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Article

Relationship between Dynamics of Modern Glaciers of the Mt. Munkhkhairkhan (Mongolian Altai) and Climate

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Abstract: Mt. Munkhkhairkhan is the most crucial region for understanding climate and glaciation changes in Mongolia. This study investigated the relationship between glacial area changes and the climate elements of Mt. Munkhkhairkhan in the Mongolian-Altai Mountains using a remote sensing approach, in-situ observations, the Mann–Kendall (MK) test, Innovative Trend Analysis Method (ITAM), Sen’s slope estimator test, and statistical analysis. The study results showed that for the last 30 years, the annual average air temperature of Mt. Munkhkhairkhan has been slightly increasing. Total annual precipitation (mainly snow) in the mountain area decreased from 1990 to 2000, but since 2000, a significant increase in precipitation levels has appeared. For the last 30 years, the glacial area has decreased by 32% to 11.7 km². Multiple regression results showed a strong correlation between Temperature, Precipitation, and Glaciers (Multiple R = 0.69, R² = 0.48). Rutherford indicated that Temperature ($t = -2.332$, $p = 0.036$) and Precipitation ($t = -3.212$, $p = 0.007$) were significant predictors in the model. Air temperature and precipitation explained 48 percent of the change in the glacier area, and R = 0.69 is a strong correlation. The glaciers and snow area in the study area have changed due to climate warming and precipitation changes and are located in arid and semi-arid regions of Central Asia. This study of Mt. Munkhkhairkhan shows that climate change significantly impacts glaciers and snow.

Keywords: global change; Mongolian Altai; temperature; precipitation; glacier area change; high mountains; remote sensing; Mt. Munkhkhairkhan

1. Introduction

Glaciers are an important indicator of climate change and are sensitive to global warming, and provide specific ecosystems with reliable freshwater resources. Glaciers, ice sheets,

and freshwater lakes are the largest reservoirs of freshwater on Earth; consequently, studying them is of vital humanitarian, ecological, economical, and ecosystemic importance [1,2]. Many researchers have shown that global warming is the main contributing factor to glacier melting, and changes in glacier-fed rivers are important issues for water management in the context of global climate change [3,4]. The global annual air temperature has increased by 0.85 °C from 1880 to the present and is likely to increase further and rapidly in the future [4,5].

Although climate change has been rapidly increasing globally, its effects differ from region to region depending on geographical and basic patterns of the climate in some regions. In particular, the high-mountain region ecosystems of Central Asia are strongly affected by climate change due to their arid climate [5,6]. The area covered by glaciers is a total of 17,950 km² in this harsh remote continental climate region [7]. The Mongolian total surface water resource has been estimated to be 599 km³/year, with 83.5% stored in lakes, 3.6% in glaciers, and 5.8% in rivers [8].

The distribution of water resources varies differently for each Mongolian territory. Western Mongolia has a dry climate and inadequate water resources [2]. The glaciers of Western Mongolia provide about 3% of the total freshwater resources and they have a vital role in the hydrographical condition of the Central Asian internal drainage basin [9,10]. Over the last few decades, due to climate change, the number of glaciers, glacial areas, and their accumulation have decreased. Associated with glacial melting, enormous changes in the neighboring hydrographical conditions and ecosystems have appeared [2,11,12]. In particular, because of ongoing glacial melting, there will be a risk of changes to glacial lakes and an increase in the water levels of rivers and lakes. Due to global warming, and the consequent permafrost thaw that has dammed pro-glacial lakes, glacial lakes risk collapsing.

