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INFLUENCE OF SALINITY BASED OSMOTIC SUCTION ON THE SHEAR STRENGTH OF A COMPACTED CLAY

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1 ABSTRACT

2 As most previous studies have neglected the positive influence of salinity (osmotic suction)
3 on most coastal soils in Australia, the design of transport infrastructure involving these soils
4 have often been overly conservative. In this study, a laboratory approach based on direct
5 shear testing is explained to determine the stress-strain behaviour of compacted coastal silty
6 clay (CL) at different levels of osmotic suction generated by varying salinity (NaCl)
7 concentrations. A broad data set for a total of 147 direct shear tests conducted on remoulded
8 and re-compacted test specimens at seven different initial matric suction conditions is
9 analysed to develop a semi-empirical model that captures the effect of osmotic suction on the
10 soil shear strength. The results suggest that greater the initial matric suction is the more
11 pronounced will be the role of osmotic suction. The proposed semi-empirical model is
12 governed by an electrical conductivity relationship with the osmotic suction generated by soil
13 salinity. A new parameter χ_2 is introduced to quantify the role of soil salinity on the apparent
14 soil shear strength corresponding to different levels of osmotic suction. When this novel
15 relationship is coupled with the conventional matric suction theory, the overall unsaturated
16 shear strength of a saline soil can be properly evaluated, as proven by the close proximity of
17 the predictions to the measurements.

18 **Keywords:** Unsaturated, matric suction, osmotic suction, shear stress

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23 INTRODUCTION

24 Most types of civil infrastructure are built on and remain under unsaturated conditions for
25 most of their service life; therefore, the stability of these structures with increased loading in
26 the future will depend on the actual shear strength of the foundation soil (sub-grade).
27 Especially in notably saline soils prevalent along the coastal belt of Australia, neglecting the
28 benefits of salinity-based osmotic suction can lead to undue design conservatism. Past studies
29 have shown that the magnitudes of both matric and osmotic suction influence the shear
30 strength of a natural or compacted soil (Graham et al., 1992, Barbour and Fredlund, 1989).
31 However, while the effect of matric suction on the shear strength is well established through
32 comprehensive testing and analysis (Khalili, 2018, Vanapalli et al., 1996, Bishop, 1960,
33 Khalili et al., 2004), only a limited number of studies have focussed on the role of osmotic
34 suction. For instance, Tiwari and Ajmera (2014), Xu (2019), and Di Maio and Scaringi
35 (2016) verified the corresponding increase in soil shear strength as the osmotic suction is
36 increased. Fredlund et al. (2012) pointed out that osmotic suction would have greater
37 influence at higher values of matric suction, hence a key reason why the authors in the
38 current study have investigated the role of osmotic suction at different levels of matric
39 suction for a coastal saline soil. Elsewhere in relation to plant morphology, Pathirage et al.
40 (2017) and Jayathilaka et al. (2019) have pointed out that the variations of soil osmotic
41 suction can also be associated with the nutrient uptake by roots; this effect is not considered
42 in this study.

43 Osmotic suction stems from the salts dissolved in the pore water of a soil, and particularly in
44 coastal areas salinity is significantly increased by sodium chloride (NaCl) that is transported
45 and deposited in inland areas over the geological time domain, i.e. in some instances, even
46 several thousands of km away from the present-day marine boundary. Arora (2017) reports

47 that around 400 million hectares, which is more than 6% of the total landmass of the earth
48 can be categorised as saline, while in Australia alone, about 30% of its landmass is
49 considered to be saline (Rengasamy, 2006).

50 The soil strength is influenced by the degree of saturation, the presence of various chemical
51 compounds, the overburden and confining ground pressure, and the fabric of the soil and pore
52 water conditions that also influence the inter-particle behaviour (Murray and Sivakumar,
53 2010). The osmotic suction induced by variations in pore water salinity can induce
54 inter-particle forces, e.g. van der Waal attraction forces, electrostatic repulsive forces, and
55 surface hydration forces (Li et al., 2013). While some of these stress components can help to
56 reduce potential swelling (Rao and Thyagaraj, 2007), as the osmotic suction changes, the
57 hydraulic and mechanical properties of a compacted soil also contribute to the thinning of the
58 diffusive double layer (DDL), cause the particles to coagulate, and increase the effective
59 stress (Di Maio et al., 2004). While the effective pressure acting on clay particles can be
60 described for instance by Derjaguin-Landau-Verwey-Overbeek (DLVO) theory of stability
61 (Liang et al., 2007), determining the magnitudes of these inter-particle forces through
62 accurate measurements is still a challenge (Fredlund et al., 2012). Electrical conductivity is
63 widely used in geophysical applications (Shevnin et al., 2010, Jiao-Jun et al., 2007, Shah and
64 Singh, 2004), while its ability to corroborate with matric suction (Hen-Jones et al., 2014,
65 Piegari and Maio, 2013) and osmotic suction (Adam et al., 2012, Read and Cameron, 1979)
66 has been widely tested and discussed. Nevertheless, the combination of electrical
67 conductivity and osmotic suction to predict the soil shear strength is still at infancy.

68 In view of the above, a novel osmotic stress parameter (χ_2) is introduced in this paper to
69 determine the stress induced by changes in osmotic suction attributed to different salt
70 concentrations in the pore water. The role that osmotic suction plays in the shear strength of a

71 soil compacted at different levels of matric suction is then investigated through a series of
72 direct shear tests.

73 OSMOTIC STRESS PARAMETER

74 A new shear strength model which can capture the influence of both the matric and osmotic
75 suction on the shear strength is introduced in this paper. Here, the shear strength of the soil
76 with pore water salinity can be partially related to the saturated shear strength parameters
77 based on the traditional effective cohesion (c') and effective friction angle (ϕ'). In addition,
78 instead of the effective stress parameter (χ_1) proposed by Khabbaz and Khalili (1998), a
79 revised osmotic stress parameter (χ_2) that depends on the level of salinity in pore water is
80 introduced as follows:

$$\sigma_{net} = (\sigma_N - u_a) + \chi_1(u_a - u_w) + \chi_2\pi \quad (1)$$

81 In the above, the term $(\sigma_N - u_a)$ is the effective normal stress, σ_N is the total normal stress,
82 the parameter $\chi_1 = \left(\frac{u_a - u_w}{AEV}\right)^{-0.55}$ is the effective stress parameter that depends on the matric
83 suction, the term $(u_a - u_w)$ is the matric suction, u_a is the pore air pressure, u_w is the pore
84 water pressure, AEV is the air entry value, and π is the osmotic suction.

85 Equation (1) conforms with the postulate that the inter-particle physio-chemical stresses can
86 be superimposed directly to the classical effective stress concept (Lu & Likos 2006; Rao &
87 Thyagaraj 2007). Also as proposed earlier by Lu and Likos (2006), the net inter-particle
88 contact forces due to physico-chemical effects can be determined by the summation of
89 chemical cementation (bond stresses), van der Waals attraction forces, and the repulsion
90 forces of DDL. Therefore, by characterising χ_2 in terms of inter-particle contact, the relevant
91 forces are more realistically acknowledged, but as noted by Fredlund et al. (2012), the

92 methods of determining or measuring the correct magnitude of the aforementioned
93 inter-particle contact forces remains a challenge.

94 EXPERIMENTAL DETERMINATION OF χ_2

95 According to Khabbaz and Khalili (1998), the shear strength of unsaturated soil can be
96 estimated in terms of the effective normal stress $(\sigma_N - u_a)$ and matric suction
97 $(u_a - u_w)$:

$$\tau'_U = [(\sigma_N - u_a) + \chi_1(u_a - u_w)] \tan \phi' + c' \quad (2)$$

98 Where, τ'_U is the unsaturated shear strength, σ_N is the total normal stress, u_a is the pore air
99 pressure, u_w is the pore water pressure, χ_1 is the effective stress parameter that depends on the
100 matric suction, ϕ' is the effective friction angle, and c' is the effective cohesion component.
101 Since the effective stress generated in an unsaturated soil element is defined by considering
102 the net stress constituting the pore water and pore air pressures, pore water salinity (hence,
103 the matric and osmotic suction, and physio-chemical pressure), it is assumed that the shear
104 strength parameters c' and ϕ' of saturated soil are independent of the matric suction
105 (Khabbaz and Khalili, 1998) and of the salinity-based osmotic suction (Lu and Likos, 2006).
106 Combining these concepts in a mathematical sense, a refined shear strength model for
107 unsaturated-osmotic conditions is now defined by the following expression:

$$\tau'_{US} = [(\sigma_N - u_a) + \chi_1(u_a - u_w) + \chi_2\pi] \tan \phi' + c' \quad (3)$$

108 Where τ'_{US} is the shear strength of osmotically induced unsaturated soil. The difference
109 between Equations (2) and (3) is given by the osmotic suction component as follows:

$$\tau'_{US} - \tau'_U = \chi_2\pi \tan \phi' \quad (4)$$

110 Hence, according to Equation (4), the only unknown parameter χ_2 can be estimated.

$$\chi_2 = \frac{\tau'_{us} - \tau'_u}{\pi \tan \phi'} \quad (5)$$

111 RELATIONSHIP BETWEEN χ_2 AND THE ELECTRICAL CONDUCTIVITY RATIO

112 Electrical conductivity is a function of the salt concentration so it can also be considered as a
 113 function of osmotic suction (Abedi-Koupai and Mehdizadeh, 2007). To investigate the
 114 behaviour of χ_2 the electrical conductivity of soil is used as an additional influencing factor.
 115 A semi-empirical model parameter (i.e. χ_2) is introduced based on the experimental results.
 116 In this model, ECR is used to represent the change of salinity in pore water and S_r is used to
 117 incorporate the influence of the degree of saturation.

$$\chi_2 \begin{cases} = 0 & \pi = 0 \\ = \frac{a}{S_r^c} (1 - \exp(-b(ECR))) & \pi \neq 0 \end{cases} \quad (6)$$

118 In the above, the electrical conductivity ratio (ECR) = $\frac{\Delta EC}{EC_i}$, $\Delta EC = (EC - EC_i)$, EC is the
 119 electrical conductivity of the saturated soil for a given salt concentration in pore water, EC_i is
 120 the initial electrical conductivity of saturated soil remoulded with distilled water, S_r is the
 121 degree of saturation, and a , b and c are empirical coefficients from the regression analysis of
 122 experimental results.

