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1	A Practical Method to Determine the Five-Parameter Orientation of Intragranular
2	Boundaries in Polycrystals
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Abstract-Intragranular boundaries are important features of polycrystalline materials and impact on 9 many physical and chemical properties. Knowledge of their physical orientation is often crucial to 10 explain such properties. However, it has proved difficult to determine the complete orientation of 11 12 intragranular boundaries, which involves the misorientation angle and axis about which the adjacent crystal lattices need to be rotated to bring them into coincidence and also the physical orientation of 13 the boundary plane, expressed by the plunge and azimuth of its normal; five parameters in total. 14 Here we present a simple and practical manual method to determine the complete intragranular 15 boundary orientation in any crystal system. The method is developed on geometrical relationships 16 exhibited between electron channelling patterns across a common boundary but then extended for 17 use with electron backscattered diffraction patterns. The method recognises the 18 channelling/diffraction band, equivalent to a crystal lattice plane, not displaced across a boundary; 19 the boundary rotation axis must be the normal to this plane. Geometrical relationships between the 20 boundary trace, the non-displaced band/plane and their respective plane normals constrain boundary 21 orientation to two alternative symmetrically equivalent solutions and are evaluated via 22 23 stereographic projection. The choice of solution is guided by comparison with the presence or absence of a similarly oriented band/plane observed in the original channelling/diffraction patterns. 24 The method therefore conforms to the low-index crystallographic lattice plane and dislocation 25 model for intragranular boundary formation and defines boundary orientation in terms of total 26 angular misorientation due to tilt and twist components and the orientation of the boundary plane. 27 Examples of intragranular boundary orientation determination using this method are provided in 28 olivine. Results are compared to and differ from those obtained via conventional misorientation 29 analysis, which only rotates adjacent crystal lattices into parallelism and does not consider boundary 30 31 plane orientation. Potential extrapolation of the new method to intergranular boundaries is also considered. 32

#### **1. INTRODUCTION**

3

34 Intragranular boundaries are important and common features of polycrystalline materials, such as metals, ceramics and rocks. Their presence impacts critically on many physical and chemical 35 properties, processes and behaviours, including: overall strength and failure, various types of 36 conductivity, diffusion and creep mechanisms, recrystallisation and recovery processes, corrosion 37 and precipitation behaviours, etc (e.g. Randle, 1993). Thus, the nature of intragranular boundaries is 38 crucial to understanding the ultimate behaviour of polycrystalline materials. However, intragranular 39 boundaries often exist as rather obscure elements within more detailed microstructural 40 41 characterisations that focus on the properties of the regions they surround, such as the crystal texture or fabric (as defined by the crystallographic preferred orientation or CPO), subgrain size, 42 43 etc. Perhaps the main reason why intragranular boundaries are often somewhat neglected is due to 44 problems associated with defining their complete physical and/or crystallographic orientations. On the one hand, intragranular boundaries are defects that perturb the crystallographic structure of 45 46 polycrystalline materials; their definition must include recognition of this impact on the crystallography. Convention therefore recognises the misorientation angle/axis ( $\theta$ /<uvw>) pair, or 47 the crystallographic direction about which one crystal lattice must be rotated by a (minimum) angle 48 to bring it into coincidence with an adjacent lattice. As this operation represents a pure rotation of 49 50 one crystal coordinate system with respect to the other (Fig. 1a), it can be described by a rotation 51 transformation matrix  $(g_m)$ , such that,

52 
$$\{C_i\} = g_m\{C_j\}$$
 (1)

where {C<sub>i</sub>} and {C<sub>j</sub>} are the respective adjacent crystal coordinate systems. Misorientation
angle/axis pairs comprise three degrees of freedom: the plunge/azimuth of the axis and the angle
(e.g. Warrington and Bufalini, 1971; Grimmer et al., 1974; Mainprice et al., 1993; Lloyd et al.,
1997; Randle, 2003). On the other hand, intragranular boundaries are also physical features (i.e.
'planes'), the orientation of which is typically crystallographically constrained but is not necessarily

crystallographically coincident. The physical orientation of the intragranular boundary plane is
defined by the plunge/azimuth of the normal to the plane, which involves two degrees of freedom
(Fig. 1b); however, it can also be represented crystallographically (i.e. <uvw>) with respect to one
and/or other of the two adjacent crystal orientations.

Thus, the complete orientation of intragranular boundaries involves five independent parameters
(e.g. Sutton and Ballufi, 1995; Kim et al., 2005; Randle, 2006; Rohrer and Randle, 2009;
Ratanaphan et al., 2014; Sutton et al., 2015; Dash et al., 2017). However, whilst it is always

65 possible to determine the misorientation angle/axis pair, particularly via electron backscattered



**Figure 1**. Intragranular boundary relationships. (a) Definition of the crystallographic orientations of two subgrains A and B and the alternative three rotations ( $\phi_1$ ,  $\Psi$ ,  $\phi_2$ ) required to bring their respective lattices XYZ and X'Y'Z' into coincidence, as represented by the misorientation angle/axis pair. (b) Regions A, B share common boundary plane (A<sub>hkl</sub>, B<sub>hkl</sub>), trace (t) and normal (h). (c) Pure tilt boundary (r, rotation axis;  $\omega$ , rotation angle). (d) Pure twist

boundary.

diffraction (EBSD) in the scanning electron microscope (SEM), the direction of the boundary plane 66 normal has proved to be notoriously difficult to establish. Although various methods, techniques 67 and approaches do exist (e.g. optical microscopy universal stage, transmission electron microscopy, 68 69 focussed ion beam, X-ray and neutron techniques, etc.), they tend to be restrictive, laborious, complex and/or expensive (e.g. Rohrer & Randle, 2009). An exception to these constraints is 70 71 provided by the grain boundary character distribution (GBCD) method (e.g. Saylor et al., 2004; 72 Rohrer et al., 2004b), which combines EBSD misorientation analysis and statistical stereology of 73 boundary traces. However, GBCD has been reserved mostly for (cubic) metals and ceramics (e.g. Saylor et al. 2004), with only a single geological application on olivine (Marquardt et al., 2015). 74 More general characterisation of boundary microstructural data has recently become more common 75 using SEM/EBSD. In part, this stems from the availability of free resource software. Perhaps the 76 77 most readily available is MTex (e.g. Bachmann et al., 2010) and in particular its Tilt and Twist Boundaries script (https://mtex-toolbox.github.io/TiltAndTwistBoundaries.html). Analytical scripts 78 based on the GBCD method for estimating the five grain boundary parameters are also available 79 (http://mimp.materials.cmu.edu/~gr20/Grain\_Boundary\_Data\_Archive/). In addition, an alternative 80 but related approach using the weighted mean Burger's vector has been suggested by Wheeler et al. 81 (2009), with a recent modification within the MTex software package (Wieser et al., 2020). 82 However, all approaches are based a priori on the recognition of tilt and twist boundaries based on 83 accurate definition of the misorientation axis; they do not consider either general boundaries or the 84 orientation of the boundary plane normal. Thus, they recognise only three of the five parameters 85 necessary to fully define an intragranular boundary. 86 87 The physical orientation of the boundary between two intragranular domains (A, B) can be

described crystallographically relative to either domain (i.e.  $A_{hkl}$ ,  $B_{hkl}$ ; see Fig. 1b). The boundary plane can also be described by its trace ( $t_{A=B}$ ) and normals ( $h_{A,B}$ ) relative to either grain. As these are perpendicular,

91 
$$h_{A,B} \cdot t_{A,B} = 0$$
 (2)

Pure tilt intragranular boundaries (Fig. 1c) develop by the progressive addition of, in principal, a
single set of edge dislocations to achieve the misorientation between adjacent domains (e.g.
Burgers, 1939a, b and 1940; Buranova, 1940; Shockley and Read, 1949; Read and Shockley, 1950).
A rotation angle (ω) and axis (r) can therefore be recognised. As the rotation axis lies within the
boundary plane and hence perpendicular to the boundary plane normal (h; Fig. 1c),

$$99 r \cdot h = 0 (3)$$

In contrast, pure twist intragranular boundaries (Fig. 1d) develop via (at least) two sets of (not
necessarily orthogonal) screw dislocations to achieve the misorientation between adjacent domains
(e.g. Bragg, 1940; Burgers, 1940; Shockley and Read, 1949; Read and Shockley, 1950). In this
case, the rotation axis is perpendicular to the boundary plane and hence parallel to its normal, such
that,

$$105 r \cdot h = 1 (4)$$

In practice most boundaries are probably combinations of tilt and twist components necessary tocreate the 'best-fit' between adjacent domains. Thus,

108 0 (pure tilt) 
$$\leq$$
 TTC  $\leq$  1 (pure twist) (5)

where TTC is the vector product of the rotation axis (r) and boundary plane normal (h) and is
known as the boundary tilt-twist component index (Amouyal et al., 2005). However, and in spite of
the apparent simplicity of Eqns 3 – 5, a boundary separating two adjacent misoriented crystal
lattices can have an effectively infinite number of physical orientations, defined by its planar
normal (Fig. 1b - d). Herein lies the inherent problem and concomitant requirement for defining the
complete five-parameter orientation of intragranular boundaries.