The Mongolian-Altai Mountain regions extend from the northwest to the southeast in the western part of Mongolia [13]. Pan and other researchers (2016) have estimated that there are 627 glaciers covering 334.0 ± 42.3 km² in the Mongolian Altai Mountains [14]. These glaciers are vitally balanced systems that maintain downstream ecosystems and provide major freshwater resources for people [9]. One of the main representatives of the glaciers is found on the Munkhkhairkhan Mountain at the southeast of the Mongolian Altai range. These glaciers are located from the Doloonuur (Nuur means the Lake) at the northwest to Ikh Turgen at the southeast of the mountain; the northwest part is called the massif of Doloonuur, the central part is called the Shuurkhai Massif, and the southeast part is called Khukhnuurs Massif (Figure 1). Among these, at the Shuurkhai Massif, five glaciers occupy 22.39% of the total area, eight glaciers occupy 40.68% of the area of the Massif of Doloonuur, and four glaciers occupy 36.92% of the total area of the Khukhnuur Massif [15]. By the estimation of Otgonbayar, D, 24 glaciers together had an area of 48.13 km² in 1970, and in 2008, 17 glaciers covering 26.57 km² remained [13,15]. This shows that between 1970 and 2008, seven glaciers completely melted. Kamp, Pan estimated a glacial area of 27.4 km² in 2011 and Ganyushkin, D, estimated 26.86 km² in 2016 [16,17]. Based on these estimates, the glacier area has decreased by 44.7% from 1970 to 2008. Some small glaciers have melted entirely.

In this mountainous area, where there is continuous permafrost and glaciers are melting, the loss of permafrost and areas of changing and collapsing small glacial lakes indicate that global warming and climate change have been accelerating [18–21]. Determining the relationship of glacier area change with various climate parameters will play a vital role in evaluating and predicting future climate change effects on mountainous areas, ecosystem stability, and water management [22]. Whether mountain glaciers, particularly glaciers at high altitudes and in arid regions, are particularly vulnerable to climate change needs to be studied further [23,24]. The purpose of our study, therefore, is to identify the effects of climate change on changes in the spatial area of glaciers. Consequently, our aims are (i) to determine the current area of and recent changes in the glaciers between 1991 and 2020 on Mt. Munkhkhairkhan; (ii) to analyze air temperature and precipitation changes in the

mountain area; (iii) to determine the relationship between climate and glacier melting, and (iv) to discuss the various implications of these changes—for example, for water use.

