Constitutive modelling of Hydro-Mechanical coupling in double porosity 1 2 media based on Mixture Coupling Theory Yue Ma¹, Jun Feng², Shangqi Ge^{3, *}, Kai Wang⁴, Xiaohui Chen⁵, Aizhong Ding⁶ 3 4 1 School of Civil Engineering, University of Leeds, Leeds, LS2 9JT, UK. Email: cnym@leeds.ac.uk 5 2 School of airport engineering, Civil Aviation Flight University of China, Guanghan, 618307, 6 CHINA. Email: sckid1987@163.com 7 3 Research Center of Coastal and Urban Geotechnical Engineering, Zhejiang University, Hangzhou, 8 310058, CHINA; Geomodelling and Artificial Intelligence Centre, School of Civil Engineering, 9 University of Leeds, Leeds, LS2 9JT, UK. Email: geshangqi@zju.edu.cn 10 4 College of Water Sciences, Beijing Normal University, Beijing, 100875, CHINA. Email: 11 wangkaik@mail.bnu.edu.cn 12 5 School of Civil Engineering, University of Leeds, Leeds, LS2 9JT, UK. Email: x.chen@leeds.ac.uk 13 6 College of Water Sciences, Beijing Normal University, Beijing, 100875, CHINA. Email: 14 ading@bnu.edu.cn 15 * Corresponding author: Shangqi Ge (geshangqi@zju.edu.cn) 16 Abstract: Modelling of fluids in deformable geoformation media has gained great attention in the 17 past decades due to significant applications such as groundwater prediction, shale gas and carbon 18 capture and storage. However, considerable research has been focused on the porous media concept, 19 and dual network (fracture and pores) multiphysics coupled modelling has remained a challenge due 20 to the lack of a systemic theory to bridge the physical deformation of the media (e.g., rocks) and the 21 interaction of water flow in pores and fractures. This paper adopts the non-equilibrium 22 thermodynamics-based approach, the Mixture Coupling Theory, to develop a thermodynamics 23 consistency constitutive model for the fully coupled Hydro-Mechanical behavior in double porosity 24 formation. The energy dispassion due to fluid flow in matrix pore and fracture is given through non-25 equilibrium thermodynamics, and the relationship between the solid and fluid is linked through 26 Helmholtz free energy. The dynamic evolution of stress, porosity change of the matrix pores and 27 fracture, are derived with respect to mechanical strain, pore pressure, and fracture pressure to account 28 for the flow-deformation interaction. The developed constitutive equations are then solved 29 numerically to show the hydraulic and mechanical behavior of double porosity formation, as well as 30 their sensitivity to parameters. 31 **Keywords**: double porosity; nonequilibrium thermodynamics; hydraulic-mechanical

32 Introduction

33 The ubiquity of double porosity media and its completely different characteristics compared to porous 34 media give rise to the importance of studies on water flow in fractured soil or rock in groundwater 35 sources evaluation, underground construction, groundwater contamination, petroleum and shale gas exploitation, underground gas storage, geothermal reservoir (Berkowitz 2002, Rutqvist and 36 37 Stephansson 2003, Gupta and Yadav 2020). Water flow in the subsurface is driven by both hydraulic 38 gradient and rock mechanical field. The interaction between water flow and the deformation of the 39 solid results in a more complex process in groundwater flow (Segura and Carol 2008, Tsang et al. 40 2015), making it difficult for mathematical modelling.

The early modelling work toward the Hydro-Mechanical coupling are the ones by Terzaghi (1943) and Biot (1962), Biot (1972), then followed extensively by many other researchers (Lewis and Schrefler 1987, Vardoulakis et al. 1996, Rutqvist and Tsang 2002, Laloui et al. 2003, Rajagopal and Tao 2005, Tarantino and Tombolato 2005, Wong and Mašín 2014, Zhou and Sheng 2015). In these studies, the porous media is assumed to be homogeneous with single porosity. However, many geomaterials have two scales of void space: the matrix pores and the fracture (Borja and Koliji 2009), as illustrated in Fig. 1.

The distinctive fluid transport and pressure distribution in the fracture and matrix pores are quite different from those in the single porosity situation, so that the classic Biot equations fail to capture the feature of the double porosity situation.

51 To describe the coupled hydro-mechanical behavior of the double porosity material, the material is 52 often viewed to be composed of two distinct but overlapping media: one consisting of the porous 53 matrix, in which there are the solid matrix and matrix pores, and the other is the fracture (Barenblatt et 54 al. 1960, Warren and Root 1963), see Fig. 1. The two media can exchange water mass as the porous 55 matrix holds a large storage capacity and low permeability, while the fracture has high permeability 56 and low storability (Song et al. 2019). Based on the above concept, various mathematical formulations 57 representing the fluid flow or hydro-mechanical coupling have been developed by different 58 approaches with different degrees of sophistication.

59 The early double porosity model (Barenblatt et al. 1960, Warren and Root 1963, Aifantis 1980) 60 explored the fluid transport behavior but failed to explore the mechanical deformation and its 61 coupling with fluid. In these research, the coupling between the fluids in the matrix pores and the 62 fracture is achieved by the fluid exchange between the two regions; the flow is assumed to be 63 independent of deformation. Later models (Wilson and Aifantis 1982, Khaled et al. 1984, Beskos and 64 Aifantis 1986, Zhang et al. 2003, Zhang and Roegiers 2005) incorporated the fluid pressure into the 65 strain equation and the strain influence on the fluid transport to achieve the coupling between flow 66 and deformation. However, these models made no progress in the coupling between the fluids in the 67 matrix pores and the fracture, as they only considered the mass exchange. The fact is that the fracture 68 fluid acting on the porous matrix must lead to the change of fracture volume and matrix pore volume, 69 and further influence the fluid transport in the matrix pore, and vice versa. Such a phenomenon is then 70 incorporated in the new fully coupled Hydro-Mechanical models proposed by Khalili (2003), Khalili 71 (2008).

72 The mathematical models are developed by different approaches. There is no certain classification of 73 the approaches for modelling the double porosity problem. Different categories of approaches can be 74 found in (Chen and Teufel 2000, Gelet et al. 2012, Boutin and Royer 2015). Among all the 75 approaches, two noticeable ones are the conventional mechanics approach and the mixture theory 76 approach. Some remarkable mathematic models have been developed by the mechanics approach 77 (Elsworth and Bai 1992, Khalili and Valliappan 1996, Pao and Lewis 2002, Khalili 2003, Khalili 78 2008) and by the mixture theory approach (Aifantis 1977, Aifantis 1979, Aifantis 1980, Wilson and 79 Aifantis 1982, Beskos and Aifantis 1986, Bai et al. 1993, Bai et al. 1993, Borja and Koliji 2009) and 80 following further work (Wilson and Aifantis 1982, Khaled et al. 1984, Beskos and Aifantis 1986, 81 Berryman and Wang 1995).