115 In this contribution, we propose a simple practical solution for the definition of the five-parameter

orientation of intragranular boundaries. Our solution arises from the recognition that all 116 intragranular boundaries involve four common geometrical parameters (Fig. 1): boundary trace (t), 117 boundary plane normal (h), rotation axis (r) and rotation angle ( $\omega$ ). In addition, the basic 118 requirement involves knowledge of the crystallographic orientation of the domains on either side of 119 a boundary, as well as the (crystallographic) orientation of the boundary plane normal direction. 120 Whilst we derive our solution from original observations made via SEM back-scattered electron 121 (BSE) electron channelling (EC), we subsequently demonstrate not only how it can be adapted for 122 EBSD data but also suggest how it may be extended to (some) intergranular boundaries. 123

124

# 2. METHOD

## 125 **2.1 Electron channelling**

SEM/EC (e.g. Hirsch et al., 1962; Venables and Harland, 1973; Joy, 1974; Goldstein and Yakowitz, 126 127 1975; Joy et al., 1982) is a related technique to SEM/EBSD (e.g. Dingley, 1989; Prior et al. 1999) and 2009) for determining crystallographic orientations via electron channelling patterns (ECP) 128 rather than EBSD patterns. It also provides images based on sample crystallographic microstructure 129 130 via BSE 'orientation contrast' (OC), also known as electron or orientation channelling contrast imaging (ECCI/OCCI) in recent literature (e.g. Zaefferer and Elhami, 2014); such images (e.g. Fig. 131 2a) are equivalent to EBSD fore-scattered electron (FSE) images. The principal differences between 132 EC and EBSD patterns are as follows. 133

(1) In EBSD, a vertical electron beam strikes a sample tilted at ~65-70°, whereas for EC a vertical
electron beam is rocked about a fixed point on a horizontal sample. (2) The angular spread of
patterns can approach ~100° in EBSD (Fig. 2b, c) but is typically only up to ~24° in EC (Fig. 2d, e),
as determined by the rocking angle (±RA) of the incident electron beam; it is much easier therefore
to identify (index) EBSD patterns, particularly via automation and in phases with low crystal
symmetry. (3) However, the smaller angular spread for ECP results in better angular resolution

compared with EBSD patterns (Fig. 2b-e). (4) In contrast, the maximum spatial resolution of ECP
formation is from 1 - 10µm due to aberrations in the SEM objective lens, which is at least an order
of magnitude greater than conventional EBSD. (5) Because ECPs are produced by rocking a
stationary electron beam about a point on the sample surface, a one-to-one relationship is



**Figure 2**. Comparison between electron backscattered diffraction (EBSD, left) and electron channelling (EC, right) patterns from pyrite. (a) EC orientation contrast (OC) and/or EBSD fore-scattered electron (FSE) image; note adjacent regions A and B separated by a distinct boundary. (b) and (c) Individual EBSD patterns from regions B and A respectively cover ~70° of the crystal structure. (d) and (e) Individual EC patterns (ECP) from regions B and A respectively cover only ~20° (i.e. electron beam 'rocking angle', RA, = ±10°) of the crystal structure. (f) The EBSD pattern from the boundary between A and B is formed by the complete superposition of patterns (b) and (c), resulting in an indistinct image. (g) The ECP from the boundary between A and B is composed of the relevant halves of patterns (d) and (e), resulting in a distinct image consisting of recognisable 'partial' patterns.

maintained between each point on the surface and in the pattern such that 'partial' ECPs
characteristic of the crystallographic orientation on each side of a boundary are produced (Fig. 2g);
this one-to-one relationship does not exist for EBSD patterns, which instead are superposed and
appear indistinct (Fig. 2f). It is this last characteristic that provides the crucial opportunity to
determine the complete orientation of intragranular boundaries.

# 149 **2.2** Boundary rotation angle(s) and axis

Figure 3a is a schematic representation of partial ECP configurations across an intragranular 150 boundary (trace t and trace normal  $h_t$ ) and illustrates how the boundary tilt angle ( $\omega$ ) and axis (r) 151 can be determined. The crucial observation is to recognise the EC (diffraction) band (or crystal 152 lattice plane) that is not displaced across the boundary trace. This band 'dips' towards the pattern 153 centre by 90- $\beta^{\circ}$ . The value of  $\beta$  is determined via the internal scaling present in ECP formation, 154 which is constrained by the specific SEM operating conditions (i.e. accelerating voltage, working 155 distance, etc.) and indicated by RA (degrees/radians) and the pattern diameter (D mm), such that, 156  $\beta^{\circ} = Bmm \times 2RA^{\circ}/Dmm$ (6a) 157

where Bmm is the orthogonal linear distance measured in the ECP between the non-displaced
channelling band and the pattern centre (Fig. 3a). Similar calculations can be defined for all other
EC bands and angular distances.

For the EC band to be non-displaced, the displacement translation vector  $(d_v)$  must be contained within the band (Fig. 3b). Thus, the normal to the non-displaced EC band must be parallel to the boundary rotation axis. As well as the non-displaced EC band, any bands parallel to the boundary trace are also recognised. The ECP configuration prior to boundary formation can be restored by translating one partial pattern relative to the other, parallel to the displacement translation vector of the non-displaced EC band until all bands become aligned (Fig. 3b). The boundary rotation angle  $(\omega_r)$  is then obtained from the displacement distance (d) in a similar manner to Eqn 6a,

168  $\omega_r^\circ = dmm \ x \ 2RA^\circ/Dmm$ 

(6b)



**Figure 3**. Determination of boundary rotation axis and angles. (a) Schematic partial ECPs across an intragranular boundary (t, trace; h<sub>t</sub>, trace normal). The normal to the non-displaced EC band, which 'dips' 90- $\beta^{\circ}$  where  $\beta$  = BxD/2RA (B, linear distance from non-displaced band centre to pattern centre; D, twice the rocking angle), defines the direction of the boundary rotation axis (r). Note also a boundary trace parallel EC band. (b) Restoration of preboundary formation ECP configuration via the displacement vector (d<sub>v</sub>) parallel to the non-displaced EC band ( $\alpha$ , acute angle between non-displaced band and boundary trace). Total rotation angle ( $\omega_r$ ) is parallel to the displacement distance (d), with tilt ( $\omega_{tilt}$ ) and twist ( $\omega_{twist}$ ) components normal and parallel to the boundary trace (determined via internal scaling relative to RA and D); the tilt component can also be represented in terms of distance (Å) via internal scaling provided by EC bandwidth ( $\delta Å_{hkl}$ ). (c) Upper hemisphere stereographic (Wulff/equal angle) and/or crystallographic projection representation: 1, boundary trace (t); 2, non-displaced EC band; 2<sub>⊥</sub>, vertical plane normal to non-displaced EC band; r, boundary rotation axis.

169	In the example shown in Fig. 3, the displacement vector is oblique to the boundary trace,	which
170	indicates a general boundary. From Fig. 1, pure tilt displacements develop normal to the b	ooundary
171	trace via rotation parallel to the trace, whereas pure twist displacements develop parallel t	to the trace
172	via rotation normal to it. The tilt ( $\omega_{tilt}$ ) and twist ( $\omega_{twist}$ ) angular components of a general b	ooundary
173	can therefore be determined by resolving $\omega_r$ into its boundary trace normal and parallel	
174	displacements respectively,	
175	$\omega_{\text{tilt}}^{\circ} = \text{dsin}\alpha \ 2\text{RA}^{\circ}/\text{Dmm}$	(6c)
176	$\omega_{\text{twist}}^{\circ} = \text{dcos}\alpha 2\text{RA}^{\circ}/\text{Dmm}$	(6d)
177	where $\alpha$ is the (acute) angle between the non-displaced EC band and the boundary trace (	t).

The displacement normal to the boundary trace can also be expressed as a length (e.g. in angstroms,
Å) via the scaling provided by the width of the diffraction bands, which depends on the SEM

accelerating voltage used, and is readily available in EC/EBSD processing software. Figure 3c summarises the boundary trace (1), non-displaced EC band (2), including its vertical normal plane ( $2_{\perp}$ ), and rotation axis (r) orientations in terms of (Wulff/equal angle) stereographic and/or crystallographic projections.

#### **184 2.3 Boundary plane orientation**

As stated previously, the main problem with defining the complete five-parameter orientation of intragranular boundaries concerns the physical orientation of the boundary plane normal direction. This is not a problem unique to intragranular boundaries. For example, the inclination ( $\psi$ ) of boundaries formed during grain boundary sliding (gbs) can be determined from a relatively simple relationship between the offsets of three mutually perpendicular directions (u, v, w) and the angle ( $\theta$ ) between the trace of the boundary and a reference direction, such that (Langdon, 2006).

191 
$$\tan \psi = (u - \operatorname{wtan} \theta)/v \tag{7}$$

However, it is impractical to use this equation (e.g. in the determination of the contribution of gbs to 192 the total strain) because of the difficulties of measuring the angles  $\theta$  and  $\psi$  at every boundary 193 (Langdon, 2006; Mohamed, 2011). Similarly, the movement characteristics of geological faults 194 relies implicitly on knowledge of the dip of the fault plane (e.g. Redmond, 1972; Yamada and 195 196 Sakaguchi, 1995 Xu et al., 2007 and 2009; Lisle and Walker, 2013); if the dip is not known, it is difficult to determine from other parameters except for specific combinations (e.g. Nieto-Fuentes et 197 al., 2014). Fortunately, the geometry of partial ECPs formed across a common intragranular 198 boundary (e.g. Fig. 3) offers a relatively simple practical solution to the problem. 199

The solution assumes implicitly that intragranular boundary orientations are constrained not only crystallographically to low-index lattice planes (i.e. Read and Shockley, 1950) but also by the origin of ECPs. Following Friedel's rule (Friedel, 1913), which introduces an effective centre of symmetry, each EC band originates from the centre of a spherical projection directly below the pattern centre defined by the rocking position (Fig. 4a). Thus, only bands 'steeper' than (90 - RA)°



Figure 4. EC spherical projection. (a) Schematic representation of trigonal α-quartz three-dimensional crystallography and ECP formation on the projection surface. (b) Quartz EC crystallographic unit triangle constructed via photo-montage over a spherical surface (Lloyd and Ferguson, 1986). (c) Multiple possible boundary plane parallel low-index lattice planes, all of which intersect at ±a<sub>2</sub> for a boundary plane trace normal to (2-1-10) and apparently parallel to and coincident with (01-11). (d) Schematic representation of parallelism but not coincidence between boundary trace (t) and EC bands (in)visible in an ECP.

- are visible in any ECP. As all bands are inclined towards the centre of the pattern, the angle of
- projection is given by Eqn 6a. The construction shown in Fig. 4a led directly to the early use of
- spherical ECP maps for crystallographic indexing purposes (e.g. Fig. 4b; Lloyd, 1987) and
- ultimately to (interactive) spherical Kikuchi maps (e.g. Day 2008 and 2009; Zhu et al., 2019;
- Hielscher et al. 2019), which we make use of later in this contribution.