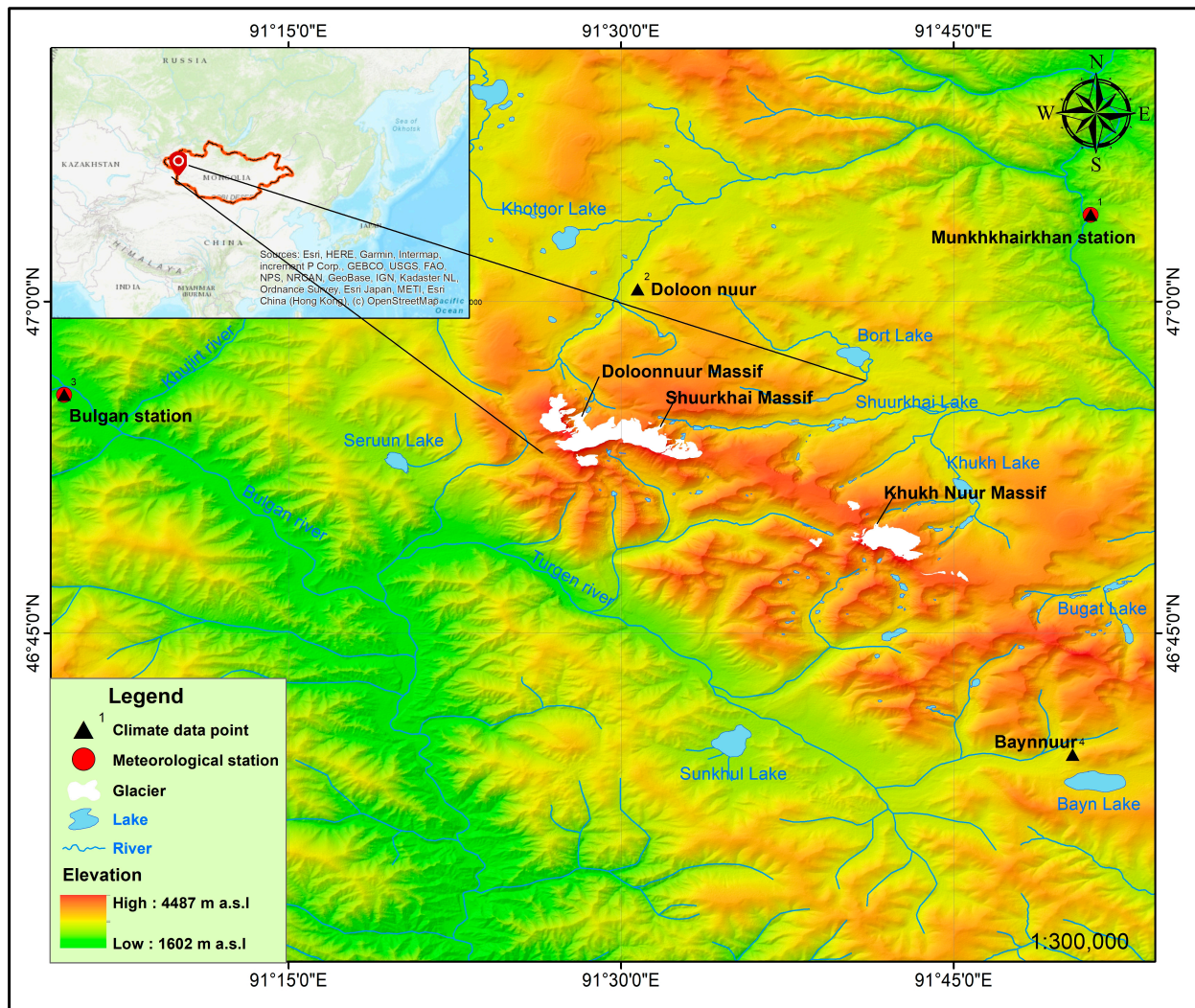


Figure 1. Geographic location of the Mt. Munkhkhairkhan (the numbers of the climate data points same to Table 2).

2. Materials and Methods

2.1. Study Area

Mt. Munkhkhairkhan stretches from the northwest to the southeast in the Mongolian Altai Mountains, from latitude $46^{\circ}56'095''$ to $46^{\circ}45'45''$ N, and longitude $90^{\circ}09'04''$ to $90^{\circ}32'48''$ E [13]. The mountain range is 3500–4300 m a.s.l., about 150 km long and 50 km wide, and borders the mountain ranges of the Munkhkhairkhan sum (a “sum” is an administrative unit) of Khovd Province and the Bulgan sum of Bayan-Ulgii Province (Figure 1).

Mt. Munkhkhairkhan is the second-highest mountain in Mongolia [16,25]. The range ends at the Deluunii River Valley in the northwest and the Baruun Khuurain River Valley in the south. Along a northeast/southwest axis, the range separates a tectonic depression from the mountains of the Mongolian-Altai. Small rivers such as the Doloon Nuur, Shuurkhai, and Bort in the north part, and the Uyench and Bodonch Rivers in the south drain the area [15]. The climate zone includes harsh winters and humid, cold summers.

The cold season (October–April) and the winter are long and cold, whereas summers (May–September) are short and warm. The Mongolian Altai is in an area well known for its extreme continental climate due to its central location on the Eurasian land mass. The

winters are dominated by a strong Asian anticyclone (the Siberian High Pressure), while the summers in the mountains are short and cool. Modern climate information has been systematically recorded in Mongolia ever since the early 1940s. However, the national climate station network is sparse, and climate data for the Mongolian Altai are limited to only six meteorological stations that are situated in the valleys between the mountain ranges—thus making it difficult to estimate the climate and weather at higher elevations. In the valleys of the Mongolian Altai, the mean temperature is -30 to -34 °C in January and less than 15 °C in July [26]. The average temperature of the coldest month is -20.8 °C, and the average summer temperature is $+18$ °C, with the warmest summer temperatures reaching 22.5 – 25 °C [15,27,28].

2.2. Data Source

2.2.1. Meteorological Data

“Prediction of Worldwide Energy Resource” or “POWER Data Access Viewer” is an open information system based on satellite observations with geo-spatial climate and solar-related data for the evaluation and planning of renewable energy systems. We used these resources for analysing precipitation and air temperature data 2 m above the ground at four locations near Mt. Munkhkhairkhan and total precipitation data from the database of NASA power (<https://power.larc.nasa.gov/data-access-viewer/> (accessed on 30 January 2022)).

