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Assessing relevant climate data for agricultural applications

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32 **Abstract**

33 Climate change is expected to substantially reduce agricultural yields, as reported in the by
34 the Intergovernmental Panel on Climate Change (IPCC). In Sub-Saharan Africa and (to a
35 lesser extent) in South Asia, limited data availability and institutional networking constrain
36 agricultural research and development. Here we performed a review of relevant aspects in
37 relation to coupling agriculture-climate predictions, and a three-step analysis of the
38 importance of climate data for agricultural impact assessment. First, using meta-data from the
39 scientific literature we examined trends in the use of climate and weather data in agricultural
40 research, and we found that despite agricultural researchers' preference for field-scale
41 weather data (50.4% of cases in the assembled literature), large-scale datasets coupled with
42 weather generators can be useful in the agricultural context. Using well-known interpolation
43 techniques, we then assessed the sensitivities of the weather station network to the lack of
44 data and found high sensitivities to data loss only over mountainous areas in Nepal and
45 Ethiopia (random removal of data impacted precipitation estimates by $\pm 1,300$ mm/year and
46 temperature estimates by $\pm 3^\circ\text{C}$). Finally, we numerically compared IPCC Fourth Assessment
47 Report climate models' representation of mean climates and interannual variability with
48 different observational datasets. Climate models were found inadequate for field-scale
49 agricultural studies in West Africa and South Asia, as their ability to represent mean climates
50 and climate variability was limited: more than 50% of the country-model combinations
51 showed $< 50\%$ adjustment for annual mean rainfall (mean climates), and there were large
52 rainfall biases in GCM outputs (1,000 to 2,500 mm/year), although this varied on a GCM
53 basis (climate variability). Temperature biases were also large for certain areas (5-10°C in the
54 Himalayas and Sahel). All this is expected to improve with IPCC's Fifth Assessment Report;
55 hence, appropriate usage of even these new climate models is still required. This improved
56 usage entails bias reduction (weighting of climate models or bias-correcting the climate
57 change signals), the implementation of methods to match the spatial scales, and the
58 quantification of uncertainties to the maximum extent possible.

59

60 **Keywords:** Sub-Saharan Africa; South Asia; climate modelling; climate model; skill;
61 uncertainty; CMIP3; CMIP5.

62

63 **1. Introduction**

64 Agriculture is expected to play an important role in the context of climate change, not only
65 because it is considered amongst the most vulnerable sectors, but also because it is part of the
66 solution (i.e. potential to mitigate greenhouse gases [GHGs] emissions) (FAO, 2009; IPCC,
67 2007). Agriculture will likely be severely affected over the next hundred years due to
68 unprecedented rates of changes in the climate system (IPCC, 2007; Jarvis et al., 2010; Lobell
69 et al., 2008; Thornton et al., 2011). Some of these impacts have already been observed
70 (Battisti and Naylor, 2009; Schlenker and Lobell, 2010). To help cope with such impacts, a
71 framework to assess the effects of climate change on agriculture and food security and to aid
72 with adaptation was established in 2008, as described by Jarvis et al. (2011): The
73 Consultative Group of International Agricultural Research (CGIAR) Research Program on
74 Climate Change, Agriculture and Food Security (CCAFS).

75

76 For adaptation to be successful, agricultural and climate data are crucial, and these are scarce
77 in their basic forms (data from research and weather stations, respectively) or not very well
78 managed and/or maintained in certain parts of the world. Most importantly, climate databases
79 and their derived products are sometimes inaccurate, or else lack the documentation
80 necessary to facilitate their use within the agricultural research community. In some
81 instances, this may be indicative of the gap between the agricultural and climate research
82 communities (Pielke et al., 2007; Thornton et al., 2011). Even when the two do collaborate,
83 agricultural researchers face critical constraints when accessing basic sources of climate data
84 (i.e. weather stations) due to a number of factors, from access to data, to weather maintenance
85 and data quality checks, to the weather itself (DeGaetano, 2006).

86

87 In the last 10 years, various datasets have been developed by different institutions, usually
88 based on either a combination of weather station data, satellite data, and numerical weather
89 prediction models in addition to interpolation methods, or on the sole application of climate
90 models. The usage of these datasets for agricultural modelling purposes is rather limited for
91 one or more of the following reasons: (1) their time step is long (monthly in the best case);
92 (2) their temporal coverage is limited to an average of several years (Hijmans et al., 2005;
93 New et al., 2002); (3) their spatial resolution is too coarse (Adler et al., 2003; Schneider et al.,
94 2010); (4) their geographic coverage is not wide enough (Di Luzio et al., 2008); and (5) only
95 certain variables (i.e. temperatures, rainfall) are reported whereas other agriculturally relevant
96 measures (e.g. potential and/or reference evapotranspiration, relative humidity, solar
97 radiation) are rarely reported (Di Luzio et al., 2008; Hijmans et al., 2005). Moreover,
98 assessments of these data (particularly climate models) have been done only under a climate-
99 science perspective (Gleckler et al., 2008; Pierce et al., 2009), for a limited number of
100 variables (Jun et al., 2008; Reifen and Toumi, 2009), or for a reduced realm (Walsh et al.,
101 2008).

102

103 In this paper, we sought to improve the general knowledge on the available climate data for
104 agricultural research using a three-step thorough analysis on fundamental aspects related to
105 agricultural modelling. First, we perform a meta-analysis on the usage of various data sources
106 for agricultural applications; second, we assess the quality and distribution of weather station

107 records by exploring both the ability of these data to fill geographic information gaps by
108 means of interpolation, and the sensitivities of the different regions to data loss; and finally,
109 we assess the accuracy of climate model outputs against different observational datasets using
110 various metrics reported in previous literature (Gleckler et al., 2008; Pierce et al., 2009). We
111 finally analyse the main implications of our findings on agricultural impact assessment.

112

113 **2. Review of knowledge and data**

114 **2.1. Understanding of processes and crop modelling**

115 Mechanisms to fix carbon in plants (i.e. photosynthesis) are affected by a number of factors
116 (El-Sharkawy, 2005; Prasad et al., 2002), although responses strongly depend on the type of
117 mechanism used by the plant to produce biomass (i.e. C₄, C₃, CAM) and on any other stresses
118 to which the plant could be subjected simultaneously. In crop production, apart from
119 appropriate plant growth it is the amount of biomass accumulated in fruits and seeds and the
120 nutrients in them that matters most (Thuzar et al., 2010). Yields are a direct consequence of
121 photosynthesis and biomass accumulation, and these are directly or indirectly affected by
122 environmental conditions [see (Challinor et al., 2009b) for a review]. Well-watered crops
123 grown under optimal temperature and solar radiation ranges develop to their full production
124 potential (van Ittersum et al., 2003), but growth potential reduces if the crop is stressed during
125 the growing season (Hew et al., 1969; Huntingford et al., 2005).

126

127 Therefore, modelling crop growth depends on (1) correct formulation of the simulation
128 model, (2) our ability to understand the effects of environmental factors on growth, and (3)
129 correct measurement of the relevant environmental factors for correct mapping of their
130 interactions (Boote et al., 1996; El-Sharkawy, 2005). Hence, crop modelling largely benefits
131 from accurate measurements of temperatures, rainfall, and solar radiation, as the main factors
132 acting on photosynthesis (Challinor and Wheeler, 2008; Hoogenboom et al., 1994), but even
133 these basic data are often unavailable, messy, or of limited quality. The more available data
134 there exists, the better calibration and evaluation of crop models can be (Adam et al., 2011;
135 Niu et al., 2009; Xiong et al., 2008).

136

137 Additionally, most crop models simulate growth of individual plants and then scale out the
138 modelling results to the plot-scale, based on management decisions such as plant and row
139 distances, and plot size (Aggarwal et al., 2006; Boote et al., 1996; Hoogenboom et al., 1994).
140 On the other hand, available weather data (when not measured in the field) is only available
141 at coarse spatial scales. Matching these two spatial scales is not an easy task [see (Challinor
142 et al., 2009a; Jagtap and Jones, 2002; Trnka et al., 2004) for a review]. The challenge is thus
143 to increase the knowledge of the interactions between atmospheric and crop-growth processes
144 (Boote et al., 1996) whilst avoiding model over-parameterisation (Challinor et al., 2009b),
145 improving the accuracy of inputs (Adam et al., 2011), and matching both spatial scales
146 (Challinor et al., 2009a). All this requires closing the gap between crop and climate scientists.

147

148 **2.2. Weather data**

149 Measurements of weather for a given site are often unavailable because (1) there is no
150 weather station; (2) weather stations are not well maintained so data are either only available

151 for a short period or contain gaps, (3) collected data are not properly stored; (4) data do not
152 pass basic quality checks; and/or (5) access to data is restricted by holding institutions (Figure
153 1). This all further constrains agricultural impact assessment, highlighting the importance of
154 making data public.

155

156

157

<Insert Figure 1 here>

158 Apart from the constraints related to access and weather station locations, probably the most
159 important issue regarding weather data is quality (Begert et al., 2008; DeGaetano, 2006)
160 (Figure 1), which also greatly affects the performance of impact models. Therefore, the
161 climate and agricultural community has partly focused on developing methods for either
162 temporal or spatial data gap filling, and on using such methods for developing global or
163 regional datasets with public access (Hijmans et al., 2005; Jones and Thornton, 1999; Soltani
164 et al., 2004).

165

166 However, uncertainties in global datasets derived from interpolation methods have been only
167 barely (if at all) estimated (Buytaert et al., 2009; Challinor and Wheeler, 2008; Soria-Auza
168 et al., 2010). Researchers using global datasets and any weather station source need to be aware
169 of these problems and ought to take this into account by testing the sensitivities of their
170 approaches to accuracy issues (i.e. inhomogeneities, discontinuities) and (if possible)
171 providing results within the range of uncertainty in input data (i.e. such as the outputs of cross
172 validated interpolation methods) (Challinor et al., 2005).

173

174

2.3. Climate model data

175 General Circulation Models (GCMs) are currently the best way to model the complex
176 processes that occur at the earth system's level (Huntingford et al., 2005; IPCC, 2007).
177 However, as CGMs are highly complex, they are computationally expensive, so they have
178 only been used for predictions at coarse spatial scales. These predictions therefore involve a
179 number of uncertainties relevant to agriculture [see (Challinor et al., 2009b; Jarvis et al.,
180 2010; Quiggin, 2008) for reviews on the topic].

181

182 In short, uncertainty in climate modelling arises from the impossibility of modelling the
183 climate system with complete determinism (Walker et al., 2003). This uncertainty can arise
184 from: context (boundaries of the system modelled), model, inputs, and parameters (Walker et
185 al., 2003). Model uncertainty can be structural or technical: structural uncertainty in models is
186 associated with our lack of understanding of the system, whereas technical uncertainty relates
187 to our inability to implement mathematical formulations in computational systems. Other
188 uncertainties in climate modelling arise from variable driving forces (greenhouse gas
189 emissions and concentrations), initial conditions and parameterised physics (Challinor et al.,
190 2009b; Walker et al., 2003). Rationalisation and quantification of all these uncertainties under
191 the context of agriculture is possible (see Challinor et al., 2009b for a review).

192

193 Crop modellers are thus challenged to understand the broad concepts of climate modelling
194 uncertainties and detect the sensitivities of crop models to them, whilst also having a basic

195 understanding of earth processes in order to identify major flaws in climate models and
196 decide the best ways to couple them with crop models.

197

198 **3. Materials and methods**

199 Throughout this paper, we built upon existing knowledge of agricultural and climate
200 modelling (Sect. 2) and:

- 201 1. Performed a meta-analysis on the usage of climate and weather data for agricultural
202 modelling purposes and summarised the desirable characteristics sought when
203 modelling crop production.
- 204 2. Analysed the robustness of the existing weather station network by assessing both the
205 ability of these data to correctly fill information gaps via interpolation methods, and
206 the network's sensitivities to information loss.
- 207 3. Assessed the accuracy of climate model outputs from the Fourth Assessment Report
208 of the IPCC (IPCC, 2007) against different observational datasets, using metrics and
209 methods reported in the climate-science literature that are also familiar to agricultural
210 researchers.

211

212 All calculations were done by means of the software packages R-2.13.1 (available at
213 <http://www.r-project.org>) and GRASS-GIS 6.4.0 (available at <http://grass.fbk.edu>) in a 64-bit
214 Red Hat Enterprise Linux 5 box.

215

216 **3.1. Study area**

217 We focused on the geographic area of Africa and South Asia, where several studies have
218 identified that significant vulnerabilities exist (Aggarwal, 2008; Aggarwal et al., 2004;
219 Barrios et al., 2008; Byjesh et al., 2010; Challinor et al., 2007a; Chipanshi et al., 2003; Jones
220 and Thornton, 2003; Lane and Jarvis, 2007; Liu et al., 2008; Lobell et al., 2008; Thornton et
221 al., 2009; Thornton et al., 2011; Washington et al., 2006). In particular, we concentrate our
222 efforts on West Africa (Senegal, Mali, Burkina Faso, Ghana and Niger), East Africa
223 (Ethiopia, Tanzania, Uganda and Kenya) and the Indo-Gangetic Plains countries (India,
224 Nepal, and Bangladesh), hereafter referred to as WAF, EAF and IGP, respectively (Figure 2).

225

226

<Insert Figure 2>

227

228 **3.2. Analysing the usage of climate data in agricultural studies**

229 **3.2.1. Meta-data from agricultural studies**

230 We gathered data from a number of publications on any topic that made use of climate data
231 for any sort of agricultural modelling. We conducted searches using various search engines
232 and downloaded only peer-reviewed publications. Review papers and the Fourth Assessment
233 report of the IPCC were particularly useful in identifying additional published studies. We
234 analysed all publications that in any way involved the usage of climate data for agricultural
235 modelling purposes. As the selection of the impact assessment model is the first decision that
236 any researcher needs to make, we focus on the driving factors of this decision. We recorded
237 different variables from the studies as follows:

- 238 (1) Problem and/or topic in question: classified in categories such as impact assessment,
239 seasonal yield forecasting, sole crop modelling, and climate attribution, among others.
240 Each study was classified into only one category by taking into account only the main
241 issue addressed by the paper;
- 242 (2) Scale of the approach: includes site, sub-national, country, regional (group of
243 countries), and global;
- 244 (3) Use of weather generators: for both present and future, we recorded whether the study
245 did or did not use a weather generator;
- 246 (4) Climate dataset (current): GCM when a GCM (regardless of which one) was used,
247 RCM when an RCM (regardless of which one) was used, weather station, satellite (no
248 further discrimination), and important datasets (i.e. CRU, WorldClim, GPCP, among
249 others);
- 250 (5) Climate dataset (future): the nature of used future projections was recorded here
251 including the downscaling method, if applicable. Classifications were: GCM “as is”
252 when studies used raw GCM outputs as inputs, pattern scaled GCMs (Mitchell et al.,
253 2004), RCMs, systematic changes to current climate data, statistical downscaling
254 (Wilby et al., 2009), and weather generator downscaled GCM (Jones et al., 2009).
255

256 For further details on the above categories the reader is referred to our supplementary
257 material (part 1). We revised a total of 205 peer-reviewed publications (See supplementary
258 material part 2), printed between the years 1983 and 2011. Most of the studies were published
259 immediately before or after the IPCC 4AR was released in 2007. When a certain study made
260 use of two different sources of present-day climate data, it was considered twice (totalling
261 247 cases).
262

263 **3.2.2. Analysing the usage of climate data in agricultural studies**

264 We analysed the recent trends in the use of climate data for agriculture: the obvious
265 constraints in the studies, the type of approaches used and the climate data inputs used to
266 drive the chosen agricultural models. By doing this, we ensured that we covered all the main
267 factors driving an agricultural researcher’s decision to select a particular approach for a given
268 problem.
269

270 **3.3. Analysis of weather station data**

271 **3.3.1. Worldwide weather station network data**

272 Long term climatological means of monthly precipitation and mean, maximum and minimum
273 temperatures were assembled, as described by Hijmans et al. (2005). However, it is important
274 to note that at the global level the sources of these data are large in number and differ in
275 coverage, availability and quality (Table 1), and thorough quality checks were done only in a
276 sub-set of the sources by original distributing institutions.
277

278 **<Insert Table 1 here>**
279

280 Additional sources such as R-Hydronet (<http://www.r-hydronet.sr.unh.edu/english/>) and
281 Oldeman (1988) database for Madagascar were also included. We discarded any weather

282 station with less than 10 years of data. The final dataset (after quality control and duplicates
283 removal, see Hijmans et al. 2005 for more details) comprised 13,141 locations with monthly
284 precipitation data, 3,744 locations with monthly mean temperature, and 2,684 locations with
285 diurnal temperature range within our study region. This dataset is hereafter referred to as
286 WCL-WS.

287

288 **3.3.2. Analysing robustness of existing weather station networks**

289 Many methods exist that allow the user to determine (interpolate) the value of a parameter
290 (e.g., monthly rainfall) in a given condition (i.e. in a given site, at a given time, or both),
291 where it had never been measured before. Some of these methods are already popular with
292 researchers using climate data (Hijmans et al., 2005; Hutchinson, 1995; Jones and Thornton,
293 1999; New et al., 2002) either on a regional or on a global basis. For climate-variable
294 interpolations, the robustness of weather records is critical for an accurate result.

295

296 We assessed the robustness of the weather station network by testing both the ability of
297 weather records to yield accurate interpolation results, and the sensitivities of the network to
298 information loss. Towards those ends, we used the WCL-WS dataset to fit a thin plate spline
299 interpolation algorithm (Hutchinson, 1995) for our study region. We investigated the effect of
300 weather station availability by using 100 cross validated folds for four variables (monthly
301 maximum, minimum and mean temperatures and total precipitation) using similar methods as
302 in Hijmans et al. (2005) and New et al. (2002) for each fold. We used longitude, latitude and
303 elevation as independent variables. We used 85% randomly selected data points for fitting the
304 splines and the remaining 15% for evaluating the result for each variable and month. For the
305 evaluation, we calculated the R^2 and the Root Mean Square Error (RMSE) and produced
306 boxplots of the 100-fold-by-12-month interpolations for each of the four variables. As the
307 number of stations considerably exceeded the amount of available memory for processing,
308 we divided the whole region of study in 5 tiles, each with an equivalent number of locations.
309 We then projected the fitted splines onto 30-arc-second gridded datasets of latitude, longitude
310 and altitude (Jarvis et al., 2008), thus producing a total of 4,800 interpolated surfaces (12
311 months times 4 variables times 100 folds). Finally, we analysed the spatial variability of
312 standard deviations and the performance of the interpolation technique as proxies for
313 sufficient distribution and geographic density of weather stations.

314

315 **3.4. Assessment of IPCC Fourth Assessment Report (4AR) model data**

316 **3.4.1. Long-term observed mean climatology from weather stations**

317 Three different long term climatology datasets were assembled: (1) the Global Historical
318 Climatology Network (GHCN, as in Sect. 3.3.1) version 2 (Peterson and Vose, 1997),
319 available at <http://www.ncdc.noaa.gov/pub/data/ghcn/v2>. We used GHCN as an independent
320 source because it is a global resource that contributed significantly to WCL-WS and also
321 because it is available at more temporally disaggregated levels (i.e. monthly), thus allowing
322 uniformity with analyses on Sect. 3.4.3 and 3.4.6. This database includes monthly historical
323 totals (1900-2010) of precipitation (20,590 stations), and means of maximum, minimum
324 (4,966) and mean (7,280) temperatures. GHCN data have been subject to quality checks and
325 to a process of “homogenisation” or “adjustment” (Peterson and Easterling, 1994); however,

326 the available data within our analysis domain consisted primarily of “unadjusted” stations.
327 For each location (6,393 stations for rainfall, 1,278 for mean temperature and 549 for
328 minimum and maximum temperature) within our study area, we averaged historical monthly
329 time series for the period 1961-1990 for maximum, minimum and mean temperatures and
330 total rainfall, resulting in a time-averaged dataset of 6,393 locations for rainfall, 1,278 for
331 mean temperature and 549 for minimum and maximum temperature. This dataset will be
332 hereafter referred to as GHCN-CL.

333

334 (2) WCL-WS (Sect. 3.3.1); and (3) the Global Surface Summary of the Day (GSOD) was
335 accessed at <http://www.ncdc.noaa.gov/cgi-bin/res40.pl>. This database contains daily data
336 from ~9,000 weather stations worldwide for 18 variables, including, mean, maximum,
337 minimum and dew point temperature, sea level and location pressure, visibility, wind speed
338 and gust, precipitation, snow depth, and specifications on the occurrence of rain, snow, fog,
339 tornado, thunder, or hail (NOAA, 2011; <ftp://ftp.ncdc.noaa.gov/pub/data/g sod/readme.txt>).
340 We selected weather stations within our study area (1,999); aggregated daily rainfall, mean,
341 maximum and minimum temperatures to a monthly time scale; and then averaged over the
342 period 1961-1990. This dataset will be hereafter referred to as GSOD-CL.