The crux of our solution is that for intragranular boundaries constrained crystallographically to low-210 index lattice planes, the trace of the boundary plane is parallel to an EC band ideally present in the 211 (partial) ECPs (e.g. Fig. 3a). However, as shown in Fig. 4c, there can be multiple possible solutions. 212 213 The problem therefore is to determine which specific lattice plane is parallel to the boundary plane. In fact, the problem is even more complex as the boundary plane, although constrained to be 214 parallel to a low-index lattice plane, does not have to be coincident with that plane (e.g. Fig. 3a). 215 This is because all EC bands originate from the centre of a spherical projection (Fig. 4a), such that 216 the distance between the band and the centre of the ECP increases with inclination (Fig. 4c). In 217 contrast, intragranular boundaries are physical features that do not originate from the centre of the 218 spherical projection. Thus, whilst EC bands dipping  $\langle (90 - RA)^{\circ} do not appear in the ECP, a$ 219 mutually parallel intragranular boundary plane can still be observed in the image. Concomitantly, it 220 follows that straight boundary traces, including those through the centre of ECPs, are not 221 constrained to be vertical. The solution to the problem of determining the physical orientation of 222 intragranular boundaries is provided by stereographic projection analysis based on the geometry of 223 partial ECPs. Figure 3c shows the basic relationships between the boundary trace and its vertical 224 normal plane, the non-displaced EC band and its vertical normal plane, and the boundary rotation 225 axis plotted in (Wulff/equal angle) stereographic and/or crystallographic projections. We now 226 develop this construction in Fig. 5. 227

We first plot (Fig. 5a) the boundary trace (1), the non-displaced EC band (2) and their respective 228 vertical normal planes  $(1_{\perp}, 2_{\perp})$ . The intersection (3) of the non-displaced EC band (2) and the 229 vertical section plane normal  $(1_{\perp})$  defines the pitch of the former on the latter. Next (Fig. 5b), we 230 construct small circles (4) with radii equal to the pitch (3) centred on the strikes (t) of the boundary 231 trace (1); these small circles define the loci of pitches of the boundary plane relative to the trace (t). 232 233 The intersections (5) of the small circles (4) with the vertical section plane normal to the nondisplaced EC band  $(2_{\perp})$  define the pitches of the boundary plane on that plane. The strikes (t) of the 234 235 boundary trace (1) and the small circle intersections (5) must lie in the boundary plane; the great



Figure 5. Determination of boundary plane orientation: red and blue colours indicate boundary plane and non-displaced EC band related elements respectively (see Fig. 3). (a) Plot of boundary trace (1), non-displaced EC band (2) and their respective vertical normal planes (1⊥, 2⊥); intersection (3) of 2 and vertical 1⊥ defines the pitch of the former on the latter. (b) Small circles (4) with radii equal to the pitch (3) constructed about strikes (t) of the

boundary trace (1); the small circles define the loci of the pitches of the boundary plane and their intersections (5) with the vertical section plane normal to the non-displaced EC band  $(2_{\perp})$  define the pitches of the boundary plane on that plane. (c) Strikes (t) of boundary trace (1) and small circle intersections (5) must lie in the boundary plane, such

that great circles (6a, b) through these points define its potential orientations; the most likely alternative is determined by coincidence between the planes and any EC band in the partial ECPs (e.g. Fig. 3a). (d) Determination of non-displaced EC band normal (r, boundary rotation axis) and boundary plane normal ( $h_{\perp}$ ) directions.

circles (6a, b) through these points therefore define its two alternative orientations (Fig. 5c). To

237 choose the most likely alternative, we check for coincidence between the predicted boundary planes

and any low-index boundary-parallel EC band (e.g. Fig. 3a); we show below that this identification

is facilitated by spherical Kikuchi maps (SKM). Finally (Fig. 5d), the non-displaced EC band

normal (r), which defines the boundary plane rotation axis, and boundary plane normal  $(h\perp)$ 

directions are plotted. The crystal indices of these positions,  $[uvw]_r$  and  $[uvw]_{h\perp}$ , are the values that should be used in Eqn. 4 to calculate the TTC component of the boundary.

In the next section, we provide actual examples of the application of the method to intragranular boundaries in olivine. The examples were imaged at the University of Leeds using a Tescan Vega3 SEM operated in EC mode using a 25kV accelerating voltage, a specimen working distance of 9mm and rocking angles of  $\pm 10-11^{\circ}$ ; beam currents were varied to obtain the best OC and ECP images.

#### 247

# **3. EXAMPLE RESULTS**

Figure 6a is an EC orientation contrast image of a single forsteritic olivine (orthorhombic, space 248 group Pbnm) grain from a sample of volcanic ejecta (Kahl et al., in prep.); note the intragranular 249 boundary comprising long, straight and short, kinked segments. Also shown are ECPs (1 and 2) 250 251 from each side of the boundary (Fig. 6b, c). Whilst the patterns are slightly different due to the misorientation introduced by the boundary (e.g. compare their respective Euler angle triplets), their 252 253 individual configurations are consistent along each side of the boundary irrespective of the (boundary) segment orientation. Figures 6d and e show partial ECPs from each side (1 and 2) of the 254 two boundary segments respectively, obtained by rocking the incident electron beam on the 255 boundary as indicated. The boundary traces are clearly observed in both images by displacement of 256 EC bands. The displacement sense is either sinistral or dextral depending on the orientation (i.e. 257 angle of intersection) of the EC band relative to the boundary trace. However, one EC band towards 258 the upper left, which is the same for both boundary segments, is continuous across the boundary. 259 Spherical geometry dictates that the inclination ('dip') of any EC band is the (angular) distance 260 between the normal to the band, measured from the band centre, and the centre of the pattern 261

subtracted from 90° (Eqn 6). All bands 'dip' towards the pattern centre (i.e. the centre of the



Figure 6. Intragranular boundary orientation method example. (a) Olivine grain with intragranular boundary comprising long, straight and short, kinked segments (electron channelling orientation contrast). (b) and (c)
Complete electron channelling patterns from points 1 and 2 either side of the boundary (note Euler angle triplets); the patterns and hence orientations do not change along the boundary segments. (d) and (e) Partial electron channelling patterns from the long, straight (d) and short, kinked (e) boundary segments (broken white lines); note displacement of all channelling bands except for one, which is the same for both segments.

spherical projection, which is therefore upper hemisphere). Scaling is provided by the rocking angle involved in pattern formation, which is  $RA = \pm 10.7^{\circ}$  for this example.

# **3.1 Boundary rotation angles and axes**

- 266 The rotation angles and axes for the two olivine intragranular boundary segments (Fig. 6) are
- determined following the procedure outlined above (Fig. 3), illustrated in Figs 7a, b and 8a, b.
- 268 Based on the orthogonal distances from the centres of the non-displaced band to the rocking
- positions (i.e. pattern centres), it is inclined at 81.8° and 82.2° towards the pattern centres
- 270 respectively for the two segments. As the normal to the non-displaced EC band defines the
- orientation of the boundary plane rotation axis, this must be the same crystallographically for both

boundary segments irrespective of the different orientations of their boundary traces and hence
planes. The rotation axes for the two segments therefore plunge 8.2° and 7.8° respectively normal to
the non-displaced band.

The rotation angles associated with formation of the two intragranular boundary segments can be determined by translating one of the partial patterns parallel to the non-displaced EC band until all of the other EC bands become continuous (Figs 7b and 8b). In other words, the effect of the intragranular boundary has been 'removed'. In the examples shown, the translation directions are not orthogonal to the trace of either boundary segment, indicating that both segments are general boundaries and that the overall rotation angle ( $\omega$ ) comprises both tilt ( $\omega_{tilt}$ ) and twist ( $\omega_{twist}$ ) components.

The overall rotation angle is defined by the (angular) separation of the partial patterns along the displacement vectors, according to the internal scaling present in the ECPs (i.e.  $RA = \pm 10.7^{\circ}$ ) and using Eqn 6. The tilt and twist components are the displacements respectively normal and parallel to the boundary trace such that,

286 
$$\omega = (\omega_{tilt}^2 + \omega_{twist}^2)^{1/2}$$
 (8)

In terms of the long, straight segment,  $\omega = 3.3^{\circ}$ , with  $\omega_{tilt} = 3.1^{\circ}$  and  $\omega_{twist} = 1.0^{\circ}$  (Fig. 7b); whilst for the short, kinked segment,  $\omega = 2.6^{\circ}$ , with  $\omega_{tilt} = 2.2^{\circ}$  and  $\omega_{twist} = 1.4^{\circ}$  (Fig. 8b). As recognised previously, both segments are general boundaries. However, by comparing the relative values of the tilt and twist components for each segment, the long, straight segment involves significantly more tilt than twist in its configuration. To quantify the boundary segments more precisely (e.g. by the TTC index) requires accurate definition of their boundary plane orientations.

# **3.2 Boundary plane orientations**

To determine the orientation of the two boundary plane segments (Fig. 6a), we follow the workflow outlined in Fig. 5. First, we plot (Figs 7c and 8c) Wulff/equal angle upper hemisphere stereographic



Figure 7. Intragranular boundary orientation determination: olivine long, straight segment (see Fig. 6). Nomenclature and colours as in Figs 3 and 5. (a) Partial ECPs from across the boundary segment (t, trend and h<sub>t</sub>, trend normal); all EC bands are displaced across the boundary except one (indicated, with 'strike/dip/sense'); normal to this band defines boundary rotation axis (r, 09/129 'plunge/azimuth'). (b) Partial ECPs 'restored' to preboundary formation positions by translation parallel to non-displaced EC band such that all bands are continuous.