For verifying this data, the climate data obtained during 1991–2020 at the Munkhkhairkhan Meteorological station (<http://irimhe.namem.gov.mn/> (accessed on 22 February 2022)) was also used. The values of the above-named data sets are different, but the general trend is consistent [29,30].

2.2.2. Satellite Data

For the estimation of Mt. Munkhkhairkhan’s glacier area change between 1991 and 2020, we used Landsat 5 TM, Landsat 8 OLI, and SRTM with a 30 m accuracy—available at USGS (Global Visualization Viewer; <http://glovis.usgs.gov/>; Table 1).

Table 1. Satellite data.

Date	Scene ID	Sensor	Resolution
1991	LT51410271991198XXX03	Landsat 5 TM	30 m
1992	LT51410271992233BJC01	Landsat 5 TM	30 m
1995	LT51400271995218BJC00	Landsat 5 TM	30 m
1997	LT51410271997246BJC00	Landsat 5 TM	30 m
2000	LT51410272000191BJC00	Landsat 5 TM	30 m
2002	LE71400282002213SGS00	Landsat 7 ETM+	30 m
2004	LT51410272004218BJC00	Landsat 5 TM	30 m
2007	LT51410272007226BJC00	Landsat 5 TM	30 m
2008	LT51410272008229BJC00	Landsat 5 TM	30 m
2010	LT51410272010234IKR02	Landsat 5 TM	30 m
2013	LC81410272013242LGN01	Landsat 8 OLI	30 m
2014	LC81410272014245 LGN01	Landsat 8 OLI	30 m
2015	LC81410272015216LGN01	Landsat 8 OLI	30 m
2017	LC81410272017205LGN01	Landsat 8 OLI	30 m
2019	LC81410272019227LGN00	Landsat 8 OLI	30 m
2020	LC81410272020214LGN00	Landsat 8 OLI	30 m

2.3. Methodology

Changes over the last 30 years in Mt. Munkhkhairkhan’s glaciers were calculated using Remote sensing analysis, the Mann–Kendall (MK) test, the Innovative Trend Analysis Method (ITAM), and Sen’s Slope Estimator Test, as well as statistical methods for determining the relationship between climate parameters (precipitation and air temperature changes; Figure 2).

