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Double-Diffusive Magnetic Layering

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ABSTRACT

Double-diffusive systems, such as thermosolutal convection, in which the density depends on two components that diffuse at different rates, are prone to both steady and oscillatory instabilities. Such systems can evolve into layered states, in which both components, and also the density, adopt a ‘staircase’ profile. Turbulent transport is enhanced significantly in the layered state. Here we exploit an analogy between magnetic buoyancy and thermosolutal convection in order to demonstrate the phenomenon of magnetic layering. We examine the long-term nonlinear evolution of a vertically-stratified horizontal magnetic field in the so-called ‘diffusive regime’, where an oscillatory linear instability operates. Motivated astrophysically, we consider the case where the viscous and magnetic diffusivities are much smaller than the thermal diffusivity. We demonstrate that diffusive layering can occur even for subadiabatic temperature gradients. Magnetic layering may be relevant for stellar radiative zones, with implications for the turbulent transport of heat, magnetic field and chemical elements.

Keywords: instabilities — magnetic buoyancy — solar tachocline — Sun: magnetic fields

1. INTRODUCTION

In an electrically conducting gas, under the influence of gravity, horizontal magnetic fields with strength varying with depth can become unstable to what is known as magnetic buoyancy instability (Newcomb 1961; Parker 1966). This instability has been studied in some detail, both in the linear and nonlinear regimes, particularly with respect to the breakup and escape of the toroidal field within the Sun (see, e.g., the review by Hughes 2007).

The simplest means of demonstrating magnetic buoyancy instability is through consideration of a plane parallel atmosphere containing a horizontal magnetic field that varies in strength with height, $\mathbf{B} = B(z)\hat{\mathbf{y}}$. **This may be regarded as a local Cartesian description of an azimuthal field in a spherical geometry.** For ideal (non-diffusive) MHD, Newcomb (1961) showed, via the energy principle, that necessary and sufficient conditions for instability are that somewhere in the plasma either

$$\frac{d\rho}{dz} > -\frac{\rho g}{a^2 + c^2}, \quad (1)$$

for modes for which the field lines do not bend (interchange modes), or

$$\frac{d\rho}{dz} > -\frac{\rho g}{c^2} \quad (2)$$

for three-dimensional (undular) perturbations with a very long wavelength in the direction of the imposed field. Here z is the vertical coordinate, increasing upwards, ρ is the density, g is the magnitude of the acceleration due to gravity, $a(z)$ and $c(z)$ are, respectively, the Alfvén speed and the adiabatic sound speed, defined by

$$a^2 = \frac{B^2}{\mu_0 \rho}, \quad c^2 = \frac{\gamma p}{\rho}, \quad (3)$$

where p is the gas pressure, μ_0 is the magnetic permeability and γ is the usual ratio of specific heats.

The instability criteria (1) and (2) can be reformulated so that the role of the magnetic gradient is more evident (Thomas & Nye 1975), yielding

$$-\frac{ga^2}{c^2} \frac{d}{dz} \ln \left(\frac{B}{\rho} \right) > N^2 \quad (4)$$

for interchange modes, or

$$-\frac{ga^2}{c^2} \frac{d}{dz} \ln(B) > N^2 \quad (5)$$

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for undular modes, where

$$N^2 = \frac{g}{\gamma} \frac{d}{dz} \ln(p\rho^{-\gamma}) \quad (6)$$

is the square of the Brünt-Väisälä frequency, a measure of the stratification. Clearly, either a sufficiently rapid decrease with height of B/ρ or B can destabilize the interchange or undular modes respectively. It is of interest to note that, despite the stabilizing influence of magnetic tension, undular modes are more readily destabilized; the physics of the instability mechanism is discussed in Hughes & Cattaneo (1987).

The instability criteria (4) and (5) are derived assuming that there are no diffusive processes present (i.e. zero viscosity, resistivity and thermal diffusivity). The criteria are modified in a significant manner by the incorporation of magnetic and thermal diffusion (Gilman 1970; Acheson 1979). For example, for interchange modes, with horizontal and vertical wavenumbers k_x and k_z , the criterion for instability becomes

$$-\frac{ga^2}{c^2} \frac{d}{dz} \ln\left(\frac{B}{\rho}\right) > \frac{\eta}{\gamma\kappa} N^2 + \frac{\eta\nu}{\gamma} \frac{k^6}{k_x^2}, \quad (7)$$

where η is the magnetic diffusivity, κ the thermal diffusivity, ν the kinematic viscosity, and $k^2 = k_x^2 + k_z^2$. The effect of the ratio η/κ can be seen clearly by comparing conditions (4) and (7). For a given background subadiabatic stratification, a weaker gradient of magnetic field is required to promote instability when $\eta/\kappa < \gamma$. Essentially, for a perturbed parcel of fluid, the stabilizing effects of the thermal stratification are eroded more quickly by diffusive effects than the destabilizing effects of the magnetic stratification, and therefore instability is more likely to occur. In astrophysical contexts, $\eta \ll \kappa$ and thus the difference between criteria (4) and (7) is significant.

Criteria (1) and (2) (or (4) and (5)) describe the onset of direct (steady) instabilities. The incorporation of diffusion not only modifies these criteria (e.g. expression (7)), but also introduces another entire class of unsteady, oscillatory instabilities. The instability mechanism (which is described below) is more subtle, but can again be distilled into a criterion relating the gradient of B/ρ to the background stratification. As shown by Hughes (1985), the criterion for instability of interchange modes is

$$-\frac{ga^2}{c^2} (\eta + \nu - \kappa(\gamma - 1)) \frac{d}{dz} \ln\left(\frac{B}{\rho}\right) > (\kappa + \nu)(\kappa + \eta)(\eta + \nu) \frac{k^6}{k_x^2} + (\kappa + \nu)N^2. \quad (8)$$

Of particular note here is that two very different forms of the field can be unstable, depending on the sign of

the factor $\eta + \nu - (\gamma - 1)\kappa$. For $\eta + \nu > (\gamma - 1)\kappa$, instability occurs if B/ρ decreases sufficiently rapidly with height, whereas for $\eta + \nu < (\gamma - 1)\kappa$, instability occurs if B/ρ increases sufficiently rapidly with height. The latter condition on the diffusion coefficients is readily satisfied in stellar interiors, and exhibits a conundrum whereby instability occurs owing to diffusive effects in a situation where both the thermodynamic and magnetic gradients individually appear to be stabilizing.

Much more is known about these types of instabilities outside of the magnetic context. Instabilities of a convectively stable state that arise owing to the disparate diffusion rates of two components that contribute to the density are generally known as *double-diffusive* instabilities. Double-diffusive systems have been studied extensively, chiefly motivated by their applications to oceanographical and astrophysical mixing; the field is reviewed in the monograph by Radko (2013). The two components contributing to the density are typically the temperature and a concentration field of some material fluid contaminant, in what are known as thermosolutal systems. The most well-studied example is that of the oceanographic thermohaline system (see, for example, the review by Schmitt 1994), where the two components are heat and salt, with cooler and saltier water being denser.

If hot, salty water overlies cold, fresh water in proportions such that the overall density stratification still has the density decreasing upwards, double-diffusive direct instability can occur, due to heat diffusing much faster than salt. A parcel of water displaced upwards loses its stabilizing (relatively cool) thermal content to its surroundings, but retains more of its destabilizing (relatively fresh) salinity content. It can thus be less dense than its surroundings — in this case, the perturbation therefore continues to move upwards (and vice versa for downward displacements). Sinusoidal perturbations elongate upwards and downwards and the direct instability in this case is often known as ‘salt fingering’ (or just ‘fingering’). The underlying mechanism of the steady magnetic buoyancy interchange instability, governed by criterion (7), may be described in an analogous fashion to that of salt fingers.

Conversely, if the situation is reversed, whereby cold, fresh water overlies hot, salty water, instability can still occur, but in an oscillatory, overstable manner. A parcel perturbed upwards is denser than its surroundings and so returns to its original position. Through diffusive exchange, it returns cooler but fresher than it was initially. If the net effect is such that the parcel is denser than its original form, then it will overshoot. The process thus repeats, leading to oscillations that grow with

time. This situation is often known as the ‘diffusive’ form of the double-diffusive instability in the oceanographic context, or ‘semi convection’ in the astrophysical context (where the mean molecular weight of the gas content replaces the effect of salt). Note, however, as intimated above with regard to equation (8), the oscillatory form of the magnetic double-diffusive instability is rather subtle. In particular, unlike for thermohaline convection, it cannot necessarily simply be considered as a reversal of the stabilizing and destabilizing roles of the components in comparison with the direct instability. Most strikingly, for $\eta + \nu < (\gamma - 1)\kappa$, instability can occur when both the B/ρ and entropy ($p\rho^{-\gamma}$) gradients are stabilizing, as measured by the diffusionless criterion (4). As shown by Hughes (1985), in these circumstances the instability can still be interpreted in terms of a fluid parcel (or flux tube) argument, but with the key ingredient being the role played by the compressive heating of the magnetic field.

Double-diffusive instabilities — of both the fingering and diffusive types — are of particular interest because the motions they engender induce mixing where it might not be expected if one merely considers convective instabilities based on the overall density gradient. Enhanced thermal and saline transport in the oceans can contribute to the existence of large-scale circulations (e.g. Rahmstorf 2006), and anomalous chemical abundances in stars can be more consistent with double-diffusively enhanced transport there (e.g. Langer 1991).

One of the most intriguing aspects of double-diffusive convection is its tendency to form layers or ‘staircases’. Importantly, further enhanced transport can result from a secondary instability leading to such layered states. These states are identified by their stepped salinity, temperature and density profiles with depth, representing well-mixed convective layers separated by steep-gradient interfaces. The formation of layers occurs for both the fingering and diffusive regimes, although there are differences in the two, particularly in the interface dynamics. Observations of both types of layering are found in the oceans: fingering staircases require overlying warm salty water, as in the tropical Atlantic (e.g. Schmitt et al. 2005), whereas diffusive staircases require overlying cold, fresh water, as found in the Arctic (e.g. Timmermans et al. 2008). For both regimes, it is found that the layered states have a measurably stronger vertical transport.

Many possible explanations have been advanced for the formation of layers. Merryfield (2000) reviewed the state of play of the theory of double-diffusive staircase formation, discussing in some detail the four main candidates proposed at that time: the collective in-

stability of salt fingers (Stern 1969); the possibility of metastable equilibria; instability via negative density diffusion (cf. Phillips 1972; Posmentier 1977); intrusions resulting from horizontal gradients (e.g. Zhurbas & Ozmidov 1983). Subsequently, a fifth candidate has emerged — the so-called γ instability of Radko (2003). This may be characterized as a mean field instability, driven by gradients in the flux ratio γ , which can operate in both the fingering and diffusive regimes¹. Somewhat of a resurgence has occurred in the theoretical understanding of layer formation, driven by the fact that computational technology has become powerful enough to accommodate the range of scales that seems necessary to simulate the process, and hence test possible theories of layer formation. The simulation of layer formation is computationally demanding for two reasons: it relies on a large range of spatial scales and, furthermore, it is, in some sense, a slow process. However, today’s most powerful computers can overcome these constraints. Consequently, persuasive numerical simulations of layer formation in both the fingering and diffusive regimes have been performed (e.g. Stellmach et al. 2011; Rosenblum et al. 2011).

Given the importance of the layering process for turbulent transport in thermosolutal convection, coupled with the recent possibility of simulating layering computationally in the astrophysical context (see the comprehensive review by Garaud 2018), this raises the interesting question of whether magnetic buoyancy instabilities (a double-diffusive system, as discussed above) can also lead to layering and, if so, what are the implications for transport in magnetized stellar interiors? Under certain constraints, there is a direct analogy between the dynamics of magnetic buoyancy instabilities (including diffusion) and the well-studied thermohaline convection (Spiegel & Weiss 1982, see also Section 2). The formal transformation between the two systems holds in the magneto-Boussinesq limit and for two-dimensional motions in which, for the magnetic system, the field remains unidirectional. This transformation, however, does not map temperature and salinity gradients in the thermohaline system directly to entropy and magnetic field gradients in the magnetic buoyancy system, as one might naively expect. Whereas salt does map directly to magnetic pressure, temperature in the thermohaline problem maps to a linear combination of temperature and magnetic pressure in the magnetic buoyancy problem. Thus the results for magnetic buoyancy are, at first sight, a little surprising, particularly, as already noted,

¹ Note that γ in this context is not the ratio of the specific heats used earlier.

the occurrence of instability for ‘stable’ gradients of both entropy and magnetic field.

The aim of this paper is to exploit the analogy between the two systems, building on the new understanding of double-diffusive dynamics revealed in the oceanographic and astrophysical contexts, in order to take a first look at the nature of the layering that can occur as a result of magnetic buoyancy. Through the analogy we know that layering will occur, at some parameter values; the question is whether such parameters, on translation to the magnetic case, might match a relevant astrophysical regime. Stellar interiors are characterized by the regime $\nu < \eta \ll \kappa$ or, equivalently, $\sigma < \tau \ll 1$, where $\sigma = \nu/\kappa$ is the Prandtl number and $\tau = \eta/\kappa$ is the ratio of magnetic to thermal diffusivity. More precisely, at the base of the solar convection zone, for example, $\sigma \approx 2 \times 10^{-6}$ and $\tau \approx 3 \times 10^{-5}$ (see, e.g., Gough 2007). This is a long way from the oceanographic regime, for which the ratio of salt to heat diffusion $\tau \approx 10^{-2}$ and $\sigma \approx 7$. Indeed, whereas it is possible to simulate the oceanographic regime at the correct parameter values, this is a forlorn hope in astrophysics. The numerical evidence suggests that layering in the fingering regime ceases to occur for small σ , but is maintained in the diffusive regime (Mirouh et al 2012; Garaud 2018). For astrophysical implications, we therefore choose to concentrate on the diffusive regime. Furthermore, this allows us to study the possibility of layering in the regime where the gradients of both B/ρ and $p/\rho^{-\gamma}$ are positive (and hence stable in the absence of diffusion).