The translation distance is the boundary rotation angle ( $\omega$ ), which can be resolved into its boundary normal tilt ( $\omega_{tilt}$ ) and boundary parallel twist ( $\omega_{twist}$ ) components for general boundaries. (c) Stereographic projection (*upper* hemisphere, Wulff/equal angle) of progressive steps involved in boundary orientation determination (see Fig. 5 and associated text for definition of numbers/symbols). Two symmetrical solutions are indicated, both 'striking' 115° and 'dipping' either 81°N or 81°S. (d) Spherical Kikuchi map (SKM) representation illustrating crystallographic indexing procedure and final selection of boundary plane orientation (115/81N) based on coincidence with (2-11) EC band/lattice plane. Note also crystallographic orientations of r and h (i.e. true boundary plane normal), allowing determination of TTC = 0.17, close to a pure tilt intragranular boundary.

projections of the orientations of the boundary trace (1), its (vertical) normal plane  $(1_{\perp})$ , the non-297 displaced EC band (2) and its (vertical) normal plane  $(2_{\perp})$ . The values are determined using Eqn 6 298  $(0 - 90^\circ)$  and the clockwise azimuthal direction around the pattern circumference  $(0 - 180^\circ)$ . Next, 299 we determine the intersection (3) of the non-displaced EC band on the (vertical) normal to the 300 boundary segment trend; the intersection defines the loci of the pitches of the boundary plane as 301 small circles (4) about the strike (t) of the boundary trace (1). The intersections (5) of the small 302 circles (4) with the vertical section plane normal to the non-displaced EC band  $(2_{\perp})$  define the 303 pitches of the boundary plane on that plane. The strikes (t) of the boundary trace (1) and the 304 intersections (5) must lie in the boundary plane; great circles through these points therefore define 305 its potential orientations (6). 306

Typically, two alternative and symmetrical grain boundary plane solutions are obtained (6a and 6b). 307 To choose the most likely alternative, we check for coincidence between the predicted boundary 308 plane orientations and an EC band by overlaying the partial ECPs onto the stereographic projection 309 (Figs 7c and 8c). This is because the dislocation model for intragranular boundary formation 310 predicts that they coincide with low index crystal planes. For the long, straight segment, alternative 311 6a oriented 115°/81°N is coincident with an EC band that is parallel to the boundary trace (see Figs 312 6b-d and 7a, b), which is therefore selected. For the short, kinked segment, alternative 6a oriented 313 158°/83°W is also coincident with an EC band that is parallel to the boundary trace (see Figs 6b-d 314 and 8a, b), which is therefore selected. Having selected the appropriate boundary plane orientations, 315 the boundary plane normal directions  $(h_{\perp})$  can be determined for both segments. 316

Figures 7c and 8c define the boundary segment orientations in terms of spherical angles. As such, they are measured and represented in terms of the sample coordinate system. In this example, the olivine grain is from a sample of volcanic ejecta and consequently has no specific specimen orientation. It would be useful therefore to express the boundary segment plane orientations in terms of (olivine) crystallography. This is possible by use of spherical Kikuchi maps (SKM; see



Figure 8. Intragranular boundary orientation determination: olivine short, kinked segment (see Fig. 6).
Nomenclature and colours as in Figs 3 and 5. (a) Partial ECPs across each boundary segment (t, trend and h<sub>t</sub>, trend normal); all EC bands are displaced across the boundary except for one (indicated, with 'strike/dip/sense'); normal to this band defines boundary rotation axis (r, 09/129 'plunge/azimuth'). (b) Partial ECPs 'restored' to their preboundary formation positions by translation parallel to the non-displaced EC band such that all bands are continuous. The translation distance is the boundary rotation angle (ω), which can be resolved into its boundary normal tilt (ω<sub>tilt</sub>) and boundary parallel twist (ω<sub>twist</sub>) components for general boundaries. (c) Stereographic projection (*upper* hemisphere, Wulff/equal angle) of progressive steps involved in boundary orientation determination (see Fig. 5 and associated text for definition of numbers/symbols). Two solutions are indicated, both 'striking' 158° and 'dipping' either 83°W or 81°E. (d) Spherical Kikuchi map (SKM) representation illustrating crystallographic indexing procedure and final selection of boundary plane orientation (158/83W) based on coincidence with (222) EC band/lattice plane. Note also crystallographic orientations of r and h (i.e. true boundary plane normal), allowing determination of TTC = 0.45, indicating a general intragranular boundary.

323 Day 2008 and 2009) and in particular the interactive versions for olivine available in the HKL

324 Channel5 and/or AZtecCrystal software packages (Fig 7d and 8d). In effect, the stereographic

325 projection construction is simply overlain onto an SKM defined by the orientation of the ECPs,

either via visual comparison or, more accurately, via their Euler angle triplets (e.g. Figs 6a, b);

however, care is required beyond 45° from the centre of projection due to increasing spherical

distortion. It is also possible to incorporate the 'restored' partial ECPs into the construction (Figs 7d

and 8d). The interactive nature of the SKM then makes it a simple matter to index

crystallographically the relevant planes and direction.

Based on the olivine SKM, the non-displaced EC band common to both boundary segments is the

332  $(1\overline{74})$  lattice plane (Figs 7d and 8d). The long, straight boundary segment is parallel to the  $(2\overline{11})$ 

lattice plane, whilst the short, kinked boundary segment is parallel to the (222) lattice plane.

However, a problem arises in the determination of the normal directions to these lattice planes due

to non-cubic (i.e. orthorhombic) symmetry relationships. We resort therefore to relationships

involving the reciprocal (orthorhombic) crystal lattice and metric tensor (e.g. Bond, 1976; Boisen
and Gibbs, 1990; De Graf and McHenry, 2012),

338 
$$M^* = \begin{bmatrix} a *^2 & 0 & 0 \\ 0 & b *^2 & 0 \\ 0 & 0 & c *^2 \end{bmatrix}$$
(9)

where  $a^* = bc/V$ ,  $b^* = ac/V$  and  $c^* = ab/V$ ;  $a = 4.76\text{\AA}$ ,  $b = 10.21\text{\AA}$  and  $c = 5.98\text{\AA}$  are the orthogonal olivine lattice parameters; and V = abc is the volume of the olivine unit cell (290.38Å<sup>3</sup>). The direction [uvw] normal to the plane (hkl) is then given by,

342 
$$\begin{bmatrix} u \\ v \\ w \end{bmatrix} = M^* \begin{pmatrix} h \\ k \\ l \end{pmatrix}$$
 (10)

Thus, the rotation axis normal to the  $(1\overline{7}\overline{4})$  non-displaced lattice plane is parallel to  $[2\overline{3}\overline{5}]$ , whilst the normal directions to the  $(2\overline{1}1)$  and (222) boundary parallel lattice planes are parallel to  $[7\overline{1}2]$ and [513] respectively. Combining the common rotation axis with the boundary plane normal

<b>Boundary Parameter</b>	long, straight segment	short, kinked segment
non-displaced EC band	(174)	$(17\overline{4})$
rotation axis plunge/trend	08°/129°	08°/129°
rotation axis (r)	[235]	[235]
boundary plane	(211)	(222)
boundary trend (t)	115°/295°	158°/328°
ooundary plane normal plunge/trend	09°/025° or 205°	10/068° or 248°
boundary plane normal (h)	[712]	[513]
rotation angle ( $\omega$ )	3.3°	2.6°
<i>tilt component</i> ( $\omega_{tilt}$ )	3.1 °	2.2 °
twist component ( $\omega_{twist}$ )	1.0 °	1.4 °
tilt-twist component (TTC)	0.17	0.45
'excess volume' nm	25.16	8.87

numbers in bold represent five-parameter definition of intragranular boundary orientations

directions yield TTC values of 0.17 for the long, straight boundary segment and 0.45 for the short,
kinked boundary segment. The former is defined as close to a pure tilt boundary whilst the latter is a
general boundary.

# 349 3.3 Summary

Table 1 summarises the results of the olivine intragranular boundary analyses (Figs 6 - 8) in terms 350 351 of both the sample (i.e. stereographic projection plane and normal orientations) and crystallographic (i.e. crystal planes and normals) coordinate systems. The former is not representative for this 352 example as the sample does not possess a rigorous spatial orientation as it is a sample of volcanic 353 354 ejecta; however, this representation may be useful in samples that are kinematically constrained. The latter represents the complete five-parameter determination of boundary plane orientations and 355 is constrained by the orthorhombic crystal structure of olivine. For example, the crystallographic 356 orientations can be represented in terms of conventional upper and/or lower hemisphere 357 orthorhombic projections (Fig. 9a, b). 358

359

#### 4. DISCUSSION

360 In this section we discuss the implications of the model derived to determine the complete

361 orientation of intragranular boundaries. In combination with the results from the example (Figs 6 -



Figure 9. (a) Upper and (b) lower hemisphere SKM orthorhombic projections of the olivine intragranular boundary parameters determined by the method developed in this contribution. Also indicated are predicted crystal slip plane and directions for boundary formation as well as the misorientation axis determined by conventional EBSD analysis. (c) Schematic representation of olivine tilt and general boundary configuration and associated crystal planes, directions and slip systems.

8), we consider: (1) intragranular boundary formation and crystal slip system determination; (2)
conventional (i.e. EBSD-based) misorientation analysis; (3) adapting the EC-based method for
EBSD analysis; and (4) the applicability of the method for intergranular boundaries. To assist in this
discussion, we have summarised Table 1 in Table 2 to compare the olivine five-parameter boundary
definitions (Figs 6 - 8) with slip system determination and EBSD misorientation analysis.