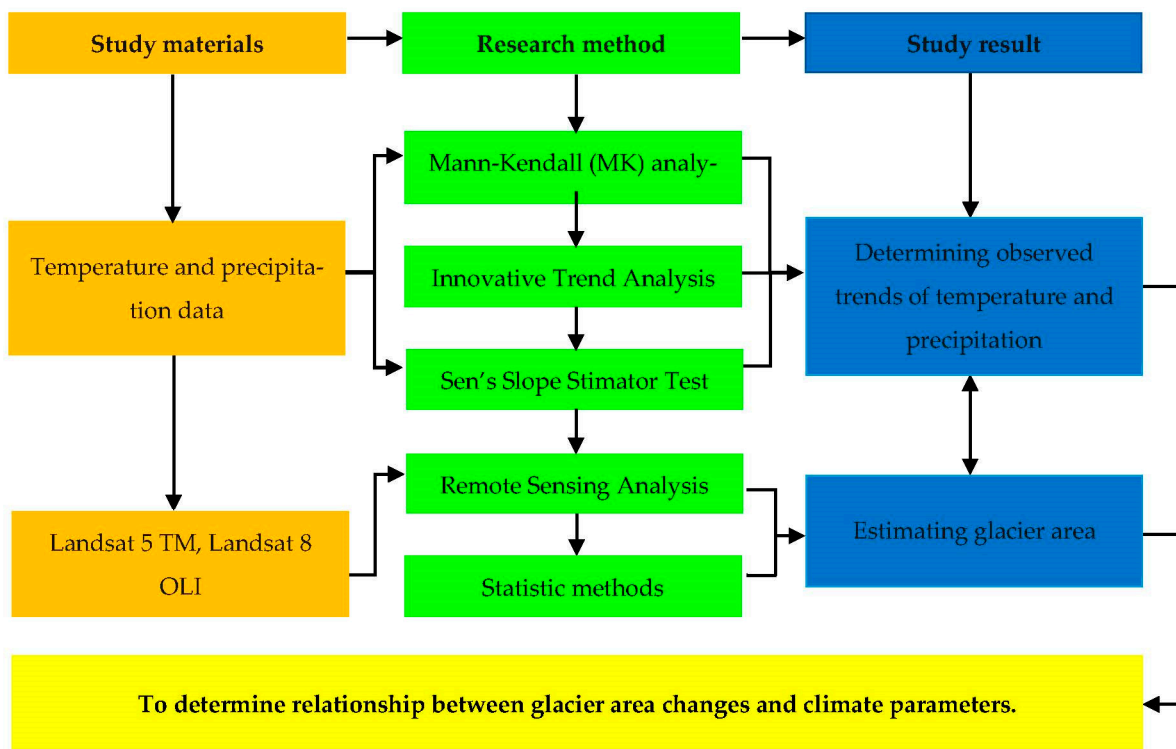


Figure 2. Scheme of the research methodology.

2.3.1. Mann–Kendall (MK) Test

The Mann–Kendall (MK) test was used for detecting the trends of variables in the hydro-meteorological observations [5,31–33]:

$$S = \sum_{i=1}^{n-1} \sum_{j=i+1}^n \text{sgn}(x_j - x_i) \tag{1}$$

Data on the change in the x_i amount of data ($i = 1, 2 \dots n-1$) and the x_j value of the data were compared. The amount of data was estimated as follows using x_i :

$$\text{sgn}(x_j - x_i) = \begin{cases} +1 & \text{if } x_j - x_i > 0 \\ 0 & \text{if } x_j - x_i = 0 \\ -1 & \text{if } x_j - x_i < 0 \end{cases} \tag{2}$$

where the values x_j and x_i and the number timeseries is greater than or equal to 10 ($n \geq 10$), the average value is equated as following the normal distribution $E(S) = 0$. Changes in the time series of the data were calculated as $Var(S)$ as follow:

$$E(S) = 0 \tag{3}$$

$$VarS = \frac{n(n-1)(2n+5) - \sum_{k=1}^m t_k(t_k-1)(2t_k+5)}{18} \tag{4}$$

where m is number of relevant groups of time series and t_k is the number of ties in relevant groups of k th. The Z statistic test is as follows:

$$Z = \begin{cases} \frac{s-1}{\delta} & \text{if } S > 0 \\ 0, & \text{if } S = 0 \\ \frac{s+1}{\delta} & \text{if } S < 0 \end{cases} \tag{5}$$

If Z is greater than 0, it shows the trend is increasing; if Z is less than 0, it indicates that the trend is decreasing.

$$A UF_K = \frac{d_k - E(d_k)}{\sqrt{Var(d_k)}} \quad k = 1, 2, \dots, n \quad (6)$$

$$UF_K = -UF_K \quad (7)$$

$$K = n + 1 - k \quad (8)$$

where UF is the change curve.

2.3.2. Innovative Trend Analysis Method

The accuracy of the Innovative Trend Analysis Method was compared to the MK test and the indicator was calculated as follows [32,33]:

$$\Phi = \frac{1}{n} \sum_{i=1}^n \frac{10(x_j - x_i)}{\mu} \quad (9)$$

where Φ is the trend indicator, n is the number of observations, x_i is the data series of the first half, x_j is the data series of second subseries, and μ is the mean of the data series.

2.3.3. Sen's Slope Estimator Test

The amount of temperature and precipitation was determined by the slope method. Q_i is the slope between 2 datapoints and is determined as follows [34,35]:

$$Q_i = \frac{x_j - x_k}{j - k}, \text{ for } i = 1, 2, \dots, N \quad (10)$$

where x_j and x_k are data points of time, and $k(j > k)$, respectively. If there is only one datum, N is the number of lasting periods: $N = \frac{n(n-1)}{2}$. If there are several data for many years, N is shown as follows: $N < \frac{n(n-1)}{2}$, where " n " is the number of observations. Values of slope calculations were sorted from minimum to maximum. The mean of β was calculated as follows:

$$\beta = \begin{cases} Q[(N+1)/2] & \text{when } N \text{ is odd} \\ Q[(N/2) + Q(N+2)/(2)/(2)] & \text{when } N \text{ is even} \end{cases} \quad (11)$$

The value of β shows whether the trend is increasing or decreasing.