123 The sensitivity of χ_2 depends mainly on the above three experimental coefficients, and their
 124 influence on χ_2 is shown in Fig. 1. The coefficient a increases to the maximum of χ_2 ,
 125 however, coefficient b does not influence the maximum value of χ_2 , but rather contributes to

126 the lowest value of ECR having the maximum χ_2 . The maximum theoretical value of χ_2 can
127 then be calculated from Equation (7).

$$\chi_{2max} = \left(\frac{a}{S_r^c}\right) \quad (7)$$

128 According to Equation (7), χ_{2max} does not depend on the electrical conductivity ratio (ECR),
129 although χ_2 is a function of ECR. This is further validated by Fig. 1(d) which shows the
130 distribution of χ_2 with respect to ECR for different $\frac{a}{S_r^c}$ ratios. According to Fig. 2, the
131 distribution of χ_{2max} with respect to the empirical coefficient c is linear at higher degrees of
132 saturation, but the distribution of χ_{2max} becomes exponential as the degree of saturation
133 decreases. Therefore, at higher matric suctions (i.e. lower degree of saturation), the influence
134 of χ_2 will be significant compared to that of a soil having a lower soil matric suction (high
135 degree of saturation).

136 The minimum ECR value where χ_2 reaches its maximum and the critical ECR (i.e. ECR_C)
137 depends on the empirical coefficient b [Fig. 1(b)], therefore, χ_{2max} increases as the soil
138 approaches a dry state. The distribution of ECR_C with respect to the empirical coefficient b
139 for different levels of saturation is shown in Fig. 3. Here, for every degree of saturation, the
140 ECR_C decreases as the coefficient b increases, following a power decay function. For
141 example, under fully saturated conditions ($S_r = 1$), when the coefficient b increases from 0.01
142 to 0.08, the ECR_C decreases from 750 to 125.

143 MATERIALS AND TESTING PROGRAM

144 **Soil type and preliminary testing**

145 The soil samples were obtained from the coastal region of Wollongong (85 km south of
146 Sydney). The particle size distribution (AS1289.3.6.1) indicates it consists of sand (48%), silt
147 (36%), and clay (16%) particles (Fig. 4). The liquid and plastic limits (AS1289.3.1.1 and
148 AS1289.3.2.1) are 46.8% and 27.7%, so this soil can be classified as a sandy, silty clay of
149 low plasticity, CL, based on the ASTM Unified soil classification (ASTM D2487 2010). The
150 modified Proctor compaction characteristics according to AS1289.5.1.1 enabled a maximum
151 dry density (MDD) of 15.58 kN/m^3 at an optimum moisture content of 27.2%, and a specific
152 gravity of the soil was determined to be 2.62 (AS1289.3.5.2).

153 The matric suction of the compacted soil was measured using the contact method and
154 Whatman No 42 filter paper approach (ASTM D5298-03). The soil water retention data was
155 only measured for test specimens remoulded with distilled and de-aired water. Thirteen
156 different samples were prepared at different moisture contents, and then they were air sealed
157 and stored for seven days in a temperature and humidity controlled room ($20 \pm 2^\circ\text{C}$, 30%RH)
158 to attain moisture equilibrium. The samples were then compacted into a 50mm diameter
159 cylindrical mould to 85% of MDD. Fig. 5 shows the soil water retention curve calibrated
160 with the Van Genuchten (1980) model (Van Genuchten 1980) with best-fit parameters:
161 $m=0.306$, $n=1.44$, and $\alpha=0.008$.

162 Although, the osmotic suction can be theoretically calculated according van Hoff's equation
163 ($\pi = \nu R^*TC$), where π is the osmotic suction, R^* is the universal gas constant, T is the
164 absolute temperature, ν is the valency, and C is the ion concentration. In this study, the actual

165 osmotic suction was measured using a WP4 Dew Point Potentiometer. Crystallised NaCl
166 were mixed with distilled water to prepare a solution with the desired salt concentrations.
167 Although the soil contains constant ion content, the pore water salinity is likely to increase
168 due to soil moisture decrease (i.e. caused by a rise in global temperature induced by climate
169 change). Therefore, it is appropriate to consider a broader range of salinity values. Therefore,
170 seven soil samples were fully saturated with solutions having NaCl concentrations of 0.0, 0.2,
171 0.4, 0.6, 0.8, 1.0 and 2.0 mol/L, where the maximum salinity of the studied soil was three
172 times more than the maximum salinity of seawater (i.e. 35 g/L). Also, the inherent salt
173 content of the soil specimen was determined by X-Ray diffraction (XRD) as less than 0.1%.
174 Therefore, for this study, the contribution of inherent salt content could be assumed as
175 negligible. The samples were then fully sealed and stored in a temperature and humidity
176 controlled room ($20\pm 2^\circ\text{C}$, 30% RH) for another 24 hours. Although the WP4C Dew Point
177 Potentiometer could be used to measure total suction, in this study, as the seven test
178 specimens remained fully saturated the matric suction was 0 kPa. On this basis, the measured
179 total suction was assumed to be equal to the osmotic suction of the specimen. The moisture
180 equilibrated samples were placed into clean plastic cups and tested with a WP4C Dew Point
181 Potentiometer (range 0 to 300 MPa). The measured values of osmotic suction are summarised
182 in Table 1.

183 Seven samples saturated with different pore water salinities were prepared and stored in the
184 temperature and humidity-controlled room for 24 hours. These moisture equilibrated samples
185 were then compacted to a dry density of 13.24 kN/m^3 (85% of MDD) and then placed into a
186 standard electrical resistivity measuring box having dimensions of 38 x 101.5 x 152.3 mm.
187 The electrical resistivity was measured using a Tinker and Rasor SR-2 soil resistivity meter
188 ($\pm 0.1 \Omega\text{cm}$ accuracy) and then converted to electrical conductivity ($=1/\text{Electrical resistivity}$).

189 The distribution of electrical conductivity with the pore water saline concentration is
190 summarised in Fig. 6.

191 **Direct shear test**

192 Seven different solutions with different osmotic suctions were prepared by mixing the
193 relevant amount of commercially available crystallised NaCl with distilled water; only NaCl
194 was used to mimic the conditions of coastal soils. Different levels of initial matric suction
195 (0, 25, 100, 200, 500, 1000 and 1500 kPa) were targeted by controlling the moisture content
196 of the specimens. The required amount of water and the relevant salt concentration were
197 added to the soil, and then the mixture was left in the temperature and humidity-controlled
198 room for seven days for chemical and moisture equilibration. The samples were then
199 compacted in a 60×60×40 mm shear box chamber to attain 85% of MDD, and then stored in
200 the temperature and humidity-controlled room for two more days.

201 A motor-driven direct shear box where the specimen carriage travels on roller bearings was
202 used to maintain a constant rate of horizontal displacement of 0.006 mm/min. A load cell and
203 two LVDT transducers accurate within ± 0.001 kN, ± 0.001 mm and ± 0.001 mm, respectively,
204 were used to determine the horizontal shear force, vertical displacement, and horizontal
205 displacement. A lever arm loading system (beam ratio 10:1) was used to apply a vertical load
206 by a top cap modified to accommodate a miniature pore water pressure transducer, so that
207 any variations in the matric suction could be monitored during shearing. An in-house coded
208 program with Lab VIEW software complemented by a National Instruments card
209 (NI USB-6009) with eight input channels was used to acquire data every 60 seconds. A
210 schematic diagram and few images of the actual test set up of the direct shear box are given
211 in Fig. 7.

212 Apart from those specimens with a matric suction of 0 kPa (fully saturated conditions), direct
213 shear tests were carried out at different levels of matric suction under constant water (CW)
214 contents. For fully saturated experiments where the matric suction = 0 kPa, the compacted
215 specimens were fully submerged in the desired saline solution for 24 h before shearing to
216 attain moisture and chemical equilibration. To ensure CW conditions, evaporation from the
217 soil specimen had to be minimised so the compression and shearing stages of the direct shear
218 tests could occur within a temperature and humidity-controlled environment. The top and
219 bottom surfaces of all test specimens were covered with a 1mm thick film of polyethylene to
220 minimise evaporation. Moreover, the space between the top and bottom sliding halves, and
221 any other gaps between the top cap and bottom plate were sealed with silicon grease. The
222 volume of air around the specimen was reduced by enclosing the direct shear box and the
223 assembly inside an airtight polythene bag which was then covered with a damp cloth to
224 reduce any variations in temperature inside the polythene bag (Fig. 7). The moisture contents
225 of soil specimens before the compaction and after the direct shear test were determined.
226 However, the average moisture content variation was found to be negligible ($< 0.1\%$).
227 Therefore, soil specimen was assumed to be at a constant water content condition during
228 compaction and shearing.

229 During the compression stage, the specimens were loaded vertically in 10, 20 and
230 40 kPa steps, where each load increment was left for one to two days until the variations of
231 vertical displacement became insignificant ($< 1\%$). The specimens were then sheared at a
232 relatively low shear strain rate of 0.006 mm/min, in order to accommodate the redistribution
233 of any variation in matric suction induced by the shearing process. Shearing continued until a
234 maximum shear strain of 25% was achieved.

235 RESULTS AND DISCUSSION

236 **Stress-strain behaviour**

237 The shear stress and strains of 147 soil samples were measured at three different normal loads
238 for given osmotic suctions and various initial matric suctions. Of those, the distributions of
239 shear stress and normal strain were plotted against the shear strain for different osmotic
240 suctions for fully saturated conditions. The stress-strain behaviour of the fully saturated soil
241 without the influence of salinity ($\pi = 0$ kPa), was determined to consider as a reference to
242 compare the stress-strain behaviour of the soil with variable osmotic suctions under
243 unsaturated conditions.

244 The saturated stress-strain behaviour of the soil for various osmotic suctions for a given
245 normal stress ($\sigma'_N = 10$ kPa) is shown in Fig. 8. The results show that the peak shear stress
246 increases gradually with the influence of osmotic suction, showing a maximum increase of
247 around 13.48 kPa. Moreover, the stress-strain behaviour of the unsaturated soil was
248 monitored for six different matric suction conditions as discussed above. Of those, the
249 stress-strain distribution for two different pressure conditions (different applied normal
250 stresses and matric suctions) is shown in Fig. 9. Similar to the behaviour of saturated soil, the
251 peak stress of unsaturated soil increases with osmotic suction. The reason for this is the
252 increased resistance for the relative movement of soil particles due to the increase in inter-
253 particle bond strength. However, the increase of peak shear stress of unsaturated soil with
254 respect to osmotic suction was higher compared to the saturated condition for a given normal
255 stress. For example, the soil specimen subjected to 1500 kPa matric suction shows an
256 increase of around 98.96 kPa of peak shear stress for the highest osmotic suction increase,
257 while the saturated soil shows only about 13.48 kPa for the same increase of osmotic suction.