# 367 4.1 Intragranular boundary formation – slip systems

368 The long, straight and short, kinked segments shown in Fig. 6a are clearly part of the same olivine

slip systems, with EBSD misorientation analysis											
	Five parameter definition Slip system EBSD misorientation										
Boundary	h(hkl)	h[uvw]	ω	r	TTC	SP	SD	ω	r		
Long, straight	(211)	[712]	3.3°	[235]	0.17	(122)	[712]				
Short, kinked	(222)	[513]	2.6°	[235]	0.45	$(12\overline{2})$ $(12\overline{2})$ (222)	$[\bar{5}1\bar{3}]$ [4 $\bar{3}\bar{1}]$ [4 $\bar{3}\bar{1}]$	6.13°	[212],		

Table 2 Comparison of aliving five-parameter boundary definitions (Figs 6 - 8) including

intragranular boundary. They share the same boundary rotation axis of  $\left| 2\overline{35} \right|$  in spite of their 369 different boundary plane orientations, which are parallel to  $(2\overline{1}1)$  and (222) respectively (Fig. 9c). 370 However, the former is close to a pure tilt boundary (TTC = 0.17), whilst the latter is a general 371 boundary (TTC = 0.45). The segments also differ in terms of their rotation angles, which is larger 372 for the long, straight segment (i.e. 3.3° compared with 2.6°). This suggests that the long, straight 373 segment represents the principal boundary and that the short, kinked segment is a minor component 374 that may well have been removed with further boundary evolution. 375

In terms of intragranular boundary formation, we have assumed the standard Reed and Shockley 376 (1950) dislocation model. Pure tilt boundaries, such as  $(2\overline{1}1)$ , form parallel to the boundary rotation 377 axis and normal to the slip plane and slip direction. The slip direction is therefore normal to the 378 boundary plane (Fig. 9c). From Eqn. 10, the boundary plane normal is determined to be  $[7\overline{12}]$ , such 379 that the slip system responsible for boundary formation is  $(122)[\overline{7}1\overline{2}]$ . In contrast, the normal to the 380 (222) short, kinked boundary plane is parallel to  $[\overline{5}\overline{1}\overline{3}]$  suggesting slip on  $(12\overline{2})[\overline{5}\overline{1}\overline{3}]$  was involved 381 in its formation. However, this is a general boundary, which require at least two slip systems to 382 form; thus,  $(12\overline{2})[\overline{5}\overline{1}\overline{3}]$  can be regarded as the tilt component. For pure twist boundaries, the slip 383 384 plane is parallel to the boundary plane, which is (222). The slip plane also contains the slip direction, although its precise orientation is not directly defined. We suggest that the twist slip 385 386 direction is parallel to the intersection of the tilt boundary and the slip plane for the pure twist boundary, which is  $[4\overline{3}\overline{1}]$  (Fig. 9a, b). Thus, the twist component of the general boundary was 387 388 accommodated by slip on  $(222)[4\overline{31}]$  (see Fig. 9c). Incidentally, the normal to the  $(12\overline{2})$  slip plane

is  $[42\overline{5}]$ , which is parallel to the intersection of the two boundary segment planes (Fig. 9).

390 Consequently, the four slip systems either share or intersect with the same slip plane,

391 namely  $(12\overline{2})$ .

# **4.2** Comparison with conventional (EBSD) misorientation analysis

The olivine sample (Figs 6 - 8) has also been analysed via EBSD, which provides Euler angle 393 triplets for each side of the intragranular boundaries (Fig. 6b, c). EBSD software typically includes 394 misorientation analysis for specified pairs of Euler angle triplets. However, as the Euler angle 395 triplets are effectively constant for each side of the boundary, only a single pair of triplets is 396 required to determine a single misorientation angle/axis pair applicable to both long, straight and 397 short, kinked segments. The angle/axis pair determined is  $6.13^{\circ}/[2\overline{1}2]$ , significantly different to the 398 rotation angles and directions determined by the five parameter method (Table 2). The EBSD 399 misorientation axis  $[2\overline{1}2]$  lies close to the non-displaced EC band  $(1\overline{7}\overline{4})$  and is oblique to both 400 boundary segment planes/traces (Fig. 9a, b). The segments are therefore general boundaries with the 401 402 same characteristics according to EBSD misorientation analysis.

It is clear from Table 2 and Fig. 9 that conventional (EBSD) misorientation angle/axis pair analysis 403 is insensitive to boundary type and physical orientation. This is not surprising as its definition 404 recognises only the axis about which one crystal lattice needs to be rotated to bring it into alignment 405 with another (Fig. 1a); the orientation of the boundary is not involved. Thus, both segments have 406 the same misorientation angle/axis pair. The only configurations for which the misorientation axis 407 could potentially define boundary plane orientation are for pure tilt and twist boundaries; the axis is 408 parallel to the boundary plane in the former and normal to the plane in the latter. However, there 409 remain a large number of possible inclinations within either pure tilt or pure twist boundary planes 410 for a given trend (Fig. 4c, d), which makes even these ideal situations undefinable via conventional 411 misorientation analysis. 412

In the initial SEM/EBSD implementation of misorientation analysis (e.g. Mainprice et al., 1993; 413 Lloyd et al., 1997), there was no requirement for two crystal lattices to be adjacent (i.e. share a 414 common boundary), although usually the calculated misorientation 'angle/axis' pair relates to 415 416 neighbouring or 'correlated' domains (Wheeler et al., 2001). Depending on crystal symmetry, there are several crystallographically related solutions for the misorientation angle/axis pair across an 417 418 intragranular boundary. Convention dictates that the solution with the smallest angle is chosen, often referred to as the 'disorientation' (MacKenzie and Thomson, 1957; Hanscomb, 1958; 419 MacKenzie, 1958; Warrington and Boon, 1975). However, an alternative approach recognises that 420 the nearest low-index boundary axis solution defines the misorientation between adjacent lattices 421 422 (e.g. Gourdet et al., 1998). The two approaches are not equivalent, which results in the distinction between coincidence site lattice (CSL) and low-order coincident axial direction (CAD) 423 visualisations of boundary geometry (e.g. Warrington and Boon, 1975). It has been suggested (e.g. 424 Cross and Randle, 2003) that an unambiguous analysis of intragranular boundary orientation 425 involves consideration of the CAD solution rather than just the disorientation. The five parameter 426 427 method derived in this contribution is most compatible with the CAD solution.

In fact, misorientation analysis angle/axis and five-parameter rotation angle/axis pairs are entirely 428 429 different parameters. In the former they bring two crystal lattices into parallelism irrespective of whether or not the lattices share a common intragranular boundary (Fig. 1a), whereas in the latter 430 they relate to the rotation between two originally identical crystal lattices due to the formation of an 431 432 intragranular boundary (Fig. 1b-d). Thus, in misorientation analysis the misorientation axis is unconstrained by the boundary plane orientation and adopts an orientation that minimises the 433 rotation angle between two adjacent crystal lattices. In contrast, in the five-parameter method not 434 435 only is the boundary plane orientation of fundamental significance but also the dislocation model for intragranular boundary formation imposes constraints on boundary plane orientations. We 436 suggest therefore that conventional EBSD misorientation analysis may involve ambiguity in terms 437 of characterising intragranular boundary orientations. 438

#### 439 4.3 Adaption to EBSD

The method developed to determine the five parameters needed to constrain the orientation of

441 intragranular boundaries is based on observations made via SEM/EC (e.g. Figs 2d-g and 3).

- 442 However, few SEM are currently capable of EC analysis and crystallographic orientation
- 443 measurements are more typically made via EBSD (e.g. Fig. 2b, c). It would be useful therefore if
- the basic method could be adapted for EBSD-based analysis and data. A solution is provided via
- 445 SKMs (e.g. Fig. 9a, b) and is illustrated using the olivine example described previously (Figs 6 8).

The adaption is as follows. (1) EBSD derived Euler angle triplets obtained from each side of a 446 boundary are input into an SKM to produce maps centred on each orientation (Fig. 10a); the maps 447 are copied into a vector drawing package. (2) Circles equivalent to ECP rocking angles (e.g.  $\pm 10^{\circ}$ ) 448 are drawn about the centre of each map to simulate ECPs from each side of the boundary and the 449 trace of the boundary is drawn across the centre of each pattern; 'cropping' tools are used to extract 450 the 'ECPs' from the SKM (Fig. 10b). (3) The extracted patterns are again 'cropped' and separated 451 along the boundary traces to form simulated 'partial ECPs' (Fig. 10c). (4) The cropped patterns are 452 joined in appropriate pairings to simulate partial ECPs obtained by rocking about a central point on 453 their common boundary (Fig. 10d). (5) The non-displaced EC/lattice band is identified, from which 454 all of the parameters needed to define the complete five- parameter orientation of the boundary can 455 456 be measured and/or determined via stereographic projection as per the basic method (Fig. 10e).

The adapted solutions in Fig. 10e compare favourably with the original solutions in Figs 7 and 8. They can therefore be used to determine the crystallographic orientations of the intragranular boundary planes as shown in Fig. 9. Thus, the original EC-based method devised to determine the five-parameter evolution of Intragranular boundaries can be easily adapted to EBSD data, making



**Figure 10**. Adaption of the EC-based five-parameter boundary orientation method to EBSD data and analysis. Nomenclature and colours as in Figs 3 and 5. (a) *HKL Channel5*<sup>©</sup> SKMs defined by EBSD derived Euler angle triplets from each side of the olivine intragranular boundaries (Fig. 6). (b) Detail of SKM-simulated ECPs. (c) Simulated ECPs 'cropped' along boundary tracings. (d) 'Cropped' patterns joined appropriately along boundary traces to simulate partial ECPs formed by rocking about a point on the common boundary. (e) SKM representation of boundary orientation determination (1 – 6 refer to steps in the stereographic projection method, see Fig. 5): *left*, long, straight segment; *right*, short, kinked segment.

the approach more generally available.

## 463 **4.4 Adaption to intergranular boundaries**

It is generally accepted (e.g. Rohrer, 2011) that the properties and behaviours of intergranular 464 boundaries are controlled more by the physical geometry of the interfacial plane rather than 465 crystallographic misorientation. This is because as misorientation increases so does dislocation 466 467 density, which reduces space between neighbouring dislocations until their cores overlap such that the ordered nature of the boundary begins to break down. Unlike intragranular boundaries, which 468 develop from a common initial crystal lattice by the progressive accumulation of dislocations (e.g. 469 Reed and Shockley, 1950), there is no conventional *physical* reason why crystallographic 470 relationships should exist across intergranular boundaries. In fact, the existence of such 471 472 relationships may have significant implications for microstructural evolution. Thus, the method developed to determine the five-parameter orientation of intragranular boundaries should not 473 necessarily be expected to apply to intergranular boundaries unless extenuating circumstances 474 475 apply.