2.3.4. Methodology to Calculate Glacial Area Changes

Glacial area was estimated by the $NDSI$ (Normalized Difference Snow Index) using Landsat TM, Landsat OLI, and Landsat ETM+ data. According to the $NDSI$, the calculated value is between -1 and 1 , and values above zero are considered glacial. It was necessary to subtract the water area because this is included in values above zero [36]:

$$NDSI = \frac{(\text{Green} - \text{SWIR})}{\text{Green} + \text{SWIR}} \quad (12)$$

3. Results

3.1. Glacial Dynamics

The glacier area of Mt. Munkhkhairkhan has decreased by 11.66 km^2 between 1991 and 2020 (and has been decreasing since at least 1970). As of 2020, it covers 24.5 km^2 . A total of 50% of this area is located at the northern parts. For the last three decades, 22% of the glaciers at the northern parts have melted, and in the eastern, south-eastern, southern, and south-western parts where glacier distribution is low, 47% on average have melted.

About 80% of the total glacial area exists above 3600 m a.s.l and the remaining 20% is below 3600 m a.s.l. From 1990 to 2020, the glacial area at and below 3600 m a.s.l. has decreased by 70% (Figure 3).

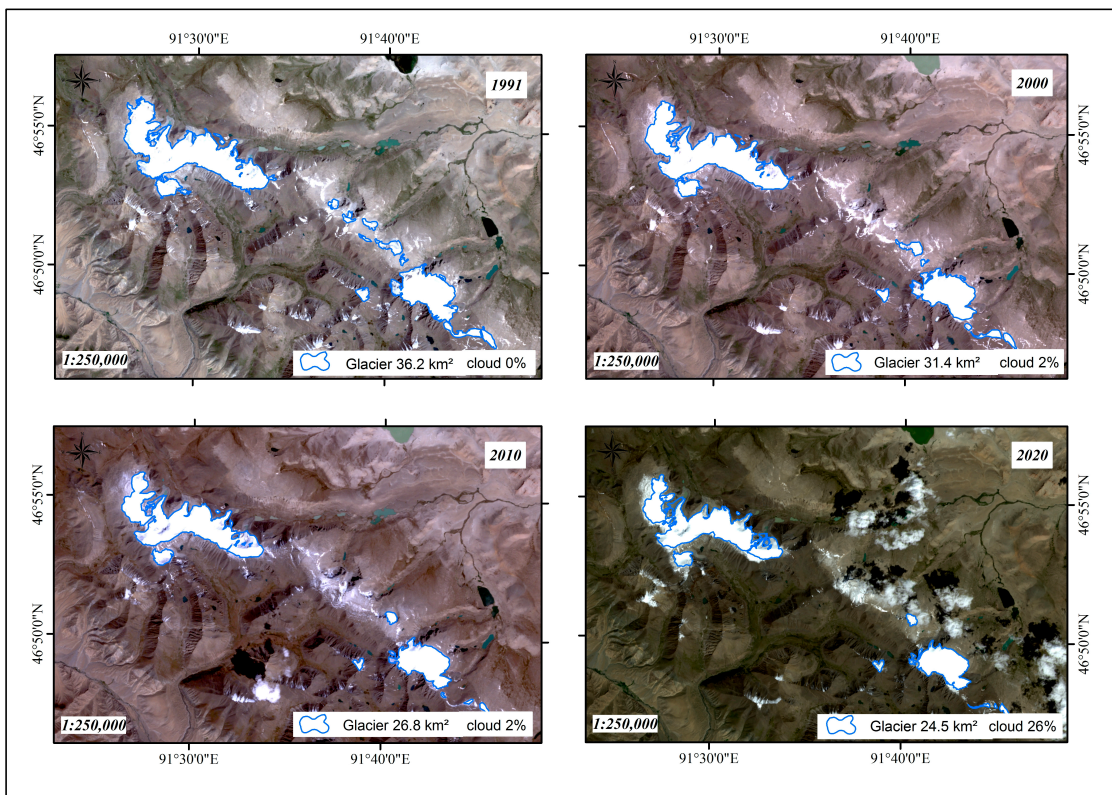


Figure 3. Changes in glacial area over the last 30 years.