258 Hence, it is evident that both matric suction (as expected) and osmotic suction has a
259 significant influence on the peak shear stress.

260 The influence of osmotic suction on normal strain response of the saturated and unsaturated
261 soil for a given normal stress is shown in Fig. 8 and 9. The results indicate that for all the
262 osmotic and matric suction conditions, the normal strain decreases, showing a contractive
263 behaviour of the specimens. Furthermore, an increase in osmotic suction results in lower
264 contraction of the specimen for both saturated and unsaturated conditions. Also, as expected,
265 at higher matric suctions the specimens exhibit a lower contraction behaviour compared to
266 saturated conditions. This contractive behaviour of soil specimens can be further elaborated
267 with maximum normal strain results. The maximum normal strain is considered as the lowest
268 achieved normal strain of the specimen. The distribution of maximum normal strain with
269 respect to osmotic suction for various matric suctions for a given normal stress
270 ($\sigma'_N = 10$ kPa) is shown in Fig. 10. For all the matric suction conditions, the maximum
271 normal strain significantly decreases with osmotic suction, showing the highest decrease of
272 change of maximum normal strain of around 4.7%. This could be because of the increased
273 resistance to the relative movement of particles due to the influence of osmotic suction.
274 Further as expected, the change of maximum normal strain also decreases with matric
275 suction. Interestingly, while the change of maximum normal strain without the influence of
276 osmotic suction is around 2.7% ($\sigma'_N = 10$ kPa), at higher matric suctions (1500 kPa) the
277 change of maximum normal strain decreases to 0.27% for the same increase of osmotic
278 suction.

279 **Model calibration and validation**

280 The peak shear strength was determined from the results of the direct shear tests. The
281 summary of all the peak stress results is given in Table 2. The experimental distribution of
282 χ_2 was calculated based on Equation (5). The saturated friction angle was calculated when the
283 soil sample became fully saturated with distilled water ($\pi = 0$ kPa). The unsaturated
284 behaviour of soil is influenced by the level of matric suction. Therefore, the influence that the
285 matric suction has on χ_2 was considered by incorporating the corresponding degree of
286 saturation into Equation (6). Due to the limited available literature for this soil suction range,
287 three independent data sets have been used for calibration (i.e. $\sigma'_N = 20$ kPa) and validation
288 (i.e. $\sigma'_N = 10$ and 40 kPa). The proposed new model for χ_2 [Equation(6)] was calibrated for
289 three major initial matric suction conditions such as $s_i = 0$ kPa (saturated), $s_i = 200$ kPa and
290 500 kPa, with respect to experimental results for a given normal stress ($\sigma'_N = 20$ kPa). The
291 distribution of χ_2 with the electrical conductivity ratio for three different levels of matric
292 suctions is shown in Fig. 11; it was used to estimate the best-fit parameters which were then
293 used to predict the unsaturated behaviour of soil in combination with the degree of saturation
294 for the other independent data sets. The fully saturated condition was used to determine the a
295 and b coefficients when the influence of c was not significant ($S_r = 1$). Then the parameter c
296 was determined based on the results from $s_i = 200$ kPa condition, and also further calibrated
297 all the three parameters with $s_i = 500$ kPa. Based on these determinations, the calibrated
298 parameters are $a = 0.003$, $b = 0.0375$ and $c = 2$. The proposed model was validated for two
299 independent loading conditions ($\sigma'_N = 10$ and 40 kPa) with the calibrated model parameters,
300 and the corresponding validation results are shown in Fig. 12. In general, the model
301 predictions match the experimental results very well, thus indicating that the proposed model
302 incorporating χ_2 is able to predict the osmotically induced shear strength of a saline soil.

303 The above results also show that χ_2 increases with an increasing ECR ($\Delta EC/EC_i$), but this
304 increase in χ_2 also decreased at high ECR values until it reached a maximum theoretical χ_2
305 value of 0.003 under saturated conditions irrespective of applied stress. The maximum
306 theoretical χ_2 increases as the initial degree of saturation decreases or the initial matric
307 suction increases. The minimum value of ECR where χ_2 reached its maximum or the critical
308 ECR (ECR_c), does not depend on the initial matric suction; hence it is evident that the ECR_c
309 is a parameter which only depends on the pore solution and surface potential of soil particles.

310 The predicted peak shear stress was calculated based on Equation (3) and the results were
311 compared with the experimental results for two independent applied normal stress conditions;
312 the corresponding distribution of model prediction and experimental results of peak shear
313 stress is shown in Fig. 13. The model predictions match the experimental results at lower
314 initial matric suctions (< 500 kPa), giving a maximum deviation of less than 5kPa. However,
315 as the initial matric suction (>500 kPa) increases, the model shows a slight deviation
316 (5 to 14.5 kPa) depending on the magnitude of osmotic suction and matric suction. Overall,
317 the model exhibits an increased deviation from the experimental results at the highest values
318 of osmotic suction and initial matric suction. The maximum deviation of model results from
319 experimental results for any condition is about 14.5kPa when the osmotic suction increases to
320 9560 kPa at the highest considered initial matric suction of 1500 kPa.

321 MODEL LIMITATIONS

322 The proposed unsaturated shear strength model was primarily based on the salinity level
323 (hence, osmotic suction) and the initial matric suction. The model was calibrated and
324 validated using the shear box results for a clayey soil of low plasticity (CL; PI=19) under
325 constant water content. The measured maximum matric suction changes upon shearing were

326 generally very small compared to the relatively high initial matric suction, as shown in
327 Fig. 14 for a typical test specimen. Therefore, while the proposed model is accurate under
328 these specific conditions, when considering its broader application to other soils, the
329 following limitations can be elucidated.

- 330 • This unsaturated shear strength model was validated only for a single soil of low
331 plasticity (CL). Therefore, the application of the model to soils of much higher
332 plasticity(e.g. CH, OH, MH) will require caution to be exercised.
- 333 • Duringshearing (constant water content) the pore structure (void ratio) can change
334 with an accompanied change in the degree of saturation. Fig. 14 shows for an initial
335 matric suction of 1500 kPa, the maximum change in matric suction upon shearing is
336 in the proximity of 35 kPa (< 2.5%). For soils of different fabric that significantly
337 dilate upon shearing(e.g. dense granular soils or highly compacted fills), the shearing-
338 induced matric suction changes may be large enough to induce notable discrepancies
339 of the proposed shear strength model.
- 340 • Under near saturation, the role of matric suction will be eliminated, hence the shear
341 strength parameter (χ_2) represented by Equation (6)becomes simplified as a sole
342 function ofsalinity. In this regard, further tests conducted at much greater osmotic
343 suction (e.g. in the proximity of say 100 MPa)will be desirable to calibrate χ_2 more
344 accurately.
- 345 • It is appreciated that remoulded soil specimens may not truly represent the actual
346 hydro-mechanical behaviour of in situ soil. However, due to technical difficulties in
347 obtaining many identical undisturbed test specimens (i.e. same microstructure and
348 pore water salinity), remoulded samples were used for this study.Undisturbed block

349 samples to fit the dimensions of the shear box apparatus will certainly be considered
350 in the future.

351 • In the field, given the climatic and environmental influences, the pore water salinity
352 can vary with time due to ion exchange. In this study, time-dependent change in
353 salinity was not considered.

354 • The proposed model was influenced by electrical conductivity measurement, where
355 only the role of NaCl was considered. However, the model can deviate from accuracy
356 if the soil solution contains other cations such as Fe^{3+} or Fe^{2+} .

357 • At very high matric suctions existing under exceedingly dry conditions, the effect
358 of salt crystallisation on the soil shear strength cannot be predicted by this model.

359 CONCLUSION

360 A series of direct shear box tests were carried out on a typical unsaturated saline soil (CL)
361 subjected to different levels of osmotic suction with known values of initial matric suction.
362 While still embracing the Mohr-Coulomb mathematical framework, this study proposed a
363 new relationship to capture the role of osmotic suction by introducing a new parameter χ_2 (i.e.
364 as an independent term to χ_1) in the original Bishop's unsaturated shear strength model
365 modified by Khabbaz and Khalili (1998). The following key conclusions can be drawn based
366 on the results of this study.

367 • The parameter χ_2 representing the role of osmotic component can be estimated
368 based on the electrical conductivity ratio (ECR). The maximum value of χ_2 and the
369 corresponding minimum (critical) value ECR_c for a given degree of saturation define
370 the appropriate bounds of χ_2 that reaches its maximum of 0.003 when $ECR_c = 900$ at
371 full-saturation ($S_r = 1$).

- 372 • At lower values of both osmotic suction and initial matric suction, the predicted value
373 of χ_2 from Equation (6) agrees with the experimental results for $a = 0.003$, $b = 0.0375$
374 and $c = 2$. However, at high levels of osmotic suction (i.e. $\pi > 4500$ kPa) and at high
375 initial matric suction (i.e. $s_i > 500$ kPa), the model deviates from accuracy.
- 376 • The results of this study prove that for a given increase in osmotic suction, the peak
377 shear stress can significantly increase for both the unsaturated and saturated soil
378 specimens. For the unsaturated soil subjected to an initial matric suction of 1500 kPa,
379 when the osmotic suction increased from 0 to 9560 kPa, the corresponding peak shear
380 stress increased significantly by about 75% from 133 kPa to 232 kPa. For the same
381 increase in osmotic suction, the corresponding increase in peak shear stress of
382 saturated test specimens from 11.5 kPa to 24.9 kPa may not seem substantial at a
383 glance, but it is noteworthy that this increase is still more than double, hence
384 demonstrating the beneficial influence of salinity even under saturated conditions.