343

344 **3.4.2. Long-term observed mean climatology from interpolated surfaces**

345 We gathered high-resolution climatology from two different sources: (1) the high resolution
346 climate surfaces in WorldClim (Hijmans et al., 2005), available at <http://www.worldclim.org>.
347 WorldClim is a 30 arc-seconds (~1km at the equator) global dataset produced from the
348 interpolation of long-term climatology as measured in weather stations. Global gridded data
349 were downloaded at the 30 arc-second resolution, then masked to our analysis domain, and
350 aggregated to 10 arc-minute using bilinear interpolation in order to reduce computational and
351 storage time; and (2) the University of East Anglia Climatic Research Unit (CRU) dataset
352 (New et al., 2002), available through <http://www.cru.uea.ac.uk/cru/data/hrg/> (CRU-CL-2.0).
353 This dataset was developed using the same interpolation method as WorldClim, with the
354 main difference that WorldClim includes many more weather stations, sometimes at the
355 expense of input data quality. CRU-CL-2.0 resolution is 10 arc-minute (~20km at the
356 equator). Data were downloaded at the global level and masked to our analysis domain.
357 WorldClim and CRU-CL-2.0 are hereafter referred to as WCL-IS and CRU-IS (interpolated
358 surfaces), respectively. We used these sources because (1) they are flag products that most
359 researchers use for impact studies; (2) they are much higher resolution than GCMs (and other
360 products such as the Global Precipitation Climatology Project [GPCP] and the Global
361 Precipitation Climatology Centre [GPCC]) and hence permit the capture of small-scale
362 weather patterns (important to agriculture) as well as a direct comparison of their within-
363 GCM-gridcell mean with the actual GCM value; (3) are based only on ground observations of
364 weather and do not incorporate side-products such as reanalysis (Uppala et al., 2005) or
365 satellite data (Huffman et al., 2007), both of whose accuracy is not as good.

366

367 **3.4.3. Long-term observed time series**

368 Two sources of weather time series were used: (1) long term (1961-1990) series of monthly
369 weather conditions were gathered from GHCN version 2 (Peterson and Vose, 1997). Again,

370 we used mainly unadjusted stations. Mean monthly temperature and total monthly historical
371 rainfall data were used without any further processing; and (2) long-term (1961-1990) series
372 of daily weather as in GSOD (NCDC, 2011). For GSOD, daily precipitation and monthly
373 temperature were aggregated to the monthly level only if all days were reported with data (for
374 rainfall) and if at least 50% of the days had data (for temperatures). This resulted in 1,999
375 stations within our analysis domain, although not all stations had data for all months and all
376 years. These two data sources are hereafter referred to as GHCN-TS and GSOD-TS,
377 respectively. Lack of data prevented us from including maximum and minimum temperatures
378 in the GHCN-TS and the GSOD-TS datasets. In contrast to GHCN-CL and GSOD-CL,
379 GHCN-TS and GSOD-TS include every month and every year, thus allowing the analysis of
380 inter-annual variability.

381

382 **3.4.4. Global climate model output**

383 The latest IPCC report (Fourth Assessment Report, 4AR) comprises the sole state-of-the-art
384 public and official source of climate data for use in impact studies (IPCC, 2007; Jarvis et al.,
385 2010). We therefore decided to use IPCC 4AR results.

386

387 We downloaded present day (1961-1990) simulations of global climate at original GCM
388 resolution (~100 km) from the CMIP3 (Coupled Model Intercomparison Project phase 3) web
389 data portal at <https://esg.llnl.gov:8443/index.jsp> (PCMDI, 2007). We downloaded monthly
390 time series of mean, maximum, minimum temperature and precipitation flux in NetCDF
391 format for 24 coupled GCMs (Table 2). Separately for each GCM, we calculated diurnal
392 temperature range for each month and year as the difference between maximum and
393 minimum temperatures and calculated total monthly rainfall as the product between the
394 precipitation rate, the water density at sea level pressure and the number of seconds in a
395 month. We used the each climate model monthly time series (GCM-TS hereafter) and also
396 calculated average 1961-1990 climatology by averaging, for each variable (mean
397 temperature, diurnal temperature range and total rainfall), every month for the whole 1961-
398 1990 period (GCM-CL hereafter). The final datasets (i.e. GCM-TS and GCM-CL,
399 respectively) consisted of three variables (mean temperature, diurnal temperature range and
400 total monthly rainfall) for 24 different GCMs.

401

402

<Insert Table 2 here>

403

404 **3.4.5. Ability to represent long-term climatology**

405 The extent to which GCM predictions are accurate has not been fully explored for some parts
406 of the world, particularly in the context of agriculture (Gleckler et al., 2008; Pierce et al.,
407 2009; Walsh et al., 2008). As previously stated (Sect. 2.1), we compared the most readily
408 available variables from both ground observations and climate models: rainfall, mean
409 temperature and diurnal temperature range. Data for other variables are not available for our
410 study regions in observational datasets. As per our stated objective (Sect. 3), we performed
411 two sets of comparisons:

412

413

- First, we compared the GCM-CL dataset with the interpolated climatology in CRU-IS,
WCL-IS (Sect. 3.4.2). We performed comparisons on a country basis in order to yield

414 country-specific results. For each GCM gridcell, the mean, maximum and minimum
415 values of all lower scale (CRU-IS, WCL-IS) cells was first calculated and then compared
416 to the GCM value through the determination coefficient (R^2) and corresponding p-value,
417 the slope of a origin-forced (so that a 1:1 relationship was sought) regression curve (S)
418 and the root mean square error (RMSE).

- 419 • Second, using the same procedure, we compared the GCM-CL dataset with observed
420 climatology in WCL-WS (Sect. 3.3.1), GHCN-CL and GSOD-CL (Sect. 3.4.1).

421

422 We analysed total rainfall, mean temperatures and diurnal temperature ranges over three
423 periods: December-January-February (DJF), June-July-August (JJA) and the whole year
424 (ANN). These months represent the most critical seasons for agriculture in our study regions,
425 and are also the most often assessed in the existing literature (Gleckler et al., 2008; Pierce et
426 al., 2009). Due to space constraints, we present only the results of comparisons between
427 GCM gridcell values and mean values within gridcells, unless otherwise stated. We do,
428 however, discuss other relevant results in more general terms.

429

430 **3.4.6. Ability to represent long-term monthly climate time series**

431 CMIP3-related GCMs are known to misrepresent certain inter-annual and/or within-decade
432 variations that are important for agricultural systems (Govindan et al., 2002). However,
433 specific aspects of these errors have not been explored in all CMIP3 models in the context of
434 agriculture. Therefore, in order to test the consistency of GCM predictions across time, we
435 compared the GCM-TS (Sect. 3.4.4) dataset against the GHCN-TS and GSOD-TS (Sect.
436 3.4.3). The comparison was done for three periods (JJA, DJF and ANN, Sect. 3.4.4) by
437 calculating the R^2 and corresponding p-value, the slope of the regression curve as forced to
438 the origin and the RMSE between the two time series (GCM-TS vs. GHCN-TS and GCM-TS
439 vs. GSOD-TS). As a GCM cell contains one or more weather stations, we averaged the
440 monthly time series as needed before comparing the two pairs of series. Finally, we compared
441 the performance of all GCMs across the geographic space of our study area.

442

443 **4. Results**

444 **4.1. Usage of climate data in agricultural studies**

445 **4.1.1. Topics of study**

446 The most addressed topic (41.4% of the studies) in our literature review was climate change
447 impact assessment (Figure 3), followed by crop growth simulation (18.5%). Water resources-
448 impact studies round out the top three topics studied (8.1%), followed by climate attribution
449 (6.9%), crop yield forecasting (6.1%), and model assessment (5.7%). Surprisingly, formal
450 studies addressing adaptation were rather scarce (3.6%). Pests and diseases, soils, abiotic
451 stresses and climate risks appeared to be a lot less addressed than impact assessment and crop
452 growth simulation studies, which together accounted for more than 50% of the total
453 publications.

454

455

<Insert Figure 3 here>

456

457 **4.1.2. Scale of studies and type of models**

458 Most of the studies performed their models at a scale less than the size of a country; site-
459 specific or sub-national level together comprised 55% of the studies. Very few (7%) of the
460 studies were performed at the global level, likely because of the type of models used: field-
461 scale mechanistic crop growth models were the most utilised overall (69.2%); followed by
462 statistical and/or empirical approaches (S/E, 21.4%), which most of the crop growth
463 modellers criticise for not being accurate enough (Lobell and Burke, 2010; Lobell et al.,
464 2008); and finally by hydrological models (10%). The frequent use of field-based crop
465 growth models suggests that the time step requirement for input data is rather high (El-
466 Sharkawy, 2005), also confirmed by the usage of weather generators (8.5 and 11.2% for
467 present and future climates, respectively).

468

469 **4.1.3. Climate data sources**

470 Unlike the model types, which were quite similar, the sources of present climate data varied
471 substantially, with a total of 32 different sources being used for present climate data (Figure
472 4A). On average, a different present-day-climate dataset was used for every 7 agricultural
473 studies. The most commonly used data source was local (non-public) weather stations (50.4%
474 of the cases), followed by University of East Anglia Climatic Research Unit (CRU) datasets
475 with 13.7% (10.9% for CRU-TS [monthly time series], and 2.8% for CRU-CL [monthly
476 climatology]). Climate model outputs were used in 14.5% of the cases: within this group,
477 10.5% used GCM data, 4% RCM [Regional Climate Model] data, 3.6% satellite imagery, and
478 2.8% WorldClim, followed by other less relevant sources. The Global Precipitation
479 Climatology Project (GPCP) (Adler et al., 2003; Huffman et al., 2009), the Global
480 Precipitation Climatology Centre (GPCC) (Schneider et al., 2010) and the Global Historical
481 Climatology Network (GHCN, (Peterson and Vose, 1997)) were rarely reported overall
482 (0.4% each).

483

484

<Insert Figure 4 here>

485

486 The future climate data used was found to be less variable overall, with only 7 different types
487 of data employed in the 125 cases citing some type of future climate data (Figure 4B). Out of
488 these 125, only one study did not clearly state which type of climate data was used. The vast
489 majority of cases (42.9%) used GCM data “as is” (AI GCM), meaning that predictions on
490 agricultural yields were based on predicted changes at coarse resolution (~100 km). All other
491 studies involved some type of downscaling, except those that employed the systematic
492 changes approach (SC variables), which can be assumed to be sensitivity analyses rather than
493 impact studies. RCMs (Regional Climate Models) were the most common way of
494 downscaling GCMs, cited in 19% of the studies, followed by statistical downscaling with
495 17.5% (SD GCM, (Tabor and Williams, 2010)), and pattern scaling with 8.7% (PS GCM,
496 (Mitchell et al., 2004)) (Figure 4B).

497

498 Uncertainty, as measured by the number of different future scenarios used (combinations of
499 emissions scenarios and climate models) was explored in only 36.5% of the studies.
500 Additionally, the average number of scenarios per study (rounded to the closest integer) was
501 3, indicating that climate uncertainties are barely (if at all) studied in agricultural science and

502 highlighting a knowledge gap in agricultural research, an issue previously raised and
503 discussed by other authors (Challinor et al., 2009b; Challinor and Wheeler, 2008), although
504 some studies addressing this aspect are underway (C. Rosenzweig, personal communication).

505

506 **4.2. Robustness of existing weather station networks**

507 The sensitivities of the network to information loss were found overall to be low.
508 Nevertheless, certain areas, variables and months were found highly sensitive. Agricultural
509 lands (Ramankutty et al., 2008), as visually inspected, are in general less sensitive to data loss
510 than non-agricultural lands. Interpolations' performance varied depending upon the variable,
511 month and parameter used to evaluate them (i.e. R^2 , RMSE, and S), but were consistent,
512 statistically significant ($p < 0.0001$) and with variability (of R^2 , RMSE, and S) between 10–
513 15% in the worst cases. Rainfall presented the lowest R^2 values (Figure 5), particularly in the
514 months of April to August, during which there was a higher variability in the R^2 value and the
515 values reached the absolute minima (0.8). Although it is possible that a high number of
516 weather stations per unit area can improve accuracy, it does not seem to happen in all
517 variables, areas and/or months.

518

519

<Insert Figure 5 here>

520

521 The DJF period presented significantly lower variability and more predictive power, probably
522 due to overall low climate variability (Cooper et al., 2008). Interestingly, maximum and
523 minimum temperatures showed different interpolation accuracies, even though they were
524 measured in the same places. Maximum RMSE for temperatures was up to 1.7°C, whilst for
525 precipitation it was up to 100 mm/year, as seen in the evaluation data. The effect of
526 geography and the difficulty of fitting unique and complex landscape features cause errors,
527 leading to high standard deviations in some areas (Figure 6). In the highlands of Eastern
528 Africa, particularly in the states of Benshangul-Gumuz, Addis Ababa and Southern Nations in
529 Ethiopia, the central areas of the Eastern and Coast States in Kenya, and the very centre of
530 Tanzania (i.e. regions of Morogoro, Dodoma and Manyara) between-fold variability was
531 found to be high (above 150 mm/year).

532

533

<Insert Figure 6 here>

534

535 Over IGP, the largest variability was found in the coastal areas of Maharashtra, Karnataka
536 and Kerala in India, where rainfall deviation was up to 600 mm/year, and in Nepal (districts
537 of Gorka, Dhawalagiri, and Lumbini), where rainfall variability can go up to 1,000 mm/year,
538 and temperature uncertainties up to 3°C, probably due to the combined effect of a more
539 complex climate in the Himalayas and low weather station density.

540

541 **4.3. Accuracy of climate model outputs**

542 **4.3.1. Ability to represent mean climate**

543 As expected, the climate models' skill varied on a variable, country and region basis, with
544 certain identifiable patterns (Figure 7, 8). The GCMs represent the observed climatology
545 from weather stations (i.e. WCL-WS, GHCN-CL and GSOD-CL) more poorly than they do

546 interpolated climatology (i.e. WCL-IS, CRU-IS), mainly because GCMs do not account for
547 local-scale variability (Boo et al., 2011). In a broad sense, we found that the more complex
548 the topography, the lower the skill of the GCMs (Gallée et al., 2004; Joubert et al., 1999). We
549 also observed that GCM skill decreased according to the complexity of the variable, with the
550 maximum skill displayed for mean temperatures, followed by temperature range and finally
551 by precipitation. These results agree with those of other studies (Gleckler et al., 2008;
552 Masson and Knutti, 2011; Pierce et al., 2009).

553

554 Annual precipitation fit in IGP and WAF was observed to dip as low as 0 in some cases, with
555 a considerable number of cases (23% for WCL-WS, 27% for GHCN-CL and 63% for GSOD-
556 CL) presenting very low adjustment ($R^2 < 0.5$) (Figure 7). In Mali, Niger, India and
557 Bangladesh, model skill in representing precipitation, compared to weather station
558 measurements, was consistently low, an issue also reported in other studies (Douglass et al.,
559 2008; Gleckler et al., 2008; Reichler and Kim, 2008). The Bergen Climate Model (BCCR-
560 BCM2.0) and the INM-CM3.0 model showed very poor performance ($R^2 < 0.5$) in more than
561 25% of the countries when compared with WCL-WS, GHCN-CL and GSOD-CL, while the
562 climate model GISS-ModelE (Hansen et al., 2007) presented the poorest performance.

563

564

<Insert Figure 7 here>

565

566 When compared with interpolated climatology (i.e. WCL-IS, CRU-IS), annual precipitation
567 R^2 values varied from 0.383 (GISS-ModelE-R in Uganda) to 0.998 (IAP-FGOALS1.0-G in
568 Burkina Faso), whilst for mean temperatures the R^2 varied from 0.195 (GISS-ModelE-R in
569 Nepal) to 0.999 (MIUB-ECHO-G in Burkina Faso), and for temperature range the values
570 were observed between 0.386 (CCCMA-CGCM3.1-T47 in Senegal) to 0.9998 (MPI-
571 ECHAM5 in Burkina Faso) (Figure 7).

572

573

<Insert Figure 8 here>

574

575 In Ethiopia, mean temperature correlations were lower compared to other countries, despite
576 the relative high density of stations in that area (data not shown). In Senegal, diurnal
577 temperature range was found to be very poorly fitted, particularly for the CCCMA models
578 (Figure 8). This result contrasts with that of other studies, which have marked the CCCMA
579 models as the most skilled (Gleckler et al., 2008; Jun et al., 2008). The ability of GCMs to
580 represent mean climate patterns over a year was neither uniform nor consistent (Table 3),
581 with the lowest performance being observed for precipitation in the DJF period (large number
582 of cases with $R^2 < 0.5$, and few cases with $R^2 > 0.8$). Performance for temperature range showed
583 almost no cases with $R^2 < 0.5$, but fewer cases with $R^2 > 0.8$ than for mean temperatures (Table
584 3).

585

586

<Insert Table 3 here>

587

588

4.3.2. Ability to represent interannual variability

589 R square values were above 0.8 in a large number of gridcells (>50%) for all GCMs for both
590 variables (rainfall, mean temperature) (data not shown); however, there were large rainfall
591 biases in GCM outputs (Figure 9, 10), in some cases between 1,000 and 2,500 mm/year,
592 depending on the GCM. These areas were located in Nepal, northern India and EAF. Most of
593 the models' biases were wet-biases (Figure 10) which were found throughout the whole
594 analysis domain, but they were particularly strong over IGP in the models CCCMA-
595 CGCM3.1-T47, CSIRO-Mk3.0 and -Mk3.5, GFDL-CM2.0, all NASA-GISS models, and
596 both UKMO-HadCM3 and -HadGEM1, whereas the opposite signal was observed over the
597 same area for the models MIROC3.2.-HIRES, NCAR-CCSM3.0, INGV-ECHAM4, CNRM-
598 CM3, and GFDL-CM2.1. Over WAF and EAF, almost all GCMs showed a dry-bias, with
599 underestimations of up to 250 mm/year in some cases. Responses varied for seasonal means
600 and totals, with the wet-season (JJA) being more sensitive to wet biases in most GCMs.

601
602 Temperature biases were also large for certain areas. In some cases, annual mean temperature
603 biases were greater than 5°C and were observed to go up to 10°C, particularly in the Sahel
604 and in the areas surrounding the Himalayas and the Tibetan Plateau in Nepal (Figure 11). The
605 most evident temperature biases were found in the NASA-GISS models (GISS-AOM, GISS-
606 ModelE-H and GISS-ModelE-R), and in INM-CM3.0, probably due to their coarse
607 resolution. The quality of higher resolution models was in general better, but geographic
608 trends were difficult to identify, as the locations with mean temperature were scant (7,280
609 locations for the whole study area). The smallest biases were observed in WAF, northern
610 EAF and central India, where temperature biases were below 1.5°C, particularly for the
611 models BCCR-BCM2.0, UKMO-HadCM3, NCAR-PCM1, CCCMA-CGCM3.1-T47 and
612 MIUB-ECHO-G, some of which have been reported to perform well in tropical areas before
613 (Gleckler et al., 2008; Jun et al., 2008). These biases were mostly concentrated in lowlands
614 and were mostly warm-biases, except for UKMO-HadCM3 (Figure 12). Cold-biased models
615 were usually the GISS-NASA models, MIROC3.2-MEDRES, UKMO-HadCM3, IPSL-CM4,
616 MRI-CGCM2.3.2A and IAP-FGOALS1.0-G both for seasons (i.e. JJA, DJF, maps not
617 shown) and for the annual mean (Figure 11, 12).

618 619 **5. Discussion**

620 **5.1. Climate data and agricultural research**

621 Although climate model data (“as is”) are often preferred for impact studies, crop modellers
622 and agricultural scientists should be cautious when developing future adaptation strategies
623 based on crop models applied using future predictions of different (and sometimes unknown)
624 nature (Jarvis et al., 2011), given the large uncertainties regarding the agricultural system and
625 plant responses, the underlying uncertainty related to parameterised processes, and the
626 differences in scales, all of which are reported in the impact-assessment literature [e.g.
627 (Challinor and Wheeler, 2008)]. This, however, does not necessarily imply that climate
628 model data cannot or should not be used, but rather means that an adequate treatment of
629 biases needs to be done before climate and crop models can be properly used together
630 (Challinor et al., 2010; Osborne et al., 2007).

631

632 Our findings demonstrate that, for regional assessments where large area process-based crop
633 models, statistical, or empirical models are to be used, products such as WorldClim (Jones
634 and Thornton, 2003; Thornton et al., 2009) and CRU (Challinor et al., 2004) coupled with
635 weather generation routines appear to be the best-bet approach (Challinor et al., 2004; Jones
636 and Thornton, 2003), although climate model data can also be used with proper bias
637 treatment (Challinor et al., 2010; Osborne et al., 2007). However, if studies are to be carried
638 out on a site-specific scale (Parry et al., 2005), weather station data is the best means by
639 which to calibrate the modelling approaches. While partnerships are constantly being built
640 and this allows researchers to share data, currently global weather station data such as GSOD
641 and GHCN seem to be good options in cases when no other data is available, particularly
642 when coupled with satellite data or other (country specific) historical weather records
643 (Álvarez-Villa et al., 2010).