476 To investigate whether the method of determining the five-parameter orientation of intragranular 477 boundaries can be adapted to intergranular boundaries, we consider the boundary relationships between a single olivine grain and its seven neighbouring olivine grains (Fig. 11a) in a garnet 478 lherzolite from the 90 Ma Thaba Putsoa kimberlite, Lesotho (Nixon and Boyd, 1973; Mercier and 479 Carter, 1985; Allsopp et al., 1989). The microstructure of this rock is characterised by a matrix of 480 dynamically recrystallised olivine with straight grain boundaries and frequent 120° triple junctions, 481 with a well-developed if somewhat unusual crystallographic preferred orientation (see Wallis et al., 482 2019). EBSD data were acquired at the University of Leeds using Oxford Instruments Aztec 2.1 483 484 software on a FEI Quanta 650 field-emission gun SEM equipped with an Oxford Instruments Nordlys-S EBSD camera. Operating conditions were: 20kV accelerating voltage, 8mm working 485 distance and 70° specimen tilt angle. 486



Figure 11. Example of potential adaption of method to intergranular boundary orientation determination. (a) EBSD 'all Euler' image of an olivine grain (0) and its immediate neighbours (1-7) from a garnet lherzolite, Thaba Putsoa, Lesotho. (b) SKM 'pairs' from each side of a boundary between grain 0 and grains 1-7; also shown are the elements of the five-parameter boundary method (see key). See Table 3 for a summary of results.

Euler angle triplets (Table 3) were measured from adjacent sides of each intragranular boundary 488 (i.e. between grains 0 and 1, 0 and 2, etc.) and used to define SKM-pairs 'fitted' along the trace 489 orientation of their common boundary (Fig. 11b), in a similar manner to partial-ECPs (e.g. Figs 7, 8 490 491 and 10). Note that the triplets and hence SKMs for Grain 0 are effectively constant. The combined partial-SKMs are then interrogated to identify the relevant relationships (see above) required to 492 493 determine the boundary orientation (Fig. 11b). However, due to the increasing distortions that accrue with distance from the centre of the projection, it is advisable to consider only an 494 approximately 45° radius small circle region about the centre. 495

Results of the analysis of the olivine intergranular boundary analysis (Fig. 11b) are summarised in 496 Table 3. Of the seven intergranular boundaries considered, only one (boundary 0:5) provides no 497 overall solution using the five-parameter method as no lattice plane is continuous across the 498 boundary trace. Another (boundary 0:2) involves the juxtaposition of opposite hemispheres, again 499 500 preventing an overall solution due to the absence of a non-displaced plane. A third (boundary 0:6) also involves the juxtaposition of opposite hemispheres but now the  $(13\overline{1})$  and  $(1\overline{3}\overline{1})$  planes are 501 'continuous' across the boundary and hence provide a potential solution but one involving two 502 possible boundary plane orientations. The validity of this solution depends on whether the 503 equivalence of different planes of the same family is appropriate in terms of the five-parameter 504 505 method. The remaining four boundaries all provide tangible solutions based on the occurrence of non-displaced lattice planes across the boundary traces. Indeed, three of these boundaries (0:1, 0:3 506 and 0:7) involve the same non-displaced families of planes and hence rotation axes, specifically 507 {025} and <019> respectively. The fourth boundary (0:4) has  $(2\overline{2}2)$  as the non-displaced plane and 508  $[5\overline{1}3]$  as its rotation axis. 509

510 Five-parameter analysis indicates that two of the intergranular boundaries (0:3 and 0:7) have the

same orientation (Table 3), parallel to (066) and with boundary plane normal parallel to [037].

512 Interestingly, these boundaries are on opposite sides of the central grain (Fig. 11a). The other two

Boundary &			Non-		Rotation angles			Boundary		EDCD
Euler angle triplets		displaced EC band	(r)	ωr	Wtilt	Øtwist	plane (hkl)/[uvw]	TTC	angle/axis	
137.6 102.6 39.5	0:1	147.8 110.8 72.6	(023)	[014]	30.2	27.4	12.7	(192)/[253]	0.06	32.7/[102]
137.0 103.1 39.5	0:2	21.5 135.4 144.0		different hemispheres – no solution				84.2/[031]		
137.7 102.8 38.5	0:3	141.9 100.3 50.7	(025)	[019]	12.5	2.5	12.3	(066)/[037]	0.90	12.3/[014]
137.5 102.6 39.8	0:4	133.9 106.3 34.9	(222)	[513]	2.5	1.1	2.2	(131)[433]	0.35	6.5/[123]
137.5 102.7 39.3	0:5	40.8 120.8 129.3		different hemispheres -no solutions					94.2/[221]	
			(121)	[42]]	non-	displace	ed EC	(122)[425]	0.49	
137.4	0.6	5.3	(131)	[433]	bana but na	l same f	amily	(344)[514]]	0.41	40/[100]
39.4	0.0	114.9	$(1\overline{2}\overline{1})$	[422]	but not also		rent	(122)[425]	0.27	49/[100]
			(131)	[433]	he	misphe	res	$(3\bar{4}\bar{4})[5\bar{1}\bar{4}]$	0.84	
137.2	0.7	133.7 97.8	(025)	[019]	26.5	5 5	25.8	(066)/[037]	0.90	190/[107]
39.4	<b>U.</b> /	56.8	$(02\overline{5})$	[019]	20.5	5.5	25.0	(000)/[057]	0.68	17.0/[102]

Table 3. Adaption to intergranular boundaries: summary of results (see Figs 11 and 12).

intergranular boundaries for which solutions have been found (i.e. 0:1 and 0: 4) are parallel to (192) and  $(1\overline{3}\overline{1})$  planes respectively, with normal directions parallel to [253] and  $[4\overline{3}\overline{3}]$ , again respectively.

516 The rotation angles for the intergranular boundaries that provide five-parameter orientation solutions (Table 3) are derived by translating one of the partial SKMs parallel to the non-displaced 517 EC/lattice plane until the SKM is reconstituted (Fig. 12). This is the same procedure used for the 518 partial ECPs (e.g. Figs 7 and 8). The three boundaries sharing a <019> rotation axis exhibit 519 increasing rotation angles from 12.5° (boundary 0:3), which is only just indicative of an 520 intergranular boundary, through 26.5° (boundary 0:7) to 30.2° (boundary 0:1). However, according 521 to their TTC values, boundaries 0:3 and 0:7, which also have the same orientation, are almost pure 522 twist boundaries, whilst 0:1 is even closer to a pure tilt boundary. The other boundary (0:4) has a 523 very small rotation angle and is in fact an intragranular boundary, which explains why it provides 524 the most rigorous five-parameter orientation solution. As none of the boundaries are either pure tilt 525

- 526 or pure twist in character, the contributions of the tilt and twist components can also be readily
- 527 determined by simple geometry (Fig. 12 and Table 3).
- As the Euler angle triplets are known for each side of the seven boundaries (Fig. 11a and Table 3), HKL Channle5 software can be used to determine the conventional misorientation angle/axis pair
- 530 for each boundary (Table 3). This approach is based only on the restoration of crystal lattices into



Figure 12. Determination of the total and tilt/twist component rotation angles for the four olivine inter/intraboundaries yielding tangible five-parameter orientation solutions via translation (broken blue arrow) parallel to the non-displaced EC band (solid blue line) of one partial-SKM (broken red outline) relative to its pair until the 'disrupted' crystal lattice configuration is restored.

parallelism and takes no account of boundary orientation or configuration; thus, all seven 531 boundaries provide a misorientation angle/axis pair. In contrast, the five-parameter method, can 532 only provide a rotation angle/axis pair if a solution to the boundary orientation exists; that is 533 534 boundaries 0:1, 0:3, 0:4 and 0:7 (see Fig. 12). A direct comparison between the two approaches is therefore only possible for these four boundaries. In general, comparisons are poor, which is not 535 surprising as the two measures are fundamentally different. The closest match is for boundary 0:3, 536 with the five-parameter method giving  $12.5^{\circ}/[019]$  against  $12.3^{\circ}/[014]$  for misorientation analysis. 537 We conclude, as before for intragranular boundaries, that conventional misorientation analysis is 538 not necessarily a good indicator of complete intergranular boundary relationships. 539

#### 540 **4.5 Related methods**

Having considered how the basic method can be adapted to both EBSD data and intergranular
boundaries, we can now turn to its relationship to existing methods for boundary orientation
determination. Although the method presented here is new *in its entirety*, there are two existing
methods that it does relate to: (1) 'plane-matching'; and (2) 'grain boundary character distribution'
(GBCD).

#### 546 (1) 'Plane-matching'

The 'plane-matching' method (Pumphrey, 1972) recognises that periodic lines observed in 547 transmission electron microscope images of high angle grain boundaries result from the mismatch 548 of either low or higher index atom planes across the grain boundary. The common factor in all 549 observations is that there is at least one set of planes of relatively high atomic density that is either 550 continuous or only slightly mismatched across the boundary. Watanabe (1983) and Watanabe et al. 551 552 (1989) reported ECP observations of grain boundaries that they argued supported the 'planematching' model. This led them to suggest that ECP provided a powerful tool for the precise 553 determination of crystallographic orientations; for example, to determine the relative orientation 554 555 relationship between adjacent grains that geometrically characterise a grain boundary. In terms of

the present model, 'plane-matching' is equivalent to the non-displaced diffraction band. However, the 'plane-matching' approach does not determine the orientation of the grain boundary; rather, it constrains the direction of growth of one grain at the expense of another grain (e.g. during recrystallisation). Nevertheless, as Watanabe (1983) and Watanabe et al. (1989) provided the ECPs for the boundaries they investigated, it seems apposite to consider examples in terms of the present model, particularly as the 'plane-matching' approach did not define the boundary orientations.