The mechanics approach is straightforward and simple, but it often requires ad hoc assumptions, and it lacks the ability of systemic self-development (Laloui et al. 2003). For the mixture theory approach, as pointed out by Heidug and Wong (1996), since it maintains the individuality of the solid and fluid phase, it highly relies on the phase interaction information that is very difficult to obtain. Physical 86 intuition and specific assumptions must be required to form the coupling between phases. This may87 bring difficulties for this approach and restrict the application.

88 In this paper, a non-equilibrium thermodynamics-based approach, the mixture coupling theory, is 89 adopted to develop the fully coupled governing equations for the hydro-mechanical behavior of 90 double porosity media saturated with single-phase flow. This theory origins from Heidug's research 91 for single porosity media with swelling effects (Heidug and Wong 1996). It is modified from the 92 mixture theory by viewing the solid-fluid mixture as a single continuum without explicitly 93 discriminating between the solid and fluid phases, therefore, this theory is more like a hybrid of the 94 Biot poroelasticity view and the mixture theory. Unlike the mixture theory adopting the momentum 95 conservation equation, mixture coupling theory directly works on the free energy conservation, 96 making it easier. This theory provides a rigorous framework to study the coupling effects between 97 multi-physics and multi-phases and has been applied to different couplings in porous media (Chen 98 2013, Chen et al. 2013, Chen et al. 2016, Chen et al. 2018, Ma et al. 2020), and it is the first attempt 99 to apply the theory to the double porosity media to develop the fully coupled Hydro-Mechanical 100 model.

By using mixture coupling theory, the very general evolution equation of stress, the porosity of the matrix pores and fracture are obtained with respect to the pore water pressure and fracture water pressure as well as coupling with mechanical strain. The final governing equations are restricted within the small strain and elastic conditions, and are the same as the model proposed by Khalili (2003) through the mechanics approach. The developed mathematical models are solved by the finite element method to illustrate the coupling phenomenon in double porosity media and the sensitivity of parameters.

108 **Balance equation**

109 Basic definitions and relationships

110 In a double porosity model, water can pass through the boundary via the porous matrix and the 111 fracture, so that two water flux, namely, porous matrix flux \mathbf{I}^{M_w} and fracture flux \mathbf{I}^{F_w} , are defined

112
$$\mathbf{I}^{M_{W}} = \rho^{M_{W}} \left(\mathbf{v}^{M_{W}} - \mathbf{v}^{s} \right), \ \mathbf{I}^{F_{W}} = \rho^{F_{W}} \left(\mathbf{v}^{F_{W}} - \mathbf{v}^{s} \right)$$
(1)

113 where the subscripts M_w , F_w , *s* represent the water in the matrix pore, water in fracture and the 114 solid phase. ρ^{M_w} and ρ^{F_w} are the density of porous water and fracture water, which are relative to 115 the volume of the whole mixture system. \mathbf{v}^{M_w} , \mathbf{v}^{F_w} , \mathbf{v}^s are the velocity of porous water, fracture 116 water and the solid.

117 ρ^{M_w} and ρ^{F_w} are related to the true mass density (relative to the volume of porous water and fracture 118 water) $\rho_t^{M_w}$ and $\rho_t^{F_w}$ through

119
$$\rho^{M_w} = \phi^{M_w} \rho_t^{M_w} , \ \rho^{F_w} = \phi^{F_w} \rho_t^{F_w}$$
(2)

120 where ϕ^{M_w} , ϕ^{F_w} are the porosity of porous matrix and fracture, and they are the volume of the matrix 121 pore and fracture against the volume of the whole mixture.

122 The Darcy velocity for porous water and fracture water are

123
$$\mathbf{u}^{M_W} = \phi^{M_W} \left(\mathbf{v}^{M_W} - \mathbf{v}^s \right), \ \mathbf{u}^{F_W} = \phi^{F_W} \left(\mathbf{v}^{F_W} - \mathbf{v}^s \right)$$
(3)

124 Balance equation

125 An arbitrary domain V with a boundary S attached to its surface is selected. The domain includes all 126 phases and the double porosity feature. Water flow can pass through the boundary while the solid 127 cannot.

128 <u>1. Balance equation for Helmholtz free energy</u>

In the double porosity model, the Helmholtz free energy change involves four parts: the mechanical energy, the energy change by porous water flow, the energy change by fracture water flow and the entropy part, the balance equation writes as

132
$$\frac{D}{Dt} \int_{V} \psi dV = \int_{S} \sigma \mathbf{n} \cdot \mathbf{v}^{s} dS - \int_{S} \mu^{M_{w}} \mathbf{I}^{M_{w}} \cdot \mathbf{n} dS - \int_{S} \mu^{F_{w}} \mathbf{I}^{F_{w}} \cdot \mathbf{n} dS - T \int_{V} \gamma dV$$
(4)

133 where: ψ is Helmholtz free energy density, σ is the Cauchy stress tensor, μ^{Mw} , μ^{Fw} are the 134 chemical potential of porous water and fracture water, respectively. *T* is the temperature, which is 135 regarded as constant in this paper; γ is the entropy produced per unit volume of the mixture.

136 The material time derivative is
$$\frac{D}{Dt} = \partial_t + \mathbf{v}^s \cdot \nabla$$
, where ∂_t is the time derivative and ∇ the gradient,

137 then the derivative version of the balance equation (4) for the Helmholtz free energy is

138
$$\dot{\psi} + \psi \nabla \cdot \mathbf{v}^{s} - \nabla \cdot \left(\boldsymbol{\sigma} \mathbf{v}^{s} \right) + \nabla \cdot \left(\mu^{Mw} \mathbf{I}^{Mw} \right) + \nabla \cdot \left(\mu^{Fw} \mathbf{I}^{Fw} \right) = -T\gamma \leq 0$$
(5)

139 <u>2. Balance equation for water mass</u>

140 Water in the matrix pore (fracture) changes in two ways: 1. Water flow \mathbf{I}^{Mw} (\mathbf{I}^{Fw}) passes through the 141 boundary and exchanges with the surroundings; 2. Water exchanges between the matrix pore and the 142 fracture. The balance equation for water in the matrix pore and the fracture are

143
$$\frac{D}{Dt} \left(\int_{V} \rho^{M_{w}} dV \right) = -\int_{S} \mathbf{I}^{M_{w}} \cdot \mathbf{n} dS - \int_{V} r_{ex} dV$$
(6)

144
$$\frac{D}{Dt} \left(\int_{V} \rho^{Fw} dV \right) = -\int_{S} \mathbf{I}^{Fw} \cdot \mathbf{n} dS + \int_{V} r_{ex} dV \tag{7}$$

145 where r_{ex} is the exchange rate of fluid mass between the fracture and matrix pore.