The layout of the paper is as follows. The mathematical problem, and specifically the link between the thermohaline and magnetic buoyancy systems, is formulated in §2, for both bounded and unbounded domains. The linear instability theory of both systems is summarized in §3. The detailed results of the nonlinear computations for two examples of the layering process are contained in §4. A summary of our results, together with their astrophysical implications, is contained in §5.

2. MATHEMATICAL FORMULATION

The aim of this section is to formulate the mathematical descriptions of both thermohaline convection and magnetic buoyancy and to describe how the systems can be mapped onto each other. As discussed in the introduction, the analogy between the two systems holds only in two dimensions; thus we shall, from the outset, restrict attention to this case. In §2.1 and §2.2 we shall consider, respectively, the general formulation of the equations of thermohaline convection and magnetic buoyancy. The case of unbounded domains, for both systems, is discussed in §2.3.

2.1. The Equations of Thermohaline Convection

The basic state for Boussinesq thermohaline convection has uniform vertical (z) gradients in temperature and salinity, \bar{T}_z and \bar{S}_z . We denote perturbations to the temperature and salinity of this state by T and S respectively; density perturbations ρ are then given by $\rho/\rho_0 = -\alpha_T T + \alpha_S S$, where ρ_0 is a representative density and α_T and α_S are (positive) expansion coefficients. We describe the two-dimensional, incompressible velocity $\mathbf{u} = (u, 0, w)$ in terms of a stream function $\psi(x, z, t)$, defined by $\mathbf{u} = \nabla \times (\psi \hat{\mathbf{y}})$; the vorticity is then $\nabla \times \mathbf{u} = \omega \hat{\mathbf{y}}$, with $\omega = -\nabla^2 \psi$. We scale lengths with d , a characteristic length of the system, times with d^2/κ , where κ is the thermal diffusivity, T with $-\bar{T}_z d$ and S with $-\bar{S}_z d$. Two-dimensional Boussinesq thermohaline convection is then governed by the following three dimensionless equations, describing, respectively, the evolution of the vorticity and the temperature and salinity perturbations (e.g. Turner 1973, noting that his ψ is the negative of our ψ):

$$\frac{1}{\sigma} \left(\frac{\partial (\nabla^2 \psi)}{\partial t} + J(\psi, \nabla^2 \psi) \right) = Ra \frac{\partial T}{\partial x} - Rs \frac{\partial S}{\partial x} + \nabla^4 \psi, \quad (9)$$

$$\frac{DT}{Dt} \equiv \frac{\partial T}{\partial t} + J(\psi, T) = \frac{\partial \psi}{\partial x} + \nabla^2 T, \quad (10)$$

$$\frac{DS}{Dt} \equiv \frac{\partial S}{\partial t} + J(\psi, S) = \frac{\partial \psi}{\partial x} + \tau \nabla^2 S, \quad (11)$$

where the Jacobian J is defined by

$$J(f, g) = \frac{\partial f}{\partial x} \frac{\partial g}{\partial z} - \frac{\partial f}{\partial z} \frac{\partial g}{\partial x}. \quad (12)$$

The problem is governed by four dimensionless parameters: the thermal and solutal Rayleigh numbers, Ra and Rs , the Prandtl number σ and the diffusivity ratio τ are defined by

$$Ra = -\frac{g\alpha_T \bar{T}_z d^4}{\kappa\nu}, \quad Rs = -\frac{g\alpha_S \bar{S}_z d^4}{\kappa\nu}, \quad \sigma = \frac{\nu}{\kappa}, \quad \tau = \frac{\kappa_s}{\kappa}, \quad (13)$$

where g is the acceleration due to gravity, κ_s the solutal diffusivity, and ν the kinematic viscosity. In an experimental setup, the characteristic length d would represent the vertical distance between the two boundaries. With the Rayleigh numbers so defined, positive (negative) Ra is thermally destabilizing (stabilizing), whereas positive (negative) Rs is solutally stabilizing (destabilizing).

2.2. The Equations of Magnetic Buoyancy

We shall consider the equations of magnetic buoyancy under the magneto-Boussinesq approximation derived by Spiegel & Weiss (1982) (see also Corfield 1984;

340 Bowker et al. 2014). Under the ‘standard’ Boussinesq
 341 approximation (Spiegel & Veronis 1960), density varia-
 342 tions appear only in the buoyancy term; in addition, to
 343 leading order, they result only from temperature varia-
 344 tions, with variations in the gas pressure assumed to be
 345 negligible. In studies of magnetoconvection (see, e.g.,
 346 the monograph by Weiss & Proctor 2014), the direct re-
 347 lation between density and temperature perturbations
 348 remains: the magnetic field plays no role in this bal-
 349 ance. By contrast, the very essence of magnetic buoy-
 350 ancy instability is the influence of the magnetic field
 351 on the pressure; this is captured within the magneto-
 352 Boussinesq approximation through an ordering in which
 353 variations in the *total* pressure (i.e. gas + magnetic) are
 354 negligible, but variations in the gas pressure and mag-
 355 netic pressure individually are not. Density variations
 356 are again taken into account only in the buoyancy term,
 357 and the velocity field is again assumed to be incompress-
 358 ible ($\nabla \cdot \mathbf{u} = 0$).

359 The basic state is taken to be magnetohydrostatic,
 360 with a horizontal, depth-dependent magnetic field
 361 $\mathbf{B} = B(z)\hat{\mathbf{y}}$, confined to a layer of depth d ; the magneto-
 362 Boussinesq approximation requires that d is much
 363 smaller than the pressure and density scale heights. We
 364 consider two-dimensional (y -independent) perturbations
 365 in which the magnetic field remains unidirectional (so-
 366 called interchange modes), and again employ a stream
 367 function $\psi(x, z, t)$ with $\mathbf{u} = \nabla \times (\psi\hat{\mathbf{y}})$.

368 Following the formulation of Spiegel & Weiss (1982),
 369 the vorticity equation may be written in dimensional
 370 terms as

$$\frac{D}{Dt} (\nabla^2 \psi) = \frac{g}{T_0} \frac{\partial \delta T}{\partial x} + \frac{g}{p_0} \frac{\partial \delta p_m}{\partial x} + \nu \nabla^4 \psi, \quad (14)$$

371 where δT and δp_m are the perturbations of temperature
 372 and magnetic pressure, and where a subscript zero de-
 373 notes a representative value. Forming the scalar product
 374 of the induction equation with the magnetic field gives

$$\frac{D\delta p_m}{Dt} = -\alpha \frac{\partial \psi}{\partial x} + \eta \nabla^2 \delta p_m, \quad (15)$$

375 where

$$\alpha = \frac{B_0^2}{\mu_0} \frac{d}{dz} \ln \left(\frac{B}{\rho} \right). \quad (16)$$

376 On adopting the ordering $\delta p_m \approx -\delta p$ (negligible vari-
 377 ation in total pressure), the energy equation becomes

$$\frac{D}{Dt} \left(\delta T + \frac{\delta p_m}{C_p \rho_0} \right) = -\beta \frac{\partial \psi}{\partial x} + \kappa \nabla^2 \delta T, \quad (17)$$

379 where C_p is the specific heat at constant pressure and

$$\beta = \frac{T_0}{\gamma} \frac{d}{dz} \ln \left(\frac{p}{\rho^\gamma} \right) \quad (18)$$

380 is the subadiabatic temperature gradient (treated as
 381 constant within the Boussinesq approximation).

382 Equation (15) is already in standard advection-
 383 diffusion form (cf. equations (10), (11)). Although
 384 equation (17) is not, equations (15) and (17) can be
 385 combined to give

$$\frac{D\delta T^*}{Dt} = -\beta^* \frac{\partial \psi}{\partial x} + \kappa \nabla^2 \delta T^*, \quad (19)$$

386 where

$$\delta T^* = \delta T - \frac{\tau \delta p_m}{C_p \rho_0 (1 - \tau)}, \quad \beta^* = \beta - \frac{\alpha}{C_p \rho_0 (1 - \tau)}, \quad (20)$$

387 and where now $\tau = \eta/\kappa$. On employing the scalings

$$\delta p_m = -S\alpha d, \quad \delta T^* = -T\beta^* d, \quad (21)$$

388 together with the standard non-dimensionalization of
 389 lengths with d and times with d^2/κ , equations (15) and
 390 (19) become (11) and (10) respectively. Since, in this
 391 formulation, the underlying gradients are β^* and α , the
 392 associated Rayleigh numbers are

$$Ra_m = -\frac{gd^4\beta^*}{\kappa\nu T_0}, \quad Rs_m = \frac{gd^4\alpha}{\kappa\nu p_0}. \quad (22)$$

393 **The dimensionless form of equation (14) is then**
 394 transformed into equation (9) through the mapping

$$Ra = Ra_m, \quad Rs = \frac{(\gamma - \tau)}{\gamma(1 - \tau)} Rs_m. \quad (23)$$

395 **For completeness, it is also instructive to**
 396 **present the set of dimensionless equations de-**
 397 **scribing magnetic buoyancy without applying**
 398 **any transformation. As discussed in the intro-**
 399 **duction, consideration solely of the magnetic**
 400 **buoyancy problem shows that instability results**
 401 **from a competition between gradients of B/ρ**
 402 **and $p\rho^{-\gamma}$ (e.g. equation (4) or (7)). With this**
 403 **in mind, for this case we therefore use α , de-**
 404 **defined by (16), and β , defined by (18), in the**
 405 **non-dimensionalization. On adopting the usual**
 406 **scalings of lengths with d , times with d^2/κ , and**
 407 **writing**

$$\delta T = -\beta d \widetilde{\delta T}, \quad \delta p_m = -\alpha d \widetilde{\delta p_m}, \quad (24)$$

408 **the dimensionless forms of equations (14), (17)**
 409 **and (15) become, on dropping tildes,**

$$\frac{1}{\sigma} \frac{D}{Dt} (\nabla^2 \psi) = Rt \frac{\partial \delta T}{\partial x} - Rb \frac{\partial \delta p_m}{\partial x} + \nabla^4 \psi, \quad (25)$$

$$\frac{D\delta T}{Dt} - \left(\frac{\gamma - 1}{\gamma} \right) \left(\frac{Rb}{Rt} \right) \frac{D\delta p_m}{Dt} = \frac{\partial \psi}{\partial x} + \nabla^2 \delta T, \quad (26)$$

$$\frac{D\delta p_m}{Dt} = \frac{\partial\psi}{\partial x} + \tau\nabla^2\delta p_m, \quad (27)$$

where, as in Hughes & Proctor (1988), the thermal and magnetic Rayleigh numbers are defined by

$$Rt = -\frac{gd^4\beta}{\kappa\nu T_0}, \quad Rb = \frac{gd^4\alpha}{\kappa\nu\rho_0}. \quad (28)$$

From expressions (22), (23) and (28), together with the definition of β^* , it follows that Rt and Rb are linked to Ra and Rs via the linear relations

$$Ra = Rt + \frac{(\gamma - 1)}{\gamma(1 - \tau)}Rb, \quad Rs = \frac{(\gamma - \tau)}{\gamma(1 - \tau)}Rb. \quad (29)$$

We have already seen how, through the transformation (20), the equations of magnetic buoyancy may be transformed into those of thermohaline convection. Alternatively, if we let

$$\delta\Sigma = \frac{\delta p_m}{\alpha}, \quad \delta\lambda = \alpha\delta T + \left(\frac{\alpha}{C_p\rho_0} - \beta\right)\delta p_m, \quad (30)$$

then (26) and (27) can be expressed as

$$\frac{D\delta\Sigma}{Dt} = -\frac{\partial\psi}{\partial x} + \eta\nabla^2\delta\Sigma, \quad (31)$$

and

$$\frac{D\delta\lambda}{Dt} = \kappa\nabla^2\delta\lambda + \tilde{\kappa}\nabla^2\delta\Sigma, \quad (32)$$

where $\tilde{\kappa} = \alpha\kappa(\beta - 1/C_p\rho_0)$. Thus within this framework, magnetic buoyancy may also be viewed as a binary fluid with cross-diffusion (see, e.g., Batiste et al. 2006). However, it is worth stressing that the transformed variables δT^* , defined by (20), and $\delta\lambda$, defined by (30), are only helpful when δT and δp_m obey the same boundary conditions.

For clarity, it is worth recapping the rationale behind the three related systems parameterized by (Ra, Rs) , (Ra_m, Rs_m) or (Rt, Rb) . Equations (9)–(11) describe thermohaline convection. The transformation (20), with associated Rayleigh numbers given by (22), leads to a set of equations of a very similar form to (9)–(11): the further transformation of Rayleigh numbers (23) recovers (9)–(11) exactly. The third system (25)–(27) results from retaining δT and δp_m as variables and scaling with the ‘natural’ gradients arising from considerations of diffusionless magnetic buoyancy, namely α and β , leading to the Rayleigh numbers given by (28). Under this formulation, expression (26) is not in standard advection-diffusion form, thereby leading to counter-intuitive behavior when interpreting this system.