**Figure 13**. Comparisons between the current and 'plane-matching' methods: (a) – (c) low-angle (intragranular) boundary between two grains (A, B) in bcc Fe-6.5 mass% Si alloy (Watanabe et al., 1989, Fig. 4); (d) – (f) high-angle grain boundary between two grains (A, B) in bcc Fe-3% Si alloy (Watanabe, 1983, Fig. 1). Colour coding and nomenclature/symbols as for Figs 3, 5, 7, 8 and 10-12. (a) and (d) SEM images of boundary microstructures (cross indicates beam incident position). (b) and (e) Partial ECPs from the beam incidence positions. (e) Interpretation of the low-angle (intragranular) boundary via the current method as a vertical pure tilt boundary (TTC = 0) parallel to (110) with 4.9° misorientation. (f) Interpretation of the high-angle grain boundary via the current method as a subvertical (87.4° dip) general boundary (TTC = 0.20) parallel to ( $\overline{230}$ ) with 65.8° misorientation comprising 64.4° tilt and 6.6° twist components.

The first example is a low-angle (intragranular) boundary between two grains in bcc Fe-6.5 mass% 562 Si alloy (Watanabe et al., 1989, Fig. 4; see Fig. 13a). Partial ECPs from across the boundary (Fig. 563 13b) clearly reveal the matched plane as (1-10), which is also the non-displaced EC band. 564 565 Superposing the partial ECPs onto a SKM for bcc iron (Fig. 13c) indicates that the matched plane/non-displaced EC band passes through the projection centre; the band must therefore be 566 vertical and hence the rotation axis is horizontal, parallel to  $[\overline{1}10]/[1\overline{1}0]$ . Furthermore, not only 567 does the boundary trace also pass through the projection centre but it is almost coincident with the 568 vertical normal to the matched plane/ non-displaced EC band. Applying the method described in 569 this contribution constrains the boundary plane to be vertical (Fig. 13c); however, the very slight 570 off-set between boundary trace and non-displaced EC band means that the boundary plane normal 571 direction is parallel to  $[881]/[\overline{88}1]$ . Combining these directions with the rotation axes via the Excel 572 spreadsheet in the Appendix yields a TCC value of zero; thus this is a pure tilt boundary. Finally, 573 using the internal scaling present in the ECPs and SKM (Fig. 13c) allows the tilt angle to be 574 determined as 4.9°, confirming this as a low-angle (intragranular) boundary. 575

The second example is a high-angle grain boundary between two grains in bcc Fe-3% Si alloy 576 (Watanabe, 1983, Fig. 4; see Fig. 13d). Partial ECPs from across the boundary (Fig. 13e) reveal the 577 matched plane and hence non-displaced EC band to be (110). Superposing the partial ECPs onto a 578 SKM for bcc iron (Fig. 13f) indicates that the matched plane/non-displaced EC band does not quite 579 pass through the projection centre; the band therefore is not quite vertical and hence the rotation 580 axis is sub-horizontal parallel to [110]. Whilst the boundary trace almost passes through the 581 projection centre, it is clearly not parallel to the matched plane/non-displaced EC band. Applying 582 583 the method described in this contribution constrains the boundary plane to be parallel to  $(\overline{230})$  and to dip at 84.7°, with the boundary plane normal direction constrained to be parallel to  $[\overline{2}30]$  (Fig. 584 13f). Combining these directions with the rotation axes via the Excel spreadsheet in the Appendix 585 yields a TCC value of 0.20, indicating a general boundary albeit with a dominant tilt component. 586 Using the internal scaling present in the ECPs and SKM (Fig. 13f) allows not only the total 587

boundary misorientation to be determined at 65.8° but also the tilt and twist components at 64.4°
and 6.6° respectively. The fact that such a large grain boundary misorientation can still exhibit a
matched plane/non-displaced EC band suggests that the two adjacent grains share an affinity (see
Watanabe et al., 1989).

592 (2) *GBCD* 

In principal, the GBCD method (e.g. Watanabe, 1979, 1984, 1986 and 1988; Watanabe et al., 1981 593 and 1986) should have much in common with our method as it targets the five parameters needed to 594 define the complete orientation of a boundary. However, in practice there are significant differences 595 between the two approaches. GBCD concerns the description of the type and frequency of grain 596 boundaries in polycrystalline materials. In particular, it seeks to characterise the distribution of 597 special types of boundaries that might impact upon material properties and behaviours. The grain-598 boundary distribution is expressed in terms of the conventional five parameters from measurements 599 of grain orientations and the orientations of the lines formed where grain boundaries intersect the 600 601 plane of observation. The fundamental aim of GBCD is to provide a quantitative description of the amount and type of (special) boundaries that are present in a polycrystalline material (e.g. CSLs). 602

603 A significant advantage that the GBCD approach has over the method presented here is that it has been automated via EBSD and stereology (e.g. Saylor and Rohrer, 2002; Saylor et al., 2004; Rohrer 604 et al., 2004b; Marquardt et al., 2015; Marquardt & Faul, 2018) and hence can consider large 605 606 populations of boundaries that potentially provide statistically significant observations. However, the method does have disadvantages. In particular, it requires either 3D EBSD data derived via 607 serial sectioning or 2D EBSD data from samples lacking CPO to ensure the approximately random 608 609 distribution required for stereological determination of grain boundary normal directions (Randle and Davies 2001; Rohrer 2007 and 2011; Brandon 2010; Marguardt et al., 2015). In addition, 610 GBCD relies on the three disorientation parameters to define the minimum misorientation 611 angle/axis pair, which we have suggested is independent of boundary orientation except in special 612

circumstances. Further comparison between GBCD and the method described here is beyond thescope of this contribution.

# 615 **4.6 Accuracy, precision and error**

The method introduced here is overtly manual and practical, even when involving Euler angle 616 triplets determined via EBSD. It is therefore subject to the accuracy, precision and error limitations 617 618 inherent in such approaches, particularly in terms of visual inspection/interpretation and manual measurements. In general, errors depend on the resolution and quality of the diffraction patterns 619 (either EC or EBSD), the size of the spherical region, boundary inclination and the precision of the 620 stereographic projection, amongst other factors. As such, it is difficult if not impossible to derive 621 precise relative error(s) or error function(s). Nevertheless, the original potential error from which all 622 623 others ultimately derive is in the recognition of the non-displaced diffraction band and the associated linear and/or angular measurements (misorientation displacements and/or angles). If the 624 non-displaced band is clearly defined and particularly if it is located within~45° of the centre of 625 projection, then measurements should be both accurate and precise. However, if the non-displaced 626 band is poorly defined, which tends to be exacerbated if it occurs towards the periphery of the 627 projection, then accuracy and indeed precision are impacted. In general, the best possible accuracy 628 is probably to within one decimal place for both angular and linear measurements. In terms of 629 630 actual boundary orientations, this means that steeper boundaries are likely to be more accurately defined than shallower boundaries. 631

Whilst the accuracy and precision in locating the boundary trace are typically high, the recognition of the actual boundary plane depends on stereographic projection. The projection involves plotting the non-displaced diffraction band and its vertical normal plane, as well as the boundary trace and normal. Initial errors in defining the non-displaced diffraction band can therefore be propagated. In general, the orientation of the boundary trace can be defined accurately for straight boundaries; however, curved boundaries need to be considered as linear segments, which must introduce some error or approximation. In addition, magnification of boundary images and especially EBSDderived 'maps' can also introduce pixilation of boundary traces, leading to lack of precision and
error. The stereographic solution also provides two alternative boundary orientations, with the final
choice depending on the presence/absence of a coincident (low-index) lattice plane, which may not
be perfectly matched and hence could add additional error.

The method presented is based on the assumption that intragranular boundaries are constrained to 643 low-index crystal lattice planes (i.e. the conventional Read-Shockley model), which can introduce 644 limitations in solutions for boundary plane orientation and hence errors. Whilst the basic 645 646 assumption can be geometrically established based on dislocation systems that usually operate with short Burger vectors, more general intragranular low angle or composite boundaries (e.g. those with 647 orientation gradients that can be solved for combinations of different slip systems) can deviate from 648 649 the conventional approach. Whilst this situation must be recognised, we would emphasise that although the basic premise of the method is the Read-Shockley model, the solutions derived are not 650 necessarily low-index planes or directions (e.g. Tables 1-3). Indeed, the converse question can be 651 posed as to how valid is the practice of 'rounding' the indices to the nearest low angle solution. 652

# 653 **4.7 Future developments**

This contribution has focussed on deriving a practical method to determine the complete (five-654 parameter) orientation of intragranular boundaries. Whilst such a method has been demonstrated, it 655 certainly needs further testing via more examples, case studies and applications. These aspects 656 represent current work in progress. In addition, as presented the method is overtly practical and 657 hence applicable to relatively small datasets; it does not compare therefore with the opportunities 658 provided by EBSD misorientation analysis, although as demonstrated such analysis may contain 659 ambiguity in terms of boundary plane orientation. Nevertheless, the practical five-parameter 660 661 boundary orientation method would certainly benefit from both quantification and automation. The former is also currently being investigated via several approaches involving definition of the 662

orientation matrix (g) in terms of crystal indices, Euler angle triplets and quaternions, although the plunge and azimuth of the boundary plane normal still have to be determined using the method proposed here. The latter requires (image) recognition of boundary traces and comparison of (quantitative) Euler angle triplet based orientations for adjacent regions across boundaries to define the necessary parameters. Hopefully, should the basic practical method prove viable, full quantification and automation will follow.

669

# **5. CONCLUSIONS**

Boundaries are important features of polycrystalline materials and influence most properties, 670 characteristics and behaviours. Accurate, reproducible and ideally rapid and efficient definition of 671 boundary configuration is central therefore to furthering understanding polycrystalline materials. 672 The complete definition of boundary orientation involves the misorientation between the crystal 673 lattices of adjacent regions across the boundary and also the physical attitude of the boundary plane 674 (e.g. the direction of the plane normal). Whilst it has become relatively easy and efficient to 675 676 determine boundary misorientation (e.g. via EBSD in the SEM), definition of the boundary plane orientation has proved difficult. 677

678 This contribution has presented a practical method to determine the complete 'five-parameter' orientation of intragranular boundaries based on matching the spherical geometries of 'partial' SEM 679 electron channelling patterns (ECPs) across boundary traces. The method assumes that intragranular 680 boundaries are constrained to low-index crystal lattice planes (i.e. Read-Shockley model) and relies 681 on recognising electron channelling bands, equivalent to crystal lattice planes, which are not 682 displaced across or are parallel to the boundary trace. The former indicates the translation or 683 displacement direction of one side of the boundary relative to the other, such that its normal 684 direction defines the boundary rotation axis. Stereographic projection considerations of the 685 686 spherical geometry relationships between the non-displaced channelling band and the boundary trace, together with the boundary trace parallel channelling bands, allow for determination of the 687

orientation of the boundary plane. Finally, one partial pattern is translated relative to the other
parallel to the non-displaced channelling band until the undistorted ECP is restored, from which the
boundary rotation angle can be measured.