146 The time derivative versions of the water balance equations are

147
$$\dot{\rho}^{Mw} + \rho^{Mw} \nabla \cdot \mathbf{v}^s + \nabla \cdot \mathbf{I}^{Mw} + r_{ex} = 0$$
(8)

$$\dot{\rho}^{F_{W}} + \rho^{F_{W}} \nabla \cdot \mathbf{v}^{s} + \nabla \cdot \mathbf{I}^{F_{W}} - r_{ex} = 0$$
⁽⁹⁾

149

148

150 Entropy production and transport law

151 Entropy production

152 During irreversible processes, such as heat and mass transfer, entropy will be produced. The entropy 153 production can be expressed in terms of thermodynamic flows and thermodynamic forces (Kondepudi 154 and Prigogine 2014). The quantification of entropy production through non-equilibrium 155 thermodynamics is the core of the Mixture Coupling Theory. During water transport in double porosity media, the entropy is generated from one mechanism, i.e., the friction between the solid and water boundary, which consists of three parts: 1. the friction of porous water \mathcal{P}_{Mw} ; 2. the friction of fracture water \mathcal{P}_{Fw} ; 3. the friction generated when water exchange between the matrix pore and the fracture networks \mathcal{P}_{ex} . From non-equilibrium thermodynamics (Katchalsky and Curran 1965), there is

161
$$\mathcal{G}_{M_{W}} = -\mathbf{I}^{M_{W}} \cdot \nabla \mu^{M_{W}} , \ \mathcal{G}_{F_{W}} = -\mathbf{I}^{F_{W}} \cdot \nabla \mu^{F_{W}}$$
(10)

162 According to Gelet et al. (2012) and Coussy (2004), \mathcal{G}_{ex} can be expressed as

163
$$\mathcal{G}_{ex} = r_{ex} \left(\mu^{Mw} - \mu^{Fw} \right) \tag{11}$$

164 Then, the overall entropy production of the mixture system can be written as

165
$$0 \le T\gamma = -\mathbf{I}^{M_w} \cdot \nabla \mu^{M_w} - \mathbf{I}^{F_w} \cdot \nabla \mu^{F_w} + r_{ex} \left(\mu^{M_w} - \mu^{F_w} \right)$$
(12)

166 In equation (12), \mathbf{I}^{M_w} , \mathbf{I}^{F_w} , r_{ex} are the thermodynamic flows, and $-\nabla \mu^{M_w}$, $-\nabla \mu^{F_w}$, $(\mu^{M_w} - \mu^{F_w})$ 167 are the corresponding thermodynamic forces that drive the transport processes.

168 The quantification of the entropy production (12) and the following substitution of $T\gamma$ term in the 169 free energy equation (5) (like what has been done leading to equation (14)) are the key features of the 170 Mixture Coupling Theory, distinguishing it from other thermodynamics approaches, such as the ones 171 by Coussy (2004), Gelet et al. (2012), Nakshatrala et al. (2018). The advantages of the two features 172 have been presented in Ma et al. (2022) when dealing with the dissolution process.

173 Transport Law

The entropy production in section 3.1, on the one hand, can be used to develop the transport law, like what has been done in Chen et al. (2018). This paper focuses on the coupling between the flow and deformation. To simplify the discussion, the fluid flow is assumed to obey the Darcy's law. The Darcy's law for porous water and the fracture water can be derived through the entropy production, Gibbs-Duhem equation and phenomenological equation, as (Chen 2013)

179
$$\mathbf{u}^{M_{W}} = -\frac{k^{M_{W}}}{\nu^{W}} \nabla p^{M_{W}} , \quad \mathbf{u}^{F_{W}} = -\frac{k^{F_{W}}}{\nu^{W}} \nabla p^{F_{W}}$$
(13)

180 where k^{M_w} , k^{F_w} are the intrinsic permeability for porous matrix and fracture, p^{M_w} and p^{F_w} 181 correspond to the porous water pressure and fracture water pressure, respectively, v^w is the water 182 viscosity.

183

184 **Constitutive equation**

185 Basic equation for deformation

Assuming the material maintains mechanical equilibrium so that there is $\nabla \cdot \boldsymbol{\sigma} = \boldsymbol{0}$. Substituting the entropy production (12) into the Helmholtz free energy balance equation (5), the Helmholtz free energy change of the mixture system can be written as

189
$$\dot{\psi} + \psi \nabla \cdot \mathbf{v}^{s} - \nabla \cdot \left(\mathbf{\sigma} \mathbf{v}^{s} \right) + \mu^{Mw} \nabla \cdot \mathbf{I}^{Mw} + \mu^{Fw} \nabla \cdot \mathbf{I}^{Fw} + r_{ex} \left(\mu^{Mw} - \mu^{Fw} \right) = 0 \tag{14}$$

Equation (14) gets rid of the entropy term in equation (5), and enables us to explore the Helmholtz free energy change through the mass flux \mathbf{I}^{Mw} , \mathbf{I}^{Fw} and r_{ex} . Multiplying μ^{Mw} , μ^{Fw} on both sides of equation (8), (9), and substituting the corresponding results into equation (14), the Helmholtz free energy change becomes

194
$$\dot{\psi} + \psi \nabla \cdot \mathbf{v}^{s} = \nabla \cdot (\mathbf{\sigma} \mathbf{v}^{s}) + \mu^{Mw} (\dot{\rho}^{Mw} + \rho^{Mw} \nabla \cdot \mathbf{v}^{s}) + \mu^{Fw} (\dot{\rho}^{Fw} + \rho^{Fw} \nabla \cdot \mathbf{v}^{s})$$
(15)

In equation (15), the term $\dot{\rho}^{Mw} + \rho^{Mw} \nabla \cdot \mathbf{v}^s$ is the pore water mass change in the mixture, and therefore $\mu^{Mw} (\dot{\rho}^{Mw} + \rho^{Mw} \nabla \cdot \mathbf{v}^s)$ represents the free energy change due to pore water mass change. Equation (15) indicating that the free energy change of the system is the result of the mechanical energy and the mass energy.

199 Next, classic continuum mechanics method is adopted to measure the deformation state. Some basic200 relationships are required (Wriggers 2008)

201
$$\mathbf{F} = \frac{\partial \mathbf{x}}{\partial \mathbf{X}} (\mathbf{X}, t) , \ \mathbf{E} = \frac{1}{2} (\mathbf{F}^{\mathrm{T}} \mathbf{F} - \mathbf{I}), \ \mathbf{T} = J \mathbf{F}^{-1} \boldsymbol{\sigma} \mathbf{F}^{-T}$$
(16)

where **X** is an arbitrary reference configuration, **x** is the position, **E** represents Green strain, **F** is the deformation gradient, **T** and σ are the second Piola-Kirchhoff stress and Cauchy stress. *J* is the

204 Jacobian of **F** (
$$J = \det \mathbf{F}$$
), and satisfies $J = \frac{dV}{dV_0}$ (V , V_0 are the volume in the current and reference

205 configuration.)