385 DATA AVAILABILITY STATEMENT

386 All data, models, and code generated or used during the study appear in the submitted article.

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492 LIST OF TABLES

493 Table 1 The osmotic suctions at different concentrations of NaCl

494	Concentration (mol/L)	Measured osmotic suction (kPa)
495	0.0	0.0
496	0.2	910
497	0.4	1790
498	0.6	2700
499	0.8	3690
500	1.0	4650
501	2.0	9560

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Table 2 Experimental summary of peak shear stress

Matric suction (kPa)	σ'_N (kPa)	Peak shear stress (kPa)						
		$\pi = 0$ kPa	$\pi = 910$ kPa	$\pi = 1790$ kPa	$\pi = 2700$ kPa	$\pi = 3690$ kPa	$\pi = 4650$ kPa	$\pi = 9560$ kPa
0	10	11.46	11.71	12.71	13.96	15.48	17.02	24.93
	20	16.49	16.89	18.01	18.99	20.54	22.55	31.05
	40	27.29	27.69	28.54	29.79	31.09	33.79	42.56
25	10	24.62	24.97	26.12	27.42	29.32	30.92	37.99
	20	29.71	30.11	31.63	32.21	34.51	35.61	45.30
	40	40.45	40.83	41.93	42.85	44.95	46.58	56.30
100	10	47.42	47.71	49.11	51.29	53.64	55.81	65.47
	20	52.30	52.89	54.30	55.96	58.74	60.05	74.51
	40	63.24	63.60	65.43	66.95	68.69	71.99	83.35
200	10	60.58	61.33	63.26	65.70	68.39	71.92	84.98
	20	65.89	66.41	68.72	71.25	73.14	77.66	93.11
	40	76.32	76.90	79.43	80.89	83.18	87.87	103.43
500	10	85.65	86.54	90.97	94.95	100.09	105.88	129.04
	20	90.75	92.17	94.30	99.62	105.63	110.97	137.16
	40	101.46	102.63	106.04	108.99	113.91	120.57	149.33
1000	10	112.81	114.37	120.08	127.35	135.70	142.78	187.60
	20	118.00	120.08	125.66	130.99	142.94	149.74	196.51
	40	128.58	130.14	136.37	145.20	150.44	160.21	204.57
1500	10	133.10	135.47	143.55	152.41	162.42	173.77	232.05
	20	137.20	139.80	147.26	153.43	168.36	178.99	239.32
	40	148.90	150.65	158.51	167.73	182.01	190.48	245.12

510 LIST OF FIGURES

511 Fig. 1 Sensitivity analysis of χ_2 (a) with varying a , (b) with varying b , (c) with varying S_r and
512 (d) with varying $\frac{a}{s_r^c}$

513 Fig. 2 Distribution of $\chi_{2_{max}}$ with respect to coefficient c for lower and higher degrees of
514 saturation

515 Fig.3 Distribution of critical ECR (ECR_c) with respect to coefficient b for different degrees of
516 saturation

517 Fig. 4 Particle size distribution of soil

518 Fig. 5 Soil water retention curve (SWRC)

519 Fig. 6 Distribution of electrical conductivity of soil specimens

520 Fig. 7 Direct shear box setup with the instrumental locations, (a) Schematic view of the shear
521 box, and (b) Images of the test setup

522 Fig. 8 Stress-strain behaviour of the saturated soil for various osmotic suctions ($\sigma'_N = 10$ kPa)

523 Fig. 9 Stress-strain behaviour of the soil under unsaturated conditions, (a) $s_i = 500$ kPa and
524 $\sigma'_N = 20$ kPa, (b) $s_i = 1500$ kPa and $\sigma'_N = 40$ kPa

525 Fig. 10 Distribution of maximum normal strain with osmotic suction for different initial
526 matric suctions ($\sigma'_N = 10$ kPa)

527 Fig. 11 Model calibration with $s_i = 0$ kPa, $s_i = 200$ kPa and $s_i = 500$ kPa ($\sigma'_N = 20$ kPa)

528 Fig. 12 Experimental and model prediction results of χ_2 for different initial matric suctions
529 with $a = 0.003$, $b = 0.0375$ and $c = 2.0$, when (a) $\sigma'_N = 10$ kPa and (b) $\sigma'_N = 40$ kPa

530 Fig. 13 Distribution of experimental and model peak shear stress for different initial matric
531 suctions and saturated osmotic suction, (a) $\sigma'_N = 10$ kPa and (b) $\sigma'_N = 40$ kPa

