fully coupled way as the pore water pressure influence is missing in the mechanical or
 398 transport equation.

399

400 **Numerical Simulation**

401 A simple numerical simulation is presented in this section to illustrate the mechanical and hydraulic
 402 behavior in the double-porosity formation. The mechanical deformation and hydraulic pressure
 403 change are given, as well as the porosity and permeability change. The sensitivity of fracture spacing
 404 and permeability is analyzed.

405 *Porosity, permeability and exchange rate*

406 1. Porosity

407 The matrix and fracture porosity change equation have been derived as equation (31) and (32), which
 408 are further reduced to equation (37) and (38) according to the assumptions in section 5.1. From the
 409 parameters identified in section 5.2, the matrix and fracture porosity equation are solved in
 410 incremental form as

$$411 \quad \Delta v^{Mw} = \zeta^{Mw} \Delta \varepsilon_{ii} + Q \Delta p^{Mw} + B \Delta p^{Fw} \quad (65)$$

$$412 \quad \Delta v^{Fw} = \zeta^{Fw} \Delta \varepsilon_{ii} + B \Delta p^{Mw} + Z \Delta p^{Fw} \quad (66)$$

413 2. Permeability

414 The permeability change is related to porosity through the Kozeny-Carman's law as (Zheng and
 415 Samper 2008)

$$416 \quad \frac{k^{Mw}}{k_0^{Mw}} = \left(\frac{v^{Mw}}{v_0^{Mw}} \right)^3 \left(\frac{1 - v_0^{Mw}}{1 - v^{Mw}} \right)^2 \quad (67)$$

417 in which the subscript '0' denotes the initial value.

418 This relationship is normally for porous matrix, for fracture, although there is no evidence to support
 419 this relationship, it is roughly assumed that the fracture permeability also follows equation (67).

420 3. Exchange rate

421 The exchange rate can be described by

422
$$r_{ex} = \frac{\chi k^{M_w}}{v^w} (p^{M_w} - p^{F_w}) \quad (68)$$

423 where χ is the shape factor. There have been some different expressions for χ proposed by different
 424 researchers (Warren and Root 1963, Kazemi et al. 1976), a summary can be found in Ranjbar and
 425 Hassanzadeh (2011). In this paper, the expression of χ developed by Warren and Root (1963) is
 426 adopted as

427
$$\chi = \frac{4N(N+2)}{L^2} \quad (69)$$

428 where $N (N = 1, 2, 3)$ represents the dimension of the porous matrix block. L is the fracture spacing.

429 ***Geometry and boundary condition***

430 A double porosity geof ormation, with 20m length and 1m height (Fig. 2), is selected. The formation
 431 initially contains water at pressure of 30MPa. At the beginning of the simulation, the pressure at the
 432 right boundary drops to 5MPa due to external disturbance while maintaining 30MPa at the left
 433 boundary. By setting the right side to be permeable, water can flow out. The formation is initially at
 434 mechanical equilibrium with no effective stress. To explore the mechanical behavior when the
 435 pressure changes, the left boundary is allowed to move while the other boundaries are constrained.

436 One observation line and three observation points A, B, C are selected.

437 Parameters adopted in this simulation are listed in Table 1 (Abousleiman and Nguyen 2005, Nair et al.
 438 2005, Gelet et al. 2012)

439 ***Numerical results***

440 ***1. Hydraulic and mechanical behavior***

441 The evolution of matrix pressure and fracture pressure along the observation line is presented in Fig. 3
 442 and Fig. 4. Because of a pressure gradient generated at the beginning, water will flow from the left to
 443 the right, as time goes by, the pressure within the domain decreases and trends to reach equilibrium
 444 (Fig. 3, 4). Comparing Fig. 3 and Fig. 4, the pressure change in the fracture is quicker thanthat in the
 445 matrix pore, this is mainly because of a greater permeability of the larger permeability in the fracture
 446 zone. Fig. 5 shows the pressure change with time at observation points A, B, C, from which it is

447 clearer that the fracture pressure changes quicker than the pore pressure, but finally, the pressure in
448 the fracture and pore reaches the same as what has been set to be the boundary condition.

449 The domain is initially at mechanical equilibrium with no effective stress, indicating that the external
450 loading is born by the water pressure. Because of the loss of water pressure through the right
451 boundary, the external loading will be taken by both water pressure and solid matrix, therefore,
452 effective stress and displacement generate correspondingly (Fig. 6 and 7). Since water pressure losses
453 more near the right boundary, effective stress generates more on the right side (Fig. 6). As the solid
454 matrix bears external loading, consolidation happens and displacement occurs on the free left
455 boundary (Fig. 7).

456 The pressure difference in the fracture and matrix results in the exchange of water mass. The
457 exchange rate is shown in Fig. 8. At the early time, the pressure difference between the fracture and
458 pore is very significant, thus the exchange rate is great. As time goes by, the pressure difference
459 becomes smaller and smaller, so that the exchange rate decreases.