206 The time derivation of J satisfies the Euler's formula

$$\dot{J} = J div v_s \tag{17}$$

208 From equation (15), with the relationships (16) and (17), the Helmholtz free energy equation (15) can

209 be switched into the reference configuration as

210
$$\dot{\Psi} = tr(\mathbf{T}\dot{\mathbf{E}}) + \mu^{Mw}\dot{m}^{Mw} + \mu^{Fw}\dot{m}^{Fw}$$
(18)

211 in which: $\Psi = J\psi$ is the Helmholtz free energy in the reference configuration, $m^{M_w} = J\rho^{M_w}$, 212 $m^{F_w} = J\rho^{F_w}$ are the mass density of porous water and fracture water in the reference configuration.

213 Helmholtz free energy density of porous/fracture water

According to classical thermodynamics, the free energy density of porous matrix water and fracture water can be written as

216 $\Psi_{porous} = -p^{M_w} + \rho_t^{M_w} \mu^{M_w}$ (19)

217
$$\Psi_{fracture} = -p^{F_W} + \rho_t^{F_W} \mu^{F_W}$$
(20)

218 Using the Gibbs-Duhem equation for porous water and fracture water, it leads to

$$\dot{p}^{Mw} = \rho_t^{Mw} \dot{\mu}^{Mw} \tag{21}$$

 $\dot{p}^{F_W} = \rho_t^{F_W} \dot{\mu}^{F_W} \tag{22}$

Invoking equation (21), (22) into the time derivation of equation(19), (20), the following relationshipscan be obtained

223
$$\dot{\psi}_{porous} = \dot{\rho}_t^{M_w} \mu^{M_w} \tag{23}$$

$$\dot{\psi}_{fracture} = \dot{\rho}_t^{Fw} \mu^{Fw} \tag{24}$$

225 Free energy density of the solid matrix

The free energy of the solid-porous-fracture mixture system consists of three parts: the free energy of the porous water, the free energy of the fracture water and the free energy of the solid matrix. By subtracting the free energy of the porous water and fracture water from the free energy of the mixture system, the free energy of the solid matrix can be obtained.

From equation (18), (23), (24) and using the density relationship (2), the free energy of the solid matrix is

232
$$\left(\Psi - J\phi^{Mw}\psi_{porous} - J\phi^{Fw}\psi_{fracture}\right)^{\Box} = tr\left(\mathbf{T}\dot{\mathbf{E}}\right) + \dot{\upsilon}^{Mw}p^{Mw} + \dot{\upsilon}^{Fw}p^{Fw}$$
(25)

where $v^{M_w} = J\phi^{M_w}$, $v^{F_w} = J\phi^{F_w}$ are the porosity of porous matrix and fracture in the reference configuration.

235 Subtracting the contribution of porous water pressure and fracture water pressure, that is

236
$$W = \left(\Psi - \phi^{M_w} \psi_{porous} - J \phi^{F_w} \psi_{fracture}\right) - \upsilon^{M_w} p^{M_w} - \upsilon^{F_w} p^{F_w}$$
(26)

Substituting equation (25) into the time derivation of the equation (26), the evolution of *W* can be obtained as below, enabling us to use the pressure \dot{p}^{Mw} and \dot{p}^{Fw} as variables

239
$$\dot{W} = tr(\mathbf{T}\dot{\mathbf{E}}) - \upsilon^{Mw}\dot{p}^{Mw} - \upsilon^{Fw}\dot{p}^{Fw}$$
(27)

- 240 where W is a function of **E**, p^{Mw} and p^{Fw} .
- From equation (27), there must be

242
$$T_{ij} = \left(\frac{\partial W}{\partial E_{ij}}\right)_{p^{M_w}, p^{F_w}}, \upsilon^{M_w} = -\left(\frac{\partial W}{\partial p^{M_w}}\right)_{E_{ij}, p^{F_w}}, \upsilon^{F_w} = -\left(\frac{\partial W}{\partial p^{F_w}}\right)_{E_{ij}, p^{M_w}}$$
(28)

243 So that

244
$$\dot{W}(\mathbf{E}, p^{M_{W}}, p^{F_{W}}) = \left(\frac{\partial W}{\partial E_{ij}}\right)_{p^{M_{W}}, p^{F_{W}}} \dot{E}_{ij} + \left(\frac{\partial W}{\partial p^{M_{W}}}\right)_{E_{ij}, p^{F_{W}}} \dot{p}^{M_{W}} + \left(\frac{\partial W}{\partial p^{F_{W}}}\right)_{E_{ij}, p^{M_{W}}} \dot{p}^{F_{W}}$$
(29)

Differentiating equation (28), with the help of equation (29), the evolution of stress, porosity matrixporosity and fracture porosity can be obtained.

247
$$\dot{T}_{ij} = L_{ijkl}\dot{E}_{kl} - M_{ij}\dot{p}^{Mw} - S_{ij}\dot{p}^{Fw}$$
(30)

248
$$\dot{\upsilon}^{M_W} = M_{ij} \dot{E}_{ij} + Q \dot{p}^{M_W} + B \dot{p}^{F_W}$$
 (31)

$$\dot{\upsilon}^{Fw} = S_{ij}\dot{E}_{ij} + B\dot{p}^{Mw} + Z\dot{p}^{Fw}$$
(32)

250 where the parameters L_{iikl} , M_{ii} , S_{ii} , H_{ii} , B, Q, Z, are as following group equations

251
$$L_{ijkl} = \left(\frac{\partial T_{ij}}{\partial E_{kl}}\right)_{p^{Mw}, p^{Fw}} = \left(\frac{\partial T_{kl}}{\partial E_{ij}}\right)_{p^{Mw}, p^{Fw}}, M_{ij} = -\left(\frac{\partial T_{ij}}{\partial p^{Mw}}\right)_{E_{ij}, p^{Fw}} = \left(\frac{\partial \upsilon^{Mw}}{\partial E_{ij}}\right)_{E_{ij}, p^{Fw}}$$

252
$$S_{ij} = -\left(\frac{\partial T_{ij}}{\partial p^{F_W}}\right)_{E_{ij}, p^{M_W}} = \left(\frac{\partial \upsilon^{F_W}}{\partial E_{ij}}\right)_{E_{ij}, p^{M_W}}, \ Z = \left(\frac{\partial \upsilon^{F_W}}{\partial p^{F_W}}\right)_{E_{ij}, p^{M_W}}$$
(33)

253
$$B = \left(\frac{\partial \upsilon^{F_W}}{\partial p^{M_W}}\right)_{E_{ij,p}^{F_W}} = \left(\frac{\partial \upsilon^{M_W}}{\partial p^{F_W}}\right)_{E_{ij,p}^{M_W}}, \ Q = \left(\frac{\partial \upsilon^{M_W}}{\partial p^{M_W}}\right)_{E_{ij,p}^{F_W}}$$

254

255 Coupled hydro-mechanical governing equations

256 Assumptions and simplifications

Equations (30), (31), (32) are the general coupled equations for stress, strain, porous/fracture pressure and porous/fracture porosity, they allow us to explore the Hydro-Mechanical coupling in a broad way including anisotropy, large deformation, et.al. This paper forms the final governing equations in a simple elastic-isotropy cases, therefore, some simplifications and assumptions are made below.