2.3. Unbounded Domains

In the analysis of oceanic thermohaline convection, it is reasonable to assume that the upper and lower boundaries play no significant role, and hence to consider a vertically unbounded fluid layer (e.g. Stern & Radko 1998). Indeed, it is within such systems that the most interesting nonlinear phenomena are observed computationally.

The basic state for thermohaline convection is again characterized by a uniform temperature gradient \bar{T}_z and a uniform salinity gradient \bar{S}_z . In the absence of any boundaries, we adopt the lengthscale d defined by $d^4 = |\kappa\nu/g\tilde{\alpha}\bar{T}_z|$, i.e. the lengthscale obtained by setting the absolute value of the local Rayleigh number equal to unity. Various possible scalings for T and S may be adopted: here we choose to scale (dimensional) T and S by $T = d|\bar{T}_z|\hat{T}$, $S = d(\tilde{\alpha}/\tilde{\beta})|\bar{T}_z|\hat{S}$. On dropping hats, the dimensionless governing equations for the perturbations of the basic state are

$$\frac{1}{\sigma} \frac{D}{Dt} (\nabla^2\psi) = \frac{\partial T}{\partial x} - \frac{\partial S}{\partial x} + \nabla^4\psi, \quad (33)$$

$$\frac{DT}{Dt} = -\text{sgn}(\bar{T}_z) \frac{\partial\psi}{\partial x} + \nabla^2 T, \quad (34)$$

$$\frac{DS}{Dt} = -\text{sgn}(\bar{S}_z) \frac{1}{R_0} \frac{\partial\psi}{\partial x} + \tau\nabla^2 S, \quad (35)$$

where the density ratio $R_0 = |\tilde{\alpha}\bar{T}_z/\tilde{\beta}\bar{S}_z| = |Ra/Rs|$. Note that by our definition, R_0 is always positive; from our choice of scalings for T and S , the signs of the temperature and salinity gradients enter explicitly in equations (34) and (35), but not in (33). It can be seen that in an unbounded domain, for given σ and τ , the dynamics is controlled simply by the ratio of the Rayleigh numbers (i.e. R_0), together with the signs of the basic state gradients. Finally, we note that scaling the dimensional density ρ with $\tilde{\alpha}d|\bar{T}_z|\rho_0$ gives the following expression for the dimensionless density:

$$\rho = \left(\frac{\text{sgn}(\bar{S}_z)}{R_0} - \text{sgn}(\bar{T}_z) \right) z + S - T. \quad (36)$$

It should be noted, e.g. from expression (36), that in an unbounded domain with linear gradients of T , S and ρ , these quantities can take negative values, depending on the range of z considered. This has no physical significance: it is only gradients that matter, not the absolute values of T , S or ρ .

In extended astrophysical systems, it is again reasonable to downplay the role of boundaries and thus, similarly, to consider infinite domains for the study of magnetic buoyancy instability. On adopting $d =$

491 $|\kappa\nu T_0/g\beta^*|^{1/4}$ as the unit of length, scaling δT^* and δp_m
 492 as

$$\delta T^* = |\beta^*| d \widetilde{\delta T^*}, \quad \delta p_m = |\beta^*| d \frac{p_0}{T_0} \widetilde{\delta p_m^*}, \quad (37)$$

493 and defining the density ratio R_1 by

$$R_1 = \frac{|\beta^*| p_0}{|\alpha| T_0}, \quad (38)$$

494 the dimensionless governing equations may be ex-
 495 pressed, after dropping the tildes, as

$$\frac{1}{\sigma} \frac{D}{Dt} (\nabla^2 \psi) = \frac{\partial \delta T^*}{\partial x} + \frac{(\gamma - \tau)}{\gamma(1 - \tau)} \frac{\partial \delta p_m}{\partial x} + \nabla^4 \psi, \quad (39)$$

$$\frac{D\delta T^*}{Dt} = -\text{sgn}(\beta^*) \frac{\partial \psi}{\partial x} + \nabla^2 \delta T^*, \quad (40)$$

$$\frac{D\delta p_m}{Dt} = -\text{sgn}(\alpha) \frac{1}{R_1} \frac{\partial \psi}{\partial x} + \tau \nabla^2 \delta p_m. \quad (41)$$

Equations (33)–(35) are thus recovered on making the identifications

$$T = \delta T^*, \quad S = -\frac{(\gamma - \tau)}{\gamma(1 - \tau)} \delta p_m, \quad \frac{1}{R_0} = \frac{1}{R_1} \frac{(\gamma - \tau)}{\gamma(1 - \tau)},$$

$$\text{sgn}(\overline{T}_z) = \text{sgn}(\beta^*), \quad \text{sgn}(\overline{S}_z) = -\text{sgn}(\alpha). \quad (42)$$

Whereas for thermosolutal convection the crucial gradients are those of T and S , for magnetic buoyancy they are $p\rho^{-\gamma}$ and B/ρ . After some manipulation, and scaling consistent with the above (see the Appendix), these quantities take the dimensionless form

$$\frac{p}{\rho^\gamma} = \text{const.} + \left(\text{sgn}(\beta^*) \gamma + \frac{\text{sgn}(\alpha)}{R_1} \frac{(\gamma - 1)}{(1 - \tau)} \right) z$$

$$+ \gamma \delta T^* + \frac{(\gamma - 1)}{(1 - \tau)} \delta p_m \quad (43)$$

498 and

$$\frac{B}{\rho} = \text{const.} + \frac{\text{sgn}(\alpha)}{R_1} z + \delta p_m. \quad (44)$$

499 In terms of perturbations from the background state, the
 500 key quantities are therefore the magnetic pressure vari-
 501 ation δp_m and the variation of $p\rho^{-\gamma}$ (related to the en-
 502 tropy variation), which we shall denote by δs and which
 503 is given by

$$\delta s = \gamma \delta T^* + \frac{(\gamma - 1)}{(1 - \tau)} \delta p_m. \quad (45)$$

Although, for thermosolutal convection, the density is related in a simple manner to the temperature and salinity fields via expression (36), for the magnetic buoyancy system the relation between the two pivotal scalar fields

and the density is not so straightforward. Indeed, as shown in the Appendix, it is possible to calculate only the deviation of the density from a reference hydrostatic state: after appropriate scaling, this deviation, $\hat{\rho}$ say, takes the dimensionless form

$$\hat{\rho} = - \left(\text{sgn}(\beta^*) + \frac{(\gamma - \tau)}{\gamma(1 - \tau)} \frac{\text{sgn}(\alpha)}{R_1} \right) z$$

$$- \delta T^* - \frac{(\gamma - \tau)}{\gamma(1 - \tau)} \delta p_m. \quad (46)$$

504 It may also be noted that an alternative, possibly more
 505 intuitive, scaling for the equations of magnetic buoyancy
 506 is to adopt $d = |\kappa\nu T_0/g\beta|^{1/4}$ as the unit of length. On
 507 scaling δT and δp_m by

$$\delta T = |\beta| d \widetilde{\delta T}, \quad \delta p_m = |\beta| d \frac{p_0}{T_0} \widetilde{\delta p_m}, \quad (47)$$

508 the dimensionless governing equations become, on drop-
 509 ping tildes,

$$\frac{1}{\sigma} \frac{D}{Dt} (\nabla^2 \psi) = \frac{\partial \delta T}{\partial x} + \frac{\partial \delta p_m}{\partial x} + \nabla^4 \psi, \quad (48)$$

$$\frac{D\delta T}{Dt} + \frac{(\gamma - 1)}{\gamma} \frac{D\delta p_m}{Dt} = -\text{sgn}(\beta) \frac{\partial \psi}{\partial x} + \nabla^2 \delta T, \quad (49)$$

$$\frac{D\delta p_m}{Dt} = -\text{sgn}(\alpha) \frac{1}{R_2} \frac{\partial \psi}{\partial x} + \tau \nabla^2 \delta p_m, \quad (50)$$

512 where $R_2 = |\beta| p_0 / |\alpha| T_0$.

513 **To summarize for unbounded systems (as we**
 514 **did for finite domains), we have derived three**
 515 **different systems governing the double-diffusive**
 516 **behavior. The three systems are parameterized,**
 517 **separately, by R_0 , R_1 and R_2 ; they are related,**
 518 **respectively, to those parameterized by (Ra, Rs) ,**
 519 **(Ra_m, Rs_m) and (Rt, Rb) , discussed in § 2.2. The**
 520 **parameter R_0 is the density ratio in thermohaline**
 521 **convection, with governing equations (33)–(35).**
 522 **The R_1 system, where R_1 is directly proportional**
 523 **to R_0 , is governed by equations (39)–(41) and**
 524 **arises from transforming the magnetic buoyancy**
 525 **equations into those of classical double-diffusive**
 526 **convection. The R_2 system, governed by equa-**
 527 **tions (48)–(50), results from retaining the stan-**
 528 **dard variables of magnetic buoyancy.**

2.4. Numerical Techniques

530 We solve the governing equations of thermohaline con-
 531 vection, (33)–(35), numerically, and then use the trans-
 532 formations (42)–(46) to translate the results into the
 533 magnetic buoyancy system in the diffusive regime. In

534 the absence of physical boundaries, we adopt periodic
 535 boundary conditions for the perturbations in both the
 536 horizontal and vertical directions. The equations are
 537 solved by a standard pseudo-spectral technique, with a
 538 2/3 de-aliasing rule; time stepping is performed by combining
 539 a second-order Adams-Bashforth scheme with exponential
 540 time differencing.

541 3. LINEAR INSTABILITIES

542 The linear instabilities resulting from thermohaline
 543 convection and magnetic buoyancy have been extensively
 544 studied (see, e.g., the reviews by Turner 1973;
 545 Hughes 2007). Here we give a brief summary, both for
 546 completeness and also as necessary background for the
 547 nonlinear results described in Section 4.

548 3.1. Bounded Domains

549 Let us first consider thermohaline convection in a
 550 bounded system, governed by the linearized versions of
 551 equations (9)–(11). Instability can occur either as a
 552 steady or an oscillatory mode (e.g. Turner 1973). Steady
 553 convection, in the fingering regime, occurs via an exchange
 554 of stabilities when

$$Ra > Ra^{(e)} = \frac{Rs}{\tau} + \frac{k^6}{k_x^2}, \quad (51)$$

555 where, as earlier, k_x and k_z are the horizontal and vertical
 556 wavenumbers and $k^2 = k_x^2 + k_z^2$. Oscillatory convection,
 557 in the diffusive regime, occurs when

$$Ra > Ra^{(o)} = \left(\frac{\sigma + \tau}{1 + \sigma} \right) Rs + \frac{(1 + \tau)(\sigma + \tau)}{\sigma} \frac{k^6}{k_x^2}, \quad (52)$$

558 provided that

$$Rs > \frac{\tau^2(1 + \sigma)}{\sigma(1 - \tau)} \frac{k^6}{k_x^2}; \quad (53)$$

559 condition (53) guarantees that there is a real frequency
 560 of oscillation when $Ra = Ra^{(o)}$. The regions of linear
 561 stability and instability in the (Ra, Rs) plane for $\tau < 1$,
 562 together with the line of neutral buoyancy, are sketched
 563 in Figure 1(a). The onset of steady convection is pre-
 564 dominantly in the third quadrant in the (Rs, Ra) plane,
 565 that of oscillatory convection entirely in the first quad-
 566 rant.

567 The stability boundaries given by (51) and (52) can be
 568 translated into the (Rb, Rt) plane via the transformation
 569 (29), to give

$$Rt^{(e)} = \frac{Rb}{\tau} + \frac{k^6}{k_x^2}, \quad (54)$$

$$Rt^{(o)} = \frac{(1 + \sigma + \tau - \gamma)}{\gamma(1 + \sigma)} Rb + \frac{(1 + \tau)(\sigma + \tau)}{\sigma} \frac{k^6}{k_x^2}. \quad (55)$$

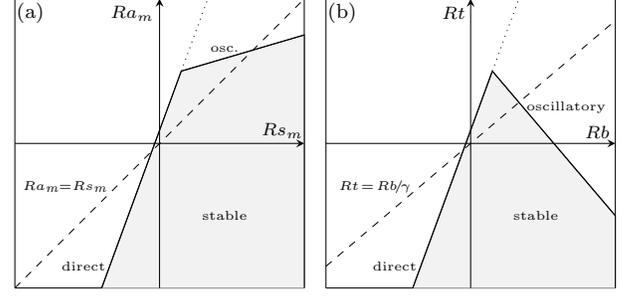


Figure 1. Sketches showing the regimes of linear instability in (a) the (Rs_m, Ra_m) plane (for $\tau < 1$) and (b) the (Rb, Rt) plane (for $\tau < \gamma - 1 - \sigma$). The lines of steady and oscillatory bifurcations are marked, together with the line of neutral buoyancy ($Ra_m = Rs_m$ in (a), $Rt = Rb/\gamma$ in (b)).