644

645 Agricultural research requires high quality and high resolution climatological data to yield
646 accurate results, but to date this has been impossible to achieve at detailed scales and with
647 sufficient coverage, partly due to the difficulty in compiling and revising field data and partly
648 due to the limited climatology knowledge of agricultural researchers (with some exceptions).
649 Large-scale datasets can be matched to certain crop models, mostly when these models can
650 be applied at large scales (Challinor et al., 2010) or do not rely on a detailed calibration of
651 varietal-level crop parameters (Lobell et al., 2011; Lobell et al., 2008). However, matching
652 different modelling scales is not a trivial matter (Baron et al., 2005; Challinor et al., 2009a).
653 Two options are available for matching two differing scales:

- 654 (1) Decreasing the resolution of the crop model from plot scale to large regions, at the
655 expense of loss of detail in some processes [see (Challinor et al., 2007b; Challinor et
656 al., 2004; Yao et al., 2007)], or
- 657 (2) Disaggregating the coarse-resolution climate data, at the expense of introducing noise
658 and possibly propagating uncertainties present in the original climate model data
659 (Tabor and Williams, 2010).

660

661 These two choices yield different results that need to be assessed and coupled. Climate data
662 can be aggregated up to any scale to match any intended use (Masson and Knutti, 2011), but
663 agricultural impacts need to be informed at an scale such that information can be used for
664 decision making and adaptation (Jarvis et al., 2011). Hence, governments and international
665 agencies should support common platforms through which data can be shared without
666 restrictions between members of the research community. Best-bet methods can then be
667 applied over such data to produce useable datasets that can be further shared, used and
668 assessed in multidisciplinary and transdisciplinary approaches.

669

670 **5.2. Robustness of existing weather station network**

671 It is tacitly acknowledged that the use of interpolated surfaces can lead to errors and biases
672 when these data are used for impact assessment (A. Jarvis, pers. comm.). However, we have
673 demonstrated here that the effects on uncertainty are actually rather low in most of the cases,
674 with very few exceptions (highlands of Ethiopia, the Himalayas, and some parts of the Sahara
675 and Southern Africa, [Figure 6](#)).

676

677 The results of this research suggest that, despite weather station density being important, it
678 may not be the only determining factor for a good ability to fill information gaps (Hijmans et
679 al., 2005). Based on our results, we suggest that, in selecting locations to measure weather,
680 the following factors be taken into account: (1) the nature of the variable (e.g. precipitation
681 might be much more difficult to monitor than temperature), (2) the area where it is measured
682 (topographically complex areas are much more variable), (3) the values of the variable in the
683 areas where it is measured (high values are subjected to larger absolute errors, assuming
684 relative errors are relatively uniform), (4) the relevance of the area for different subjects (i.e.
685 the Sahara might be irrelevant for agriculture but can be of high relevance for other fields
686 such as climate science, ecology or biodiversity and conservation), (5) possible errors in
687 measurements and other underlying factors that can influence the measurability or
688 correctness of estimates of a particular variable, and (6) possible political or social constraints
689 on access to the site. Improving weather station distribution and status, as well as improving
690 the cross-checking, correction and validation of data collected at the different sites, is
691 fundamental for improving climate data for agricultural impact assessment.

692

693 **5.3.Global climate model accuracy and performance**

694 **5.3.1. CMIP3 climate model skill**

695 GCM performance is highly reliant on the type of comparisons performed, on the GCM
696 formulation and on the nature of climate conditions in the analysed areas (Gleckler et al.,
697 2008; Masson and Knutti, 2011). Underlying factors driving GCM performance are indeed
698 difficult to track, given the complexity of the models. IPCC 4AR (CMIP3) models showed
699 varied performance, with a high tendency to being wet-biased and no general trend for
700 temperature. These responses reportedly have their origin in different factors: first, some
701 GCMs have weak forcing on sea surface temperatures (SSTs), whereas climate in Africa and
702 Asia is strongly coupled with the Atlantic and Indian Ocean and with inland water bodies
703 (Gallée et al., 2004; Lebel et al., 2000); second, models do not properly account for the
704 relation between inter-annual variability, ENSO and the monsoonal winds (Gallée et al.,
705 2004; Hulme et al., 2001); third, the resolution of the models prevents acknowledgement of
706 local-scale land use, orographic patterns and small water bodies (Hudson and Jones, 2002);
707 fourth, cloud thickness and latent heat and moisture flux between clouds has not been
708 properly resolved in the models (Gallée et al., 2004); and fifth, convective parameterisations
709 produce an early onset of the seasonal rains and over-prediction of wet days and high-rainfall
710 events (Gallée et al., 2004).

711

712 The NASA models GISS-ModelE (-R and -H) consistently presented very low predictive
713 ability, mainly because of the models' coarse spatial resolution in conjunction with the
714 reasons mentioned above (Hansen et al., 2007). These results agree with those of Gleckler et
715 al. (2008), who reported that NCAR-PCM1, GISS-ModelE (-R and -H) and GISS-AOM
716 models are the worst performing in the 24 GCMs of the CMIP3 ensemble. Similar results are
717 reported by other authors that have assessed this or similar model ensembles (Jun et al., 2008;
718 Pierce et al., 2009). Lack of detail in land use and land use changes (Eltahir and Gong, 1996),
719 monsoon winds (Eltahir and Gong, 1996; Gallée et al., 2004), and sea surface temperature

720 anomalies (SSTs) of the Atlantic and the Indian Oceans (Lebel et al., 2000; Sun et al., 1999)
721 also causes the scales at which climate model information is robust to be varied (Masson and
722 Knutti, 2011), and prevents local scale seasonal weather patterns from being modelled
723 consistently (Douglass et al., 2008; Hansen et al., 2007).
724

725 **5.3.2. Plugging climate model data into agricultural research**

726 GCMs do not provide realistic representations of climate conditions in a particular site, but
727 rather provide estimated conditions for a large area. Our results, in agreement with those from
728 the agricultural community (Baron et al., 2005; Challinor et al., 2003) and the climate
729 community (Jun et al., 2008; Masson and Knutti, 2011), indicate that climate model outputs
730 cannot be input directly into plot-scale (agricultural) models, but support the idea that higher
731 resolution climate modelling largely improves results. Either the CMIP3 (assessed here) or
732 the upcoming CMIP5 (being released at the moment) (Moss et al., 2010) climate model
733 outputs can be adequately used in agricultural modelling if: (1) the scales between the models
734 are matched (see Sect. 5.1), (2) skill of models is assessed and ways to create robust model
735 ensembles are defined, (3) uncertainty and model spread are quantified in a robust way, and
736 (4) decision making in the context of uncertainty is fully understood.
737

738 Producing robust (i.e. skilled and certain) ensembles for agriculture is not an easy task,
739 mainly because of the scales at which these have been found to be robust (Masson and
740 Knutti, 2011). Opinions are contrasting: some authors support sub-selecting models based
741 upon performance under present conditions (Matsueda and Palmer, 2011; Pierce et al., 2009),
742 calculating a mean ensemble by weighting models based on skill (Matsueda and Palmer,
743 2011; Walsh et al., 2008), while others advocate using all available models with no-weighting
744 at all (Reifen and Toumi, 2009). We suggest that until sensitivities of agricultural models to
745 ensemble spread are fully explored (Baigorria et al., 2007), the full CMIP3 (or CMIP5)
746 ensembles should be used.
747

748 Strategies for combining plot-scale and large-scale models and for optimising the overall
749 result (including estimation of uncertainties derived from the scale-matching process) need to
750 be further researched. The potential of high-quality and less uncertain climate predictions of
751 current and future climate conditions for agricultural research is expected to have a direct
752 impact on decision-making at different levels and for different purposes: to improve yields on
753 the farm, to direct country level policies and investment, to define research foci, to direct
754 international agencies' investments, and to clarify global greenhouse emissions limits and
755 commitments (Challinor et al., 2009a; Funke and Paetz, 2011; IPCC, 2007).
756

757 **6. Conclusions**

758 A thorough analysis of different aspects of climate data for agricultural applications was
759 performed. All topics addressed here are of high relevance to agricultural applications,
760 particularly in the global tropics. Several important points were raised: (1) spatial scale is the
761 most important issue for agricultural researchers, as they prefer to use monthly products with
762 higher resolution rather than daily products with very low spatial resolution, or else limit their
763 areas of study to field plots; (2) the sensitivities of Sub-Saharan African and Southeast Asian

764 climate to data loss and poor availability were found to not be limiting factors for the region,
765 with the exceptions of mountainous areas in Nepal and Ethiopia; and (3) climate modelling,
766 although constantly improving and useful, still requires considerable future development.

767

768 As such, CMIP3 GCMs can be used with a certain degree of confidence to represent large-
769 area climate conditions for some areas and periods. In areas where predictions lack enough
770 skill for agricultural modelling, models can be bias-corrected using different methods [see
771 (Challinor et al., 2009a; Hawkins et al., 2011; Reifen and Toumi, 2009)]. Whilst model skill
772 is expected to improve with the upcoming IPCC Fifth Assessment Report, climate model
773 ensembles as well as different methods for ‘calibrating’ (i.e. pre-processing for input into
774 crop models) climate model data both need to be used, as uncertainties go beyond those
775 derived from emissions scenarios (Hawkins et al., 2011). The proper usage of climate
776 projections for agricultural impact assessment is of paramount importance in order to
777 properly inform adaptation.

778

779 Finally, it is critical to understand the implications of all this to agriculture. Crops are
780 sensitive to shortages in water and heat stresses during key periods during their development
781 (i.e. flowering, fruit filling). Therefore, lack of skill in representing seasonal and inter-annual
782 variability is expected to produce a significant obstacle to agricultural impact assessment of
783 climate change; several examples in the literature exist that illustrate this (Baigorria et al.,
784 2008; Baigorria et al., 2007). The importance of this factor depends on the strength of the
785 climate signal on yields and the variables that drive this signal. Future impact assessments
786 need to take into account input data and climate model data inaccuracies, sensitivities and
787 uncertainties; make their own assessments of the inaccuracies and uncertainties; and
788 comprehensively quantify and report uncertainties in the impact assessment process.

789

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798

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1094 **Figure captions**

1095

1096 **Figure 1** Cascade of constraints to climate data, as normally observed in agricultural impact
1097 assessment

1098

1099 **Figure 2** Areas of study. Bold-outlined areas indicate the areas on which the study focused
1100 (SN: Senegal, ML: Mali, NE: Niger, BF: Burkina Faso, GH: Ghana, UG: Uganda, ET:
1101 Ethiopia, KE: Kenya, TZ: Tanzania, NP: Nepal, BD: Bangladesh, IN: India)

1102

1103 **Figure 3** Topics treated in the analysed agricultural studies. WG: weather generators.

1104

1105 **Figure 4** Frequency of use of the different data sources in agricultural studies. A. Present-day
1106 climates. B. Future climates. Datasets acronyms are as follows: CRU-TS: Climatic Research
1107 Unit monthly time series product at 0.5 degree, GCM: global climate model output, RCM:
1108 regional climate model, CRU-CL: CRU monthly climatology product at 10 arc-minute,
1109 MARS: Data from the MARS European project, GSOD: Global summary of the day,
1110 ARTES: Africa rainfall and temperature evaluation system, VEMAP: United States
1111 comprehensive dataset, ATEAM: Advanced Terrestrial Ecosystem Analysis and Modelling,
1112 PRISM: United States dataset, GPCP: Global Precipitation Climatology Project, GPCC:
1113 Global Precipitation Climatology Centre, GHCN: Global Historical Climatology Network, AI
1114 GCM: GCM data “as is”, SD GCM: statistically downscaled GCM, PS GCM: pattern scaled
1115 GCM, WG GCM: GCM data through a weather generator, SC Variables: systematic changes
1116 in target key variables, Unclear: not specified clearly in study, ARPEGE: the ARPEGE
1117 Atmospheric GCM (Déqué et al., 1994).

1118

1119 **Figure 5** Performance of the interpolations for all variables and months as measured by the
1120 R-square value. A. Rainfall, B. Mean temperature, C. Maximum temperature, D. Minimum
1121 temperature

1122

1123 **Figure 6** Uncertainties in WorldClim expressed as standard deviations from the mean of the
1124 100 cross-validated folds for (A) total annual rainfall (in mm), and (B) annual mean
1125 temperature (in °C).

1126

1127 **Figure 7** Comparison (R-square based) of observed climatology (CL-WS [w], GHCN-CL [g]
1128 and GSOD-CL [o]) and each of the GCMs (GCM-CL) for each of the countries in the study
1129 area for mean temperature (top), temperature range (middle) and precipitation (bottom), for
1130 the annual and two seasonal (DJF, JJA) means or totals. All R^2 values were statistically
1131 significant at $p < 0.0001$

1132

1133 **Figure 8** Comparison (R-square based) of interpolated climatology (i.e. CRU-IS [c], WCL-IS
1134 [w]), and each of the GCMs (GCM-CL) for each of the countries in the study area for mean
1135 temperature (top), temperature range (middle) and precipitation (bottom) for the annual mean
1136 or total and two seasons (DJF, JJA). All R^2 values were statistically significant at $p < 0.001$.

1137

1138 **Figure 9** Root mean squared error (RMSE), in millimetres, between observed (GHCN-TS)
1139 and GCM (GCM-TS) time series, for the 24 GCMs in Table 2, for annual total rainfall
1140 between the years 1961-1990.

1141

1142 **Figure 10** Mean bias of GCM (GCM-TS) time series compared to observed time series
1143 (GHCN-TS), for the 24 GCMs in Table 2, for annual total rainfall between the years 1961-
1144 1990. Bias is expressed as the slope of the regression curve between observed and climate-
1145 model series. Values below 1 (light grey areas) indicate that GCMs are wet-biased, whereas
1146 values above 1 (dark grey areas) indicate that GCMs are dry-biased.

1147

1148 **Figure 11** Root mean squared error (RMSE), in Celsius degree, between observed (GHCN-
1149 TS) and GCM (GCM-TS) time series, for the 24 GCMs in Table 2, for annual mean
1150 temperature between the years 1961-1990

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1152 **Figure 12** Mean bias of GCM (GCM-TS) time series compared to observed time series
1153 (GHCN-TS), for the 24 GCMs in Table 2, for annual mean temperature between the years
1154 1961-1990. Bias is expressed as the slope of the regression curve between observed and
1155 climate-model series. Values below 1 (light grey areas) indicate that GCMs are warm-biased,
1156 whereas values above 1 (dark grey areas) indicate that GCMs are cold-biased.

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Table 1 Number of locations per data source (global)

Source*	Precipitation stations	Mean temperature stations	Min., Max. temperature stations	Period
GHCN v2	20,590	7,280	4,966	1950-2000
WMO CLINO	4,261	3,084	2,504	1961-1990
FAOCLIM 2.0	27,372	20,825	11,543	1960-1990
CIAT	18,895	13,842	5,321	1950-2000

1166 *GHCN v2: Global Historical Climatology Network version 2 (Peterson and Vose, 1997);
1167 WMO CLINO: World Meteorological Organization Climatology Normals; FAOCLIM 2.0:
1168 Food and Agriculture Organization of the United Nations Agro-Climatic database (FAO,
1169 2001); CIAT: Database assembled by Peter J. Jones at the International Center for Tropical
1170 Agriculture (CIAT).
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Table 2 Available GCMs, resolutions, and main references

Model	Country	Atmosphere	Ocean	Reference
BCCR-BCM2.0	Norway	T63, L31	1.5x0.5, L35	(Furevik et al., 2003)
CCCMA-CGCM3.1 (T47)	Canada	T47 (3.75x3.75), L31	1.85x1.85, L29	(Scinocca et al., 2008)
CCCMA-CGCM3.1 (T63)	Canada	T63 (2.8x2.8), L31	1.4x0.94, L29	(Scinocca et al., 2008)
CNRM-CM3	France	T63 (2.8x2.8), L45	1.875x(0.5-2), L31	(Salas-Méla et al., 2005)
CSIRO-Mk3.0	Australia	T63, L18	1.875x0.84, L31	(Gordon et al., 2002)
CSIRO-Mk3.5	Australia	T63, L18	1.875x0.84, L31	(Gordon et al., 2002)
GFDL-CM2.0	USA	2.5x2.0, L24	1.0x(1/3-1), L50	(Delworth et al., 2006)
GFDL-CM2.1	USA	2.5x2.0, L24	1.0x(1/3-1), L50	(Delworth et al., 2006)
GISS-AOM	USA	4x3, L12	4x3, L16	(Russell et al., 1995)
GISS-MODEL-EH	USA	5x4, L20	5x4, L13	(Schmidt et al., 2006)
GISS-MODEL-ER	USA	5x4, L20	5x4, L13	(Schmidt et al., 2006)
IAP-FGOALS1.0-G	China	2.8x2.8, L26	1x1, L16	(Yongqiang et al., 2004)
INGV-ECHAM4	Italy	T42, L19	2x(0.5-2), L31	(Gualdi et al., 2008)
INM-CM3.0	Russia	5x4, L21	2.5x2, L33	(Diansky and Zalensky, 2002)
IPSL-CM4	France	2.5x3.75, L19	2x(1-2), L30	(Marti et al., 2005)
MIROC3.2-HIRES	Japan	T106, L56	0.28x0.19, L47	(Hasumi and Emori, 2004)
MIROC3.2-MEDRES	Japan	T42, L20	1.4x(0.5-1.4), L43	(Hasumi and Emori, 2004)
MIUB-ECHO-G	Germany/Korea	T30, L19	T42, L20	(Grötznér et al., 1996)
MPI-ECHAM5	Germany	T63, L32	1x1, L41	(Jungclaus et al., 2006)
MRI-CGCM2.3.2A	Japan	T42, L30	2.5x(0.5-2.0)	(Yukimoto et al., 2001)
NCAR-CCSM3.0	USA	T85L26, 1.4x1.4	1x(0.27-1), L40	(Collins et al., 2006)
NCAR-PCM1	USA	T42 (2.8x2.8), L18	1x(0.27-1), L40	(Washington et al., 2000)
UKMO-HADCM3	UK	3.75x2.5, L19	1.25x1.25, L20	(Gordon et al., 2000)
UKMO-HADGEM1	UK	1.875x1.25, L38	1.25x1.25, L20	(Johns et al., 2006)

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Table 3 Summarised performance of all GCMs with available data for each of the variables and periods in the study countries for different ranges of the R^2 skill evaluation parameter.

Variable	Period	Dataset*	$R^2 < 0.5$ (%)*	$0.5 < R^2 < 0.7$ (%)*	$R^2 > 0.8$ (%)*	$R^2 > 0.9$ (%)*
Rainfall	Annual	IS	2.8	6.6	77.8	54.3
		WS	37.5	19.4	30.8	17.0
		ALL	23.6	14.3	49.6	31.9
	DJF	IS	17.7	19.3	49.1	25.9
		WS	38.1	17.2	31.4	15.7
		ALL	29.9	18.1	38.5	19.8
	JJA	IS	12.8	17.2	58.9	40.1
		WS	15.2	19.1	52.1	34.5
		ALL	14.2	18.3	54.8	36.7
Diurnal temperature range	Annual	IS	0.4	2.2	81.8	73.1
		WS	0.4	1.2	54.5	46.1
		ALL	0.4	1.7	68.1	59.6
	DJF	IS	0.4	2.2	80.4	71.2
		WS	0.4	2.4	53.1	47.7
		ALL	0.4	2.3	66.8	59.4
	JJA	IS	0.4	2.0	80.7	67.2
		WS	0.4	1.2	54.5	46.1
		ALL	0.4	1.6	67.6	56.6
Mean temperature	Annual	IS	0.7	1.2	96.4	95.7
		WS	2.4	1.9	93.5	91.0
		ALL	1.7	1.6	94.7	92.8
	DJF	IS	3.5	1.9	93.2	91.5
		WS	2.3	2.3	93.9	91.2
		ALL	2.8	2.2	93.6	91.3
	JJA	IS	0.0	0.0	100.0	98.8
		WS	0.0	0.1	99.8	98.5
		ALL	0.0	0.1	99.9	98.6

1183 * Values are expressed as percent of country-GCM combinations for comparisons of GCM-
 1184 CL and different observational datasets: interpolated surfaces (IS), namely, WCL-IS and
 1185 CRU-IS; weather stations (WS), namely, GHCN-CL, WCL-WS, GSOD-CL; and as the
 1186 average of IS and WS (ALL)

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Assessing relevant climate data for agricultural applications

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32 **Abstract**

33 Climate change is expected to substantially reduce agricultural yields, as reported in the by
34 the Intergovernmental Panel on Climate Change (IPCC). In Sub-Saharan Africa and (to a
35 lesser extent) in South Asia, limited data availability and institutional networking constrain
36 agricultural research and development. Here we performed a review of relevant aspects in
37 relation to coupling agriculture-climate predictions, and a three-step analysis of the
38 importance of climate data for agricultural impact assessment. First, using meta-data from the
39 scientific literature we examined trends in the use of climate and weather data in agricultural
40 research, and we found that despite agricultural researchers' preference for field-scale
41 weather data (50.4% of cases in the assembled literature), large-scale datasets coupled with
42 weather generators can be useful in the agricultural context. Using well-known interpolation
43 techniques, we then assessed the sensitivities of the weather station network to the lack of
44 data and found high sensitivities to data loss only over mountainous areas in Nepal and
45 Ethiopia (random removal of data impacted precipitation estimates by $\pm 1,300$ mm/year and
46 temperature estimates by $\pm 3^\circ\text{C}$). Finally, we numerically compared IPCC Fourth Assessment
47 Report climate models' representation of mean climates and interannual variability with
48 different observational datasets. Climate models were found inadequate for field-scale
49 agricultural studies in West Africa and South Asia, as their ability to represent mean climates
50 and climate variability was limited: more than 50% of the country-model combinations
51 showed $< 50\%$ adjustment for annual mean rainfall (mean climates), and there were large
52 rainfall biases in GCM outputs (1,000 to 2,500 mm/year), although this varied on a GCM
53 basis (climate variability). Temperature biases were also large for certain areas (5-10°C in the
54 Himalayas and Sahel). All this is expected to improve with IPCC's Fifth Assessment Report;
55 hence, appropriate usage of even these new climate models is still required. This improved
56 usage entails bias reduction (weighting of climate models or bias-correcting the climate
57 change signals), the implementation of methods to match the spatial scales, and the
58 quantification of uncertainties to the maximum extent possible.