261 1. The mechanical behavior is restricted to small strain condition, therefore the Green Strain tensor

262 E_{ij} and Piola-Kirchhoff stress T_{ij} can be replaced by strain tensor ε_{ij} and Cauchy stress σ_{ij} .

263 2. Although many double porosity formations show the features of anisotropic and heterogeneous, 264 following some other research (e.g. Berryman and Wang (1995)), it is roughly assumed that the 265 material is isotropic. Therefore, material-dependent constants M_{ij} , S_{ij} can be substituted by a form of 266 scalar multiplied by Kronecker delta.

267
$$M_{ij} = \zeta^{M_w} \delta_{ij} , \ S_{ij} = \zeta^{F_w} \delta_{ij}$$
(34)

and considering only the elasticity, the elastic stiffness L_{ijkl} can be a fourth-order isotropic tensor

269
$$L_{ijkl} = G\left(\delta_{ik}\delta_{jl} + \delta_{il}\delta_{jk}\right) + \left(K - \frac{2G}{3}\right)\delta_{ij}\delta_{kl}$$
(35)

in which G and K denote the shear modulus and bulk modulus of the material.

According to equations (34), (35) and assumption (1), the stress evolution equation (30) can be simplified to

273
$$\dot{\sigma}_{ij} = \left(K - \frac{2G}{3}\right)\dot{\varepsilon}_{kk}\delta_{ij} + 2G\dot{\varepsilon}_{ij} - \zeta^{Mw}\dot{p}^{Mw}\delta_{ij} - \zeta^{Fw}\dot{p}^{Fw}\delta_{ij}$$
(36)

With assumption 1, 2, the porosity evolution equations (31) and (32) can be simplified as

275
$$\dot{\upsilon}^{M_w} = \zeta^{M_w} \dot{\varepsilon}_{ii} + Q \dot{p}^{M_w} + B \dot{p}^{F_w}$$
(37)

276
$$\dot{\upsilon}^{F_W} = \zeta^{F_W} \dot{\varepsilon}_{ii} + B \dot{p}^{M_W} + Z \dot{p}^{F_W}$$
(38)

277 Parameter identification

278 Identification of ζ^{Mw} and ζ^{Fw}

279 Consider a situation where the porous matrix block is blocked from the fracture, which means that 280 there is no mass exchange between the matrix pores and the fracture. In this situation, the porous 281 matrix block can be viewed as a non-porous material. The stress/strain change is purely induced by 282 fracture water. This situation is the same as the traditional single porosity research as the porous block 283 can be viewed as solid grain and the fracture become the pores.

284 In this situation, the incremental relationship between stress and pore fluid pressure is

$$\dot{\sigma}_{ij} = -\dot{p}^{Fw} \delta_{ij} \tag{39}$$

strain rate is related to the fracture water pressure through

$$\dot{\varepsilon}_{ij} = -\frac{\dot{p}^{Fw}}{3K_{pb}}\delta_{ij} \tag{40}$$

288 where K_{pb} is the bulk modulus of the porous matrix block.

289 Substituting equation (39) and (40) into equation (36) for the above situation

290
$$-\dot{p}^{Fw}\delta_{ij} = \left(K - \frac{2G}{3}\right)\left(-\frac{\dot{p}^{Fw}}{K_{pb}}\delta_{ij}\right) + 2G\left(-\frac{\dot{p}^{Fw}}{3K_{pb}}\delta_{ij}\right) - \zeta^{Fw}\dot{p}^{Fw}\delta_{kl}$$
(41)

291 From equation (41), the expression of ζ^{F_w} can be obtained as

$$\zeta^{F_w} = 1 - \frac{K}{K_{pb}}$$
(42)

Next, to identify the parameter ζ^{Mw} , a second situation is assumed where the porous water pressure and fracture water pressure change in the same way by \dot{p}^{Δ} . In this situation, there is no discrimination between matrix pores and fracture, they are all viewed as the void space in the system. The incremental relationship between the stress and fluid pressure, and the strain response are

297
$$\dot{\sigma}_{ij} = -\dot{p}^{\Delta} \delta_{ij} , \dot{\varepsilon}_{ij} = -\frac{\dot{p}^{\Delta}}{3K_s} \delta_{ij}$$
(43)

Invoking equation (43) into equation (36) for the second situation, it's easy to obtain

$$\zeta^{Mw} + \zeta^{Fw} = 1 - \frac{K}{K_s}$$
(44)

300 Then, with equation (42), the expression of ζ^{M_W} can be obtained

$$\zeta^{Mw} = \frac{K}{K_{pb}} - \frac{K}{K_s}$$
(45)

302 Identification of Q, B and Z

In the first situation considered in section 5.2.1, the fracture volume fraction change is related tofracture water pressure through

$$\dot{\upsilon}^{F_W} = -\phi^{F_W} \frac{\dot{p}^{F_W}}{K_{pb}} \tag{46}$$

306 Invoking (46), (40) and the expression for ζ^{F_w} into equation (38) for the first situation leads to

$$-\phi^{F_{W}}\frac{\dot{p}^{F_{W}}}{K_{pb}} = \left(1 - \frac{K}{K_{pb}}\right)\left(-\frac{\dot{p}^{F_{W}}}{K_{pb}}\right) + Z\dot{p}^{F_{W}}$$
(47)

308 The expression of Z can be derived as

309
$$Z = \frac{1}{K_{pb}} \left(1 - \frac{K}{K_{pb}} - \phi^{F_W} \right)$$
(48)

310 Then, in the second situation in section 5.2.1, the fracture volume change is

$$\dot{\upsilon}^{F_W} = -\phi^{F_W} \frac{\dot{p}^{\Delta}}{K_s} \tag{49}$$

312 With equation (49), (43), (48), equation (38) for the second situation can be written as

313
$$-\phi^{F_w} \frac{\dot{p}^{\Delta}}{K_s} = \left(1 - \frac{K}{K_{pb}}\right) \left(-\frac{\dot{p}^{\Delta}}{K_s}\right) + B\dot{p}^{\Delta} + \frac{1}{K_{pb}} \left(1 - \frac{K}{K_{pb}} - \phi^{F_w}\right) \dot{p}^{\Delta}$$
(50)