532 Fig. 14 Variation of matric suction during shearing ($\sigma'_N = 10$ kPa and $\pi = 0$ kPa)

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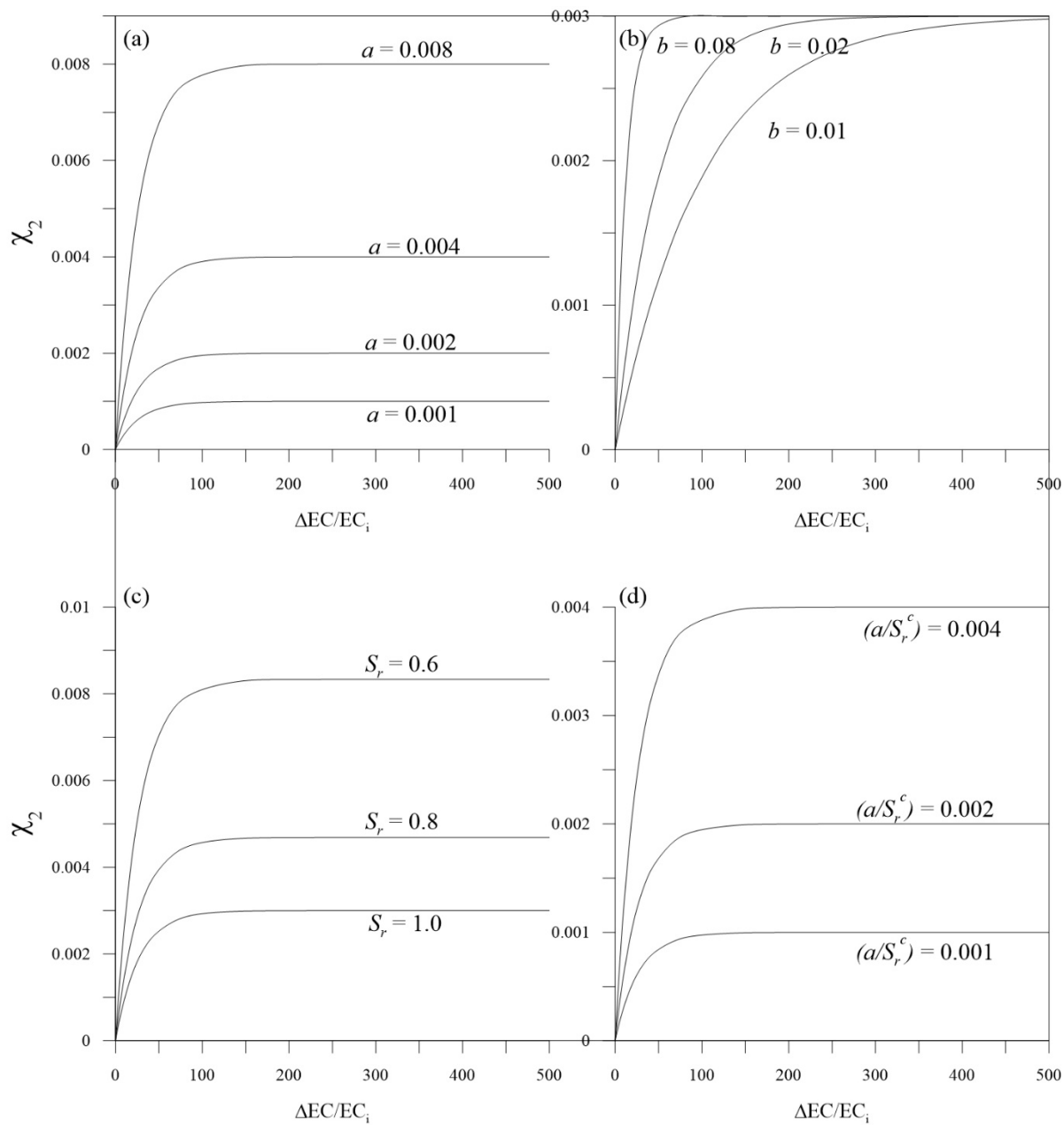
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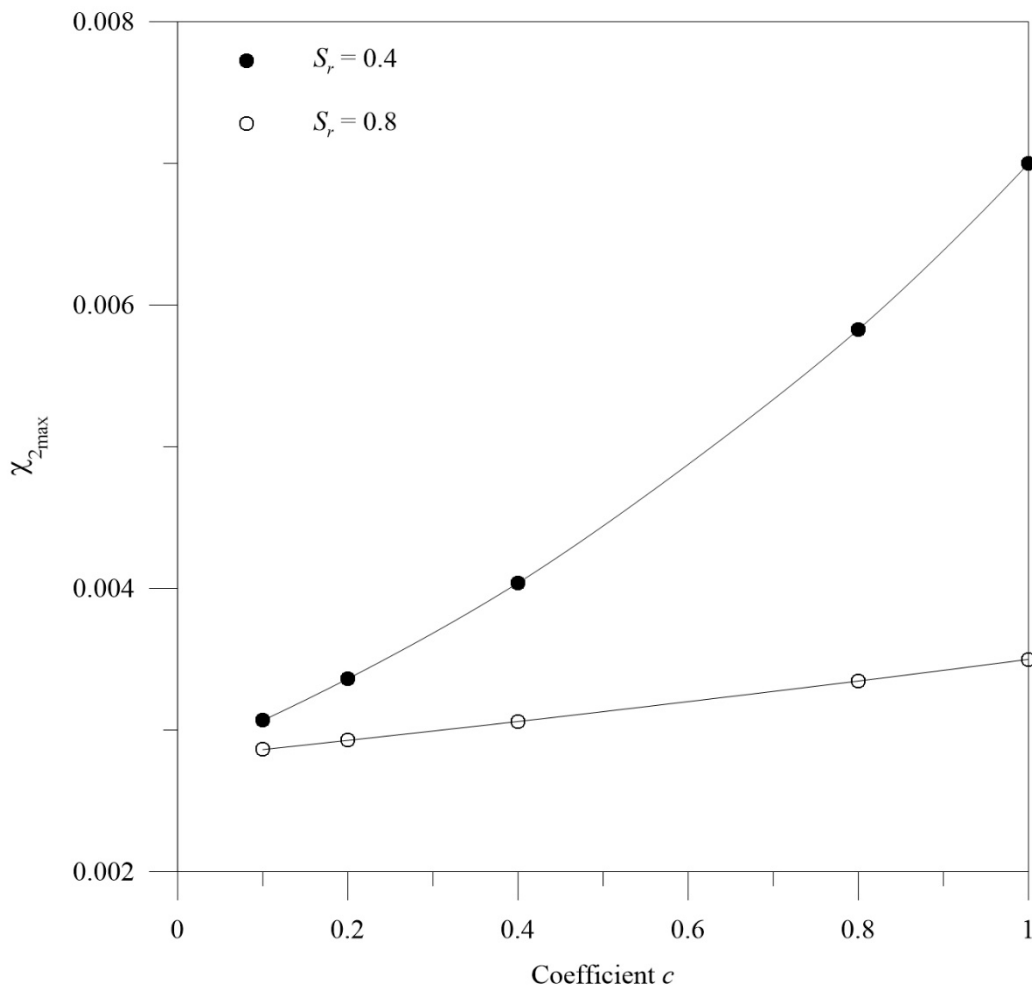
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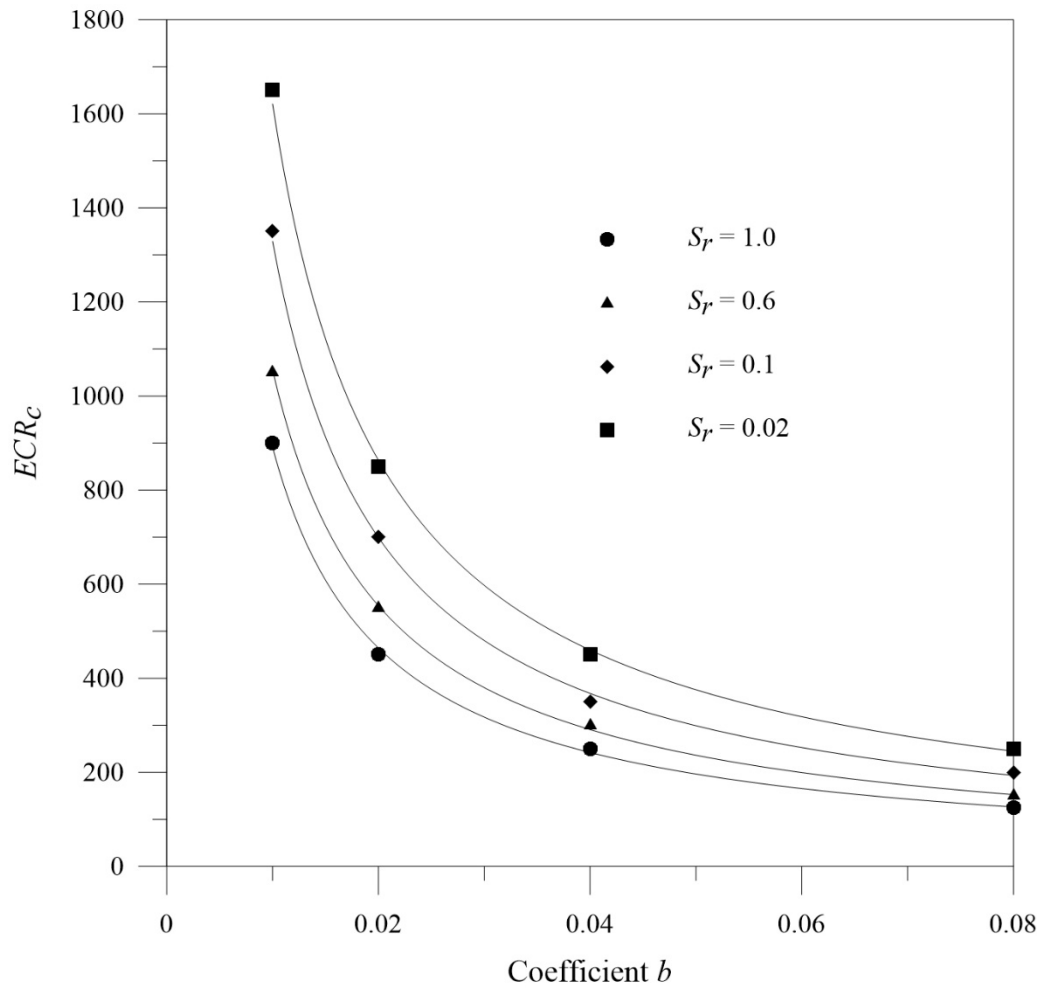
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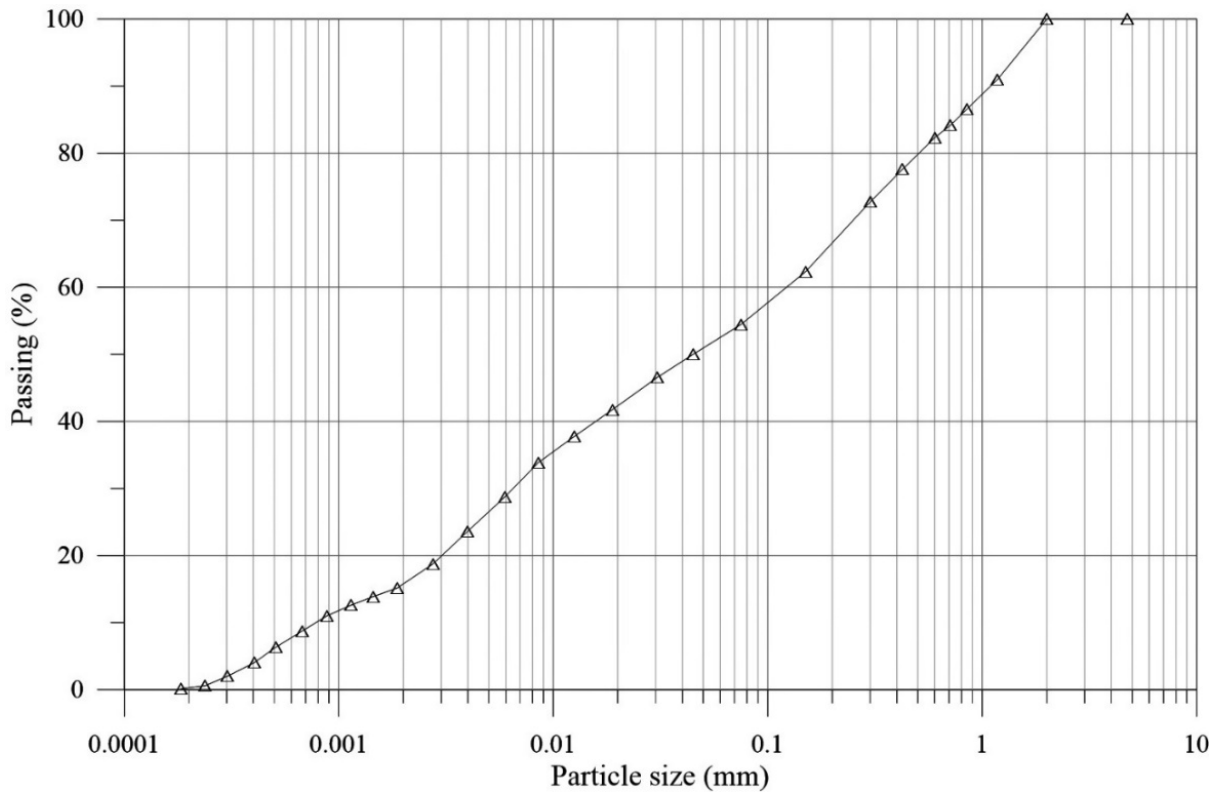