571 Note that the criteria for direct (steady) instability, (51)
 572 and (54), are identical in the (Ra, Rs) and (Rt, Rb) sys-
 573 tems. The criteria for oscillatory instability, (52) and
 574 (55), however differ significantly. Expression (55) is the
 575 dimensionless form of (8): as discussed in the Introduc-
 576 tion, and shown in Figure 1(b), it describes the appear-
 577 ance, for $\sigma + \tau < \gamma - 1$, of instability when Rt is neg-
 578 ative and Rb positive — i.e. instability in the quadrant
 579 in which, in the absence of diffusion, both the thermal
 580 and magnetic (B/ρ) gradients are stabilizing.

581 3.2. Unbounded Domains

582 In an unbounded system it is natural to seek pertur-
 583 bations that are periodic in both the horizontal and verti-
 584 cal directions. It is convenient to express the stability
 585 criteria in terms of R_0^{-1} . The critical value of R_0^{-1} for
 586 steady convection is given by

$$(R_0^{-1})^{(e)} = \text{sgn}(\bar{S}_z) \tau \left(\text{sgn}(\bar{T}_z) + \frac{k^6}{k_x^2} \right), \quad (56)$$

587 and that for oscillatory convection by

$$(R_0^{-1})^{(o)} = \text{sgn}(\bar{S}_z) \left(\text{sgn}(\bar{T}_z) \left(\frac{1 + \sigma}{\sigma + \tau} \right) + \frac{(1 + \sigma)(1 + \tau)}{\sigma} \frac{k^6}{k_x^2} \right), \quad (57)$$

588 provided also that

$$\frac{k^6}{k_x^2} < \text{sgn}(\bar{T}_z) \frac{\sigma(1 - \tau)}{(\sigma + \tau)}. \quad (58)$$

589 For $\tau < 1$, steady convection (fingering) chiefly occurs
 590 when the temperature gradient is stabilizing and the
 591 solutal gradient destabilizing. In this case, from (56),
 592 there is instability when

$$R_0^{-1} > \tau \left(1 + \frac{k^6}{k_x^2} \right). \quad (59)$$

593 Periodic boundary conditions allow modes that are in-
 594 dependent of height z — so-called ‘elevator modes’, with

595 $k_z = 0$. From (59), these are the most readily destabilized
596 modes, with $k_x \rightarrow 0$.

597 Oscillatory (diffusive) modes, on the other hand, occur
598 when the temperature gradient is destabilizing and
599 the solutal gradient stabilizing, leading, from (57), to
600 instability when

$$R_0^{-1} < \frac{1 + \sigma}{\sigma + \tau} - \frac{(1 + \sigma)(1 + \tau)}{\sigma} \frac{k^6}{k_x^2}. \quad (60)$$

601 As for the steady mode, the most readily destabilized
602 oscillatory mode is an elevator mode ($k_z = 0$) with
603 $k_x \rightarrow 0$. It should though be noted that in the un-
604 stable regime, away from marginal stability, the mode
605 of maximum growth rate — for steady and oscillatory
606 modes — takes a finite value of k_x .

607 Through the use of the transformation (29), it is
608 straightforward to show that the region of instability
609 in the $Rt < 0$, $Rb > 0$ quadrant is delineated by the
610 following inequalities:

$$\frac{\gamma - \tau}{\gamma - 1} < R_0^{-1} < \frac{1 + \sigma}{\sigma + \tau}. \quad (61)$$

611 The condition for there to be a finite range of R_0^{-1} sat-
612 isfying these inequalities (i.e. that the left hand side of
613 inequality (61) is less than the right hand side) may be
614 expressed as

$$(1 - \tau)(\gamma - (1 + \sigma + \tau)) > 0, \quad (62)$$

615 in accord with expression (55).

616 The instability criteria may alternatively be expressed
617 in terms of R_1^{-1} or R_2^{-1} . For $\alpha < 0$ and $\beta^* > 0$, steady
618 convection occurs if

$$R_1^{-1} > \frac{\gamma\tau(1 - \tau)}{(\gamma - \tau)} \left(1 + \frac{k^6}{k_x^2} \right), \quad (63)$$

619 and, with $\alpha > 0$ and $\beta^* < 0$, oscillatory convection
620 occurs if

$$R_1^{-1} < \frac{\gamma(1 - \tau)}{(\gamma - \tau)} \left(\frac{1 + \sigma}{\sigma + \tau} - \frac{(1 + \sigma)(1 + \tau)}{\sigma} \frac{k^6}{k_x^2} \right). \quad (64)$$

621 Expressions (63) and (64) are straightforward scalings
622 of (59) and (60).

623 In terms of R_2^{-1} , for $\alpha < 0$ and $\beta > 0$, steady convec-
624 tion occurs if

$$R_2^{-1} > \tau \left(1 + \frac{k^6}{k_x^2} \right), \quad (65)$$

625 and, with $\alpha > 0$ and $\beta > 0$ (both gradients ‘stabilizing’),
626 oscillatory convection occurs if

$$R_2^{-1} > \frac{\gamma(1 + \sigma)}{(\gamma - (1 + \sigma + \tau))} \left(1 + \frac{(1 + \tau)(\sigma + \tau)}{\sigma} \frac{k^6}{k_x^2} \right), \quad (66)$$

627 provided that $\sigma + \tau < \gamma - 1$. In transforming between the
628 formally identical criteria (63) and (65) (or between (64)
629 and (66)), it should be noted that k^6/k_x^2 is scaled with
630 β^* for the expressions involving R_1 and with β for those
631 involving R_2 . **We note also the somewhat counter-**
632 **intuitive difference in the inequalities (64) and**
633 **(66); this arises since, in the unstable region of**
634 **the $Rt < 0$, $Rb > 0$ quadrant,**

$$R_1 = \frac{(\gamma - 1)}{\gamma(1 - \tau)} - R_2. \quad (67)$$

4. MAGNETIC LAYERING

635 In this section, we discuss two representative cases of
636 layer formation in the diffusive regime, in a domain of
637 width 100π and height 200π . We solve the governing
638 equations of thermohaline convection in an unbounded
639 domain, (33)–(35), and relate these to the magnetic
640 buoyancy system via the transformations (42). Moti-
641 vated astrophysically, we adopt the smallest values of
642 the Prandtl number σ and the diffusivity ratio τ com-
643 patible with long-time runs in large domains; as such,
644 we set $\sigma = \tau = 0.01$. To accommodate the fine-scale
645 structure, 2048×4096 spectral modes are used. Note,
646 from (60), that oscillatory instability then occurs for
647 $R_0^{-1} < 50.5$ ($R_1^{-1} < 50.297$), and, from (61), that insta-
648 bility in the fourth (i.e. ‘stable-stable’) quadrant in the
649 (R_b, R_t) plane occurs in the range $2.485 < R_0^{-1} < 50.5$
650 ($2.475 < R_1^{-1} < 50.297$). We consider in detail two par-
651 ticular values of the background stratification parame-
652 ter: $R_0^{-1} = 1.5$, which lies in the first quadrant of Fig-
653 ure 1(b), and $R_0^{-1} = 4$, which lies in the fourth. Since
654 the underlying system that we solve computationally is
655 that of thermohaline convection, we quote nice round
656 numbers for R_0^{-1} ; since τ is small, R_1^{-1} is very slightly
657 smaller than R_0^{-1} . The initial condition for both sets of
658 simulations consists of 20 elevator modes, together with
659 a small random perturbation. These modes are essen-
660 tially the fastest growing modes for $R_0^{-1} = 4$ and are
661 about 1.5 times narrower than the fastest growing mode
662 for $R_0^{-1} = 1.5$; the long-term evolution though is not
663 dependent on the precise form of the initial conditions.
664
665

4.1. The case of $R_0^{-1} = 1.5$ ($R_1^{-1} = 1.494$)

666 Figure 2 shows the early evolution of the kinetic en-
667 ergy. Initially the kinetic energy grows in an oscilla-
668 tory fashion, representative of the linear instability in
669 the diffusive regime. The exponentially growing eleva-
670 tor modes are also exact solutions to the fully nonlinear
671 equations, since the Jacobian terms in equations (9)–
672 (11) vanish identically. They are though unstable once
673

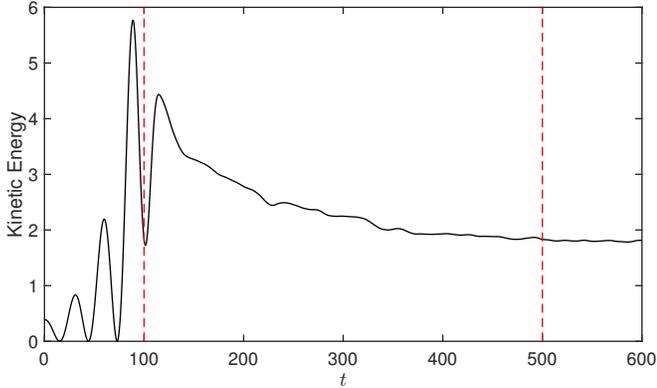


Figure 2. Kinetic energy as a function of time, during the initial instability phase: $R_0^{-1} = 1.5$. The dashed lines denote the times of the snapshots in Figures 3(a,b).

they attain sufficiently large amplitude; snapshots of ω , δs and δp_m during the break-up of the elevators are displayed in Figure 3(a). The nonlinear evolution of this secondary instability causes a rapid and total disruption of the elevators, as shown by the snapshots in Figure 3(b). All memory of the elevators is lost and the system is characterized by small-scale vorticity, together with small-scale entropy and magnetic pressure perturbations; at this stage, the mean profiles of p/ρ^γ , B/ρ and ρ remain essentially linear, as shown in Figure 4.

Following a period of equilibration ($100 \lesssim t \lesssim 1000$), in which the kinetic energy is stationary, the trend for the kinetic energy (although subject to sizeable short-term fluctuations) is an inexorable gradual increase, as shown in Figure 5. Associated with this rise in kinetic energy is the gradual emergence of a layered state from the homogeneous turbulence, and its subsequent evolution. Figure 3(c) shows the four-layered states in δs and δp_m at $t = 3600$, with the latter more pronounced, owing to the small value of τ . The layering in the vorticity ω is much less distinct. The interfaces between the layers are highly turbulent and mobile, with jets of fluid penetrating the interfaces and eventually leading to their destruction.

Figure 6(a) shows ω , δs and δp_m at $t = 8050$, where now only three layers remain; indeed, it can be seen from the figure that the upper layer is already disintegrating under turbulent erosion. Figure 6(b) shows the corresponding plots at $t = 11700$, at which point only two layers remain. The staircase structure in p/ρ^γ , B/ρ and ρ associated with the layers can be seen clearly in Figure 7(a), which plots the horizontal averages of these quantities at the times corresponding to Figures 3(c), 6(a,b).

One of the most significant features of the layering process in double-diffusive convection is the increase in

the vertical turbulent flux of the two diffusing components; this feature is illustrated by the thermohaline simulations of Stellmach et al. (2011) in the fingering regime and those of Mirouh et al (2012) in the diffusive regime. As shown in Figure 8, for the thermohaline problem, the fluxes of both heat and salt are positive (upwards) in the diffusive regime. It thus follows immediately that the flux of δp_m will be negative. It is though not obvious *a priori* what the sign of the entropy flux will be, since, from (42) and (45),

$$\langle w \delta s \rangle = \gamma \langle w T \rangle - \frac{\gamma(\gamma - 1)}{(\gamma - \tau)} \langle w S \rangle, \quad (68)$$

with $\langle w T \rangle$ and $\langle w S \rangle$ both positive, and where $\langle \cdot \rangle$ denotes a global average. In the turbulent regime considered here, $\langle w T \rangle$ and $\langle w S \rangle$ are of comparable magnitude, as shown in Figure 8, and hence, since τ is small, $\langle w \delta s \rangle \approx \langle w T \rangle \approx \langle w S \rangle$. As the layering proceeds, the fluxes increase in magnitude, whilst becoming much more noisy. Between the quasi-stationary phase when the flow is homogeneous (following the instability of the initial finger modes) and $t = 11700$ (corresponding to Figure 6(b)) there is an approximate five-fold increase in the turbulent fluxes. **Note that the ratio of these two average turbulent fluxes defines the “ γ ” of the “ γ -instability” in the mean-field theory of the cause of layering in thermohaline double-diffusive convection (Radko 2003): $\gamma = \langle w T \rangle / \langle w S \rangle$. A study of the variation of this γ with R_0 (or rather γ^{-1} with R_0^{-1}) is required to determine whether the theory fits our results, but this effort lies beyond the scope of this paper.**

4.2. The case of $R_0^{-1} = 4$ ($R_1^{-1} = 3.984$)

Here we consider the evolution from an equilibrium state for which both the magnetic field and the entropy gradient may, at least in the absence of diffusion, be considered to be stable (i.e. B/ρ and p/ρ^γ both increasing with height). In terms of the natural parameters for the magnetic problem (Rt and Rb), the presence of instability is somewhat surprising. However, viewed in terms of the transformed parameters (22), there is no particular significance to the line $Rt = 0$; as discussed above, it simply corresponds to $R_0^{-1} = (\gamma - \tau)/(\gamma - 1) = 2.485$ here.