314 Then, the expression *B* can be obtained as

315
$$B = \left(1 - \frac{K}{K_{pb}} - \phi^{F_W}\right) \left(\frac{1}{K_s} - \frac{1}{K_{pb}}\right)$$
(51)

316 Applying the second situation to equation (37), with the strain and porosity evolution equations

317
$$\dot{\varepsilon}_{ij} = -\frac{\dot{p}^{\Delta}}{3K_s} \delta_{ij}, \ \dot{\upsilon}^{Mw} = -\phi^{Mw} \frac{\dot{p}^{\Delta}}{K_s}$$
(52)

318 It leads to

319
$$-\phi^{M_w}\frac{\dot{p}^{\Delta}}{K_s} = \zeta^{M_w} \left(-\frac{\dot{p}^{\Delta}}{K_s}\right) + Q\dot{p}^{\Delta} + \left(1 - \frac{K}{K_{pb}} - \phi^{F_w}\right) \left(\frac{1}{K_s} - \frac{1}{K_{pb}}\right) \dot{p}^{\Delta}$$
(53)

320 So that the expression for Q can be obtained as

321
$$Q = \left(\frac{1}{K_{pb}} - \frac{1}{K_s}\right) \left(1 - \frac{K}{K_{pb}} + \frac{K}{K_s} - \phi^{F_w}\right) - \frac{\phi^{M_w}}{K_s}$$
(54)

322 Governing equations

323 Effective stress and mechanical equation

324 With the expression of ζ^{M_w} and ζ^{F_w} , equation (36) can be written as

325
$$\dot{\sigma}_{ij} = \left(K - \frac{2G}{3}\right)\dot{\varepsilon}_{kk}\delta_{ij} + 2G\dot{\varepsilon}_{ij} - \zeta^{Mw}\dot{p}^{Mw}\delta_{ij} - \zeta^{Fw}\dot{p}^{Fw}\delta_{ij}$$
(55)

326 From which the effective stress can be written as

$$\dot{\sigma}_{ij} = \dot{\sigma}'_{ij} - \zeta^{Mw} \dot{p}^{Mw} \delta_{ij} - \zeta^{Fw} \dot{p}^{Fw} \delta_{ij}$$
(56)

328 where σ'_{ii} is the effective stress.

327

The proposed effective stress includes the influence of matrix water pressure p^{Mw} and fracture water pressure p^{Fw} through the coefficient ζ^{Mw} and ζ^{Fw} . It is the same as the one proposed by Callari and Federico (2000) and the one by Khalili et al. (2005), Khalili (2008) reduced for saturated condition. When there is no fracture, the effective stress can be reduced to the effective stress in porous media by regarding $K = K_{pb}$.

Assuming the mechanical equilibrium condition $\partial \sigma_{ii} / \partial x_i = 0$, and using displacement variables

335
$$d_i (i = 1, 2, 3)$$
 through $\varepsilon_{ij} = \frac{1}{2} (d_{i,j} + d_{j,i})$, it leads to

336
$$G\nabla^{2}\dot{\mathbf{d}} + \left(\frac{G}{1-2\theta}\right)\nabla\left(\nabla\cdot\dot{\mathbf{d}}\right) - \zeta^{Mw}\nabla\dot{p}^{Mw} - \zeta^{Fw}\nabla\dot{p}^{Fw} = 0$$
(57)

337 in which θ is Poisson's ratio

338 Hydraulic behavior

According to the parameters identified in section 5.2, the porosity change equation (37) and(38) can now be quantitated as

341
$$\dot{\upsilon}^{M_{W}} = \zeta^{M_{W}} \dot{\varepsilon}_{ii} + \left[\frac{\zeta^{M_{W}}}{K} \left(1 - \zeta^{M_{W}} - \phi^{F_{W}}\right) - \frac{\phi^{M_{W}}}{K_{s}}\right] \dot{p}^{M_{W}} + \left[-\frac{\zeta^{M_{W}}}{K} \left(\zeta^{F_{W}} - \phi^{F_{W}}\right)\right] \dot{p}^{F_{W}}$$
(58)

342
$$\dot{\upsilon}^{F_{W}} = \zeta^{F_{W}} \dot{\varepsilon}_{ii} + \left(\zeta^{F_{W}} - \phi^{F_{W}}\right) \left(\frac{1}{K_{s}} - \frac{1}{K_{pb}}\right) \dot{p}^{M_{W}} + \frac{1}{K_{pb}} \left(\zeta^{F_{W}} - \phi^{F_{W}}\right) \dot{p}^{F_{W}}$$
(59)

The two equations represent the pore and fracture porosity change with respect to the volumetric strain, the porous water pressure and the fracture water pressure in a fully coupled way, which are the same as the porosity change equation in Khalili (2003).

346 <u>1. Porous matrix water</u>

From water partial mass equation (8), water density relationship (2), Darcy velocity equation (3),
water flux equation (1) and Euler identity, the conservation equation of water can be written as

349
$$\left(\boldsymbol{\upsilon}^{M_{W}}\boldsymbol{\rho}_{t}^{M_{W}}\right)^{\Box} + \nabla \cdot \left(\boldsymbol{\rho}_{t}^{M_{W}}\mathbf{u}^{M_{W}}\right) + r_{ex} = 0 \tag{60}$$

350 Expanding the first term in equation (60) and considering the variation of true density as

351
$$\dot{\rho}_{t}^{M_{w}} = \rho_{t}^{M_{w}} \left(\frac{1}{\rho_{t}^{M_{w}}} \frac{\partial \rho_{t}^{M_{w}}}{\partial p^{M_{w}}}\right) \frac{\partial p^{M_{w}}}{\partial t} = \rho_{t}^{M_{w}} \frac{1}{K_{w}} \dot{p}^{M_{w}}$$
(61)

352 where K_w is the compressibility of water.