59

60 **Keywords:** Sub-Saharan Africa; South Asia; climate modelling; climate model; skill;
61 uncertainty; CMIP3; CMIP5.

62

1. Introduction

Agriculture is expected to play an important role in the context of climate change, not only because it is considered amongst the most vulnerable sectors, but also because it is part of the solution (i.e. potential to mitigate greenhouse gases [GHGs] emissions) (FAO, 2009; IPCC, 2007). Agriculture will likely be severely affected over the next hundred years due to unprecedented rates of changes in the climate system (IPCC, 2007; Jarvis et al., 2010; Lobell et al., 2008; Thornton et al., 2011). Some of these impacts have already been observed (Battisti and Naylor, 2009; Schlenker and Lobell, 2010). To help cope with such impacts, a framework to assess the effects of climate change on agriculture and food security and to aid with adaptation was established in 2008, as described by Jarvis et al. (2011): The Consultative Group of International Agricultural Research (CGIAR) Research Program on Climate Change, Agriculture and Food Security (CCAFS).

For adaptation to be successful, agricultural and climate data are crucial, and these are scarce in their basic forms (data from research and weather stations, respectively) or not very well managed and/or maintained [in certain parts of the world](#). Most importantly, climate databases and their derived products are sometimes inaccurate, or else lack the documentation necessary to facilitate their use within the agricultural research community. In some instances, this may be indicative of the gap between the agricultural and climate research communities (Pielke et al., 2007; Thornton et al., 2011). Even when the two do collaborate, agricultural researchers face critical constraints when accessing basic sources of climate data (i.e. weather stations) due to a number of factors, from access to data, to weather maintenance and data quality checks, to the weather itself (DeGaetano, 2006).

In the last 10 years, various datasets have been developed by different institutions, usually based on either a combination of weather station data, satellite data, and numerical weather prediction models in addition to interpolation methods, or on the sole application of climate models. The usage of these datasets for agricultural modelling purposes is rather limited for one or more of the following reasons: (1) their time step is long (monthly in the best case); (2) their temporal coverage is limited to an average of several years (Hijmans et al., 2005; New et al., 2002); (3) their spatial resolution is too coarse (Adler et al., 2003; Schneider et al., 2010); (4) their geographic coverage is not wide enough (Di Luzio et al., 2008); and (5) only certain variables (i.e. temperatures, rainfall) are reported whereas other agriculturally relevant measures (e.g. potential and/or reference evapotranspiration, relative humidity, solar radiation) are rarely reported (Di Luzio et al., 2008; Hijmans et al., 2005). Moreover, assessments of these data (particularly climate models) have been done only under a climate-science perspective (Gleckler et al., 2008; Pierce et al., 2009), for a limited number of variables (Jun et al., 2008; Reifen and Toumi, 2009), or for a reduced realm (Walsh et al., 2008).

In this paper, we sought to improve the general knowledge on the available climate data for agricultural research using a three-step thorough analysis on fundamental aspects related to agricultural modelling. First, we perform a meta-analysis on the usage of various data sources for agricultural applications; second, we assess the quality and distribution of weather station

107 records by exploring both the ability of these data to fill geographic information gaps by
108 means of interpolation, and the sensitivities of the different regions to data loss; and finally,
109 we assess the accuracy of climate model outputs against different observational datasets using
110 various metrics reported in previous literature (Gleckler et al., 2008; Pierce et al., 2009). We
111 finally analyse the main implications of our findings on agricultural impact assessment.

112

113 **2. Review of knowledge and data**

114 **2.1. Understanding of processes and crop modelling**

115 Mechanisms to fix carbon in plants (i.e. photosynthesis) are affected by a number of factors
116 (El-Sharkawy, 2005; Prasad et al., 2002), although responses strongly depend on the type of
117 mechanism used by the plant to produce biomass (i.e. C₄, C₃, CAM) and on any other stresses
118 to which the plant could be subjected simultaneously. In crop production, apart from
119 appropriate plant growth it is the amount of biomass accumulated in fruits and seeds and the
120 nutrients in them that matters most (Thuzar et al., 2010). Yields are a direct consequence of
121 photosynthesis and biomass accumulation, and these are directly or indirectly affected by
122 environmental conditions [see (Challinor et al., 2009b) for a review]. Well-watered crops
123 grown under optimal temperature and solar radiation ranges develop to their full production
124 potential (van Ittersum et al., 2003), but growth potential reduces if the crop is stressed during
125 the growing season (Hew et al., 1969; Huntingford et al., 2005).

126

127 Therefore, modelling crop growth depends on (1) correct formulation of the simulation
128 model, (2) our ability to understand the effects of environmental factors on growth, and (3)
129 correct measurement of the relevant environmental factors for correct mapping of their
130 interactions (Boote et al., 1996; El-Sharkawy, 2005). Hence, crop modelling largely benefits
131 from accurate measurements of temperatures, rainfall, and solar radiation, as the main factors
132 acting on photosynthesis (Challinor and Wheeler, 2008; Hoogenboom et al., 1994), but even
133 these basic data are often unavailable, messy, or of limited quality. The more available data
134 there exists, the better calibration and evaluation of crop models can be (Adam et al., 2011;
135 Niu et al., 2009; Xiong et al., 2008).

136

137 Additionally, most crop models simulate growth of individual plants and then scale out the
138 modelling results to the plot-scale, based on management decisions such as plant and row
139 distances, and plot size (Aggarwal et al., 2006; Boote et al., 1996; Hoogenboom et al., 1994).
140 On the other hand, available weather data (when not measured in the field) is only available
141 at coarse spatial scales. Matching these two spatial scales is not an easy task [see (Challinor
142 et al., 2009a; Jagtap and Jones, 2002; Trnka et al., 2004) for a review]. The challenge is thus
143 to increase the knowledge of the interactions between atmospheric and crop-growth processes
144 (Boote et al., 1996) whilst avoiding model over-parameterisation (Challinor et al., 2009b),
145 improving the accuracy of inputs (Adam et al., 2011), and matching both spatial scales
146 (Challinor et al., 2009a). All this requires closing the gap between crop and climate scientists.

147

148 **2.2. Weather data**

149 Measurements of weather for a given site are often unavailable because (1) there is no
150 weather station; (2) weather stations are not well maintained so data are either only available

151 for a short period or contain gaps, (3) collected data are not properly stored; (4) data do not
152 pass basic quality checks; and/or (5) access to data is restricted by holding institutions (Figure
153 1). This all further constrains agricultural impact assessment, highlighting the importance of
154 making data public.

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<Insert Figure 1 here>

158 Apart from the constraints related to access and weather station locations, probably the most
159 important issue regarding weather data is quality (Begert et al., 2008; DeGaetano, 2006)
160 (Figure 1), which also greatly affects the performance of impact models. Therefore, the
161 climate and agricultural community has partly focused on developing methods for either
162 temporal or spatial data gap filling, and on using such methods for developing global or
163 regional datasets with public access (Hijmans et al., 2005; Jones and Thornton, 1999; Soltani
164 et al., 2004).

165

166 However, uncertainties in global datasets derived from interpolation methods have been only
167 barely (if at all) estimated (Buytaert et al., 2009; Challinor and Wheeler, 2008; Soria-Auza
168 et al., 2010). Researchers using global datasets and any weather station source need to be aware
169 of these problems and ought to take this into account by testing the sensitivities of their
170 approaches to accuracy issues (i.e. inhomogeneities, discontinuities) and (if possible)
171 providing results within the range of uncertainty in input data (i.e. such as the outputs of cross
172 validated interpolation methods) (Challinor et al., 2005).

173

174

2.3.Climate model data

175 General Circulation Models (GCMs) are currently the best way to model the complex
176 processes that occur at the earth system's level (Huntingford et al., 2005; IPCC, 2007).
177 However, as CGMs are highly complex, they are computationally expensive, so they have
178 only been used for predictions at coarse spatial scales. These predictions therefore involve a
179 number of uncertainties relevant to agriculture [see (Challinor et al., 2009b; Jarvis et al.,
180 2010; Quiggin, 2008) for reviews on the topic].

181

182 In short, uncertainty in climate modelling arises from the impossibility of modelling the
183 climate system with complete determinism (Walker et al., 2003). This uncertainty can arise
184 from: context (boundaries of the system modelled), model, inputs, and parameters (Walker et
185 al., 2003). Model uncertainty can be structural or technical: structural uncertainty in models is
186 associated with our lack of understanding of the system, whereas technical uncertainty relates
187 to our inability to implement mathematical formulations in computational systems. Other
188 uncertainties in climate modelling arise from variable driving forces (greenhouse gas
189 emissions and concentrations), initial conditions and parameterised physics (Challinor et al.,
190 2009b; Walker et al., 2003). Rationalisation and quantification of all these uncertainties under
191 the context of agriculture is possible (see Challinor et al., 2009b for a review).

192

193 Crop modellers are thus challenged to understand the broad concepts of climate modelling
194 uncertainties and detect the sensitivities of crop models to them, whilst also having a basic

195 understanding of earth processes in order to identify major flaws in climate models and
196 decide the best ways to couple them with crop models.

197

198 **3. Materials and methods**

199 Throughout this paper, we built upon existing knowledge of agricultural and climate
200 modelling (Sect. 2) and:

- 201 1. Performed a meta-analysis on the usage of climate and weather data for agricultural
202 modelling purposes and summarised the desirable characteristics sought when
203 modelling crop production.
- 204 2. Analysed the robustness of the existing weather station network by assessing both the
205 ability of these data to correctly fill information gaps via interpolation methods, and
206 the network's sensitivities to information loss.
- 207 3. Assessed the accuracy of climate model outputs from the Fourth Assessment Report
208 of the IPCC (IPCC, 2007) against different observational datasets, using metrics and
209 methods reported in the climate-science literature that are also familiar to agricultural
210 researchers.

211

212 All calculations were done by means of the software packages R-2.13.1 (available at
213 <http://www.r-project.org>) and GRASS-GIS 6.4.0 (available at <http://grass.fbk.edu>) in a 64-bit
214 Red Hat Enterprise Linux 5 box.

215

216 **3.1. Study area**

217 We focused on the geographic area of Africa and South Asia, where several studies have
218 identified that significant vulnerabilities exist (Aggarwal, 2008; Aggarwal et al., 2004;
219 Barrios et al., 2008; Byjesh et al., 2010; Challinor et al., 2007a; Chipanshi et al., 2003; Jones
220 and Thornton, 2003; Lane and Jarvis, 2007; Liu et al., 2008; Lobell et al., 2008; Thornton et
221 al., 2009; Thornton et al., 2011; Washington et al., 2006). In particular, we concentrate our
222 efforts on West Africa (Senegal, Mali, Burkina Faso, Ghana and Niger), East Africa
223 (Ethiopia, Tanzania, Uganda and Kenya) and the Indo-Gangetic Plains countries (India,
224 Nepal, and Bangladesh), hereafter referred to as WAF, EAF and IGP, respectively (Figure 2).

225

226

<Insert Figure 2>

227

228 **3.2. Analysing the usage of climate data in agricultural studies**

229 **3.2.1. Meta-data from agricultural studies**

230 We gathered data from a number of publications on any topic that made use of climate data
231 for any sort of agricultural modelling. We conducted searches using various search engines
232 and downloaded only peer-reviewed publications. Review papers and the Fourth Assessment
233 report of the IPCC were particularly useful in identifying additional published studies. We
234 analysed all publications that in any way involved the usage of climate data for agricultural
235 modelling purposes. As the selection of the impact assessment model is the first decision that
236 any researcher needs to make, we focus on the driving factors of this decision. We recorded
237 different variables from the studies as follows:

- 238 (1) Problem and/or topic in question: classified in categories such as impact assessment,
239 seasonal yield forecasting, sole crop modelling, and climate attribution, among others.
240 Each study was classified into only one category by taking into account only the main
241 issue addressed by the paper;
- 242 (2) Scale of the approach: includes site, sub-national, country, regional (group of
243 countries), and global;
- 244 (3) Use of weather generators: for both present and future, we recorded whether the study
245 did or did not use a weather generator;
- 246 (4) Climate dataset (current): GCM when a GCM (regardless of which one) was used,
247 RCM when an RCM (regardless of which one) was used, weather station, satellite (no
248 further discrimination), and important datasets (i.e. CRU, WorldClim, GPCP, among
249 others);
- 250 (5) Climate dataset (future): the nature of used future projections was recorded here
251 including the downscaling method, if applicable. Classifications were: GCM “as is”
252 when studies used raw GCM outputs as inputs, pattern scaled GCMs (Mitchell et al.,
253 2004), RCMs, systematic changes to current climate data, statistical downscaling
254 (Wilby et al., 2009), and weather generator downscaled GCM (Jones et al., 2009).

255
256 For further details on the above categories the reader is referred to our supplementary
257 material (part 1). We revised a total of 205 peer-reviewed publications (See supplementary
258 material part 2), printed between the years 1983 and 2011. Most of the studies were published
259 immediately before or after the IPCC 4AR was released in 2007. When a certain study made
260 use of two different sources of present-day climate data, it was considered twice (totalling
261 247 cases).

262 **3.2.2. Analysing the usage of climate data in agricultural studies**

264 We analysed the recent trends in the use of climate data for agriculture: the obvious
265 constraints in the studies, the type of approaches used and the climate data inputs used to
266 drive the chosen agricultural models. By doing this, we ensured that we covered all the main
267 factors driving an agricultural researcher’s decision to select a particular approach for a given
268 problem.

269 **3.3. Analysis of weather station data**

270 **3.3.1. Worldwide weather station network data**

272 Long term climatological means of monthly precipitation and mean, maximum and minimum
273 temperatures were assembled, as described by Hijmans et al. (2005). However, it is important
274 to note that at the global level the sources of these data are large in number and differ in
275 coverage, availability and quality (Table 1), and thorough quality checks were done only in a
276 sub-set of the sources by original distributing institutions.

277
278 **<Insert Table 1 here>**

279
280 Additional sources such as R-Hydronet (<http://www.r-hydronet.sr.unh.edu/english/>) and
281 Oldeman (1988) database for Madagascar were also included. We discarded any weather

282 station with less than 10 years of data. The final dataset (after quality control and duplicates
283 removal, [see Hijmans et al. 2005 for more details](#)) comprised 13,141 locations with monthly
284 precipitation data, 3,744 locations with monthly mean temperature, and 2,684 locations with
285 diurnal temperature range within our study region. This dataset is hereafter referred to as
286 WCL-WS.

287

288 **3.3.2. Analysing robustness of existing weather station networks**

289 Many methods exist that allow the user to determine (interpolate) the value of a parameter
290 (e.g., monthly rainfall) in a given condition (i.e. in a given site, at a given time, or both),
291 where it had never been measured before. Some of these methods are already popular with
292 researchers using climate data (Hijmans et al., 2005; Hutchinson, 1995; Jones and Thornton,
293 1999; New et al., 2002) either on a regional or on a global basis. For climate-variable
294 interpolations, the robustness of weather records is critical for an accurate result.

295

296 We assessed the robustness of the weather station network by testing both the ability of
297 weather records to yield accurate interpolation results, and the sensitivities of the network to
298 information loss. Towards the ~~as~~ ends, we used the WCL-WS dataset to fit a thin plate
299 spline interpolation algorithm (Hutchinson, 1995) for our study region. We investigated the
300 effect of weather station availability by using 100 cross validated folds for four variables
301 (monthly maximum, minimum and mean temperatures and total precipitation) using similar
302 methods as in Hijmans et al. (2005) and New et al. (2002) for each fold. We used longitude,
303 latitude and elevation as independent variables. We used 85% randomly selected data points
304 for fitting the splines and the remaining 15% for evaluating the result for each variable and
305 month. For the evaluation, we calculated the R^2 and the Root Mean Square Error (RMSE)
306 and produced boxplots of the 100-fold-by-12-month interpolations for each of the four
307 variables. As the number of stations considerably exceeded the amount of available memory
308 for processing, we divided the whole region of study in 5 tiles, each with an equivalent
309 number of locations. We then projected the fitted splines onto 30-arc-second gridded datasets
310 of latitude, longitude and altitude (Jarvis et al., 2008), thus producing a total of 4,800
311 interpolated surfaces (12 months times 4 variables times 100 folds). Finally, we analysed the
312 spatial variability of standard deviations and the performance of the interpolation technique
313 as proxies for sufficient distribution and geographic density of weather stations.

314

315 **3.4. Assessment of IPCC Fourth Assessment Report (4AR) model data**

316 **3.4.1. Long-term observed mean climatology from weather stations**

317 Three different long term climatology datasets were assembled: (1) the Global Historical
318 Climatology Network (GHCN, [as in Sect. 3.3.1](#)) version 2 (Peterson and Vose, 1997),
319 available at <http://www.ncdc.noaa.gov/pub/data/ghcn/v2>. [We used GHCN as an independent](#)
320 [source because it is a global resource that contributed significantly to WCL-WS and also](#)
321 [because it is available at more temporally disaggregated levels \(i.e. monthly\), thus allowing](#)
322 [uniformity with analyses on Sect. 3.4.3 and 3.4.6.](#) This database includes monthly historical
323 totals (1900-2010) of precipitation (20,590 stations), and means of maximum, minimum
324 (4,966) and mean (7,280) temperatures. GHCN data have been subject to quality checks and
325 to a process of “homogenisation” or “adjustment” (Peterson and Easterling, 1994); however,

326 the available data within our analysis domain consisted primarily of “unadjusted” stations.
327 For each location (6,393 stations for rainfall, 1,278 for mean temperature and 549 for
328 minimum and maximum temperature) within our study area, we averaged historical monthly
329 time series for the period 1961-1990 for maximum, minimum and mean temperatures and
330 total rainfall, resulting in a time-averaged dataset of 6,393 locations for rainfall, 1,278 for
331 mean temperature and 549 for minimum and maximum temperature. This dataset will be
332 hereafter referred to as GHCN-CL.

333

334 (2) WCL-WS (Sect. 3.3.1); and (3) the Global Surface Summary of the Day (GSOD) was
335 accessed at <http://www.ncdc.noaa.gov/cgi-bin/res40.pl>. This database contains daily data
336 from ~9,000 weather stations worldwide for 18 variables, including, mean, maximum,
337 minimum and dew point temperature, sea level and location pressure, visibility, wind speed
338 and gust, precipitation, snow depth, and specifications on the occurrence of rain, snow, fog,
339 tornado, thunder, or hail (NOAA, 2011; <ftp://ftp.ncdc.noaa.gov/pub/data/g sod/readme.txt>).
340 We selected weather stations within our study area (1,999); aggregated daily rainfall, mean,
341 maximum and minimum temperatures to a monthly time scale; and then averaged over the
342 period 1961-1990. This dataset will be hereafter referred to as GSOD-CL.

343

344 **3.4.2. Long-term observed mean climatology from interpolated surfaces**

345 We gathered high-resolution climatology from two different sources: (1) the high resolution
346 climate surfaces in WorldClim (Hijmans et al., 2005), available at <http://www.worldclim.org>.
347 WorldClim is a 30 arc-seconds (~1km at the equator) global dataset produced from the
348 interpolation of long-term climatology as measured in weather stations. Global gridded data
349 were downloaded at the 30 arc-second resolution, then masked to our analysis domain, and
350 aggregated to 10 arc-minute using bilinear interpolation in order to reduce computational and
351 storage time; and (2) the University of East Anglia Climatic Research Unit (CRU) dataset
352 (New et al., 2002), available through <http://www.cru.uea.ac.uk/cru/data/hrg/> (CRU-CL-2.0).
353 This dataset was developed using the same interpolation method as WorldClim, with the
354 main difference that WorldClim includes many more weather stations, sometimes at the
355 expense of input data quality. CRU-CL-2.0 resolution is 10 arc-minute (~20km at the
356 equator). Data were downloaded at the global level and masked to our analysis domain.
357 WorldClim and CRU-CL-2.0 are hereafter referred to as WCL-IS and CRU-IS (interpolated
358 surfaces), respectively. We used these sources because (1) they are flag products that most
359 researchers use for impact studies; (2) they are much higher resolution than GCMs (and other
360 products such as the Global Precipitation Climatology Project [GPCP] and the Global
361 Precipitation Climatology Centre [GPCC]) and hence permit the capture of small-scale
362 weather patterns (important to agriculture) as well as a direct comparison of their within-
363 GCM-gridcell mean with the actual GCM value; (3) are based only on ground observations of
364 weather and do not incorporate side-products such as reanalysis (Uppala et al., 2005) or
365 satellite data (Huffman et al., 2007), both of whose accuracy is not as good.