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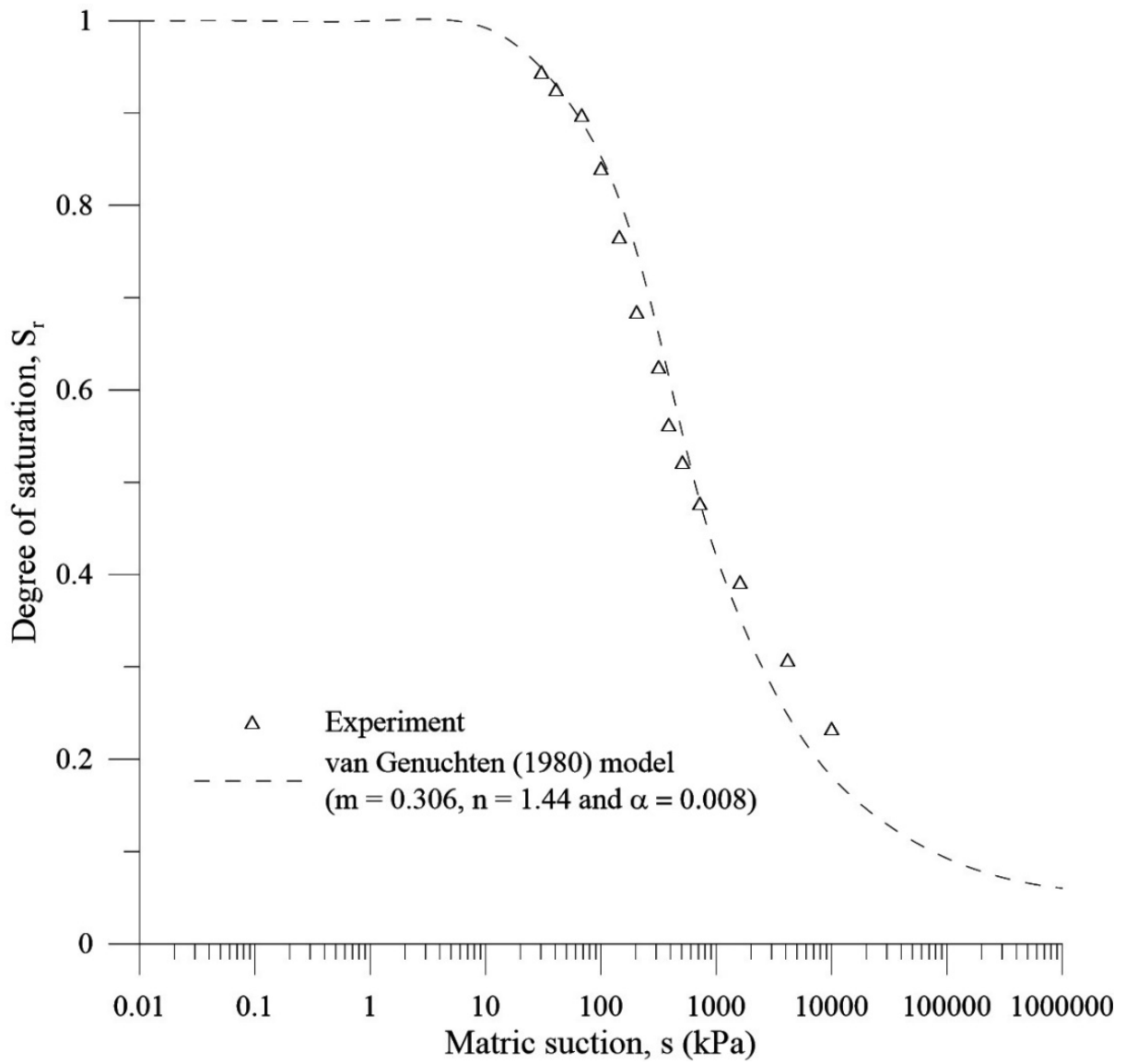
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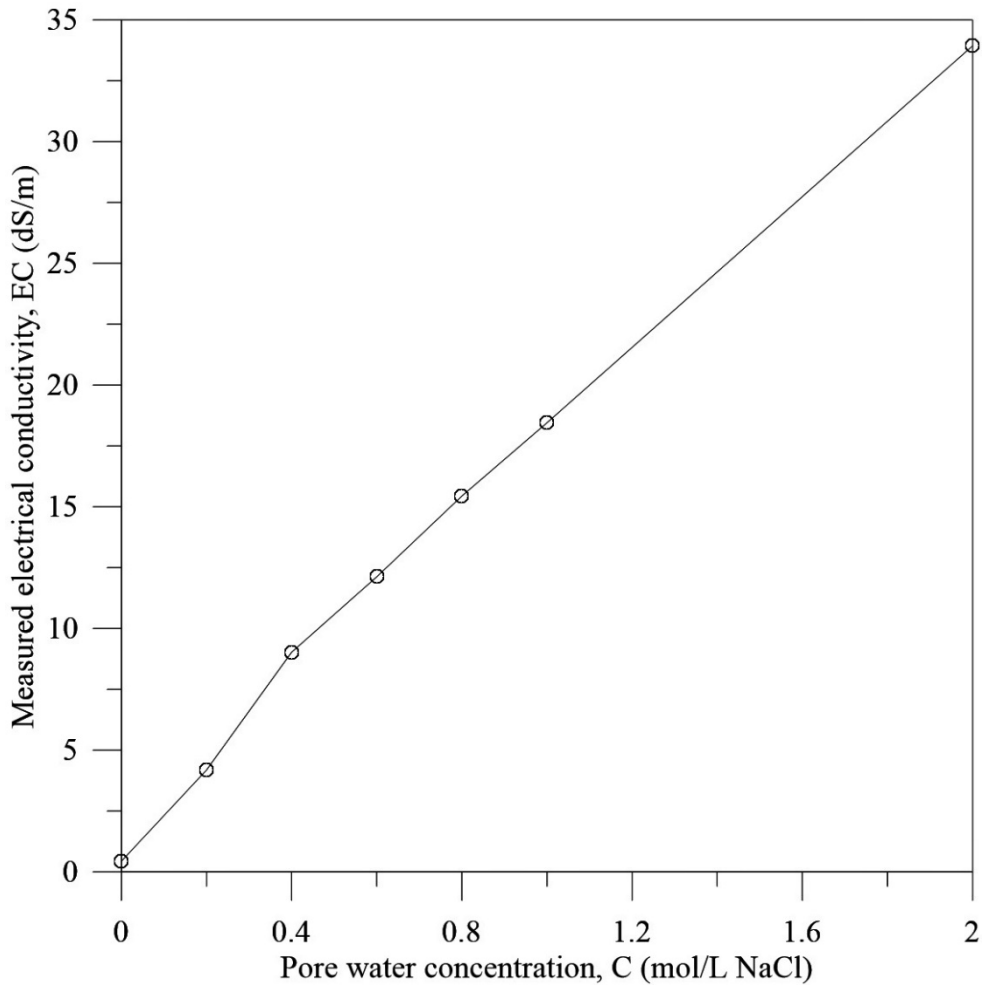
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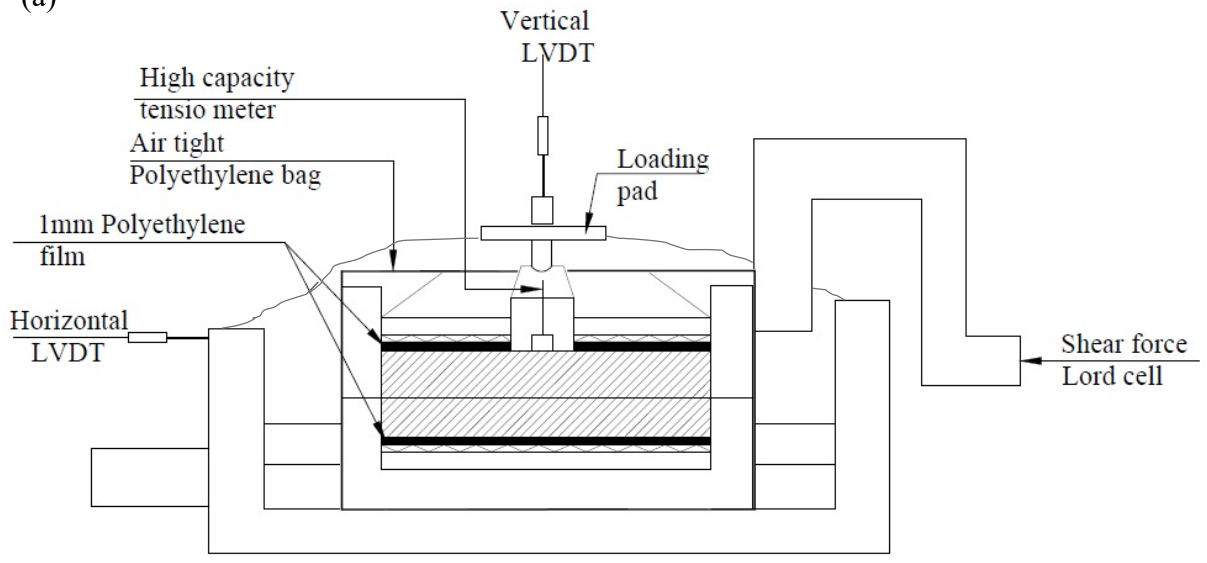
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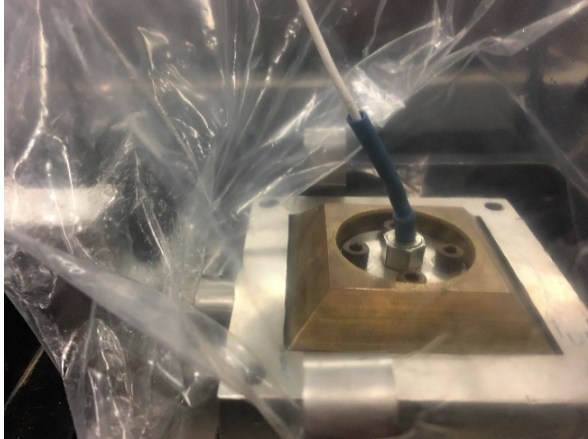
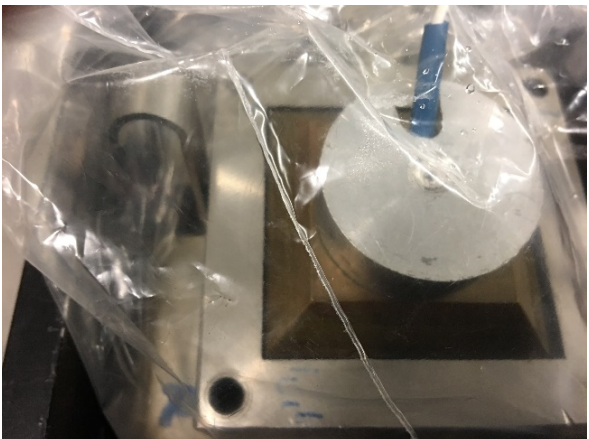
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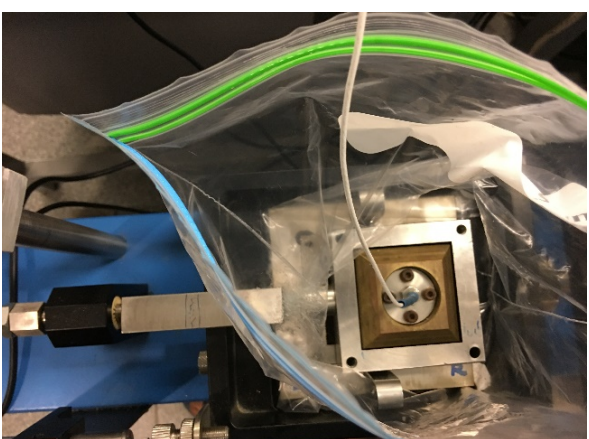