In its broad aspects, the evolution for $R_0^{-1} = 4$ is similar to that of $R_0^{-1} = 1.5$. Following the growth and saturation of the linear instability, there is a persistent yet noisy increase of kinetic energy, as shown in Figure 9. Comparison of Figures 5 and 9 shows that the increased stratification leads to a reduced growth rate

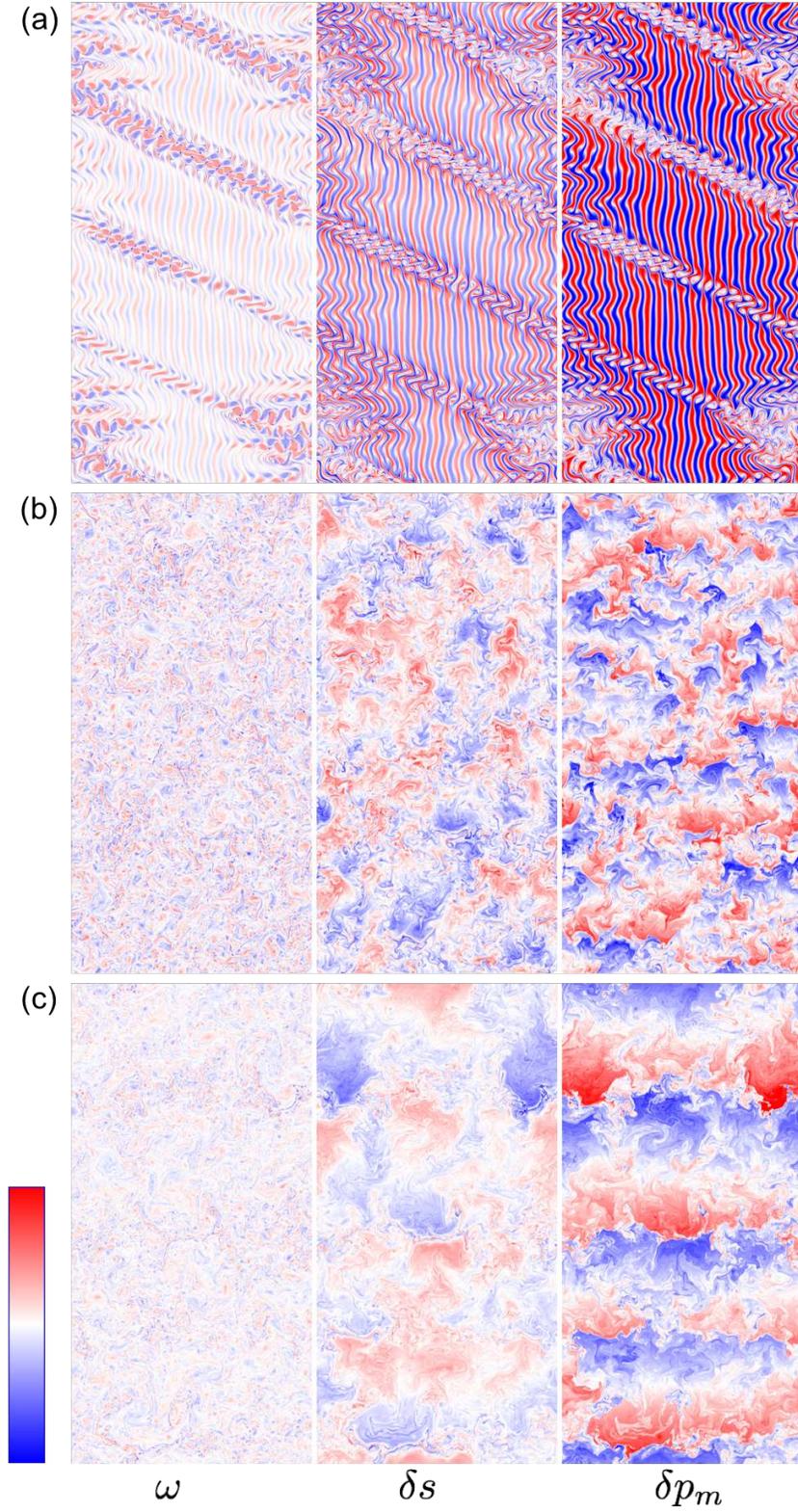


Figure 3. Snapshots at (a) $t = 100$, (b) $t = 500$ and (c) $t = 3600$ of ω , δs and δp_m (all scaled independently) for the case of $R_0^{-1} = 1.5$. The color table ranges from blue (largest negative value) to red (largest positive value), with white denoting the zero value.

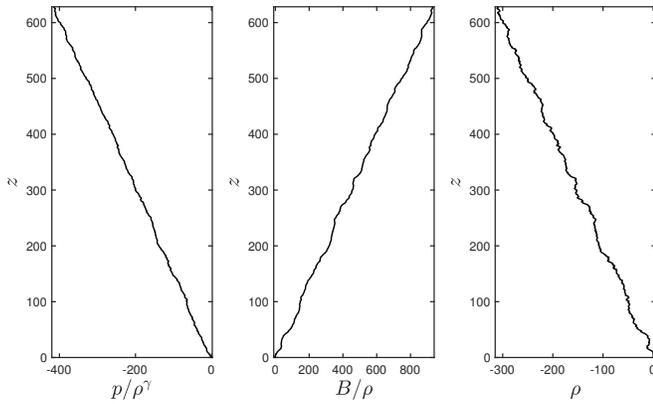


Figure 4. Horizontally averaged profiles of p/ρ^γ , B/ρ and ρ versus height at $t = 500$; $R_0^{-1} = 1.5$.

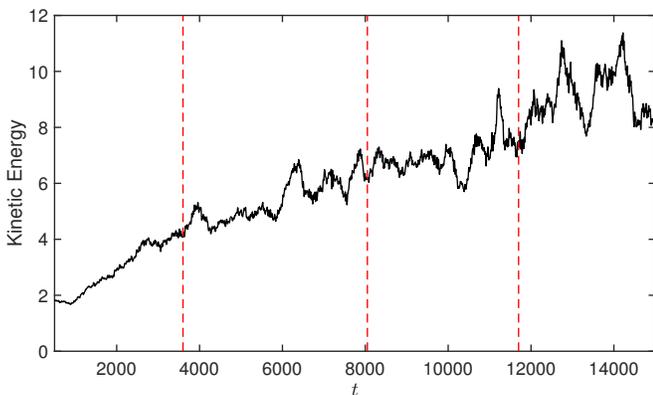


Figure 5. Kinetic energy as a function of time, for $t > 500$: $R_0^{-1} = 1.5$. The dashed lines denote the times of the snapshots in Figures 3c, 6a,b.

of the kinetic energy. Figure 10 shows that, after the initial linear instability of the elevators and the subsequent breakup into smaller-scale turbulence, layers form by eroding the regions of turbulence, as for the evolution for $R_0^{-1} = 1.5$, albeit now on a slower timescale. Figure 10(c) shows the evolution shortly after the layering process has extended across the entire domain. In comparison with the first distinct staircase to emerge for the case of $R_0^{-1} = 1.5$ (see Figure 3), the staircase for $R_0^{-1} = 4$ has shallower steps; furthermore, the flow between the interfaces is less turbulent, leading to a more coherent staircase structure.

Of particular note is the structure of the staircase and its relation to that of the equivalent thermohaline system. Figure 11 shows, for the same time as shown in Figure 10(c), the horizontally averaged profiles of p/ρ^γ , B/ρ and ρ , together with the corresponding profiles of T , S and ρ for the thermohaline problem. In the latter, the background temperature gradient is destabilizing, whereas the salinity gradient is stabilizing. As expected,

the staircase structure is more sharply defined in S than T , owing to the small value of the diffusivity ratio τ . For the magnetic buoyancy problem, the profile of B/ρ is related to that of S via the transformations (42) and (44); indeed, for small τ , B/ρ is essentially $-S$. The most striking feature of Figure 11 is the sharpness and structure of the profile in p/ρ^γ . Since here the mean entropy gradient is ‘stabilizing’ (p/ρ^γ increasing with height), convection can occur only through a local reversal (or reversals) of this gradient. Hughes & Weiss (1995) examined this phenomenon in their explanation of steady convection in the regime with $Rt < 0$ and $Rb < 0$ for a fluid confined by rigid boundaries. In that case, the role of the boundary layers is paramount, with the strong field in the magnetic boundary layers leading to an exceptionally stable entropy gradient in the boundary layers. To compensate, there is of necessity a negative (destabilizing) entropy gradient across the remainder of the cell; this is such as to drive steady convection. Here, in an unbounded domain, a staircase is formed in which there are weakly unstable entropy gradients between the interfaces and strongly stable gradients across the interfaces themselves.

Figure 12 shows that the layering process again leads to a marked overall increase in turbulent transport, but with significant short-term fluctuations; the fluxes are about one fifth of their values for $R_0^{-1} = 1.5$. It is of interest to note that, as for the case of $R_0^{-1} = 1.5$, the balance $\langle w \delta s \rangle \approx \langle wT \rangle \approx \langle wS \rangle > 0$ still holds, even though the background entropy gradient is now positive. Although we see no layer merger in the very long run we have performed, we envisage, based on the results from other simulations of thermohaline convection, that at yet longer times the layers would eventually merge, ultimately giving only one step.

Three-dimensional thermohaline simulations of the two cases we have considered have been performed by Mirouh et al (2012), who find the existence of layers for $R_0^{-1} = 1.5$ but not for $R_0^{-1} = 4$. This discrepancy between our results and theirs for the case of $R_0^{-1} = 4$ may be a genuine difference between two and three dimensions, or may be a result of the facts that (a) our two-dimensional simulations were performed at much higher resolution than the three-dimensional cases, and (b) our two-dimensional cases were integrated for much longer than was possible for the three-dimensional runs, thus allowing layered states eventually to emerge.

5. DISCUSSION

The purpose of this paper has been to expand ideas related to the mixing processes available in stars. Often, the inferred information about the interiors of stars

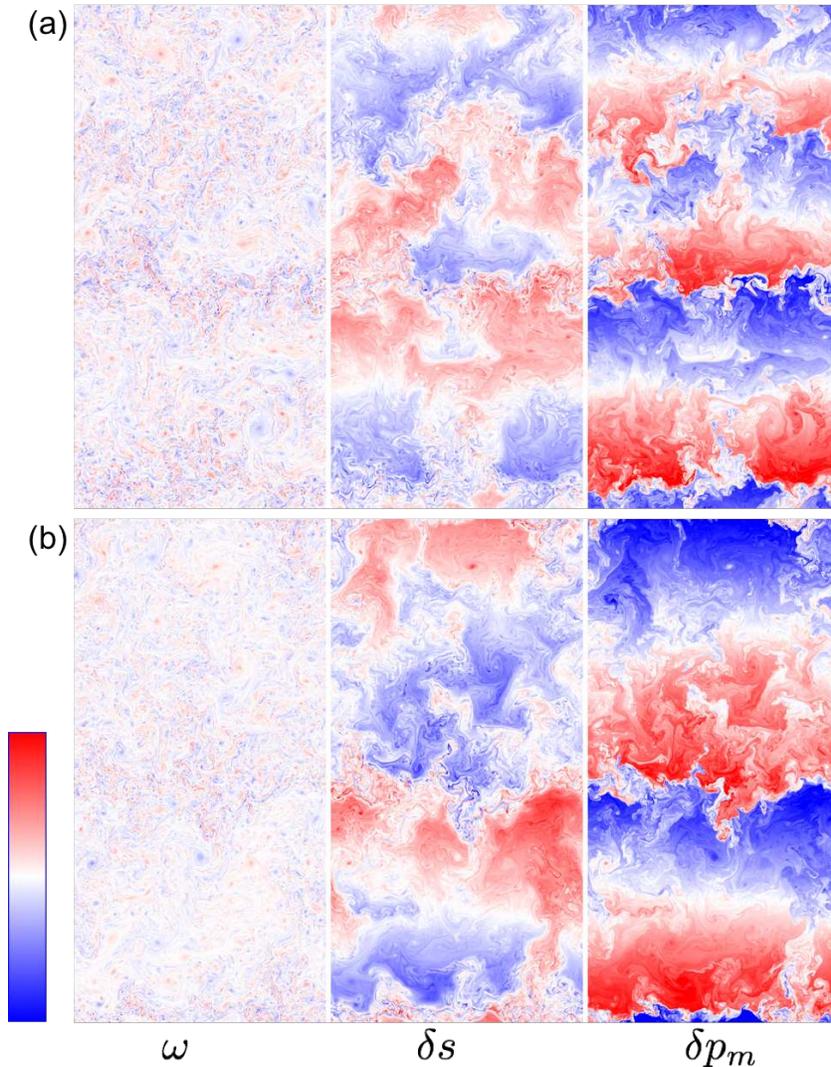


Figure 6. Snapshots at (a) $t = 8050$ and (b) $t = 11700$ of ω , δs and δp_m (all scaled independently) for the case of $R_0^{-1} = 1.5$. The color table ranges from blue (largest negative value) to red (largest positive value), with white denoting the zero value.

829 gleaned from observations reveals that our knowledge
 830 of mixing processes therein is incomplete. For example,
 831 when helioseismology probed the interior rotation profile
 832 of the Sun, it revealed the tachocline, raising new chal-
 833 lenges to our understanding of solar angular momentum
 834 transport. Similarly, the long-standing issue regarding
 835 solar lithium abundances challenges our understanding
 836 of chemical mixing processes. Such issues are clearly
 837 not confined to the Sun, with similar questions arising
 838 for many astrophysical bodies.