366

367 **3.4.3. Long-term observed time series**

368 Two sources of weather time series were used: (1) long term (1961-1990) series of monthly
369 weather conditions were gathered from GHCN version 2 (Peterson and Vose, 1997). Again,

370 we used mainly unadjusted stations. Mean monthly temperature and total monthly historical
371 rainfall data were used without any further processing; and (2) long-term (1961-1990) series
372 of daily weather as in GSOD (NCDC, 2011). For GSOD, daily precipitation and monthly
373 temperature were aggregated to the monthly level only if all days were reported with data (for
374 rainfall) and if at least 50% of the days had data (for temperatures). This resulted in 1,999
375 stations within our analysis domain, although not all stations had data for all months and all
376 years. These two data sources are hereafter referred to as GHCN-TS and GSOD-TS,
377 respectively. Lack of data prevented us from including maximum and minimum temperatures
378 in the GHCN-TS and the GSOD-TS datasets. In contrast to GHCN-CL and GSOD-CL,
379 GHCN-TS and GSOD-TS include every month and every year, thus allowing the analysis of
380 inter-annual variability.

381

382 **3.4.4. Global climate model output**

383 The latest IPCC report (Fourth Assessment Report, 4AR) comprises the sole state-of-the-art
384 public and official source of climate data for use in impact studies (IPCC, 2007; Jarvis et al.,
385 2010). We therefore decided to use IPCC 4AR results.

386

387 We downloaded present day (1961-1990) simulations of global climate at original GCM
388 resolution (~100 km) from the CMIP3 (Coupled Model Intercomparison Project phase 3) web
389 data portal at <https://esg.llnl.gov:8443/index.jsp> (PCMDI, 2007). We downloaded monthly
390 time series of mean, maximum, minimum temperature and precipitation flux in NetCDF
391 format for 24 coupled GCMs (Table 2). Separately for each GCM, w~~We~~ calculated diurnal
392 temperature range for each month and year as the difference between maximum and
393 minimum temperatures and calculated total monthly rainfall as the product between the
394 precipitation rate, the water density at sea level pressure and the number of seconds in a
395 month. We used the each climate model monthly time series (GCM-TS hereafter) and also
396 calculated average 1961-1990 climatology by averaging, for each variable (mean
397 temperature, diurnal temperature range and total rainfall), every month for the whole 1961-
398 1990 period (GCM-CL hereafter). The final datasets (i.e. GCM-TS and GCM-CL,
399 respectively) consisted of three variables (mean temperature, diurnal temperature range and
400 total monthly rainfall) for 24 different GCMs.

401

402

<Insert Table 2 here>

403

404 **3.4.5. Ability to represent long-term climatology**

405 The extent to which GCM predictions are accurate has not been fully explored for some parts
406 of the world, particularly in the context of agriculture (Gleckler et al., 2008; Pierce et al.,
407 2009; Walsh et al., 2008). As previously stated (Sect. 2.1), we compared the most readily
408 available variables from both ground observations and climate models: rainfall, mean
409 temperature and diurnal temperature range. Data for other variables are not available for our
410 study regions in observational datasets. As per our stated objective (Sect. 3), we performed
411 two sets of comparisons:

412

413

- First, we compared the GCM-CL dataset with the interpolated climatology in CRU-IS,
WCL-IS (Sect. 3.4.2). We performed comparisons on a country basis in order to yield

414 country-specific results. For each GCM gridcell, the mean, maximum and minimum
415 values of all lower scale (CRU-IS, WCL-IS) cells was first calculated and then compared
416 to the GCM value through the determination coefficient (R^2) and corresponding p-value,
417 the slope of a origin-forced (so that a 1:1 relationship was sought) regression curve (S)
418 and the root mean square error (RMSE).

- 419 • Second, using the same procedure, we compared the GCM-CL dataset with observed
420 climatology in WCL-WS (Sect. 3.3.1), GHCN-CL and GSOD-CL (Sect. 3.4.1).

421

422 We analysed total rainfall, mean temperatures and diurnal temperature ranges over three
423 periods: December-January-February (DJF), June-July-August (JJA) and the whole year
424 (ANN). These months represent the most critical seasons for agriculture in our study regions,
425 and are also the most often assessed in the existing literature (Gleckler et al., 2008; Pierce et
426 al., 2009). Due to space constraints, we present only the results of comparisons between
427 GCM gridcell values and mean values within gridcells, unless otherwise stated. We do,
428 however, discuss other relevant results in more general terms.

429

430 **3.4.6. Ability to represent long-term monthly climate time series**

431 CMIP3-related GCMs are known to misrepresent certain inter-annual and/or within-decade
432 variations that are important for agricultural systems (Govindan et al., 2002). However,
433 specific aspects of these errors have not been explored in all CMIP3 models in the context of
434 agriculture. Therefore, in order to test the consistency of GCM predictions across time, we
435 compared the GCM-TS (Sect. 3.4.4) dataset against the GHCN-TS and GSOD-TS (Sect.
436 3.4.3). The comparison was done for three periods (JJA, DJF and ANN, Sect. 3.4.4) by
437 calculating the R^2 and corresponding p-value, the slope of the regression curve as forced to
438 the origin and the RMSE between the two time series (GCM-TS vs. GHCN-TS and GCM-TS
439 vs. GSOD-TS). As a GCM cell contains one or more weather stations, we averaged the
440 monthly time series as needed before comparing the two pairs of series. Finally, we compared
441 the performance of all GCMs across the geographic space of our study area.

442

443 **4. Results**

444 **4.1. Usage of climate data in agricultural studies**

445 **4.1.1. Topics of study**

446 The most addressed topic (41.4% of the studies) in our literature review was climate change
447 impact assessment (Figure 3), followed by crop growth simulation (18.5%). Water resources-
448 impact studies round out the top three topics studied (8.1%), followed by climate attribution
449 (6.9%), crop yield forecasting (6.1%), and model assessment (5.7%). Surprisingly, formal
450 studies addressing adaptation were rather scarce (3.6%). Pests and diseases, soils, abiotic
451 stresses and climate risks appeared to be a lot less ~~important-addressed~~ than impact
452 assessment and crop growth simulation studies, which together accounted for more than 50%
453 of the total publications.

454

455

<Insert Figure 3 here>

456

457 **4.1.2. Scale of studies and type of models**

458 Most of the studies performed their models at a scale less than the size of a country; site-
459 specific or sub-national level together comprised 55% of the studies. Very few (7%) of the
460 studies were performed at the global level, likely because of the type of models used: field-
461 scale mechanistic crop growth models were the most utilised overall (69.2%); followed by
462 statistical and/or empirical approaches (S/E, 21.4%), which most of the crop growth
463 modellers criticise for not being accurate enough (Lobell and Burke, 2010; Lobell et al.,
464 2008); and finally by hydrological models (10%). The frequent use of field-based crop
465 growth models suggests that the time step requirement for input data is rather high (El-
466 Sharkawy, 2005), also confirmed by the usage of weather generators (8.5 and 11.2% for
467 present and future climates, respectively).

468

469 **4.1.3. Climate data sources**

470 Unlike the model types, which were quite similar, the sources of present climate data varied
471 substantially, with a total of 32 different sources being used for present climate data (Figure
472 4A). On average, a different present-day-climate dataset was used for every 7 agricultural
473 studies. The most commonly used data source was local (non-public) weather stations (50.4%
474 of the cases), followed by University of East Anglia Climatic Research Unit (CRU) datasets
475 with 13.7% (10.9% for CRU-TS [monthly time series], and 2.8% for CRU-CL [monthly
476 climatology]). Climate model outputs were used in 14.5% of the cases: within this group,
477 10.5% used GCM data, 4% RCM [Regional Climate Model] data, 3.6% satellite imagery, and
478 2.8% WorldClim, followed by other less relevant sources. The Global Precipitation
479 Climatology Project (GPCP) (Adler et al., 2003; Huffman et al., 2009), the Global
480 Precipitation Climatology Centre (GPCC) (Schneider et al., 2010) and the Global Historical
481 Climatology Network (GHCN, (Peterson and Vose, 1997)) were rarely reported overall
482 (0.4% each).

483

484

<Insert Figure 4 here>

485

486 The future climate data used was found to be less variable overall, with only 7 different types
487 of data employed in the 125 cases citing some type of future climate data (Figure 4B). Out of
488 these 125, only one study did not clearly state which type of climate data was used. The vast
489 majority of cases (42.9%) used GCM data “as is” (AI GCM), meaning that predictions on
490 agricultural yields were based on predicted changes at coarse resolution (~100 km). All other
491 studies involved some type of downscaling, except those that employed the systematic
492 changes approach (SC variables), which can be assumed to be sensitivity analyses rather than
493 impact studies. RCMs (Regional Climate Models) were the most common way of
494 downscaling GCMs, cited in 19% of the studies, followed by statistical downscaling with
495 17.5% (SD GCM, (Tabor and Williams, 2010)), and pattern scaling with 8.7% (PS GCM,
496 (Mitchell et al., 2004)) (Figure 4B).

497

498 Uncertainty, as measured by the number of different future scenarios used (combinations of
499 emissions scenarios and climate models) was explored in only 36.5% of the studies.
500 Additionally, the average number of scenarios per study (rounded to the closest integer) was
501 3, indicating that climate uncertainties are barely (if at all) studied in agricultural science and

502 highlighting a knowledge gap in agricultural research, an issue previously raised and
503 discussed by other authors (Challinor et al., 2009b; Challinor and Wheeler, 2008), although
504 some studies addressing this aspect are underway (C. Rosenzweig, personal communication).
505

506 **4.2. Robustness of existing weather station networks**

507 The sensitivities of the network to information loss were found overall to be low.
508 Nevertheless, certain areas, variables and months were found highly sensitive. Agricultural
509 lands (Ramankutty et al., 2008), as visually inspected, are in general less sensitive to data loss
510 than non-agricultural lands. Interpolations' performance varied depending upon the variable,
511 month and parameter used to evaluate them (i.e. R^2 , RMSE, and S), but were consistent,
512 statistically significant ($p < 0.0001$) and with variability (of R^2 , RMSE, and S) between 10–
513 15% in the worst cases. Rainfall presented the lowest R^2 values (Figure 5), particularly in the
514 months of April to August, during which there was a higher variability in the R^2 value and the
515 values reached the absolute minima (0.8). Although it is possible that a high number of
516 weather stations per unit area can improve accuracy, it does not seem to happen in all
517 variables, areas and/or months.

518
519 **<Insert Figure 5 here>**
520

521 The DJF period presented significantly lower variability and more predictive power, probably
522 due to overall low climate variability (Cooper et al., 2008). Interestingly, maximum and
523 minimum temperatures showed different interpolation accuracies, even though they were
524 measured in the same places. Maximum RMSE for temperatures was up to 1.7°C, whilst for
525 precipitation it was up to 100 mm/year, as seen in the evaluation data. The effect of
526 geography and the difficulty of fitting unique and complex landscape features cause errors,
527 leading to high standard deviations in some areas (Figure 6). In the highlands of Eastern
528 Africa, particularly in the states of Benshangul-Gumuz, Addis Ababa and Southern Nations in
529 Ethiopia, the central areas of the Eastern and Coast States in Kenya, and the very centre of
530 Tanzania (i.e. regions of Morogoro, Dodoma and Manyara) between-fold variability was
531 found to be high (above 150 mm/year).

532
533 **<Insert Figure 6 here>**
534

535 Over IGP, the largest variability was found in the coastal areas of Maharashtra, Karnataka
536 and Kerala in India, where rainfall deviation was up to 600 mm/year, and in Nepal (districts
537 of Gorka, Dhawalagiri, and Lumbini), where rainfall variability can go up to 1,000 mm/year,
538 and temperature uncertainties up to 3°C, probably due to the combined effect of a more
539 complex climate in the Himalayas and low weather station density.

541 **4.3. Accuracy of climate model outputs**

542 **4.3.1. Ability to represent mean climate**

543 As expected, the climate models' skill varied on a variable, country and region basis, with
544 certain identifiable patterns (Figure 7, 8). The GCMs represent the observed climatology
545 from weather stations (i.e. WCL-WS, GHCN-CL and GSOD-CL) more poorly than they do

546 interpolated climatology (i.e. WCL-IS, CRU-IS), mainly because GCMs do not account for
547 local-scale variability (Boo et al., 2011). In a broad sense, we found that the more complex
548 the topography, the lower the skill of the GCMs (Gallée et al., 2004; Joubert et al., 1999). We
549 also observed that GCM skill decreased according to the complexity of the variable, with the
550 maximum skill displayed for mean temperatures, followed by temperature range and finally
551 by precipitation. These results agree with those of other studies (Gleckler et al., 2008;
552 Masson and Knutti, 2011; Pierce et al., 2009).

553

554 Annual precipitation fit in IGP and WAF was observed to dip as low as 0 in some cases, with
555 a considerable number of cases (23% for WCL-WS, 27% for GHCN-CL and 63% for GSOD-
556 CL) presenting very low adjustment ($R^2 < 0.5$) (Figure 7). In Mali, Niger, India and
557 Bangladesh, model skill in representing precipitation, compared to weather station
558 measurements, was consistently low, an issue also reported in other studies (Douglass et al.,
559 2008; Gleckler et al., 2008; Reichler and Kim, 2008). The Bergen Climate Model (BCCR-
560 BCM2.0) and the INM-CM3.0 model showed very poor performance ($R^2 < 0.5$) in more than
561 25% of the countries when compared with WCL-WS, GHCN-CL and GSOD-CL, while the
562 climate model GISS-ModelE (Hansen et al., 2007) presented the poorest performance.

563

564

<Insert Figure 7 here>

565

566 When compared with interpolated climatology (i.e. WCL-IS, CRU-IS), annual precipitation
567 R^2 values varied from 0.383 (GISS-ModelE-R in Uganda) to 0.998 (IAP-FGOALS1.0-G in
568 Burkina Faso), whilst for mean temperatures the R^2 varied from 0.195 (GISS-ModelE-R in
569 Nepal) to 0.999 (MIUB-ECHO-G in Burkina Faso), and for temperature range the values
570 were observed between 0.386 (CCCMA-CGCM3.1-T47 in Senegal) to 0.9998 (MPI-
571 ECHAM5 in Burkina Faso) (Figure 7).

572

573

<Insert Figure 8 here>

574

575 In Ethiopia, mean temperature correlations were lower compared to other countries, despite
576 the relative high density of stations in that area (data not shown). In Senegal, diurnal
577 temperature range was found to be very poorly fitted, particularly for the CCCMA models
578 (Figure 8). This result contrasts with that of other studies, which have marked the CCCMA
579 models as the most skilled (Gleckler et al., 2008; Jun et al., 2008). The ability of GCMs to
580 represent mean climate patterns over a year was neither uniform nor consistent (Table 3),
581 with the lowest performance being observed for precipitation in the DJF period (large number
582 of cases with $R^2 < 0.5$, and few cases with $R^2 > 0.8$). Performance for temperature range showed
583 almost no cases with $R^2 < 0.5$, but fewer cases with $R^2 > 0.8$ than for mean temperatures (Table
584 3).

585

586

<Insert Table 3 here>

587

588

4.3.2. Ability to represent interannual variability

589 R square values were above 0.8 in a large number of gridcells (>50%) for all GCMs for both
590 variables (rainfall, mean temperature) (data not shown); however, there were large rainfall
591 biases in GCM outputs (Figure 9, 10), in some cases between 1,000 and 2,500 mm/year,
592 depending on the GCM. These areas were located in Nepal, northern India and EAF. Most of
593 the models' biases were wet-biases (Figure 10) which were found throughout the whole
594 analysis domain, but they were particularly strong over IGP in the models CCCMA-
595 CGCM3.1-T47, CSIRO-Mk3.0 and -Mk3.5, GFDL-CM2.0, all NASA-GISS models, and
596 both UKMO-HadCM3 and -HadGEM1, whereas the opposite signal was observed over the
597 same area for the models MIROC3.2.-HIRES, NCAR-CCSM3.0, INGV-ECHAM4, CNRM-
598 CM3, and GFDL-CM2.1. Over WAF and EAF, almost all GCMs showed a dry-bias, with
599 underestimations of up to 250 mm/year in some cases. Responses varied for seasonal means
600 and totals, with the wet-season (JJA) being more sensitive to wet biases in most GCMs.

601

602 Temperature biases were also large for certain areas. In some cases, annual mean temperature
603 biases were greater than 5°C and were observed to go up to 10°C, particularly in the Sahel
604 and in the areas surrounding the Himalayas and the Tibetan Plateau in Nepal (Figure 11). The
605 most evident temperature biases were found in the NASA-GISS models (GISS-AOM, GISS-
606 ModelE-H and GISS-ModelE-R), and in INM-CM3.0, probably due to their coarse
607 resolution. The quality of higher resolution models was in general better, but geographic
608 trends were difficult to identify, as the locations with mean temperature were scant (7,280
609 locations for the whole study area). The smallest biases were observed in WAF, northern
610 EAF and central India, where temperature biases were below 1.5°C, particularly for the
611 models BCCR-BCM2.0, UKMO-HadCM3, NCAR-PCM1, CCCMA-CGCM3.1-T47 and
612 MIUB-ECHO-G, some of which have been reported to perform well in tropical areas before
613 (Gleckler et al., 2008; Jun et al., 2008). These biases were mostly concentrated in lowlands
614 and were mostly warm-biases, except for UKMO-HadCM3 (Figure 12). Cold-biased models
615 were usually the GISS-NASA models, MIROC3.2-MEDRES, UKMO-HadCM3, IPSL-CM4,
616 MRI-CGCM2.3.2A and IAP-FGOALS1.0-G both for seasons (i.e. JJA, DJF, maps not
617 shown) and for the annual mean (Figure 11, 12).

618

619 **5. Discussion**

620 **5.1. Climate data and agricultural research**

621 Although climate model data (“as is”) are often preferred for impact studies, crop modellers
622 and agricultural scientists should be cautious when developing future adaptation strategies
623 based on crop models applied ~~over-using~~ future predictions of different (and sometimes
624 unknown) nature (Jarvis et al., 2011), given the large uncertainties regarding the agricultural
625 system and plant responses, the underlying uncertainty related to parameterised processes,
626 and the differences in scales, all of which are reported in the impact-assessment literature
627 [e.g. (Challinor and Wheeler, 2008)]. This, however, does not necessarily imply that climate
628 model data cannot or should not be used, but rather means that an adequate treatment of
629 biases needs to be done before climate and crop models can be properly used together
630 (Challinor et al., 2010; Osborne et al., 2007).

631

632 Our findings demonstrate that, for regional assessments where large area process-based crop
633 models, statistical, or empirical models are to be used, products such as WorldClim (Jones
634 and Thornton, 2003; Thornton et al., 2009) and CRU (Challinor et al., 2004) coupled with
635 weather generation routines appear to be the best-bet approach (Challinor et al., 2004; Jones
636 and Thornton, 2003), although climate model data can also be used with proper bias
637 treatment (Challinor et al., 2010; Osborne et al., 2007). However, if studies are to be carried
638 out on a site-specific scale (Parry et al., 2005), weather station data is the best means by
639 which to calibrate the modelling approaches. While partnerships are constantly being built
640 and this allows researchers to share data, Currently global weather station data such as
641 GSOD and GHCN seem to be good options in such cases when no other data is available,
642 particularly when coupled with satellite data or other (country specific) historical weather
643 records (Álvarez-Villa et al., 2010).

644
645 Agricultural research requires high quality and high resolution climatological data to yield
646 accurate results, but to date this has been impossible to achieve at detailed scales and with
647 sufficient coverage, partly due to the difficulty in compiling and revising field data and partly
648 due to the limited climatology knowledge of agricultural researchers (with some exceptions).
649 Large-scale datasets can be matched to certain crop models, mostly when these models can
650 be applied at large scales (Challinor et al., 2010) or do not rely on a detailed calibration of
651 varietal-level crop parameters (Lobell et al., 2011; Lobell et al., 2008). However, matching
652 different modelling scales is not a trivial matter (Baron et al., 2005; Challinor et al., 2009a).
653 Two options are available for matching two differing scales:

- 654 (1) Decreasing the resolution of the crop model from plot scale to large regions, at the
655 expense of loss of detail in some processes [see (Challinor et al., 2007b; Challinor et
656 al., 2004; Yao et al., 2007)], or
- 657 (2) Disaggregating the coarse-resolution climate data, at the expense of introducing noise
658 and possibly propagating uncertainties present in the original climate model data
659 (Tabor and Williams, 2010).

660
661 These two choices yield different results that need to be assessed and coupled. Climate data
662 can be aggregated up to any scale to match any intended use (Masson and Knutti, 2011), but
663 agricultural impacts need to be informed at an scale such that information can be used for
664 decision making and adaptation (Jarvis et al., 2011). Hence, governments and international
665 agencies should support common platforms through which data can be shared without
666 restrictions between members of the research community. Best-bet methods can then be
667 applied over such data to produce useable datasets that can be further shared, used and
668 assessed in multidisciplinary and transdisciplinary approaches.

669 670 **5.2. Robustness of existing weather station network**

671 It is tacitly acknowledged that the use of interpolated surfaces can lead to errors and biases
672 when these data are used for impact assessment (A. Jarvis, pers. comm.). However, we have
673 demonstrated here that the effects on uncertainty are actually rather low in most of the cases,
674 with very few exceptions (highlands of Ethiopia, the Himalayas, and some parts of the Sahara
675 and Southern Africa, **Figure 6**).