460 2. Porosity and permeability

461 The matrix and fracture porosity distribution along the observation line are presented in Fig. 9 and Fig.
462 10. The Figs show up to a 5% decrease in matrix porosity and a 8% decrease in fracture porosity.
463 Since the pressure change and strain change on the right part of the domain are more significant than
464 those on the left part, and the porosity change shows similar distribution. As the permeability is
465 related to porosity through equation (67), permeability change has a similar trend with porosity
466 change, as shown in Fig. 11 and 12.

467 ***Sensitivity analysis of fracture spacing***

468 An important part of the double porosity model is the exchange of water from the porous matrix to
469 fracture. A greater exchange rate will help the porous water pressure change quicker. From equation
470 (69), the exchange rate is inversely proportional to the fracture spacing L . Different values of L
471 ($L=0.1, 0.5, 1, 5, 10$) are chosen to explore the sensitivity dependence of hydraulic and
472 mechanical behavior on L . The results are presented in Fig. 13-16.

473 From Fig. 13, the exchange rate in low L value is much higher than that in big L value. A lower L
474 value means a high ‘fracture density’, so that water can be transferred quicker between porous block
475 and fracture. In the simulated case, at a certain time or space, the matrix pressure is mostly greater
476 than fracture pressure, so that matrix water flows into the fracture space. The lower L value is, the
477 quicker matrix water lose to fracture, therefore, the matrix pressure changes quicker (Fig. 14).
478 Similarly, fracture pressure changes slower with a small L value (Fig. 15). The difference in pressure
479 change consequently leads to a difference in mechanical behavior (Fig. 16).

480 *Sensitivity analysis of fracture permeability*

481 The hydraulic transport is highly affected by permeability. In this section, the permeability sensitivity
482 is explored by setting the fracture permeability as $k^{Fw} = 5*10^{-19} m^2, 5*10^{-18} m^2, 5*10^{-17} m^2$ while
483 keeping the pore permeability as $k^{Mw} = 5*10^{-20} m^2$ to represent the permeability difference at
484 different magnitudes. Other parameters remain the same as those listed in Table 1.

485 The fracture water pressure under different permeabilities is shown in Fig. 17, from which it is clear
486 that the fracture water pressure drops much quicker with a higher permeability. Because the water
487 exchange rate between the fracture and matrix pore depends on the pressure difference, hence, the
488 exchange rate under high fracture permeability is quicker (Fig. 18), which further promotes the matrix
489 pore water pressure drop (Fig. 19).

490 The pressure change trend under different fracture permeability is similar to that in section 6.3, but the
491 porosity change under different fracture permeability is quite different, as shown in Fig. 20. When
492 $k^{Mw} = 5*10^{-19} m^2$, the fracture porosity drops most on the right boundary, but when
493 $k^{Mw} = 5*10^{-18} m^2$, or $k^{Mw} = 5*10^{-17} m^2$, the maximum porosity is not at the right boundary but at a
494 point closed to the right boundary. According to equation (66), the change in fracture porosity Δv^{Fw}
495 comes from three parts: 1. the change of strain $\zeta^{Fw} \Delta \varepsilon_{ii}$; 2. The change of pore pressure $B \Delta p^{Mw}$; 3.
496 The change of fracture pressure $Z \Delta p^{Fw}$. The contribution of the three parts under different
497 permeabilities is presented in Fig. 21. The consolidation (strain change) and fracture pressure drop
498 would decrease the fracture porosity while the matrix pressure change trends to increase the fracture

499 porosity, leading to the overall decrease of fracture porosity. It can be found that the main difference
500 among the three permeabilities is the fracture pressure contribution part: under all permeabilities, the
501 strain and pore pressure changes slow and their contribution to the fracture porosity are similar.
502 However, when permeability is higher, the fracture pressure drops quickly, leading to a quick and
503 significant decrease of fracture porosity; when permeability is lower, fracture pressure drops slowly
504 and fracture porosity changes slowly. The combined influence of the strain, matrix pressure and
505 fracture pressure results in the dramatic trend in Fig. 20. Owing to the permeability difference, the
506 matrix porosity also changes in a dramatic trend, as presented in Fig. 22 and Fig. 23.

507 **Conclusion**

508 This paper derives the constitutive equations for double porosity formation under a coupled hydro-
509 mechanical situation by using the Mixture Coupling Theory. The cross-coupling relations between
510 stress, strain, porous/fracture water pressure and porous/fracture porosity are obtained, allowing us to
511 explore the Hydro-Mechanical response in a broad way. The final governing equations are formed for
512 the elastic condition, leading to the same equations developed by other approaches.

513 The constitutive models are solved by the finite element method, and the results show the hydraulic
514 and mechanical deformation, as well as the porosity and permeability change. The sensitivity analysis
515 shows that the smaller fracture spacing greatly increases the exchange rate and facilitates the matrix
516 water pressure change, the fracture permeability sensitivity analysis shows that the greater
517 permeability significantly accelerates the pressure change and affects the porosity change.

518 **Data Availability statement**

519 Some or all data, models, or code that support the findings of this study are available from the
520 corresponding author upon reasonable request.

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