839 In situations where there is a clearly dominant mech-
 840 anism (such as convection), theories of the transport
 841 (such as mixing length theory) have readily emerged.
 842 On the other hand, when extra mixing is required to
 843 explain observations in stellar radiative zones, where no
 844 dominant transport mechanism is apparent, a taxonomy
 845 of potential mixing processes has more gradually been

846 unveiled. Perhaps the most notable recent advances
 847 concern processes introduced by rotational effects (ro-
 848 tational mixing), shear turbulence, double-diffusive con-
 849 vection, overshooting convection and gravity wave trans-
 850 port (see, e.g., Zahn 2008). The late Jean-Paul Zahn
 851 and collaborators published extensively on such work,
 852 and a perspective of the complexity of this taxonomy is
 853 afforded in Figure 1 of Mathis & Zahn (2005).

854 The role of the magnetic field in the dynamics of mix-
 855 ing can be two-fold. First, magnetic fields can poten-
 856 tially affect many of the proposed non-magnetic mix-
 857 ing processes. Magnetic fields often inhibit instabilities
 858 (see, e.g., Chandrasekhar 1961) and also therefore their
 859 transport and mixing properties. The effect of magnetic
 860 fields is often therefore considered in this constraining
 861 context. Second, magnetic fields can be a source of in-

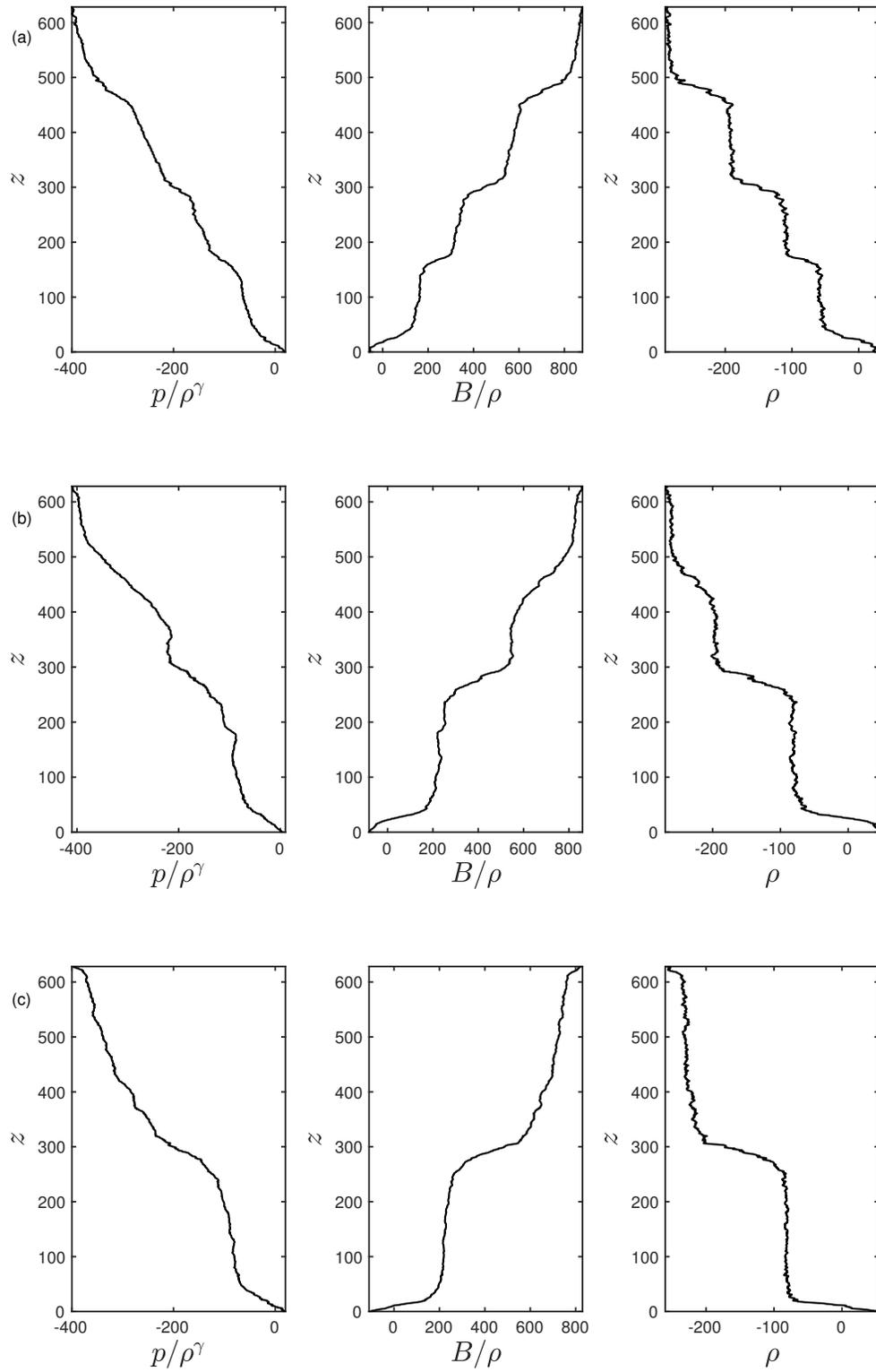


Figure 7. Horizontally averaged profiles of p/ρ^γ , B/ρ and ρ versus height at (a) $t = 3600$, (b) $t = 8050$, (c) $t = 11700$ (corresponding to the snapshots in Figures 3(c), 6a,b): $R_0^{-1} = 1.5$.

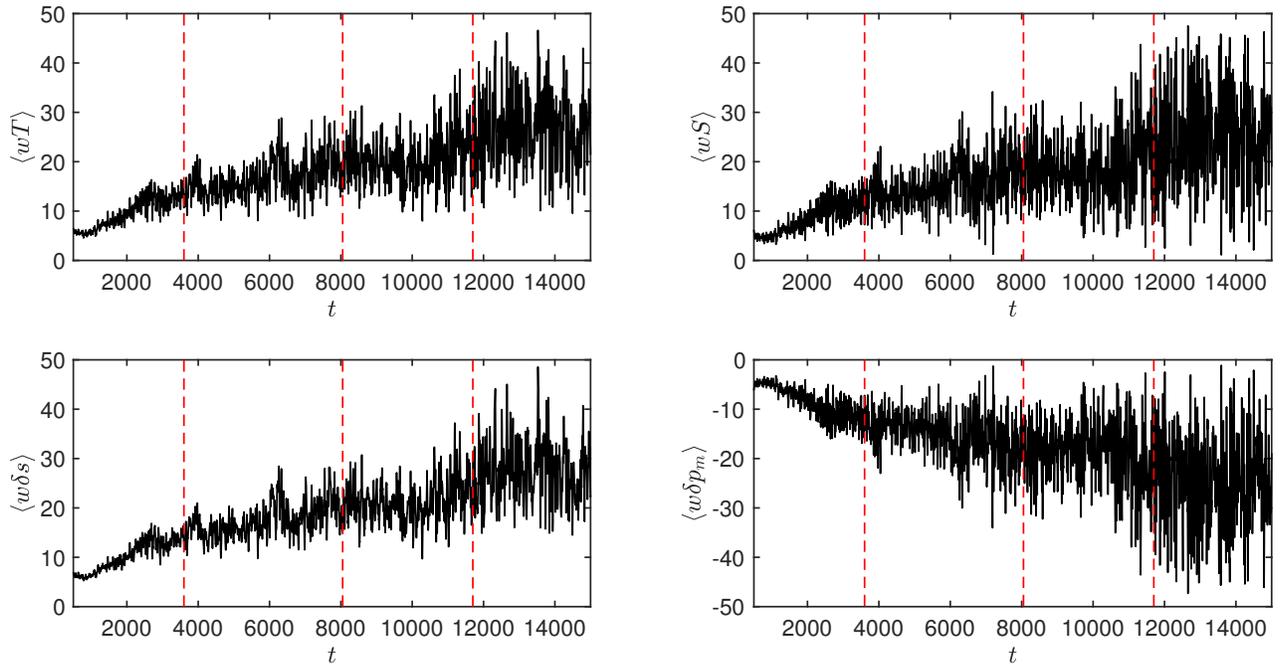


Figure 8. The top row shows the averaged vertical fluxes of T and S as a function of time, for all but the initial stages of the evolution, for the thermohaline problem with $R_0^{-1} = 1.5$. The bottom row shows the averaged vertical fluxes of δs and δp_m for the equivalent magnetic buoyancy problem. The dashed lines denote the times of the snapshots in Figures 3(c), 6(a,b).

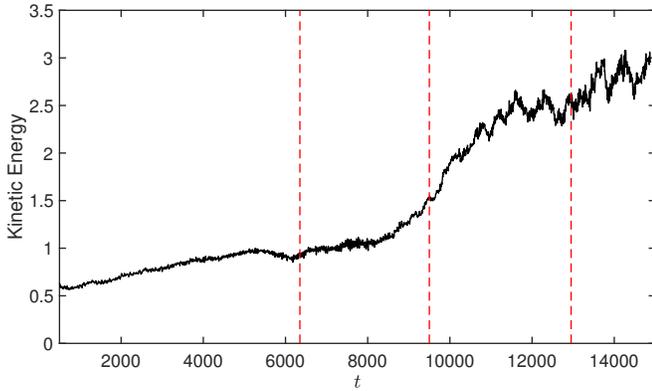


Figure 9. Kinetic energy as a function of time, for $t > 500$: $R_0^{-1} = 4$. The dashed lines denote the times of the snapshots in Figure 10.

2003). Another example that is more directly relevant to our study here is the work of Busso et al. (2007), which invokes magnetic buoyancy instabilities as a potential source of extra mixing (known in this context as ‘cool bottom processing’) to explain certain observed anomalies in low-mass red giant branch (RGB) and asymptotic giant branch (AGB) stars. Busso et al. (2007) ascribe the vertical transport required to explain the observations to the rise of thin buoyant flux tubes; they then infer the interior field strength necessary to produce these tubes at just the right rate to create the desired transport. Although only a phenomenological approach, this work ultimately imposes requirements on the interior fields that would lead to the necessary mixing. Our work is clearly directly relevant to this type of transport.

The overall aim of this paper has been to demonstrate that mixing induced by magnetic buoyancy instabilities could be a powerful and far more prevalent dynamical process than is currently widely recognized, and hence that it should be added to the overall catalogue of mixing processes. We have demonstrated that not only do magnetic buoyancy instabilities initiate extra mixing of magnetic and thermodynamic properties, but secondary instabilities to layered states can also significantly enhance that mixing. Furthermore, such mixing can be engendered under conditions that appear to be very stable in terms of the individual components.

In more detail, by exploiting the analogy derived by Spiegel & Weiss (1982), we have shown how the phenomenon of layering in thermosolutal convection implies the formation of layers in a system driven by magnetic buoyancy. In the astrophysically relevant regime in which the Prandtl number σ and diffusivity ratio τ are both small, the numerical evidence to date reveals that layering seems to occur more readily in the diffusive regime than in the fingering regime (Garaud 2018). In our translation of these results to the magnetic buoyancy case, we have thus concentrated on the diffusive regime. Although necessarily restricted to two-dimensional motions in pursuing the analogy, this has allowed us to conduct high resolution simulations in order to explore the regime of small σ and τ (though of course these are still much larger than the true astrophysical values). When both σ and τ are small, this leads to an enlarged region of instability (see expression (60)) and hence the potential for layering at higher values of R_0^{-1} .

In this paper, we have demonstrated, through numerical simulations, the formation of layers for $R_0^{-1} = 1.5$ and $R_0^{-1} = 4$ with $\sigma = \tau = 0.01$. From a thermohaline perspective, the two cases are very similar in many regards. However, when viewed from the magnetic perspective, the latter case is significant in that it falls in

stability themselves, thereby generating a further means of dynamical transport.

The latter context is the theme of this paper. Much of the research performed in this area has been directed at explaining the existence and geometry of observed magnetic fields. That is, the evolution and transport of the magnetic field itself has been the chief characteristic of interest, rather than any other induced mixing. For example, the stability and therefore the ultimate configuration of large-scale poloidal and toroidal fields in the deep interiors of stars has been investigated in a similar manner to the exploration of field configurations in plasma confinement devices (see, e.g., Markey & Tayler 1973; Pitts & Tayler 1985). Another substantial avenue of research has been devoted to dynamo instabilities (see, e.g., Moffatt & Dormy 2019). Generally driven by some combination of turbulence, rotation and shear, such instabilities can explain the initial generation of both small-scale and large-scale (mean) fields from weak seed fields, rather than the re-organization of large-scale fields as in Tayler instabilities. Magnetic buoyancy, the instability mechanism that is the subject of this paper, has been studied in the context of the creation and transport of compact magnetic flux structures from large-scale fields, in an effort to seek the origin of solar magnetic active regions and sunspots.