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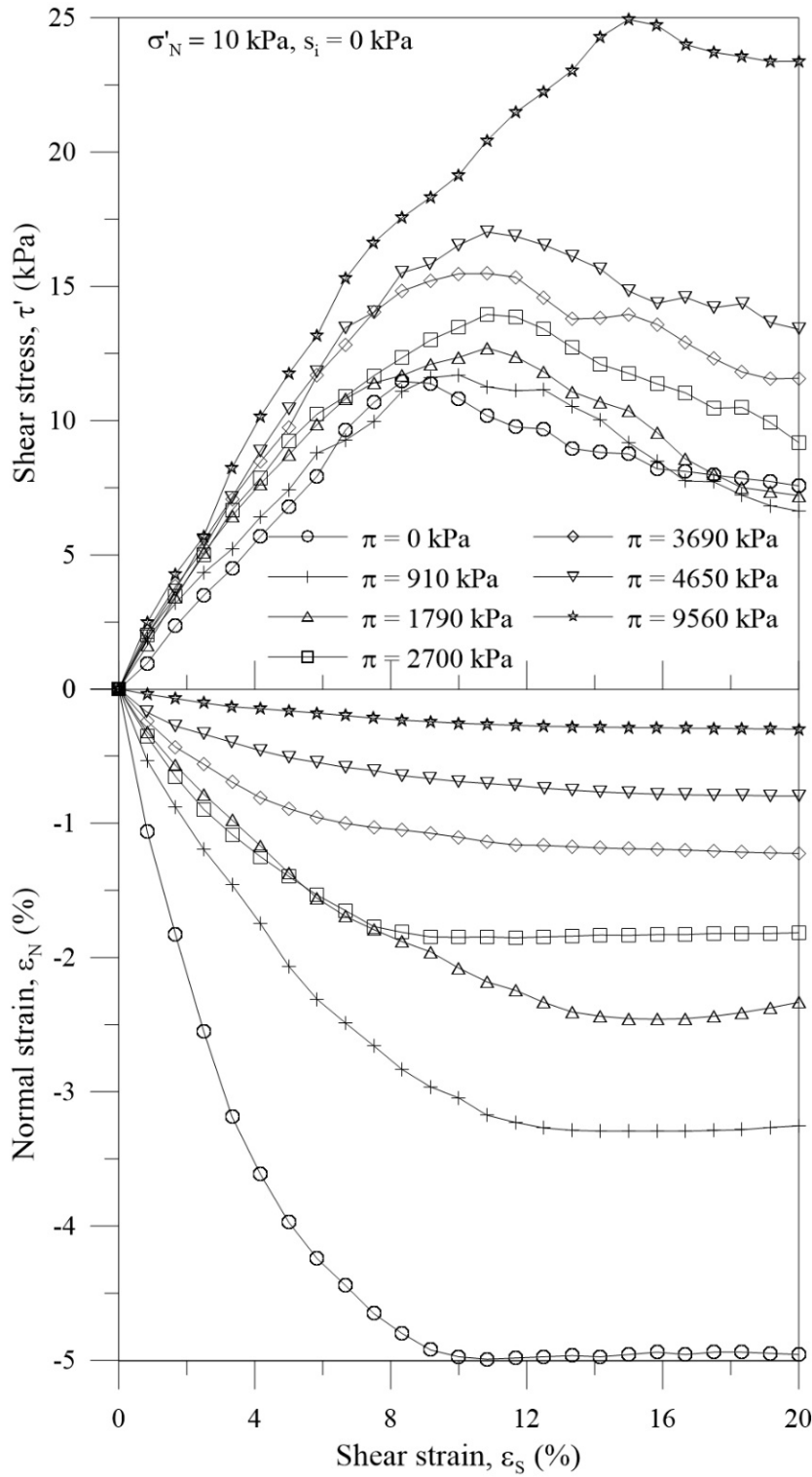
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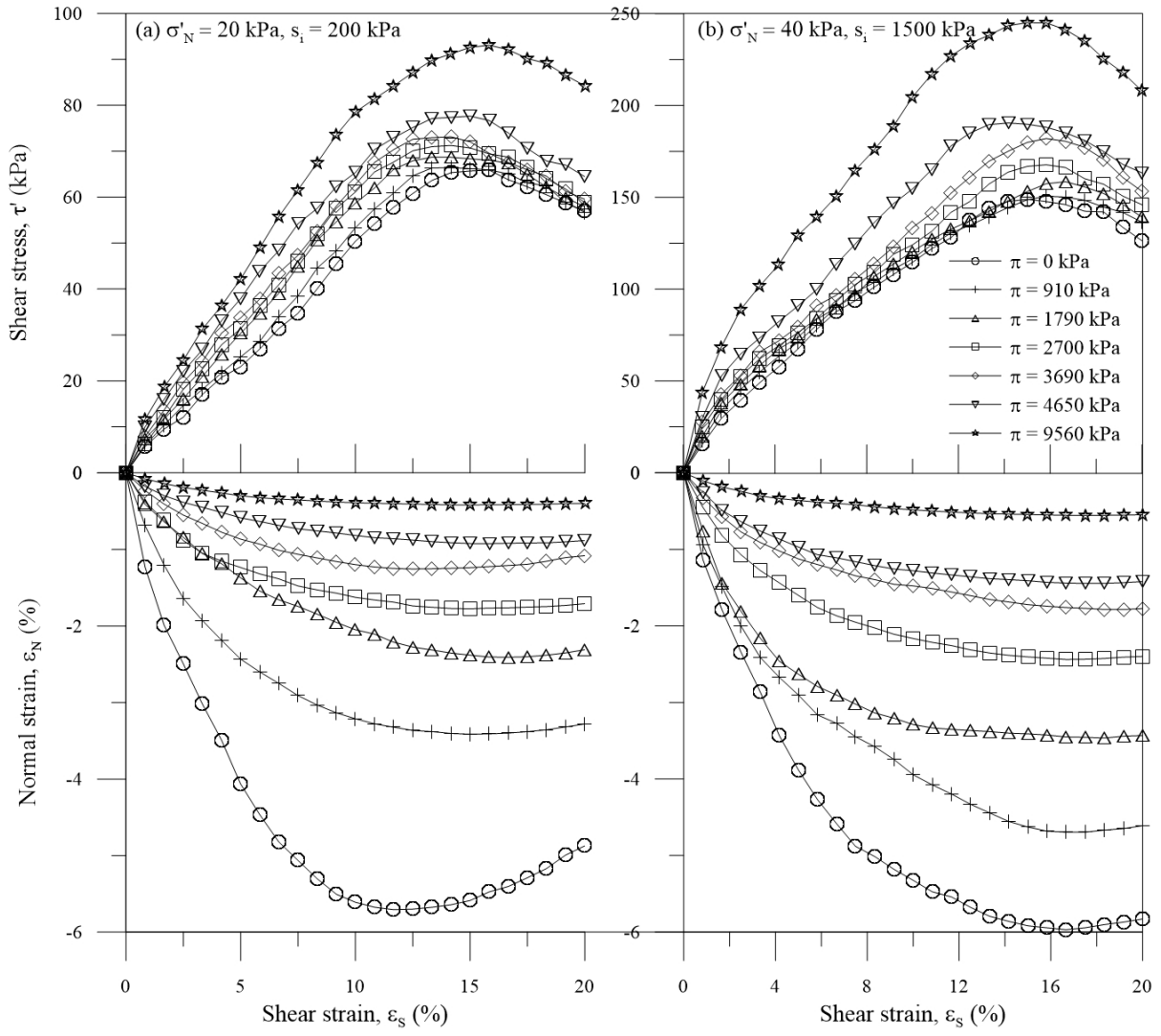
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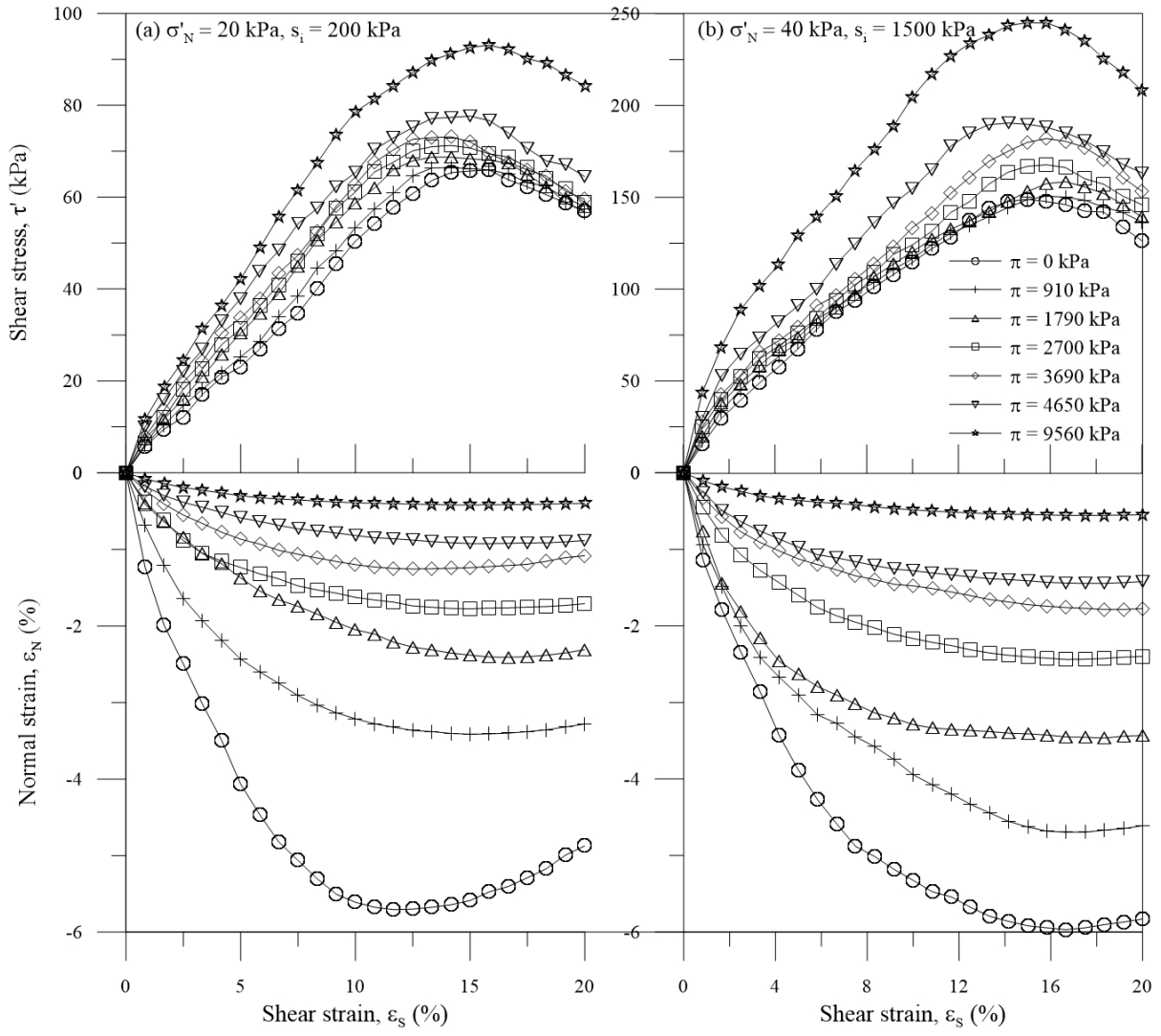