676

677 The results of this research suggest that, despite weather station density being important, it
678 may not be the only determining factor for a good ability to fill information gaps (Hijmans et
679 al., 2005). Based on our results, we suggest that, in selecting locations to measure weather,
680 the following factors be taken into account: (1) the nature of the variable (e.g. precipitation
681 might be much more difficult to monitor than temperature), (2) the area where it is measured
682 (topographically complex areas are much more variable), (3) the values of the variable in the
683 areas where it is measured (high values are subjected to larger absolute errors, assuming
684 relative errors are relatively uniform), (4) the relevance of the area for different subjects (i.e.
685 the Sahara might be irrelevant for agriculture but can be of high relevance for other fields
686 such as climate science, ecology or biodiversity and conservation), (5) possible errors in
687 measurements and other underlying factors that can influence the measurability or
688 correctness of estimates of a particular variable, and (6) possible political or social constraints
689 on access to the site. Improving weather station distribution and status, as well as improving
690 the cross-checking, correction and validation of data collected at the different sites, is
691 fundamental for improving climate data for agricultural impact assessment.

692

693 **5.3.Global climate model accuracy and performance**

694 **5.3.1. CMIP3 climate model skill**

695 GCM performance is highly reliant on the type of comparisons performed, on the GCM
696 formulation and on the nature of climate conditions in the analysed areas (Gleckler et al.,
697 2008; Masson and Knutti, 2011). Underlying factors driving GCM performance are indeed
698 difficult to track, given the complexity of the models. IPCC 4AR (CMIP3) models showed
699 varied performance, with a high tendency to being wet-biased and no general trend for
700 temperature. These responses reportedly have their origin in different factors: first, some
701 GCMs have weak forcing on sea surface temperatures (SSTs), whereas climate in Africa and
702 Asia is strongly coupled with the Atlantic and Indian Ocean and with inland water bodies
703 (Gallée et al., 2004; Lebel et al., 2000); second, models do not properly account for the
704 relation between inter-annual variability, ENSO and the monsoonal winds (Gallée et al.,
705 2004; Hulme et al., 2001); third, the resolution of the models prevents acknowledgement of
706 local-scale land use, orographic patterns and small water bodies (Hudson and Jones, 2002);
707 fourth, cloud thickness and latent heat and moisture flux between clouds has not been
708 properly resolved in the models (Gallée et al., 2004); and fifth, convective parameterisations
709 produce an early onset of the seasonal rains and over-prediction of wet days and high-rainfall
710 events (Gallée et al., 2004).

711

712 The NASA models GISS-ModelE (-R and -H) consistently presented very low predictive
713 ability, mainly because of the models' coarse spatial resolution in conjunction with the
714 reasons mentioned above (Hansen et al., 2007). These results agree with those of Gleckler et
715 al. (2008), who reported that NCAR-PCM1, GISS-ModelE (-R and -H) and GISS-AOM
716 models are the worst performing in the 24 GCMs of the CMIP3 ensemble. Similar results are
717 reported by other authors that have assessed this or similar model ensembles (Jun et al., 2008;
718 Pierce et al., 2009). Lack of detail in land use and land use changes (Eltahir and Gong, 1996),
719 monsoon winds (Eltahir and Gong, 1996; Gallée et al., 2004), and sea surface temperature

720 anomalies (SSTs) of the Atlantic and the Indian Oceans (Lebel et al., 2000; Sun et al., 1999)
721 also causes the scales at which climate model information is robust to be varied (Masson and
722 Knutti, 2011), and prevents local scale seasonal weather patterns from being modelled
723 consistently (Douglass et al., 2008; Hansen et al., 2007).
724

725 **5.3.2. Plugging climate model data into agricultural research**

726 GCMs do not provide realistic representations of climate conditions in a particular site, but
727 rather provide estimated conditions for a large area. Our results, in agreement with those from
728 the agricultural community (Baron et al., 2005; Challinor et al., 2003) and the climate
729 community (Jun et al., 2008; Masson and Knutti, 2011), indicate that climate model outputs
730 cannot be input directly into plot-scale (agricultural) models, but support the idea that higher
731 resolution climate modelling largely improves results. Either the CMIP3 (assessed here) or
732 the upcoming CMIP5 (being released at the moment) (Moss et al., 2010) climate model
733 outputs can be adequately used in agricultural modelling if: (1) the scales between the models
734 are matched (see Sect. 5.1), (2) skill of models is assessed and ways to create robust model
735 ensembles are defined, (3) uncertainty and model spread are quantified in a robust way, and
736 (4) decision making in the context of uncertainty is fully understood.
737

738 Producing robust (i.e. skilled and certain) ensembles for agriculture is not an easy task,
739 mainly because of the scales at which these have been found to be robust (Masson and
740 Knutti, 2011). Opinions are contrasting: some authors support sub-selecting models based
741 upon performance under present conditions (Matsueda and Palmer, 2011; Pierce et al., 2009),
742 calculating a mean ensemble by weighting models based on skill (Matsueda and Palmer,
743 2011; Walsh et al., 2008), while others advocate using ~~as many as all~~ available models with
744 no-weighting at all (Reifen and Toumi, 2009). We suggest that until sensitivities of
745 agricultural models to ensemble spread are fully explored (Baigorria et al., 2007), the full
746 CMIP3 (or CMIP5) ensembles should be used.
747

748 Strategies for combining plot-scale and large-scale models and for optimising the overall
749 result (including estimation of uncertainties derived from the scale-matching process) need to
750 be further researched. The potential of high-quality and less uncertain climate predictions of
751 current and future climate conditions for agricultural research is expected to have a direct
752 impact on decision-making at different levels and for different purposes: to improve yields on
753 the farm, to direct country level policies and investment, to define research foci, to direct
754 international agencies' investments, and to clarify global greenhouse emissions limits and
755 commitments (Challinor et al., 2009a; Funke and Paetz, 2011; IPCC, 2007).
756

757 **6. Conclusions**

758 A thorough analysis of different aspects of climate data for agricultural applications was
759 performed. All topics addressed here are of high relevance to agricultural applications,
760 particularly in the global tropics. Several important points were raised: (1) spatial scale is the
761 most important issue for agricultural researchers, as they preferred to use monthly products
762 with higher resolution rather than daily products with very low spatial resolution, or else
763 limited their areas of study to field plots; (2) the sensitivities of Sub-Saharan African and

764 Southeast Asian climate to data loss and poor availability were found to not be limiting
765 factors for the region, with the exceptions of mountainous areas in Nepal and Ethiopia; and
766 (3) climate modelling, although constantly improving and useful, still requires considerable
767 future development.

768

769 As such, CMIP3 GCMs can be used with a certain degree of confidence to represent large-
770 area climate conditions for some areas and periods. In areas where predictions lack enough
771 skill for agricultural modelling, models can be bias-corrected using different methods [see
772 (Challinor et al., 2009a; Hawkins et al., 2011; Reifen and Toumi, 2009)]. Whilst model skill
773 is expected to improve with the upcoming IPCC Fifth Assessment Report, climate model
774 ensembles as well as different methods for ‘calibrating’ (i.e. pre-processing for input into
775 crop models) climate model data both need to be used, as uncertainties go beyond those
776 derived from emissions scenarios (Hawkins et al., 2011). The proper usage of climate
777 projections for agricultural impact assessment is of paramount importance in order to
778 properly inform adaptation.

779

780 Finally, it is critical to understand the implications of all this to agriculture. Crops are
781 sensitive to shortages in water and heat stresses during key periods during their development
782 (i.e. flowering, fruit filling). Therefore, lack of skill in representing seasonal and inter-annual
783 variability is expected to produce a significant obstacle to agricultural impact assessment of
784 climate change; several examples in the literature exist that illustrate this (Baigorria et al.,
785 2008; Baigorria et al., 2007). The importance of this factor depends on the strength of the
786 climate signal on yields and the variables that drive ~~such~~-this signal. Future impact
787 assessments need to take into account input data and climate model data inaccuracies,
788 sensitivities and uncertainties; make their own assessments of the inaccuracies and
789 uncertainties; and comprehensively quantify and report uncertainties in the impact assessment
790 process.

791

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800

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1096 **Figure captions**

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1098 **Figure 1** Cascade of constraints to climate data, as normally observed in agricultural impact
1099 assessment

1100

1101 **Figure 2** Areas of study. Bold-outlined areas indicate the areas on which the study focused
1102 (SN: Senegal, ML: Mali, NE: Niger, BF: Burkina Faso, GH: Ghana, UG: Uganda, ET:
1103 Ethiopia, KE: Kenya, TZ: Tanzania, NP: Nepal, BD: Bangladesh, IN: India)

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1105 **Figure 3** Topics treated in the analysed agricultural studies. WG: weather generators.

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1107 **Figure 4** Frequency of use of the different data sources in agricultural studies. A. Present-day
1108 climates. B. Future climates. Datasets acronyms are as follows: CRU-TS: Climatic Research
1109 Unit monthly time series product at 0.5 degree, GCM: global climate model output, RCM:
1110 regional climate model, CRU-CL: CRU monthly climatology product at 10 arc-minute,
1111 MARS: Data from the MARS European project, GSOD: Global summary of the day,
1112 ARTES: Africa rainfall and temperature evaluation system, VEMAP: United States
1113 comprehensive dataset, ATEAM: Advanced Terrestrial Ecosystem Analysis and Modelling,
1114 PRISM: United States dataset, GPCP: Global Precipitation Climatology Project, GPCC:
1115 Global Precipitation Climatology Centre, GHCN: Global Historical Climatology Network, AI
1116 GCM: GCM data “as is”, SD GCM: statistically downscaled GCM, PS GCM: pattern scaled
1117 GCM, WG GCM: GCM data through a weather generator, SC Variables: systematic changes
1118 in target key variables, Unclear: not specified clearly in study, ARPEGE: the ARPEGE
1119 Atmospheric GCM (Déqué et al., 1994).

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1121 **Figure 5** Performance of the interpolations for all variables and months as measured by the
1122 R-square value. A. Rainfall, B. Mean temperature, C. Maximum temperature, D. Minimum
1123 temperature

1124

1125 **Figure 6** Uncertainties in WorldClim expressed as standard deviations from the mean of the
1126 100 cross-validated folds for (A) total annual rainfall (in mm), and (B) annual mean
1127 temperature (in °C).

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1129 **Figure 7** Comparison (R-square based) of observed climatology (CL-WS [w], GHCN-CL [g]
1130 and GSOD-CL [o]) and each of the GCMs (GCM-CL) for each of the countries in the study
1131 area for mean temperature (top), temperature range (middle) and precipitation (bottom), for
1132 the annual and two seasonal (DJF, JJA) means or totals. All R^2 values were statistically
1133 significant at $p < 0.0001$

1134

1135 **Figure 8** Comparison (R-square based) of interpolated climatology (i.e. CRU-IS [c], WCL-IS
1136 [w]), and each of the GCMs (GCM-CL) for each of the countries in the study area for mean
1137 temperature (top), temperature range (middle) and precipitation (bottom) for the annual mean
1138 or total and two seasons (DJF, JJA). All R^2 values were statistically significant at $p < 0.001$.

1139

1140 **Figure 9** Root mean squared error (RMSE), in millimetres, between observed (GHCN-TS)
1141 and GCM (GCM-TS) time series, for the 24 GCMs in Table 2, for annual total rainfall
1142 between the years 1961-1990.

1143

1144 **Figure 10** Mean bias of GCM (GCM-TS) time series compared to observed time series
1145 (GHCN-TS), for the 24 GCMs in Table 2, for annual total rainfall between the years 1961-
1146 1990. Bias is expressed as the slope of the regression curve between observed and climate-
1147 model series. Values below 1 (light grey areas) indicate that GCMs are wet-biased, whereas
1148 values above 1 (dark grey areas) indicate that GCMs are dry-biased.

1149

1150 **Figure 11** Root mean squared error (RMSE), in Celsius degree, between observed (GHCN-
1151 TS) and GCM (GCM-TS) time series, for the 24 GCMs in Table 2, for annual mean
1152 temperature between the years 1961-1990

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1154 **Figure 12** Mean bias of GCM (GCM-TS) time series compared to observed time series
1155 (GHCN-TS), for the 24 GCMs in Table 2, for annual mean temperature between the years
1156 1961-1990. Bias is expressed as the slope of the regression curve between observed and
1157 climate-model series. Values below 1 (light grey areas) indicate that GCMs are warm-biased,
1158 whereas values above 1 (dark grey areas) indicate that GCMs are cold-biased.

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Table 1 Number of locations per data source ([global](#))

Source*	Precipitation stations	Mean temperature stations	Min., Max. temperature stations	Period
GHCN v2	20,590	7,280	4,966	1950-2000
WMO CLINO	4,261	3,084	2,504	1961-1990
FAOCLIM 2.0	27,372	20,825	11,543	1960-1990
CIAT	18,895	13,842	5,321	1950-2000

1168 *GHCN v2: Global Historical Climatology Network version 2 (Peterson and Vose, 1997);
1169 WMO CLINO: World Meteorological Organization Climatology Normals; FAOCLIM 2.0:
1170 Food and Agriculture Organization of the United Nations Agro-Climatic database (FAO,
1171 2001); CIAT: Database assembled by Peter J. Jones at the International Center for Tropical
1172 Agriculture (CIAT).
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Table 2 Available GCMs, resolutions, and main references

Model	Country	Atmosphere	Ocean	Reference
BCCR-BCM2.0	Norway	T63, L31	1.5x0.5, L35	(Furevik et al., 2003)
CCCMA-CGCM3.1 (T47)	Canada	T47 (3.75x3.75), L31	1.85x1.85, L29	(Scinocca et al., 2008)
CCCMA-CGCM3.1 (T63)	Canada	T63 (2.8x2.8), L31	1.4x0.94, L29	(Scinocca et al., 2008)
CNRM-CM3	France	T63 (2.8x2.8), L45	1.875x(0.5-2), L31	(Salas-Méla et al., 2005)
CSIRO-Mk3.0	Australia	T63, L18	1.875x0.84, L31	(Gordon et al., 2002)
CSIRO-Mk3.5	Australia	T63, L18	1.875x0.84, L31	(Gordon et al., 2002)
GFDL-CM2.0	USA	2.5x2.0, L24	1.0x(1/3-1), L50	(Delworth et al., 2006)
GFDL-CM2.1	USA	2.5x2.0, L24	1.0x(1/3-1), L50	(Delworth et al., 2006)
GISS-AOM	USA	4x3, L12	4x3, L16	(Russell et al., 1995)
GISS-MODEL-EH	USA	5x4, L20	5x4, L13	(Schmidt et al., 2006)
GISS-MODEL-ER	USA	5x4, L20	5x4, L13	(Schmidt et al., 2006)
IAP-FGOALS1.0-G	China	2.8x2.8, L26	1x1, L16	(Yongqiang et al., 2004)
INGV-ECHAM4	Italy	T42, L19	2x(0.5-2), L31	(Gualdi et al., 2008)
INM-CM3.0	Russia	5x4, L21	2.5x2, L33	(Diansky and Zalensky, 2002)
IPSL-CM4	France	2.5x3.75, L19	2x(1-2), L30	(Marti et al., 2005)
MIROC3.2-HIRES	Japan	T106, L56	0.28x0.19, L47	(Hasumi and Emori, 2004)
MIROC3.2-MEDRES	Japan	T42, L20	1.4x(0.5-1.4), L43	(Hasumi and Emori, 2004)
MIUB-ECHO-G	Germany/Korea	T30, L19	T42, L20	(Grötzner et al., 1996)
MPI-ECHAM5	Germany	T63, L32	1x1, L41	(Jungclaus et al., 2006)
MRI-CGCM2.3.2A	Japan	T42, L30	2.5x(0.5-2.0)	(Yukimoto et al., 2001)
NCAR-CCSM3.0	USA	T85L26, 1.4x1.4	1x(0.27-1), L40	(Collins et al., 2006)
NCAR-PCM1	USA	T42 (2.8x2.8), L18	1x(0.27-1), L40	(Washington et al., 2000)
UKMO-HADCM3	UK	3.75x2.5, L19	1.25x1.25, L20	(Gordon et al., 2000)
UKMO-HADGEM1	UK	1.875x1.25, L38	1.25x1.25, L20	(Johns et al., 2006)

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Table 3 Summarised performance of all GCMs with available data for each of the variables and periods in the study countries for different ranges of the R^2 skill evaluation parameter.

Variable	Period	Dataset*	$R^2 < 0.5$ (%)*	$0.5 < R^2 < 0.7$ (%)*	$R^2 > 0.8$ (%)*	$R^2 > 0.9$ (%)*
Rainfall	Annual	IS	2.8	6.6	77.8	54.3
		WS	37.5	19.4	30.8	17.0
		ALL	23.6	14.3	49.6	31.9
	DJF	IS	17.7	19.3	49.1	25.9
		WS	38.1	17.2	31.4	15.7
		ALL	29.9	18.1	38.5	19.8
	JJA	IS	12.8	17.2	58.9	40.1
		WS	15.2	19.1	52.1	34.5
		ALL	14.2	18.3	54.8	36.7
Diurnal temperature range	Annual	IS	0.4	2.2	81.8	73.1
		WS	0.4	1.2	54.5	46.1
		ALL	0.4	1.7	68.1	59.6
	DJF	IS	0.4	2.2	80.4	71.2
		WS	0.4	2.4	53.1	47.7
		ALL	0.4	2.3	66.8	59.4
	JJA	IS	0.4	2.0	80.7	67.2
		WS	0.4	1.2	54.5	46.1
		ALL	0.4	1.6	67.6	56.6
Mean temperature	Annual	IS	0.7	1.2	96.4	95.7
		WS	2.4	1.9	93.5	91.0
		ALL	1.7	1.6	94.7	92.8
	DJF	IS	3.5	1.9	93.2	91.5
		WS	2.3	2.3	93.9	91.2
		ALL	2.8	2.2	93.6	91.3
	JJA	IS	0.0	0.0	100.0	98.8
		WS	0.0	0.1	99.8	98.5
		ALL	0.0	0.1	99.9	98.6

1185 * Values are expressed as percent of country-GCM combinations for comparisons of GCM-
 1186 CL and different observational datasets: interpolated surfaces (IS), namely, WCL-IS and
 1187 CRU-IS; weather stations (WS), namely, GHCN-CL, WCL-WS, GSOD-CL; and as the
 1188 average of IS and WS (ALL)

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Figure 1
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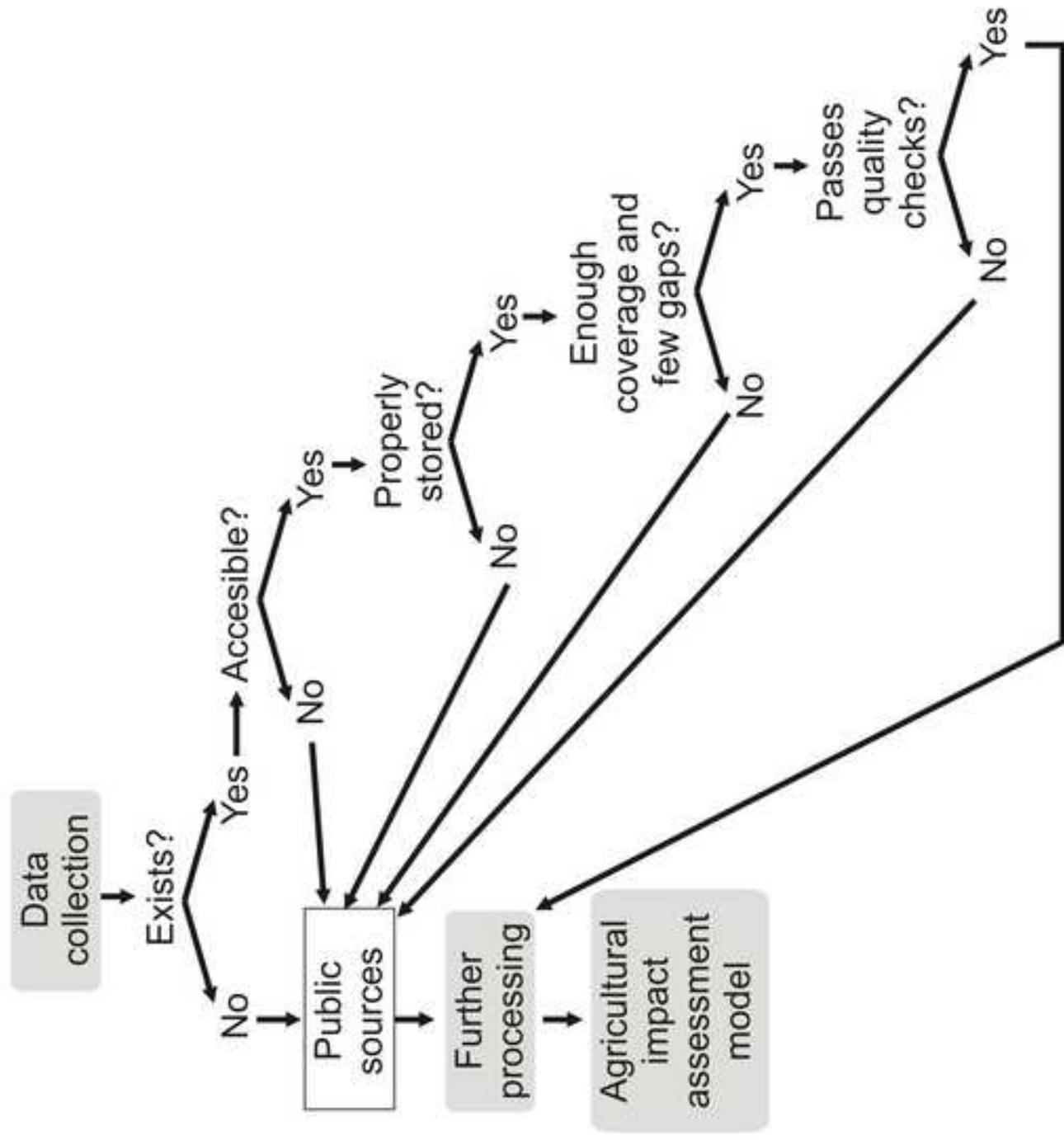