These examples all address the origin of certain magnetic field configurations rather than any induced transport of other ingredients, such as heat, angular momentum or chemical species. Perhaps the most well-known example in which magnetic field drives an instability that generates transport of another important quantity is the magneto-rotational instability (MRI), where the presence of a magnetic field instigates turbulence in an accretion disk, thereby allowing the turbulent transport of angular momentum and material (see, e.g., Balbus

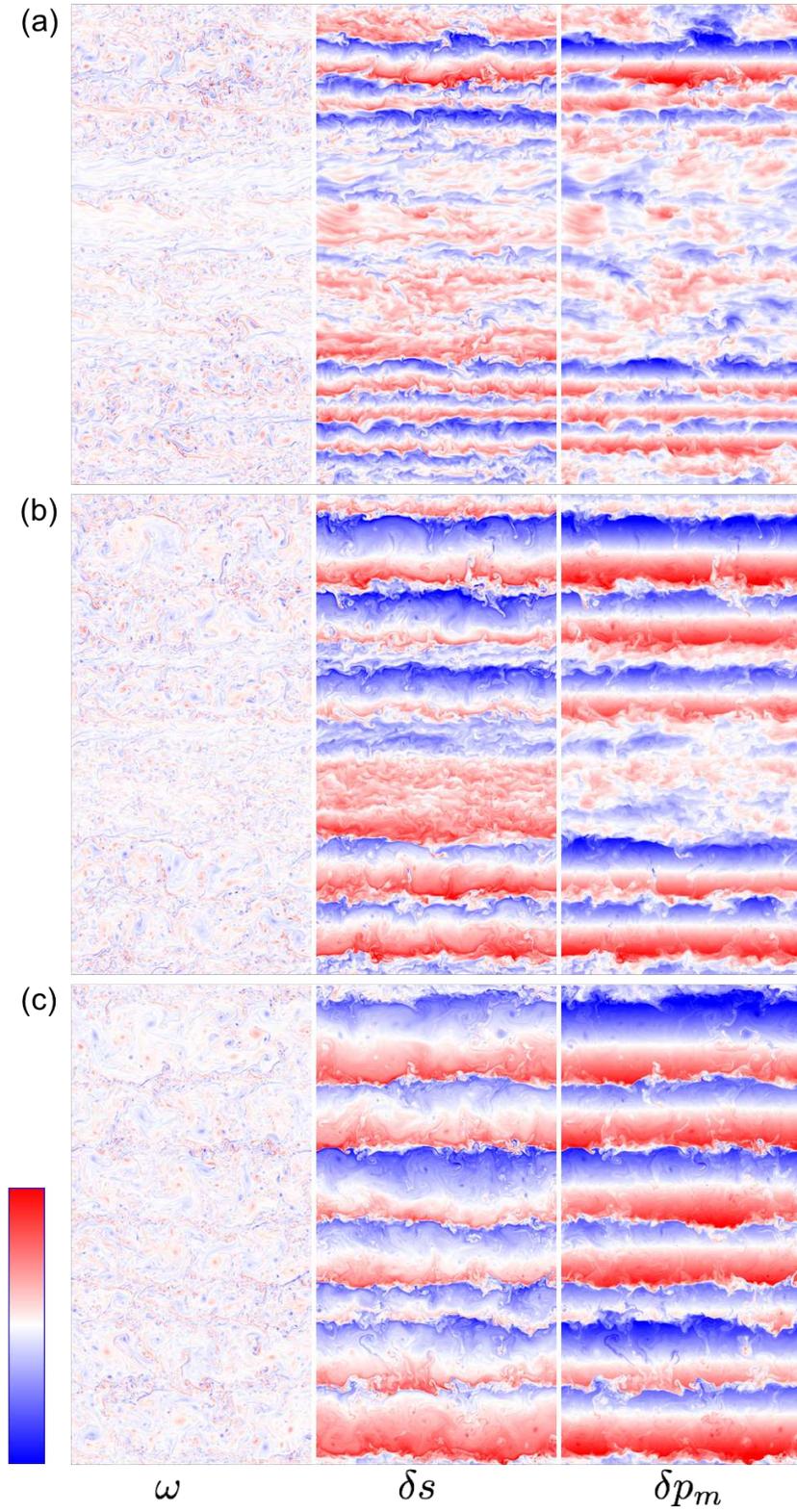


Figure 10. Snapshots at (a) $t = 6350$, (b) $t = 9500$ and (c) $t = 12950$ of ω , δs and δp_m (all scaled independently) for the case of $R_0^{-1} = 4$. The color table ranges from blue (largest negative value) to red (largest positive value), with white denoting the zero value.

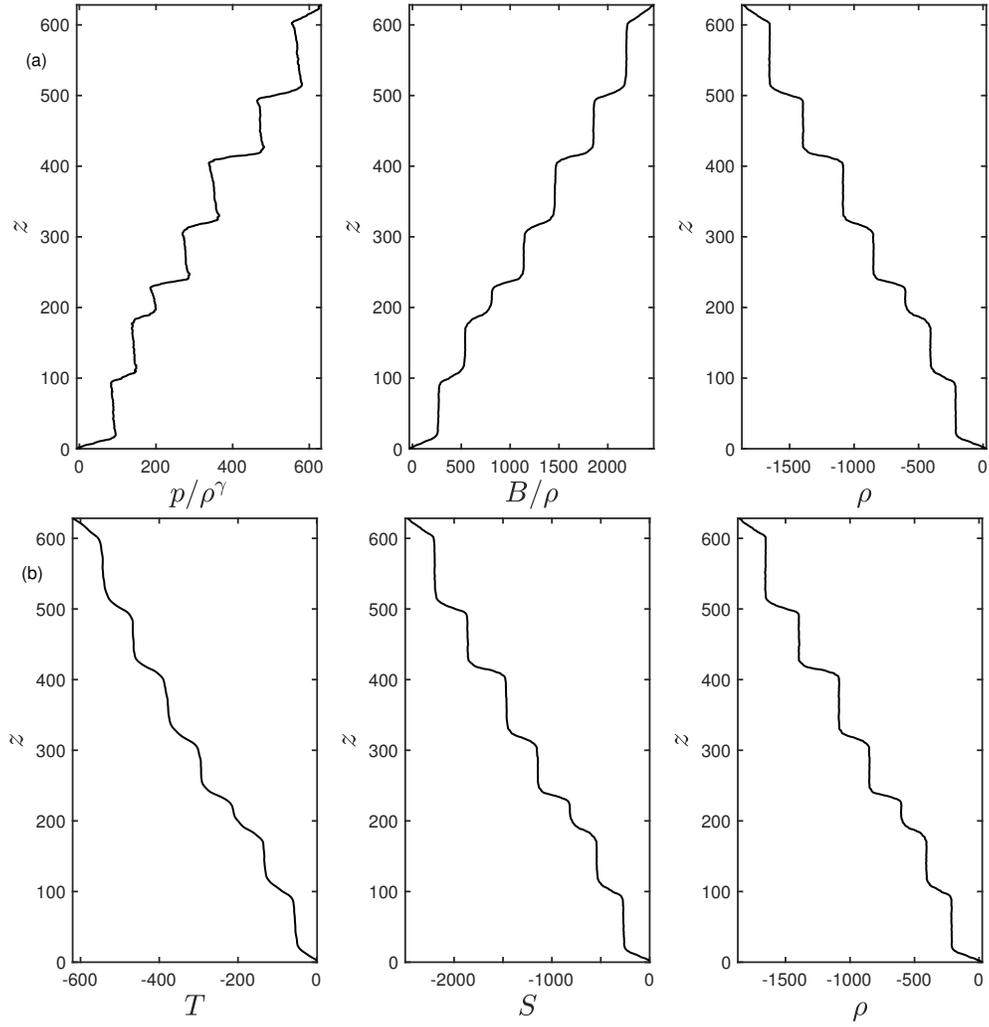


Figure 11. (a) Horizontally averaged profiles of p/ρ^γ , B/ρ and ρ versus height at $t = 12950$ (corresponding to the snapshots in Figure 10(c)): $R_0^{-1} = 4$. (b) The horizontally averaged profiles of T , S and ρ for the equivalent thermohaline problem.

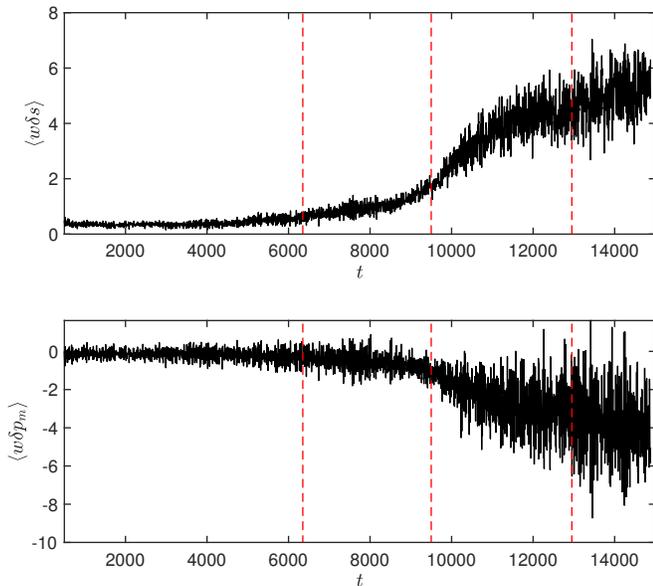


Figure 12. Averaged vertical fluxes of δs and δp_m as functions of time, for all but the initial stages of the evolution: $R_0^{-1} = 4$. The dashed lines denote the times of the snapshots in Figure 10.

what one might regard as the ‘stable-stable’ quadrant of magnetic buoyancy instabilities, where the system is described by a subadiabatic entropy gradient and by B/ρ decreasing with height (see Figure 1(b)). Thus, at first glance, one might not even expect instability in this regime, let alone layering. In each of these two cases, we have shown that the initial instability to simple vertical ‘elevator’ modes quickly gives way to turbulence, which subsequently evolves to layered states (see Figures 3, 6, 10). These then slowly merge to form wider and wider layers. As time progresses, each new scenario possesses stronger transport properties, with the layered states being significantly (5-6 times) more efficient than the more homogeneous state that emerges after the initial instability (see Figures 8, 12).

It is important to consider where magnetic layer formation and the associated enhanced transport may be of significance astrophysically. As shown in Figure 1, diffusive magnetic buoyancy instabilities are found in both the first and fourth quadrants in the (R_b, R_t) plane; furthermore, as we have demonstrated, the parameter regime of layer formation extends from close to the line of neutral stability in the first quadrant all the way into the fourth quadrant. In stellar convective zones, with strongly supercritical turbulent convection, magnetic buoyancy will presumably not be a major player. However, in radiative zones, where conditions are more quiescent and timescales much longer, there is the opportunity for magnetic buoyancy to act as the agent for

layer formation. For small values of σ and τ , and considering elevator modes, inequality (66) shows that with $\alpha > 0$ and $\beta > 0$, oscillatory convection occurs for

$$R_2^{-1} \gtrsim \frac{\gamma}{(\gamma - 1)}, \quad (69)$$

or, in dimensional terms,

$$(\gamma - 1) \frac{ga^2}{c^2} \frac{d}{dz} \ln \left(\frac{B}{\rho} \right) > N^2 \quad (70)$$

(cf. inequality (8)). The magneto-Boussinesq approximation holds under the assumption that $a^2 \ll c^2$. Inequality (70) thus requires that for instability the subadiabatic gradient is smaller than the gradient of B/ρ . As explained by Corfield (1984) and Bowker et al. (2014), this condition is indeed necessary for asymptotic consistency of the magneto-Boussinesq approximation². Application of inequality (70) to the Sun suggests that the most favorable conditions for oscillatory instability with $\alpha > 0$ and $\beta > 0$ will be towards the top of the radiative zone, where the subadiabatic gradient is indeed weak, and where the strong magnetic field in the overlying tachocline will inevitably lead, locally, to a magnetic field that increases with height. Although, it is hard to be definitive in asserting that layering will occur — owing to the impossibility of simulations in the astrophysical regime or the lack of a rigorous theory — we have shown unambiguously in § 4.2 that for small σ and τ (though not astrophysically small), pronounced layering does indeed occur with $\alpha > 0$ and $\beta > 0$ ($Rb > 0$, $Rt < 0$).

A non-solar application (but still drawing phenomenologically from our understanding of the solar radiative zone and tachocline dynamics) is provided by the study of Busso and collaborators (Busso et al. 2007), discussed above. Our work shows that the transport in a layered context is substantially greater than the simple advective transport associated with the small-scale magnetic structures that initially emerge, and hence may well be very different from that of highly conceptual flux tubes.

We conclude by considering future directions for the study of magnetic layering. Our approach in this study has been to exploit the analogy of Spiegel &

² To be precise, the magneto-Boussinesq approximation is based on an expansion in two small parameters: $\varepsilon_1 = d/H$ (H is scale height) and $\varepsilon_2 = \delta\rho/\rho_0$, with $\varepsilon_1 \gtrsim \varepsilon_2$. The ratio a^2/c^2 is $O(\varepsilon_2/\varepsilon_1)$ and the subadiabatic gradient is $O(\varepsilon_2)$, not $O(\varepsilon_1)$ as one might naively suppose.

Weiss (1982), thus allowing us to translate between two-dimensional thermohaline convection and interchange modes of magnetic buoyancy instability under the magneto-Boussinesq approximation. It is though important to look beyond the constraints imposed by the analogy. For thermohaline convection, the nature of two-dimensional versus three-dimensional simulations was investigated in some detail by Garaud & Brummell (2015). It was found that two-dimensional simulations of the diffusive case (as performed here) were reasonably representative of the full three-dimensional dynamics (although more care is definitely needed for the fingering case). More significantly, for the magnetic system in general, we expect similarities between the two- and three-dimensional evolutions since the basic state field imposes a preferred horizontal direction. The most readily excited linear mode of magnetic buoyancy instability, although three-dimensional, has a very long wavelength in the direction of the imposed field (see, e.g., Acheson 1979). Furthermore, the nonlinear evolution of three-dimensional magnetic buoyancy instabilities (Matthews et al 1995) has many features in common with that of two-dimensional (interchange) modes (Cattaneo & Hughes 1988). The most important consequence of relaxing the Boussinesq approximation (through considering either the anelastic approximation or the full compressible equations) is the introduction of a preferred lengthscale into the system, through, for example, the pressure scale height. It is clearly important to understand the influence of this scale on the layering and transport processes. **Also, in seeking more realism, it should be noted that, even in the extension to 3D, the current problem as set up only examines the instability of an initially**

unidirectional field, and more complex initial field geometries should be studied.