353 With equation (61), equation (60) becomes

354
$$\dot{\upsilon}^{Mw}\rho_t^{Mw} + \upsilon^{Mw}\rho_t^{Mw}\frac{1}{K_w}\dot{p}^{Mw} + \nabla \cdot \left(\rho_t^{Mw}\mathbf{u}^{Mw}\right) + r_{ex} = 0$$
(62)

Invoking the porous volume fraction evolution equation (37) and Darcy's law (13), the governing
equation for porous matrix water transport is

357
$$\zeta^{Mw} \nabla \cdot \dot{\mathbf{d}} + \left(Q + \frac{\nu^{Mw}}{K_w}\right) \dot{p}^{Mw} + B \dot{p}^{Fw} + r_{ex} = \frac{k^{Mw}}{\nu^w} \nabla^2 p^{Mw}$$
(63)

358 with
$$\zeta^{M_w} = \frac{K}{K_{pb}} - \frac{K}{K_s}, \ Q = \frac{\zeta^{M_w}}{K} \left(1 - \zeta^{M_w} - \phi^{F_w}\right) - \frac{\phi^{M_w}}{K_s}, \ B = -\frac{\zeta^{M_w}}{K} \left(1 - \frac{K}{K_{pb}} - \phi^{F_w}\right).$$

359

360 <u>2. Fracture water</u>

361 Similar to the above steps, the governing equation for the fracture water transport can be obtained as

362
$$\zeta^{Fw} \nabla \cdot \dot{\mathbf{d}} + B \dot{p}^{Mw} + \left(Z + \frac{\upsilon^{Fw}}{K_w} \right) \dot{p}^{Fw} - r_{ex} = \frac{k^{Fw}}{\upsilon^w} \nabla^2 p^{Fw}$$
(64)

363 with
$$\zeta^{F_W} = 1 - \frac{K}{K_{pb}}, B = \left(\zeta^{F_W} - \phi^{F_W}\right) \left(\frac{1}{K_s} - \frac{1}{K_{pb}}\right), Z = \frac{1}{K_{pb}} \left(\zeta^{F_W} - \phi^{F_W}\right).$$

364 Verification and discussion

Equations (30)-(32) provide very general coupled formulations for the coupled evolution of PK-2 stress, porous matrix porosity and fracture porosity, along with the dynamic change of porous water pressure and fracture water pressure, and Green strain. These formulations are for general cases like large deformation, isotropic and anisotropic. With the assumption of small deformation and isotropic, the governing equations (57), (63), (64) derived in this paper are the same as the ones proposed based on the mechanics approach, such as Khalili et al. (1999), Khalili (2003) or the ones by Pao and Lewis (2002), Khalili (2008) reduced for saturated condition which have been verified by comparing with experimental data (Khalili 2003) and further verified through comparing with different double porosity models (Ashworth and Doster 2019). This indicates that the governing equations developed by mechanics approaches are only specific cases of the general constitutive equations by using mixture coupling theory.

376 Additionally, the mixture coupling theory framework developed in this paper is more rigorous and 377 realistic, with the least assumptions. For example, in the mechanics approach, the assumptions (e.g. 378 isotropic) were made at the very beginning of the derivation process, therefore, restricting its 379 following derivation process, limiting any possibility of extension to other conditions, for example, 380 anisotropic. In other words, the mechanics approach can only obtain the constitutive relations for a 381 specific case. However, through the mixture coupling theory adopted in this paper, no specific 382 assumptions are required at the beginning, so that the general constitutive relations (i.e. equation 383 (30),(31),(32)) can be obtained. Such a relationship applies to many conditions, e.g., large 384 deformation, isotropic/anisotropic. Following the general constitutive relations, elasticity and isotropy 385 conditions are selected to obtain the final governing equations. It can thus be concluded that the 386 mixture coupling theory approach is more flexible and widely applicable.

387 The key research object of Mixture Coupling Theory is the Helmholtz free energy change of the 388 system, i.e., equation (15) and (25), which are achieved through the mass and energy balance, as well 389 as the entropy production and the Gibbs-Duhem equation. Since many dynamics processes can be described through energy dispassion, Mixture Coupling Theory can be used in a lot of fields, such as 390 391 thermo-hydro-mechanical-chemical coupling (Ma et al. 2022), swelling (Chen et al. 2016), dissolution 392 (Ma et al. 2022), or potentially wave propagation as wave is mainly a movement of energy through a 393 medium (Kumar et al. 2021, Rajak et al. 2022), or other fields like biological tissue. However, all the 394 aforementioned research topics are mainly for porous media. This paper is the first attempt to develop 395 the fully coupled equations for the double porosity media. It is noticed that Aghighi et al. (2021) used 396 a similar approach to study the sorption in double porosity media, but their equations are not

397 presented in a fully coupled way as the pore water pressure influence is missing in the mechanical or398 transport equation.

399

400 Numerical Simulation

401 A simple numerical simulation is presented in this section to illustrate the mechanical and hydraulic 402 behavior in the double-porosity formation. The mechanical deformation and hydraulic pressure 403 change are given, as well as the porosity and permeability change. The sensitivity of fracture spacing 404 and permeability is analyzed.

405 *Porosity, permeability and exchange rate*

406 <u>1.Porosity</u>

407 The matrix and fracture porosity change equation have been derived as equation (31) and (32), which 408 are further reduced to equation (37) and (38) according to the assumptions in section 5.1. From the 409 parameters identified in section 5.2, the matrix and fracture porosity equation are solved in 410 incremental form as

411
$$\Delta \upsilon^{M_w} = \zeta^{M_w} \Delta \varepsilon_{ii} + Q \Delta p^{M_w} + B \Delta p^{F_w}$$
(65)

412
$$\Delta \upsilon^{F_W} = \zeta^{F_W} \Delta \varepsilon_{ii} + B \Delta p^{M_W} + Z \Delta p^{F_W}$$
(66)

413 <u>2.Permeability</u>

The permeability change is related to porosity through the Kozeny-Carman's law as (Zheng andSamper 2008)

416
$$\frac{k^{M_w}}{k_0^{M_w}} = \left(\frac{\nu^{M_w}}{\nu_0^{M_w}}\right)^3 \left(\frac{1 - \nu_0^{M_w}}{1 - \nu^{M_w}}\right)^2$$
(67)

417 in which the subscript '0' denotes the initial value.

418 This relationship is normally for porous matrix, for fracture, although there is no evidence to support

419 this relationship, it is roughly assumed that the fracture permeability also follows equation (67).

- 420 <u>3. Exchange rate</u>
- 421 The exchange rate can be described by

422
$$r_{ex} = \frac{\chi k^{Mw}}{\nu^{W}} \left(p^{Mw} - p^{Fw} \right)$$
(68)

423 where χ is the shape factor. There have been some different expressions for χ proposed by different 424 researchers (Warren and Root 1963, Kazemi et al. 1976), a summary can be found in Ranjbar and 425 Hassanzadeh (2011). In this paper, the expression of χ developed by Warren and Root (1963) is 426 adopted as

427
$$\chi = \frac{4N(N+2)}{L^2} \tag{69}$$

428 where N(N=1,2,3) represents the dimension of the porous matrix block. L is the fracture spacing.

429 Geometry and boundary condition

A double porosity geoformation, with 20m length and 1m height (Fig. 2), is selected. The formation initially contains water at pressure of 30MPa. At the beginning of the simulation, the pressure at the right boundary drops to 5MPa due to external disturbance while maintaining 30MPa at the left boundary. By setting the right side to be permeable, water can flow out. The formation is initially at mechanical equilibrium with no effective stress. To explore the mechanical behavior when the pressure changes, the left boundary is allowed to move while the other boundaries are constrained.