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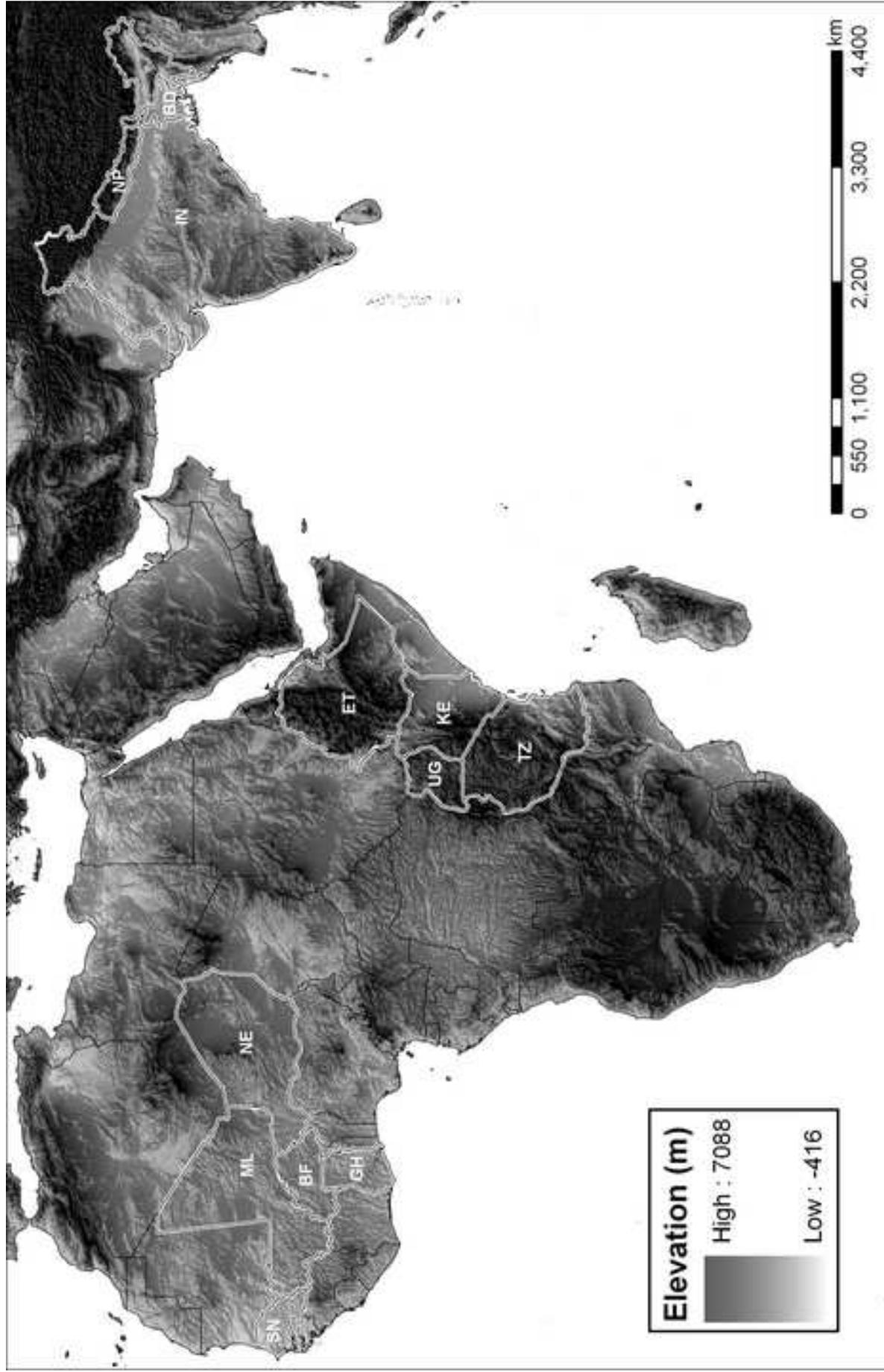


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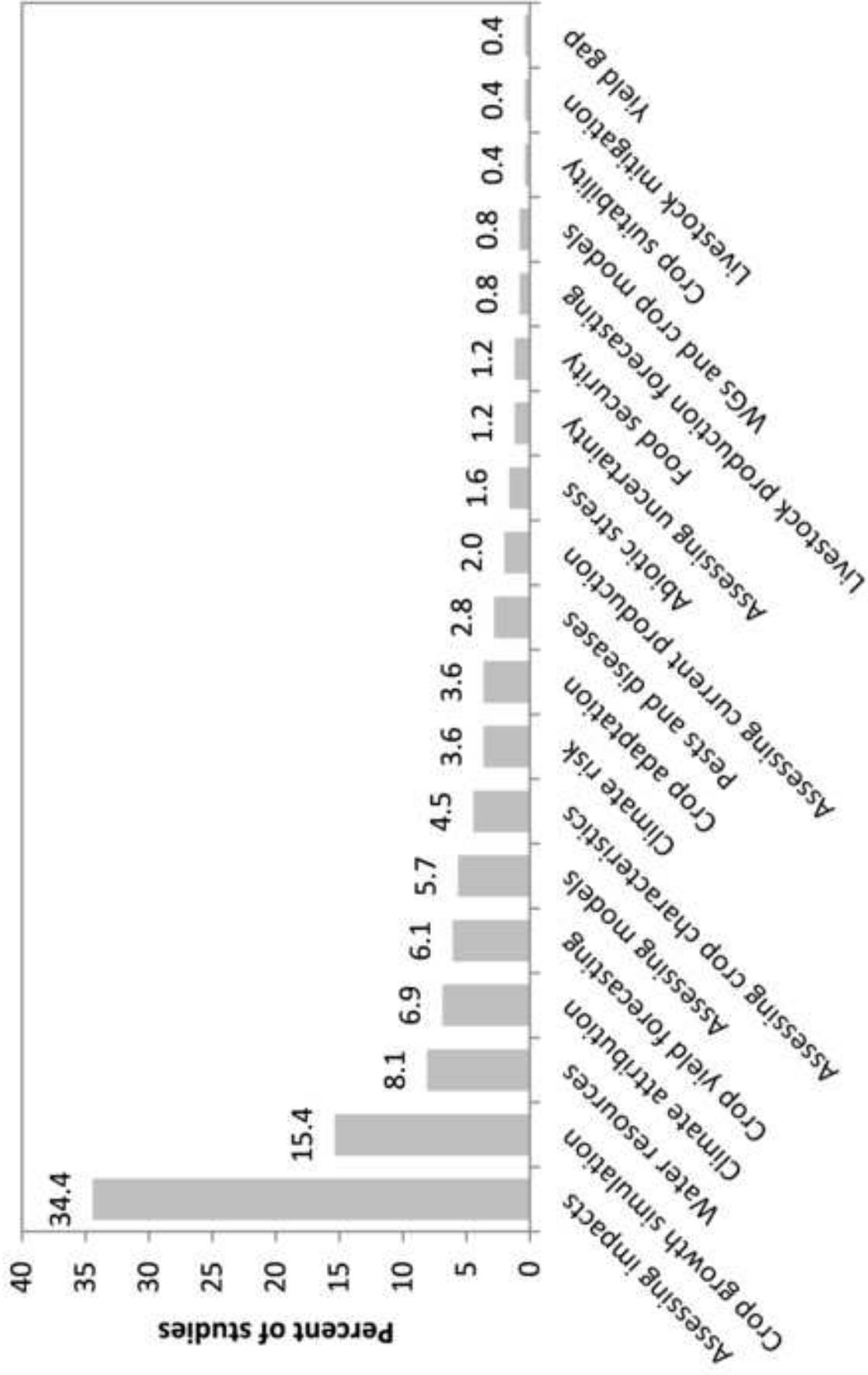


Figure 4a

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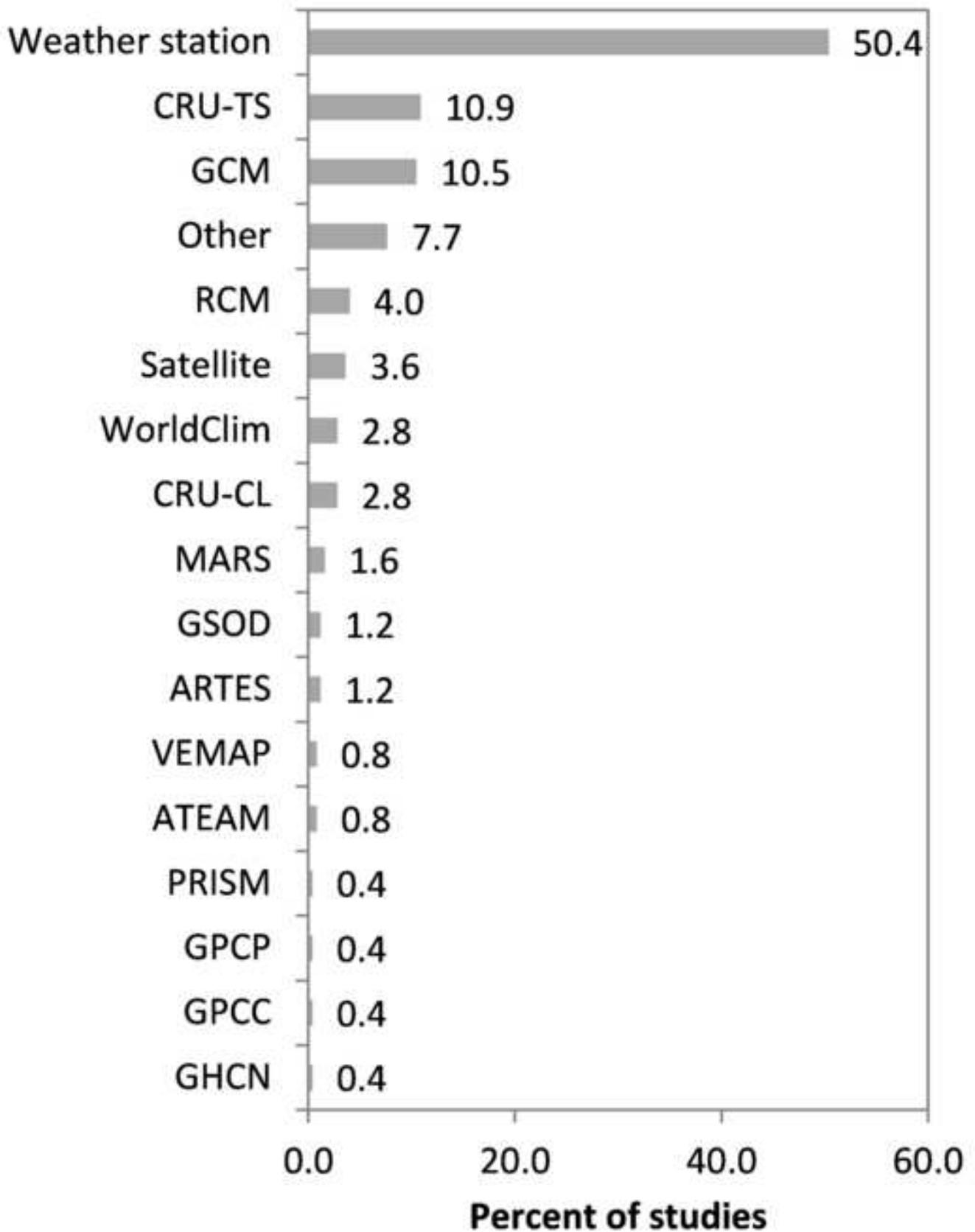


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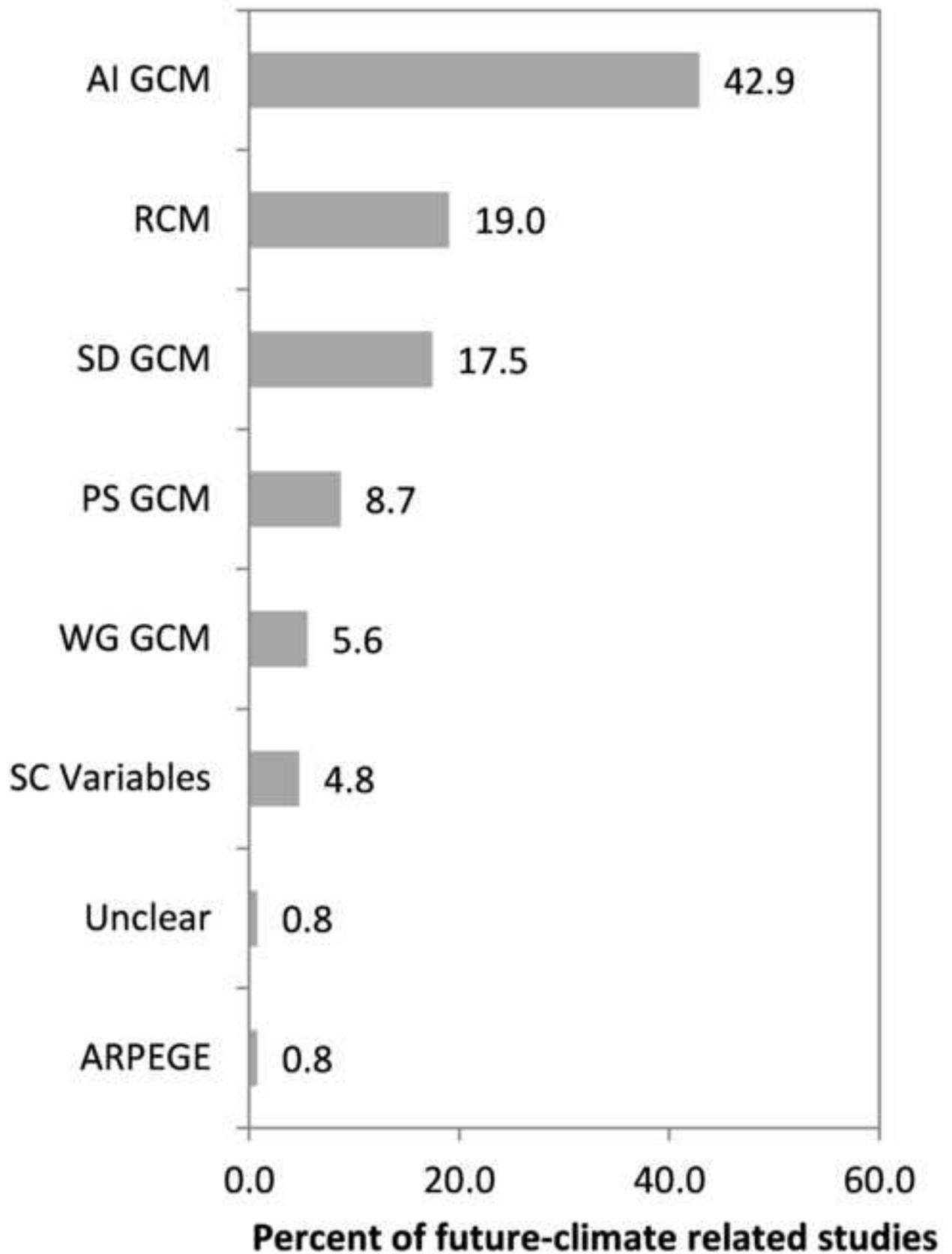


Figure 5a

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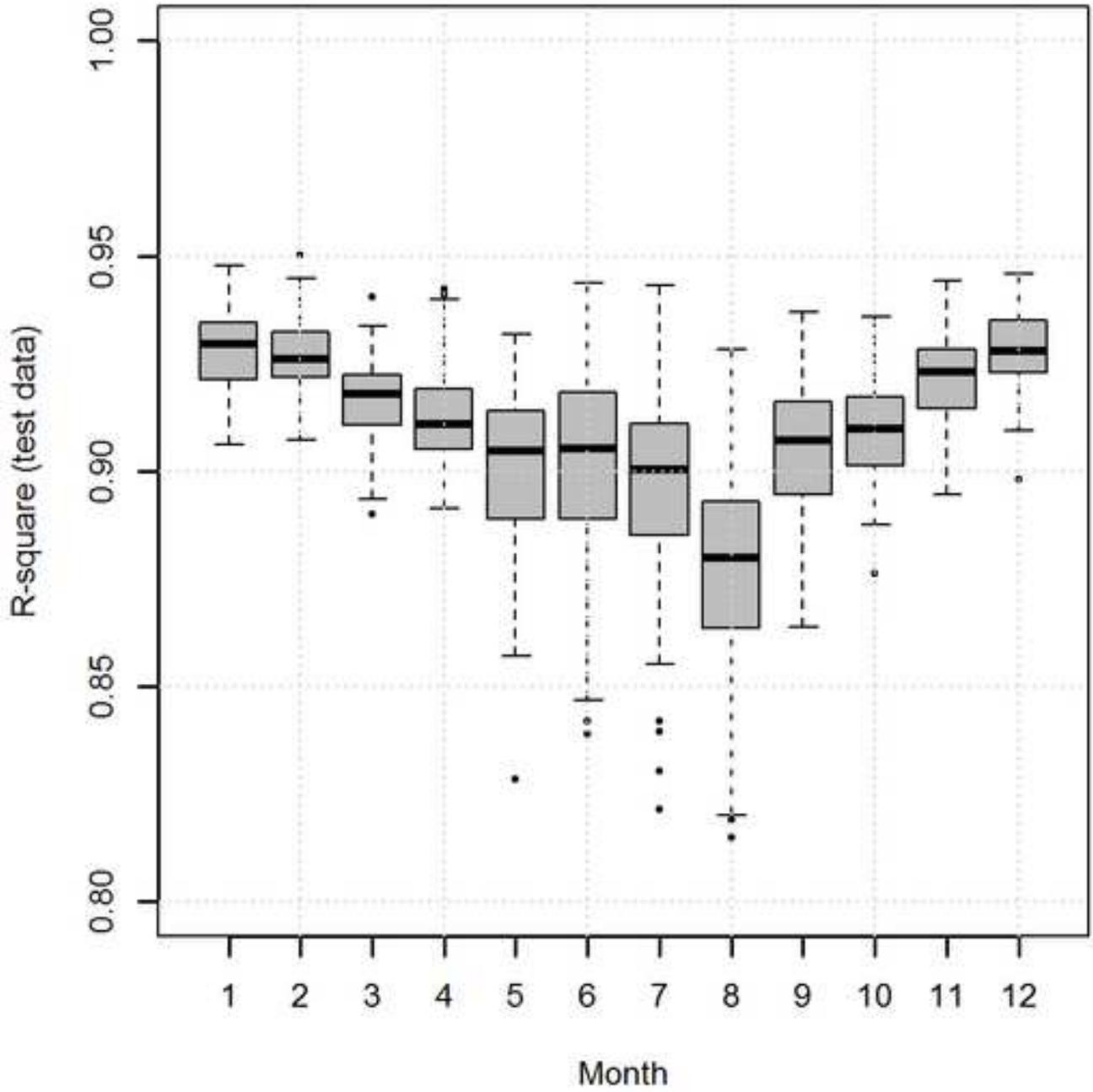