Although developed from EC theory for intragranular boundaries, it has been shown how the 691 method can be adapted relatively simply to both EBSD data (i.e. Euler angle triplets), involving the 692 use of interactive spherical Kikuchi maps, and intergranular boundaries. However, not all 693 intergranular boundaries may possess the necessary relationships between adjacent grains. In 694 addition, the new method not only readily distinguishes between tilt and twist boundaries but also 695 allows the relative contributions of these two end-member forms to the statistically more common 696 general boundary configurations to be accurately determined and represented, both in terms of 697 component rotation angles and also the tilt-twist component (TTC) index. 698

Whilst this contribution is principally concerned with the derivation of a basic method to determine 699 accurately the complete five-parameter orientation of intra- and some inter- granular boundaries, the 700 701 examples included suggest that conventional (i.e. EBSD) misorientation analysis may not necessarily provide an accurate and/or reliable representation of (intragranular) boundary 702 configurations. Misorientation analysis is based only on matching crystal lattices via rotations about 703 a specific axis (i.e. angle/axis pairs) and is independent of boundary configuration. We anticipate 704 therefore that our method could potentially offer a new field of analysis based on data that are 705 readily available and stimulate discussions on the nature and significance of (intragranular) 706 boundaries. 707

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917

#### **APPENDIX:**

5	1	n
3	l	J

# 918 Excel spreadsheet to assist crystallographic calculations involved in the five-parameter based 919 determination of the orientation of intragranular boundaries

920 We have developed a simple Excel spreadsheet to perform the crystallographic calculations

921 involved in the five-parameter based determination of the orientation of intragranular boundaries

922 (see below). This Appendix explains the calculations, which are all standard crystallography

relationships; all user input cells are indicated by bold red text. In general, the determination of the

924 five-parameters involves crystal directions, [uvw], rather than planes, (hkl). The description is based

925 on olivine; modifications for other phases and in particular crystal symmetries are indicated where

necessary. Note that the calculations involve rounding, whilst some indices can be user-simplified.

# 927 **0.** Miller/Miller-Bravais conversions (hexagonal and trigonal lattices)

As the crystal operations involved in boundary determination assume Miller indices and notation
(i.e. (hkl) or [uvw]), it is necessary to convert between these and Miller-Bravais indices applicable
to hexagonal and trigonal lattices (i.e. (HKIL) or [UVTW]).

#### 931 Miller-Bravais to Miller

932	Directions:	u = U - T	v = V - T	w = W	(Ala)
933	Planes:	h = H	$\mathbf{k} = \mathbf{K}$	l = L	(A1b)

934 Miller to Miller-Bravais

935	Directions:	U = (2u - v)/3	V = (2v - u)/3	T = (u + v)/3	W = w (A1c)
936	Planes:	H = h	K = k	$\mathbf{I} = \mathbf{h} + \mathbf{k}$	L = 1 (A1d)

#### 937 **1. Input phase lattice constants**

These are the physical dimensions (a, b, c) of the unit cell in a crystal lattice. The number of unique

terms depends on crystal symmetry but three input values are required.

# 940 **2. Input rotation axis**

941 The five-parameter boundary orientation determination method requires the crystallographic

orientation of the rotation axis and boundary plane normal directions. The former is input here.

However, the value may involve calculations in Sections 4 and/or 5.

#### 944 **3. Calculate normalised rotation axis**

The calculation of TTC involves the normalised rotation axis (e.g. Amouyal et al., 2005; Jhang et

946 al., 2018). Normalisation,  $[\bar{u}\bar{v}\bar{w}]$  is achieved via,

947 
$$\bar{u} = u/\Sigma$$
  $\bar{v} = v/\Sigma$   $\bar{w} = w/\Sigma$  (A2)

948 where  $\Sigma = (a x u + b x v + c x w)^{1/2}$ .

#### 949 **4.** Calculate normal [uvw] to plane (hkl) - orthorhombic lattices

In general, the normal direction to a crystal plane in non-cubic crystal lattices does not consist of the same indices due to the different lengths of the lattice constants, although there are symmetrydependent exceptions. For example, in the orthorhombic olivine lattice, the normal to the plane (111) is not [111] but [413]. The main text explains how the reciprocal lattice concept can be used to solve this problem (i.e. Eqns 9 and 10).

#### 955 **5.** Calculate plane (hkl) from two directions <u<sub>1</sub>v<sub>1</sub>w<sub>1</sub>>, <u<sub>2</sub>v<sub>2</sub>w<sub>2</sub>>,

When using Spherical Kikuchi Maps, lattice planes are not always displayed, particularly as their indices increase. Fortunately, any plane can be determined from two known directions (i.e. zone axes) through which it passes. First, the two zone axes are written twice, one beneath the other:

959  $u_4$   $v_1$   $w_1$   $u_1$   $v_1$   $w_4$ 960  $u_2$   $v_2$   $w_2$   $u_2$   $v_2$   $w_2$ 

961 Next, the first and last pairs are excluded. Finally, the lattice plane is calculated from,

962  $h = v_1 x w_2 - v_2 x w_1$   $k = w_1 x u_2 - u_1 x w_2$   $l = u_1 x v_2 - v_1 x u_2$  (A3)

# 963 **6.** Calculate tilt/twist component (TTC)

This is calculated simply from Eqn. 3 in the main text using the normalised rotation axis and boundary plane normal indices. The former is automatically copied into the calculation; the latter is input manually after appropriate use of sections 4 and/or 5.

Excel spr	eadsheet	to assist c	calculation	n of five-p	arameter	based orie	entation of	f intra/inte
Input valu	ues: BOL	D RED; o	utput valu	ies BOLI	) BLACK			
0 Miller/M	iller Brev	ala Camua	alana (hay					
U. Miller/N	illier-Brava	als Conver	sions (nex	agonal ar	id trigonal	lattices)		
Miller-Brav	ais to Miller							
Direction	<u< td=""><td>V</td><td>Т</td><td>W&gt;</td><td>to</td><td><u< td=""><td>v</td><td>W&gt;</td></u<></td></u<>	V	Т	W>	to	<u< td=""><td>v</td><td>W&gt;</td></u<>	v	W>
	1	1	-2	0		3	3	0
Plane	{H	K		L}	to	{h	k	}
	1	1	-2	0		1	1	0
Miller to M	iller-Bravais	N		to	- 11	V	т	
Direction	<u 3</u 	v 3	w>	10	<0 1	v 1	-2	0
Plane	ل له	k	B	to	۲. ۲H	ĸ	- <u>-</u>	
	1	1	0		1	1	-2	0
1. Input P	hase Lattic	e Constan	ts - Olivin	e				
	a	b	C					
	4.756	10.21	5.98					
2. Innut R	otation Av	is						
z. mpar n	u	v	w					
	0	0	1					
3. Calcula	te Normal	ised Rotat	ion Axis					
Products o	f lattice cor	nstants and	l crystal dir	ection				
	0.000	0.000	5.980					
SQRT OF SC	quares of pi	roducis =	5.980 ation:					
Calculate	0 000		1 000					
	0.000	0.000	1.000					
4. Calcula	te normals	s [uvw] to	planes (hk	l) in ortho	rhombic la	ttices		
Calculate r	eciprocal la	attce excep	t for [100],	[010], [001]	] // ⊥ (100),	(010), (001	)	
Lattice volu	ime =	290.381						
a* =	0.210	b* =	0.098	C* =	0.167			
Conversion		o* o*	0	0		h	,	
u v	_	a.a 0	0 b* b*	0	•	li k		
w	_	0	0	c*.c*	•			
Input plane	(hkl)		-					
		0.04421	0	0		0		
	=	0	0.009593	0	•	0		
		0	0	0.027964		1		
		0		_		•		
		0	v 100 -	0		0		
	=	0 027964	x 100 =	2 796389	-	3		
		0.027001		2.700000				
5. Calulate	e plane (hl	kl) contain	ing two di	rections <	uvw>			
Input first c	lirection <u< td=""><td>1, v1 w1&gt;</td><td></td><td>Input seco</td><td>nd direction</td><td>i <u2, td="" v2="" w2<=""><td>2&gt;</td><td></td></u2,></td></u<>	1, v1 w1>		Input seco	nd direction	i <u2, td="" v2="" w2<=""><td>2&gt;</td><td></td></u2,>	2>	
u1	v1	w1		u2	v2	w2		
1 Comuranti	0	0		0	0	3		
Conversion	expression	ר ר	- 1	0	0			
0	0	 	0	0	0 3			
Calculate o	common pla	ane		U	5			
h	k	I						
0	-3	0						
6. Calcula	te tilt/twis	t compone	nt (tilt = 0	< TTC < 1	=twist)			
Rotation a	kis (A <sub>[uvw]</sub> )				Grain bour	idary norma	l (B[uw])	
u	V	w	θ		u	V	w	θ
0 Drativit	0	1	0		0	0	3	4.32
Products 0	i lattice par	ameters ar	IU AXIS INDI	ces			14/	0
0.000	v 0.000	5 980	0,000		0,000	v 0.000	w 17 940	4 320
SQRT of se	quares of in	idices=	5.980		SQRT of s	quares of in	dices=	17.940
Calculate r	ormalised	rotation axi	s:		Calculate r	normalised	grain bound	lary normal
A* <sub>[uvw]</sub> =	0.000	0.000	1.000		N* <sub>[uvw]</sub> =	0.000	0.000	1.000
TTC =	A*,,,,,,,,]●N*,	[	1.000					