436 One observation line and three observation points A, B, C are selected.

437 Parameters adopted in this simulation are listed in Table 1 (Abousleiman and Nguyen 2005, Nair et al.
438 2005, Gelet et al. 2012)

439 Numerical results

440 <u>1. Hydraulic and mechanical behavior</u>

The evolution of matrix pressure and fracture pressure along the observation line is presented in Fig. 3 and Fig. 4. Because of a pressure gradient generated at the beginning, water will flow from the left to the right, as time goes by, the pressure within the domain decreases and trends to reach equilibrium (Fig. 3, 4). Comparing Fig. 3 and Fig. 4, the pressure change in the fracture is quicker thanthat in the matrix pore, this is mainly because of a greater permeability of the larger permeability in the fracture zone. Fig. 5 shows the pressure change with time at observation points A, B, C, from which it is clearer that the fracture pressure changes quicker than the pore pressure, but finally, the pressure inthe fracture and pore reaches the same as what has been set to be the boundary condition.

The domain is initially at mechanical equilibrium with no effective stress, indicating that the external loading is born by the water pressure. Because of the loss of water pressure through the right boundary, the external loading will be taken by both water pressure and solid matrix, therefore, effective stress and displacement generate correspondingly (Fig. 6 and 7). Since water pressure losses more near the right boundary, effective stress generates more on the right side (Fig. 6). As the solid matrix bears external loading, consolidation happens and displacement occurs on the free left boundary (Fig. 7).

The pressure difference in the fracture and matrix results in the exchange of water mass. The exchange rate is shown in Fig. 8. At the early time, the pressure difference between the fracture and pore is very significant, thus the exchange rate is great. As time goes by, the pressure difference becomes smaller and smaller, so that the exchange rate decreases.

460 <u>2. Porosity and permeability</u>

The matrix and fracture porosity distribution along the observation line are presented in Fig. 9 and Fig. 10. The Figs show up to a 5% decrease in matrix porosity and a 8% decrease in fracture porosity. Since the pressure change and strain change on the right part of the domain are more significant than those on the left part, and the porosity change shows similar distribution. As the permeability is related to porosity through equation (67), permeability change has a similar trend with porosity change, as shown in Fig. 11 and 12.

467 Sensitivity analysis of fracture spacing

An important part of the double porosity model is the exchange of water from the porous matrix to fracture. A greater exchange rate will help the porous water pressure change quicker. From equation (69), the exchange rate is inversely proportional to the fracture spacing L. Different values of L(L=0.1, 0.5, 1, 5, 10) are chosen to explore the sensitivity dependence of hydraulic and mechanical behavior on L. The results are presented in Fig. 13-16. From Fig. 13, the exchange rate in low L value is much higher than that in big L value. A lower Lvalue means a high 'fracture density', so that water can be transferred quicker between porous block and fracture. In the simulated case, at a certain time or space, the matrix pressure is mostly greater than fracture pressure, so that matrix water flows into the fracture space. The lower L value is, the quicker matrix water lose to fracture, therefore, the matrix pressure changes quicker (Fig. 14). Similarly, fracture pressure changes slower with a small L value (Fig. 15). The difference in pressure change consequently leads to a difference in mechanical behavior (Fig. 16).

480 Sensitivity analysis of fracture permeability

The hydraulic transport is highly affected by permeability. In this section, the permeability sensitivity is explored by setting the fracture permeability as $k^{Fw} = 5*10^{-19}m^2$, $5*10^{-18}m^2$, $5*10^{-17}m^2$ while keeping the pore permeability as $k^{Mw} = 5*10^{-20}m^2$ to represent the permeability difference at different magnitudes. Other parameters remain the same as those listed in Table 1.

The fracture water pressure under different permeabilities is shown in Fig. 17, from which it is clear that the fracture water pressure drops much quicker with a higher permeability. Because the water exchange rate between the fracture and matrix pore depends on the pressure difference, hence, the exchange rate under high fracture permeability is quicker (Fig. 18), which further promotes the matrix pore water pressure drop (Fig. 19).

490 The pressure change trend under different fracture permeability is similar to that in section 6.3, but the 491 porosity change under different fracture permeability is quite different, as shown in Fig. 20. When $k^{Mw} = 5 * 10^{-19} m^2$, the fracture porosity drops most on the right boundary, but when 492 $k^{Mw} = 5 * 10^{-18} m^2$, or $k^{Mw} = 5 * 10^{-17} m^2$, the maximum porosity is not at the right boundary but at a 493 494 point closed to the right boundary. According to equation (66), the change in fracture porosity Δv^{Fw} comes from three parts: 1. the change of strain $\zeta^{F_w} \Delta \varepsilon_{ii}$; 2. The change of pore pressure $B \Delta p^{M_w}$; 3. 495 The change of fracture pressure $Z\Delta p^{F_W}$. The contribution of the three parts under different 496 497 permeabilities is presented in Fig. 21. The consolidation (strain change) and fracture pressure drop 498 would decrease the fracture porosity while the matrix pressure change trends to increase the fracture 499 porosity, leading to the overall decrease of fracture porosity. It can be found that the main difference 500 among the three permeabilities is the fracture pressure contribution part: under all permeabilities, the 501 strain and pore pressure changes slow and their contribution to the fracture porosity are similar. 502 However, when permeability is higher, the fracture pressure drops quickly, leading to a quick and 503 significant decrease of fracture porosity; when permeability is lower, fracture pressure drops slowly 504 and fracture porosity changes slowly. The combined influence of the strain, matrix pressure and 505 fracture pressure results in the dramatic trend in Fig. 20. Owing to the permeability difference, the 506 matrix porosity also changes in a dramatic trend, as presented in Fig. 22 and Fig. 23.

507 Conclusion

This paper derives the constitutive equations for double porosity formation under a coupled hydromechanical situation by using the Mixture Coupling Theory. The cross-coupling relations between stress, strain, porous/fracture water pressure and porous/fracture porosity are obtained, allowing us to explore the Hydro-Mechanical response in a broad way. The final governing equations are formed for the elastic condition, leading to the same equations developed by other approaches.

The constitutive models are solved by the finite element method, and the results show the hydraulic and mechanical deformation, as well as the porosity and permeability change. The sensitivity analysis shows that the smaller fracture spacing greatly increases the exchange rate and facilitates the matrix water pressure change, the fracture permeability sensitivity analysis shows that the greater permeability significantly accelerates the pressure change and affects the porosity change.

518 Data Availability statement

519 Some or all data, models, or code that support the findings of this study are available from the 520 corresponding author upon reasonable request.

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