Explaining the formation, maintenance and transport properties of layers is an area of intense current research, not just for double-diffusive systems, but in the contexts of forced stratified turbulence, of planetary jet formation (where the jets are manifestations of a potential vorticity staircase) and of the corrugated shear flow in fusion plasmas (the $\mathbf{E} \times \mathbf{B}$ staircase). There is still considerable debate over the underlying physical mechanisms and, indeed, whether there is a common thread between the different systems that exhibit layering. In terms of magnetic buoyancy layering, the important challenge ahead is to build upon the numerical simulations to devise a theoretical model valid in the regime $R_1^{-1} \gg 1, \sigma \ll 1, \tau \ll 1$ that can provide an estimate of where layering is to be expected and, in such cases, how the turbulent transport depends on the parameters of the problem.

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APPENDIX

A. DERIVATION OF EXPRESSIONS FOR p/ρ^γ , B/ρ AND $\hat{\rho}$

Here we provide the derivations of expressions (43), (44), (46) for p/ρ^γ , B/ρ and the density deviation $\hat{\rho}$. Under the Boussinesq approximation, in which scale heights of the basic state are large, we may approximate the basic state as being linear in z . Thus, to first order in small quantities, we may write, before any rescaling or non-dimensionalization,

$$\frac{p}{\rho^\gamma} = \frac{p_0}{\rho_0^\gamma} \left(1 + z \frac{d}{dz} \ln \left(\frac{p}{\rho^\gamma} \right) + \frac{\delta p}{p_0} - \gamma \frac{\delta \rho}{\rho_0} \right). \quad (\text{A1})$$

Recall that, under the magneto-Boussinesq approximation, it is the variation in *total* pressure that is small, and, crucially, much smaller than the individual variations of the gas and magnetic pressures; i.e. $\delta p \approx -\delta p_m$. On using

this result, together with the perfect gas law, expression (A1) can be written as

$$\frac{p}{\rho^\gamma} = \frac{p_0}{\rho_0^\gamma} \left(1 + z \frac{d}{dz} \ln \left(\frac{p}{\rho^\gamma} \right) + \gamma \frac{\delta T}{T_0} + (\gamma - 1) \frac{\delta p_m}{p_0} \right) \quad (\text{A2})$$

$$= \frac{p_0}{\rho_0^\gamma} \left(1 + z \frac{d}{dz} \ln \left(\frac{p}{\rho^\gamma} \right) + \gamma \frac{\delta T^*}{T_0} + \frac{(\gamma - 1) \delta p_m}{(1 - \tau) p_0} \right), \quad (\text{A3})$$

1096 on using the definition for δT^* , given by (20).

1097 The consistent scaling is to scale p/ρ^γ with $(p_0/\rho_0^\gamma)(d|\beta^*|/T_0)$, thus giving the dimensionless expression (43):

$$\frac{p}{\rho^\gamma} = \text{const.} + \left(\text{sgn}(\beta^*)\gamma + \frac{\text{sgn}(\alpha)(\gamma - 1)}{R_1(1 - \tau)} \right) z + \gamma \frac{\delta T^*}{T_0} + \frac{(\gamma - 1) \delta p_m}{(1 - \tau) p_0}. \quad (\text{A4})$$

1098 Thus for the *diffusive regime*, in which $\text{sgn}(\alpha) = +1$ and $\text{sgn}(\beta^*) = -1$, we obtain

$$\frac{p}{\rho^\gamma} = \text{const.} - \left(\gamma - \frac{1}{R_1} \frac{(\gamma - 1)}{(1 - \tau)} \right) z + \gamma \frac{\delta T^*}{T_0} + \frac{(\gamma - 1) \delta p_m}{(1 - \tau) p_0}. \quad (\text{A5})$$

1099 Similarly, before any rescaling or non-dimensionalization, we can write,

$$\frac{B}{\rho} = \frac{B_0}{\rho_0} \left(1 + z \frac{d}{dz} \ln \left(\frac{B}{\rho} \right) + \frac{\delta p_m}{B_0^2/\mu_0} \right). \quad (\text{A6})$$

1100 Note that in the magneto-Boussinesq approximation, the term involving $\delta \rho$ is formally smaller (by a ratio of the square
1101 of the Alfvén speed to the square of the sound speed) and hence is neglected (see Hughes & Weiss 1995). Since we have
1102 chosen to scale δp_m with $|\beta^*|dp_0/T_0$, it is consistent to scale B/ρ with $|\beta^*|dp_0\mu_0/T_0B_0\rho_0$. This leads to the following
1103 expression for (dimensionless) B/ρ :

$$\frac{B}{\rho} = \text{const.} + \frac{\text{sgn}(\alpha)}{R_1} z + \delta p_m. \quad (\text{A7})$$

1104 In the diffusive regime, this becomes

$$\frac{B}{\rho} = \text{const.} + \frac{1}{R_1} z + \delta p_m. \quad (\text{A8})$$

1105 In thermohaline convection, determining an expression for the overall density is straightforward. For magnetic
1106 buoyancy, it is a little more involved since the basic state density profile involves quantities that are not used in the
1107 scaling (unlike in thermohaline convection where the basic state density depends on the temperature and salinity
1108 profiles, which then go into R_0). To first order in small quantities, we may write

$$\rho = \rho_0 \left(1 + z \frac{d}{dz} \ln \rho - \frac{\delta T}{T_0} - \frac{\delta p_m}{p_0} \right). \quad (\text{A9})$$

1109 We may express the equation for the magneto-hydrostatic basic state,

$$\frac{d}{dz} \left(p + \frac{B^2}{2\mu_0} \right) = -\rho g, \quad (\text{A10})$$

1110 as

$$\left(\gamma + \frac{B^2}{2\mu_0 p} \right) \frac{d \ln \rho}{dz} = -\frac{\gamma \beta}{T_0} - \frac{\alpha}{p} - \frac{\rho g}{p}, \quad (\text{A11})$$

1111 where α and β are as defined in § 2.2. The second term in the bracket can be neglected (Alfvén speed \ll sound speed).

1112 Thus expression (A9) becomes

$$\frac{\rho}{\rho_0} = 1 - \left(\frac{\beta}{T_0} + \frac{\alpha}{\gamma p} + \frac{\rho g}{\gamma p} \right) z - \frac{\delta T}{T_0} - \frac{\delta p_m}{p_0}. \quad (\text{A12})$$

1113 There is no straightforward way of dealing with the third term in the bracket, since it brings in quantities that are not
1114 used in the scalings of the variables. We therefore consider deviations away from the reference state defined by

$$\rho = \text{const.} - \frac{\rho_0^2 g}{\gamma p_0} z. \quad (\text{A13})$$

1115 If we denote the deviation from this state as $\hat{\rho}$ then

$$\frac{\hat{\rho}}{\rho_0} = - \left(\frac{\beta}{T_0} + \frac{\alpha}{\gamma p_0} \right) z - \frac{\delta T}{T_0} - \frac{\delta p_m}{p_0}. \quad (\text{A14})$$

1116 On substituting for β^* and δT^* , and scaling $\hat{\rho}$ with $d\rho_0|\beta^*|/T_0$, we obtain the following dimensionless expression for
1117 $\hat{\rho}$:

$$\hat{\rho} = - \left(\text{sgn}(\beta^*) + \frac{(\gamma - \tau)}{\gamma(1 - \tau)} \frac{\text{sgn}(\alpha)}{R_1} \right) z - \delta T^* + \frac{(\gamma - \tau)}{\gamma(1 - \tau)} \delta p_m. \quad (\text{A15})$$

REFERENCES

- 1118 Acheson, D. J. 1979, *SoPh*, 62, 23
 1119 Balbus, S. A. 2003, *Ann. Rev. Astron. Astrophys.*, 41:
 1120 555597
 1121 Batiste, O., Knobloch, E., Alonso, A. & Mercader, I. 2006,
 1122 *JFM*, 560, 149
 1123 Bowker, J.A., Hughes, D.W. & Kersalé, E. 2014, *GApFD*,
 1124 108, 553
 1125 Busso, M., Wasserburg, G.J., Noll, K.M & Calandra, A.
 1126 2007, *ApJ*, 671:802
 1127 Cattaneo, F. & Hughes, D. W. 1988, *JFM*, 196, 323
 1128 Chandrasekhar, S. 1961, *Hydrodynamic and*
 1129 *Hydromagnetic Stability* (Oxford: Clarendon Press)
 1130 Corfield, C. N. 1984, *GApFD*, 29, 19
 1131 Garaud, P. 2018, *Ann. Rev. Fluid Mech*, 50, 275
 1132 Garaud, P. & Brummell, N.H. 2015, *ApJ*, 815, 42
 1133 Gilman, P. A. 1970, *ApJ*, 162, 1019
 1134 Gough, D. O. 2007, in *The Solar Tachocline*, ed. D. W.
 1135 Hughes, R. Rosner, & N. O. Weiss (Cambridge:
 1136 Cambridge Univ. Press), 3
 1137 Hughes, D.W. 1985, *GApFD*, 32, 273
 1138 Hughes, D. W. 2007, in *The Solar Tachocline*, ed. D. W.
 1139 Hughes, R. Rosner, & N. O. Weiss (Cambridge:
 1140 Cambridge Univ. Press), 275
 1141 Hughes, D. W., & Cattaneo, F. 1987, *GApFD* 39, 65
 1142 Hughes, D. W., & Proctor, M. R. E. 1988, *AnRFM*, 20, 187
 1143 Hughes, D. W., & Weiss, N. O. 1995, *JFM*, 301, 383
 1144 Langer, N. 1991, *A&A*, 252, 669
 1145 Markey, P. & Tayler, R.J. 1973, *MNRAS*, 173, 77
 1146 Mathis, S. & Zahn, J.-P. 2005, *A&A*, 440, 653
 1147 Matthews, P. C., Hughes, D. W., Proctor, M. R. E. 1995,
 1148 *ApJ*, 448, 938
 1149 Merryfield, W. J. 2000, *JPO*, 30, 1046
 1150 Mirouh, G. M., Garaud, P., Stellmach, S., Traxler, A. L.
 1151 Wood, T. S. 2012, *ApJ*, 750:61
 1152 Moffatt, H. K. & Dormy, E. 2019 *Self-Exciting Fluid*
 1153 *Dynamos* (Cambridge: Cambridge University Press)
 1154 Newcomb, W. A. 1961, *PhFl*, 4, 391
 1155 Parker, E. N. 1966, *ApJ*, 145, 811
 1156 Phillips, O. M. 1972, *DSR*, 19, 79
 1157 Pitts, E. & Tayler, R.J. 1985, *MNRAS*, 216, 139
 1158 Posmentier, E. S. 1977, *JPO*, 7, 298
 1159 Radko, T. 2003, *JFM*, 497, 365
 1160 Radko, T. 2013, *Double-Diffusive Convection* (Cambridge:
 1161 Cambridge University Press)
 1162 Rahmstorf, S. 2006, in *Encyclopedia of Quaternary*
 1163 *Sciences* (ed. S. A. Elias) (Elsevier)
 1164 Rosenblum, E., Garaud, P., Traxler, A. & Stellmach, S.
 1165 2011, *ApJ*, 731:66
 1166 Schmitt, R. W. 1994, *AnRFM*, 26, 255
 1167 Schmitt, R. W., Ledwell, J.R., Montgomery, E.T., Polzin,
 1168 K.L., Toole, J.M. 2005 *Sci*, 308, 685
 1169 Spiegel, E. A., & Veronis, G. 1960 *ApJ*, 131, 442
 1170 Spiegel, E. A., & Weiss, N. O. 1982, *GApFD*, 22, 219
 1171 Stellmach, S., Traxler, A., Garaud, P., Brummell, N. &
 1172 Radko, T. 2011, *JFM*, 677, 554
 1173 Stern, M. E. 1969, *JFM*, 35, 209
 1174 Stern, M. E., & Radko, T. 1998, *JMR*, 56, 157
 1175 Thomas, J. H., & Nye, A. H. 1975, *PhFl*, 18, 490
 1176 Timmermans, M.-L., Toole, J., Proshutinsky, A.,
 1177 Krishfield, R. & Plueddemann, A. 2008, *JPO*, 38, 133
 1178 Turner, J. S. 1973, *Buoyancy Effects in Fluids* (Cambridge:
 1179 Cambridge University Press)
 1180 Weiss, N. O. & Proctor, M. R. E. 2014, *Magnetoconvection*
 1181 (Cambridge: Cambridge University Press)
 1182 Zahn, J.-P. 2008, in *The Art of Modelling Stars in the 21st*
 1183 *Century*, *Proc. IAU Symp.* 252, ed. L. Deng & K.L. Chan
 1184 Zhurbas, V. M., & Ozmidov, R. V. 1983, *IzAOP*, 19, 977