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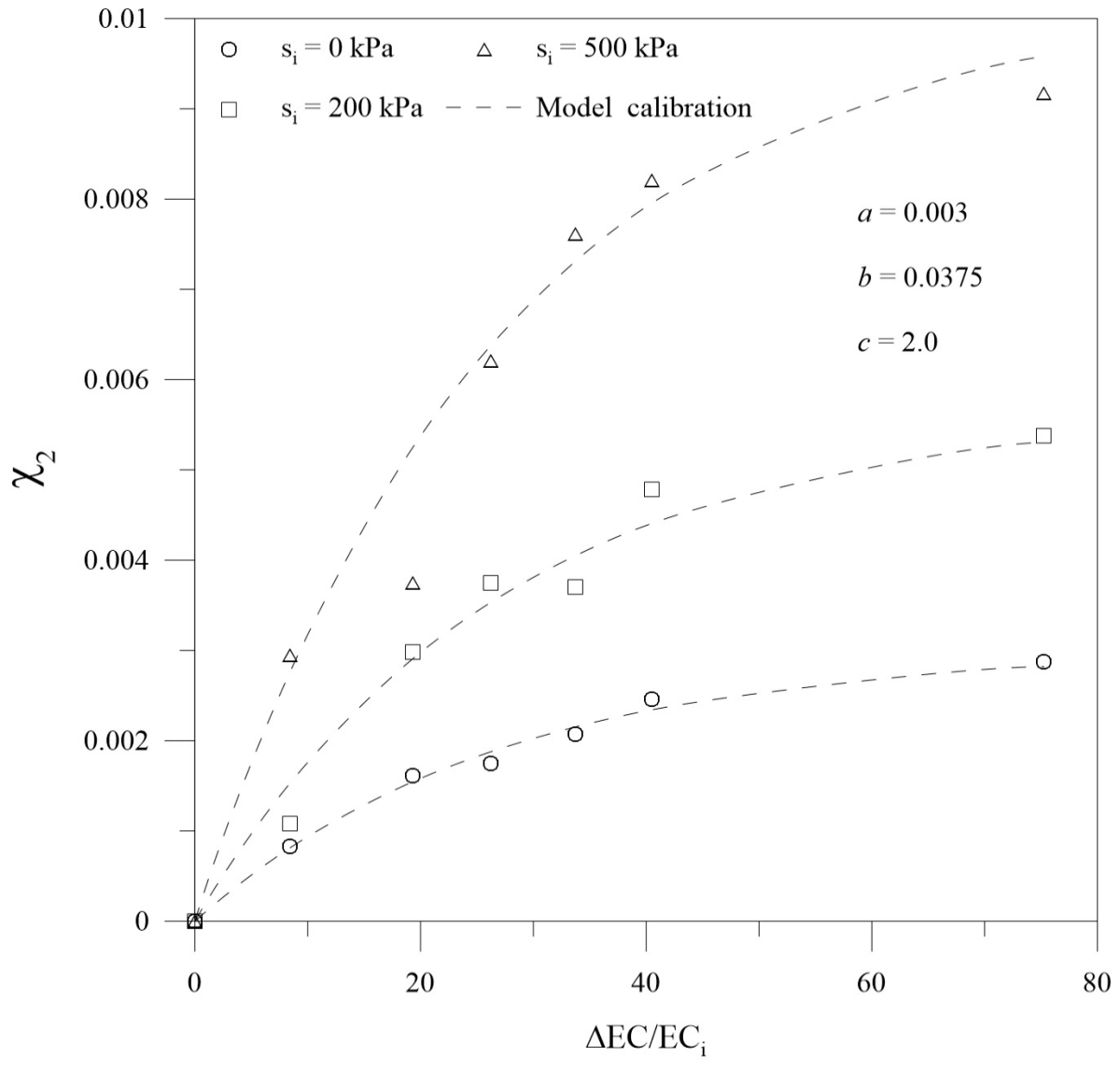
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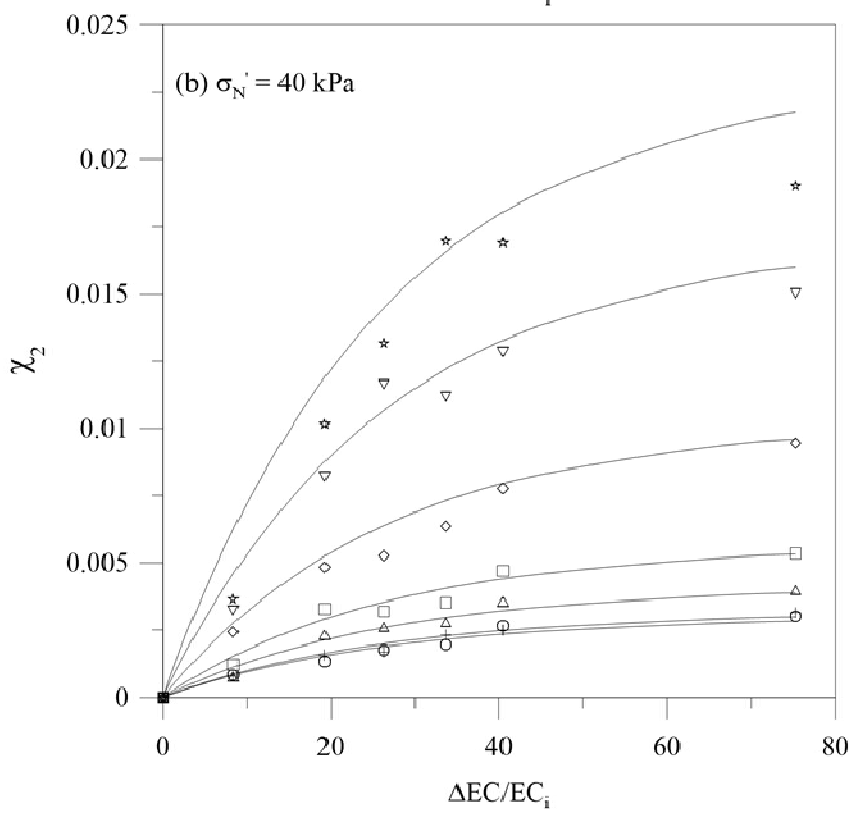
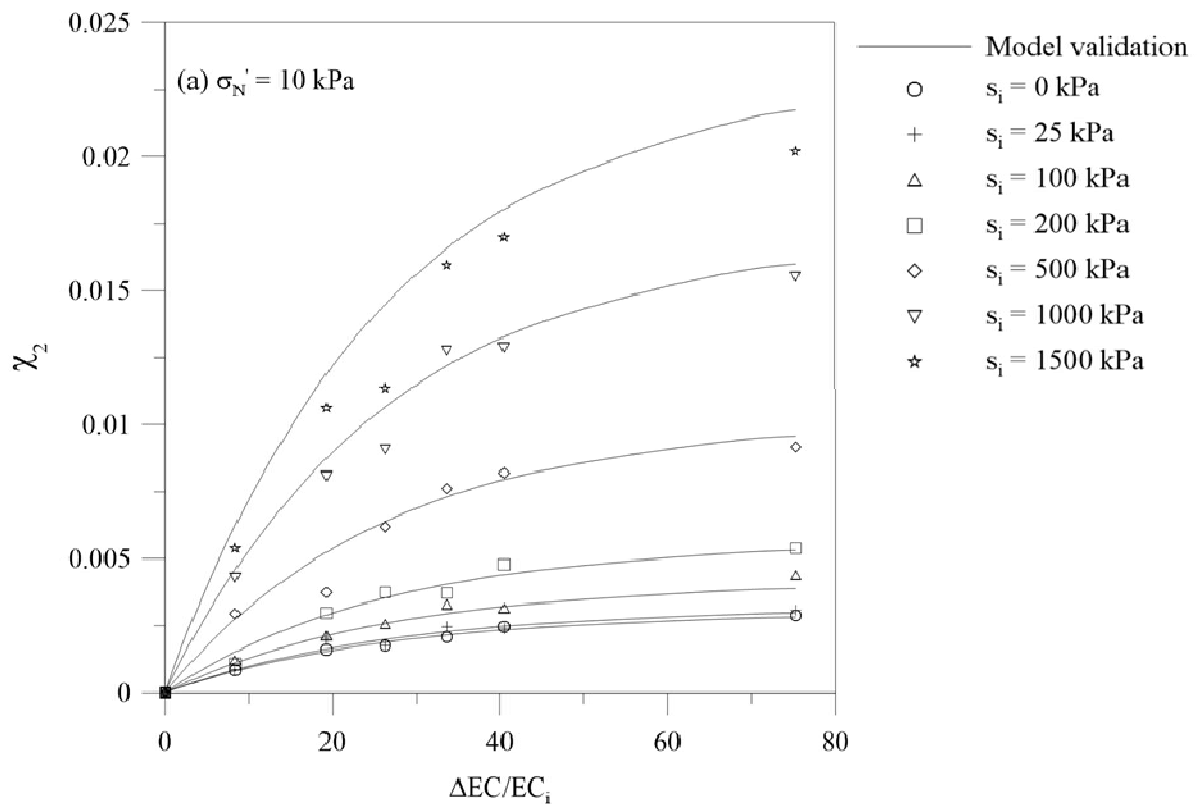
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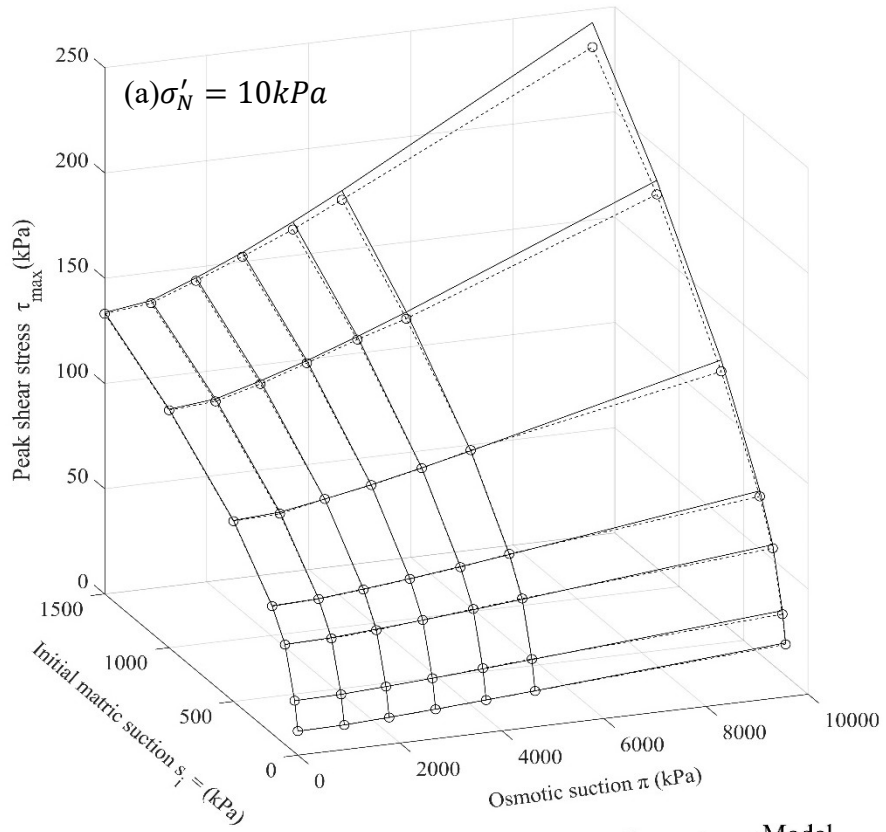
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