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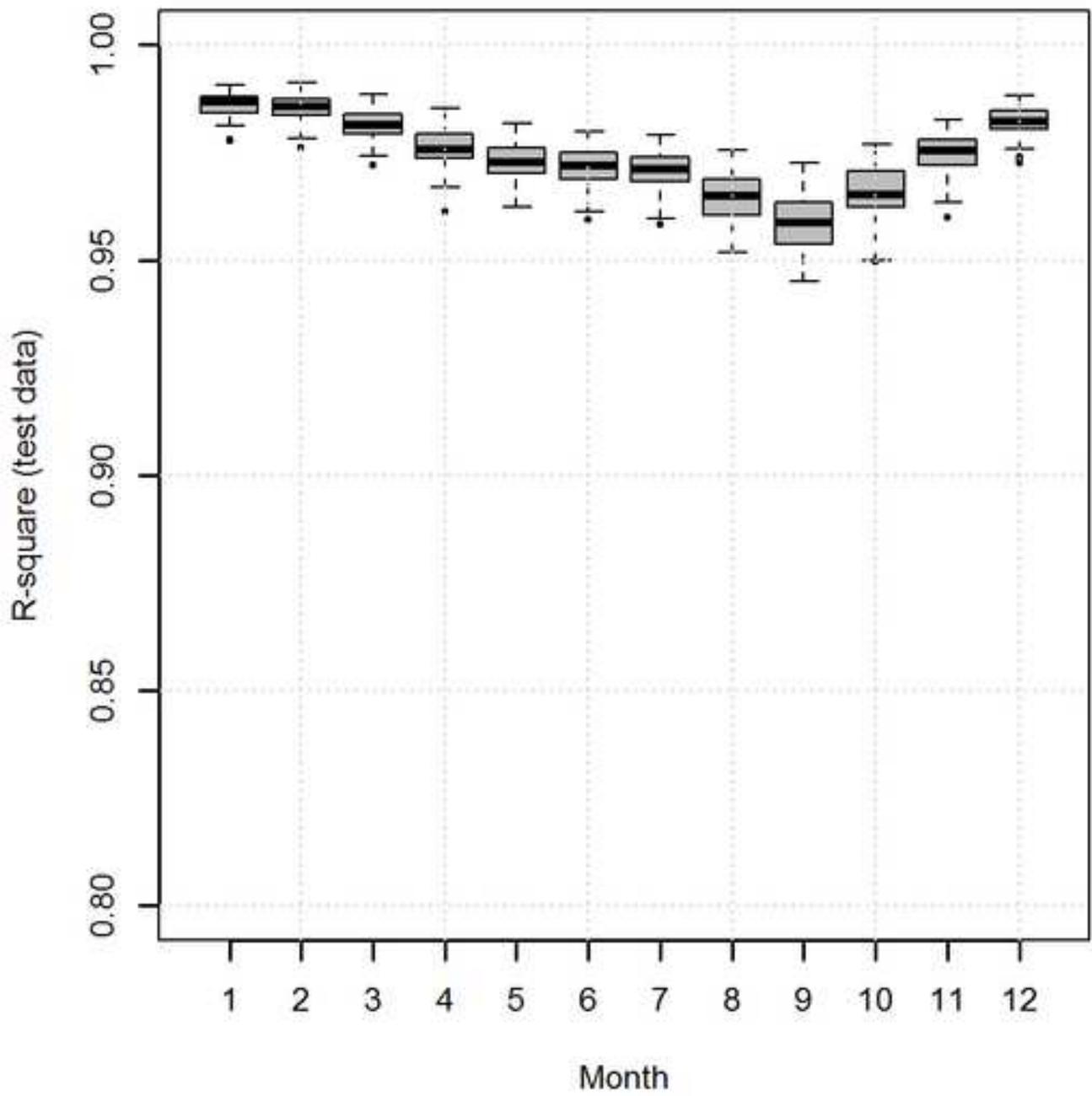


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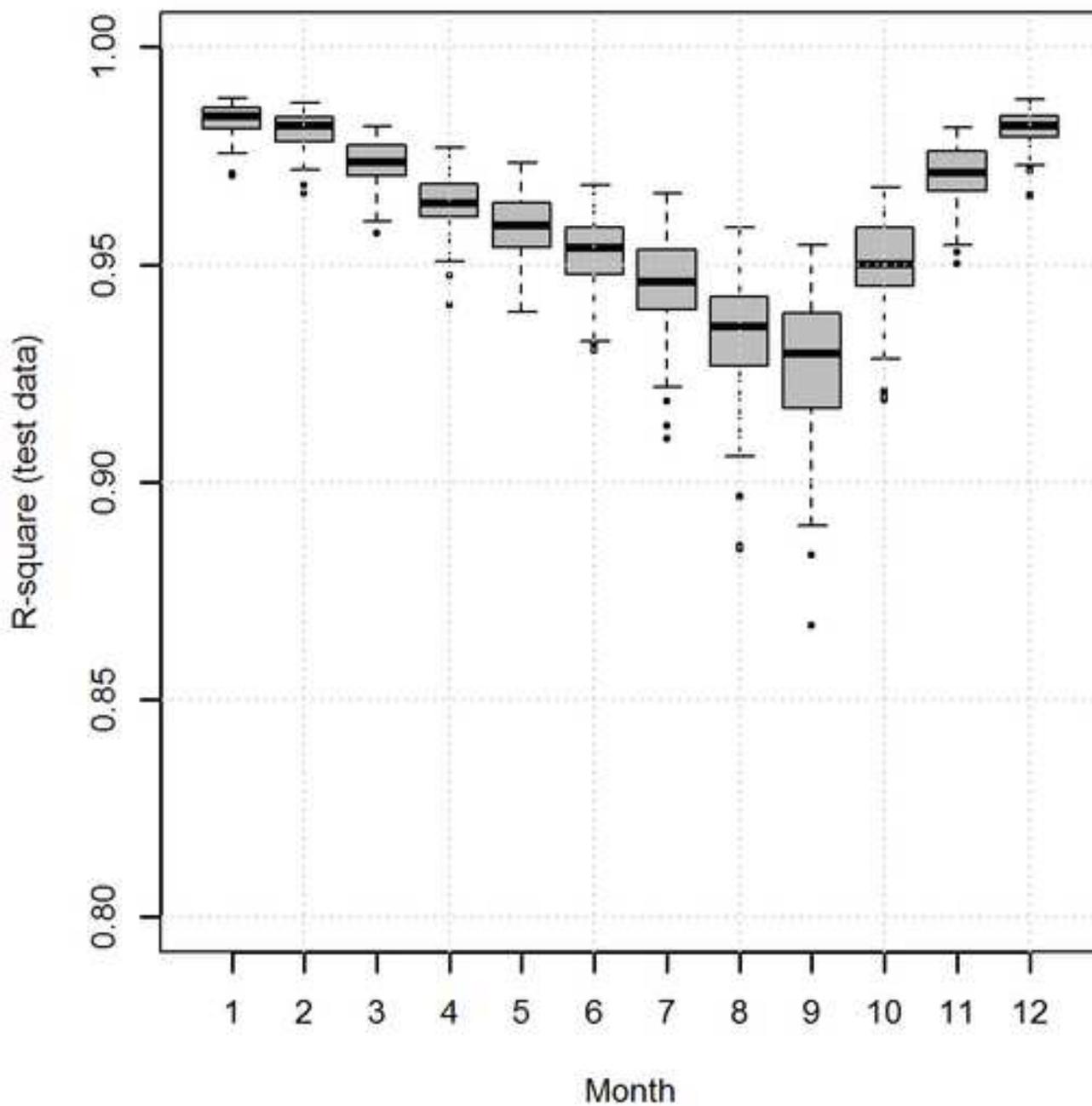


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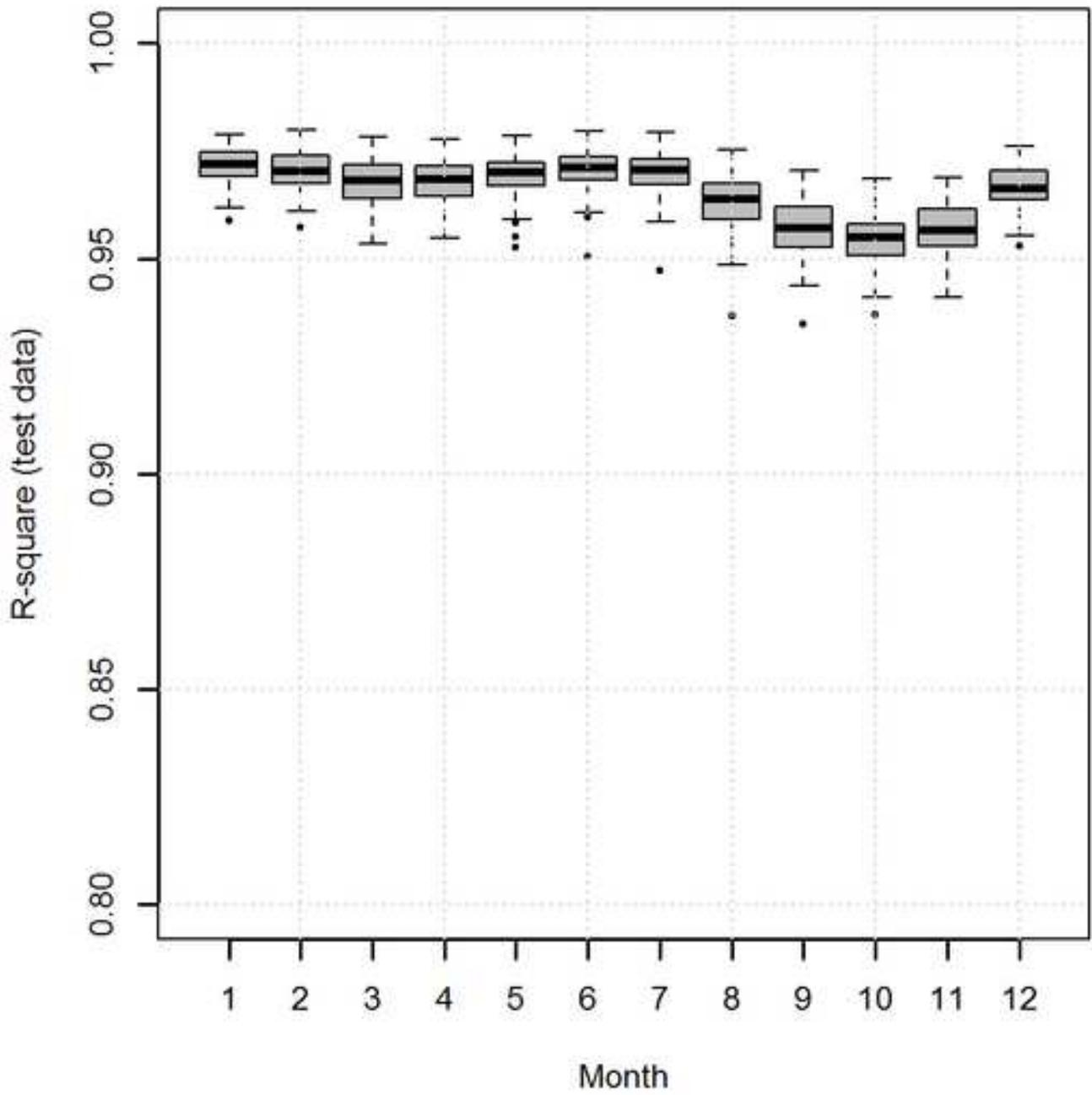


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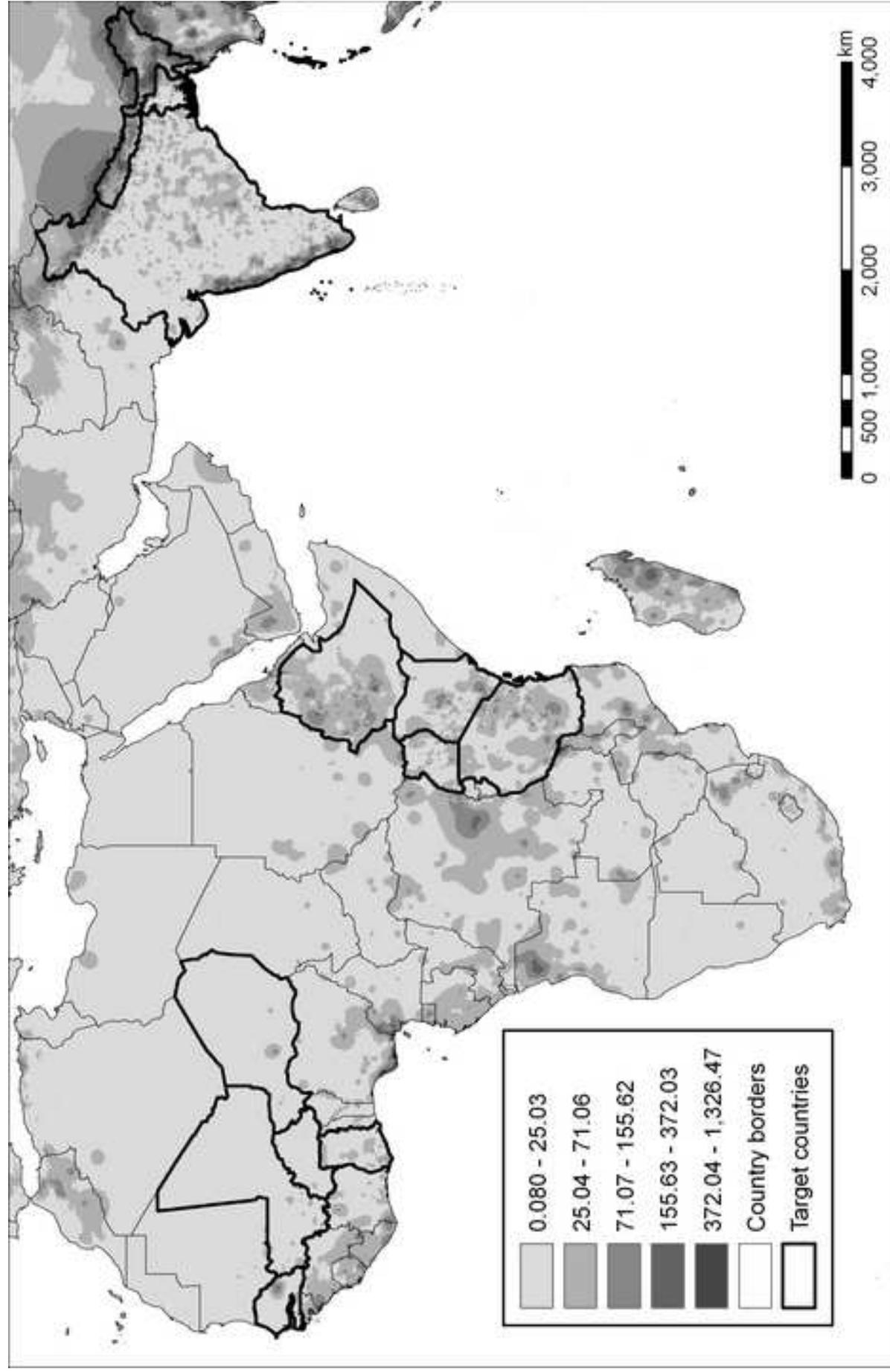


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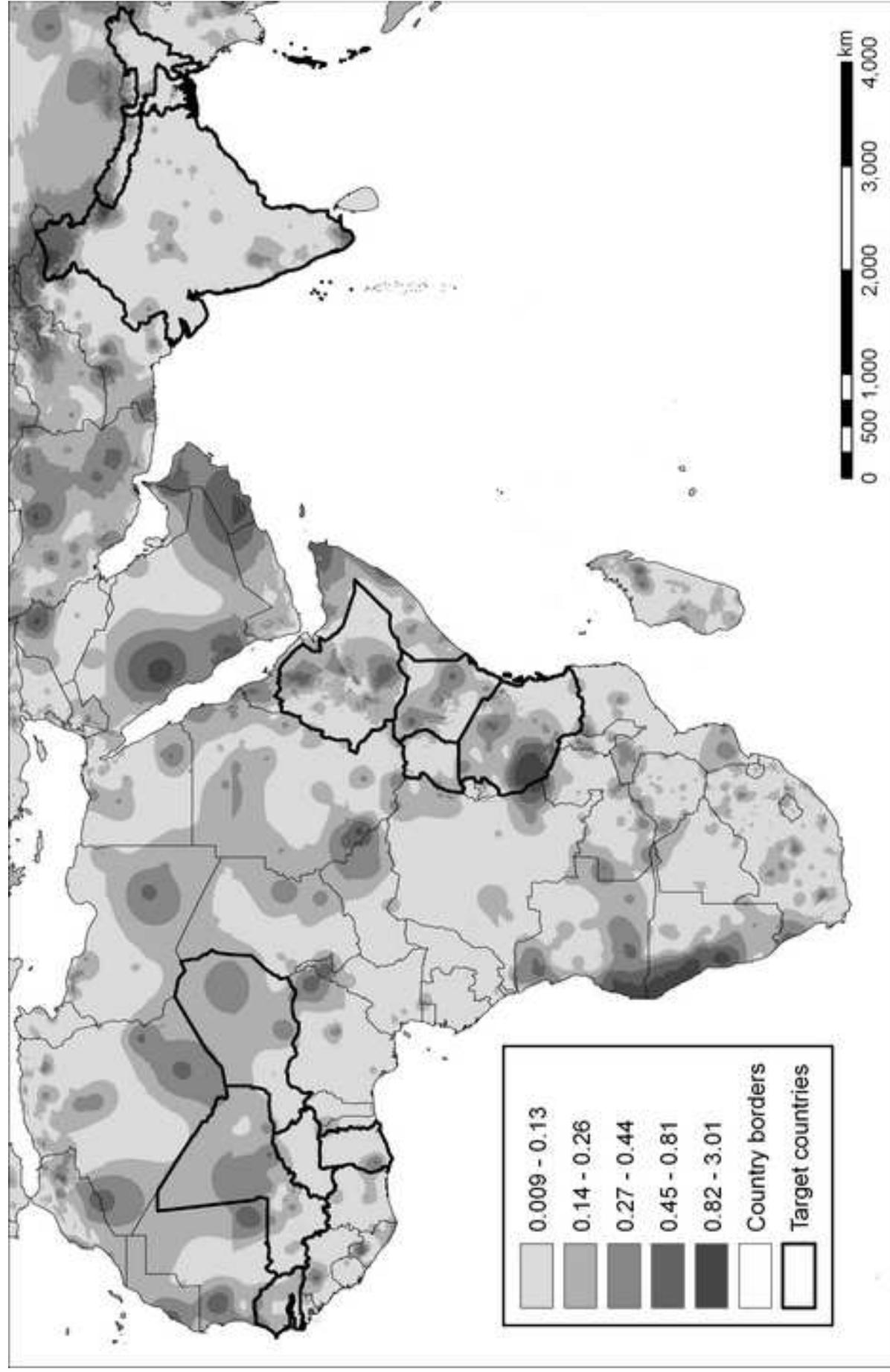


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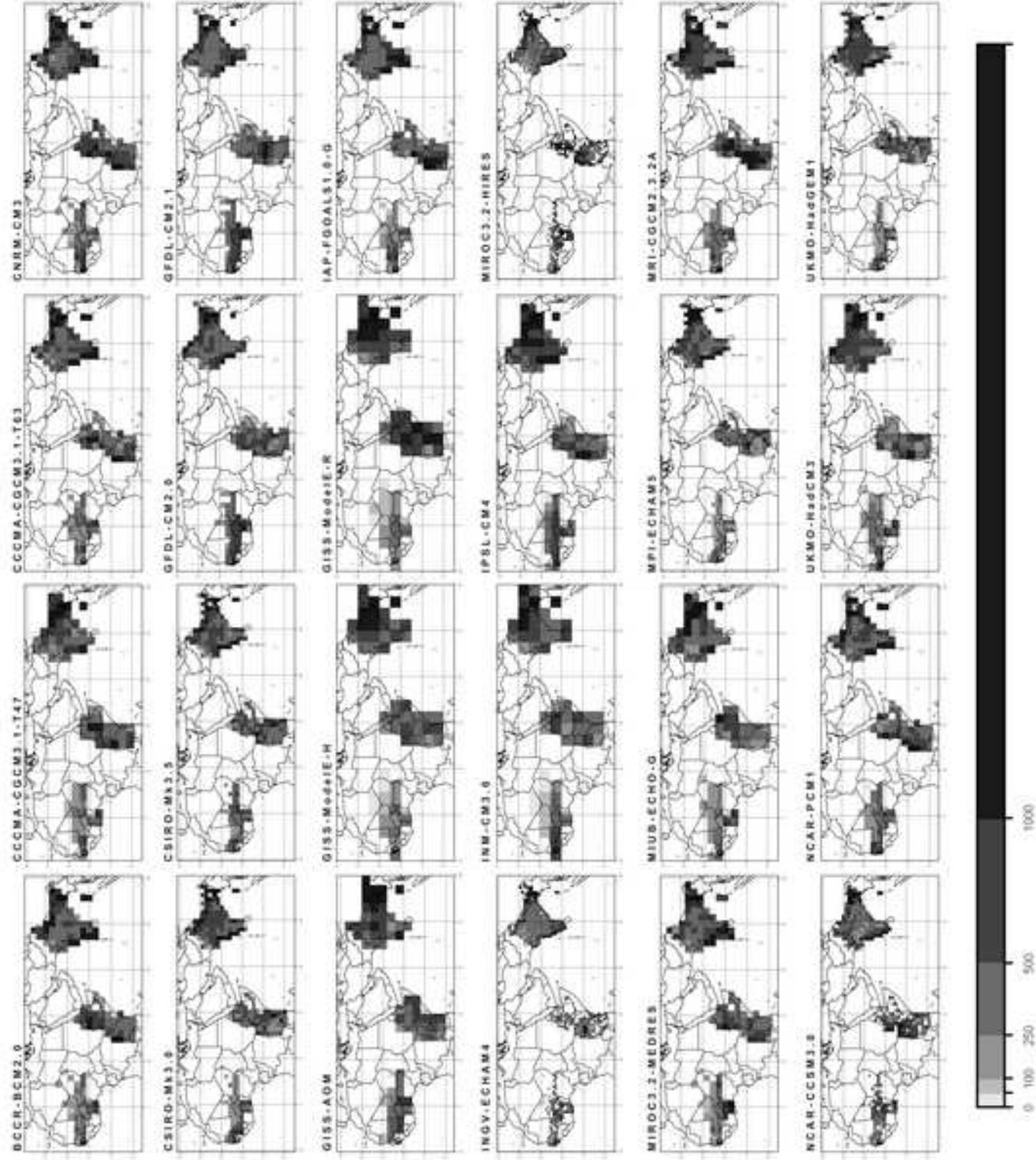


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