The total glacial area decreased by 13.8% from 1991 to 2000, by 16.02% from 2000 to 2010, and by 6.48%, from 2010 to 2020. For the last 30 years, the area has declined by 11.7 km² (32.3%), from 1991 to 2020. This result shows that the glacial area has melted at an average rate of 1.07% per year (Figure 4).

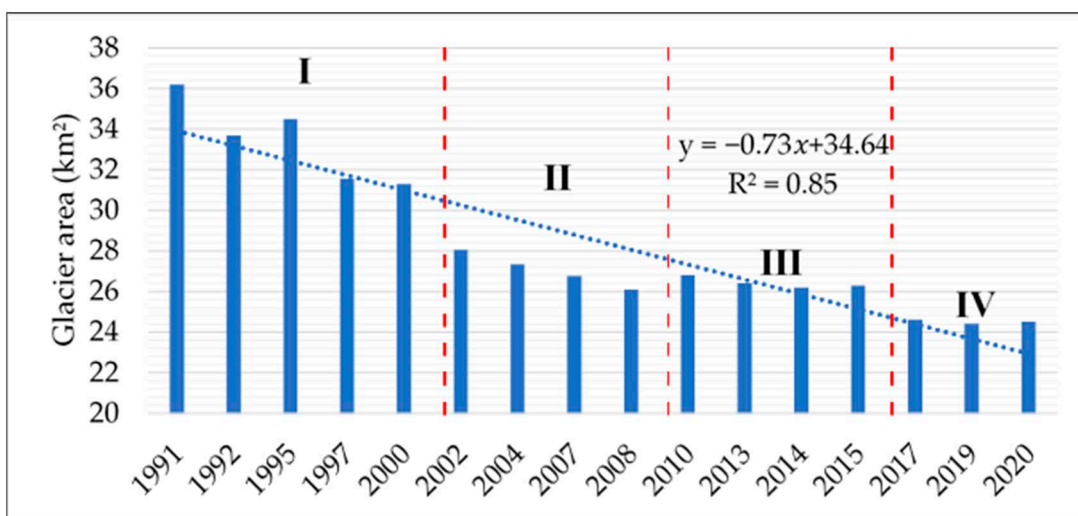


Figure 4. Overall decrease in glacial area over the last 30 years (4 stage of changes glacier area).

Figure 4 shows the four stages of changes in the glacier area. Of these;

- I. 1990–2000: The first-decade glacier melting was intense (0.4 km^2 per year). The total area of glaciers decreased from 36.2 to 31.3 km^2 — $4.9 \text{ km}^2/10$ years.
- II. 2000–2010: A period of steadily increasing glacier melting and decreasing glacier area. Compared to 2008, there was a sharp decrease in area in 2010, but there was a period with a steady decrease up to 2002 (31.3 – 28.5 km^2). Over 10 years the glacier area decreased by 4.5 km^2 (0.45 km^2 per year).
- III. 2010–2015: Glacier area changes were relatively stable, with low glacier melting (0.3 km^2 per year). The period 2010–2014 corresponds to a period of increased precipitation and the greatest decrease in air temperature (Figure 5). In 2008, the rainfall was 126.8 mm , and in 2010, it almost doubled to 231.9 mm ; this situation was maintained until 2015. However, the air temperature dropped sharply from $-3.2 \text{ }^\circ\text{C}$ to $-5.52 \text{ }^\circ\text{C}$, and the average temperature was $-3.3 \text{ }^\circ\text{C}$ until 2015 (Figure 5).
- IV. 2015–2020: Glacier area changes from 2015 to 2017 showed a dramatic reduction (26.3 – 24.2 km^2), but from 2017–2020 was relatively stable (24.6 – 24.5 km^2 ; 0.1 km^2 per year). The period 2010–2014 corresponds to the period of increased precipitation and the greatest decrease in air temperature (Figure 5). In 2008, the rainfall was 126.8 mm , and in 2010, it almost doubled to 231.9 mm ; this situation was maintained until 2015. However, the air temperature dropped sharply from $-3.2 \text{ }^\circ\text{C}$ to $-5.52 \text{ }^\circ\text{C}$, and the average temperature was $-3.3 \text{ }^\circ\text{C}$ until 2015 (Figure 5).

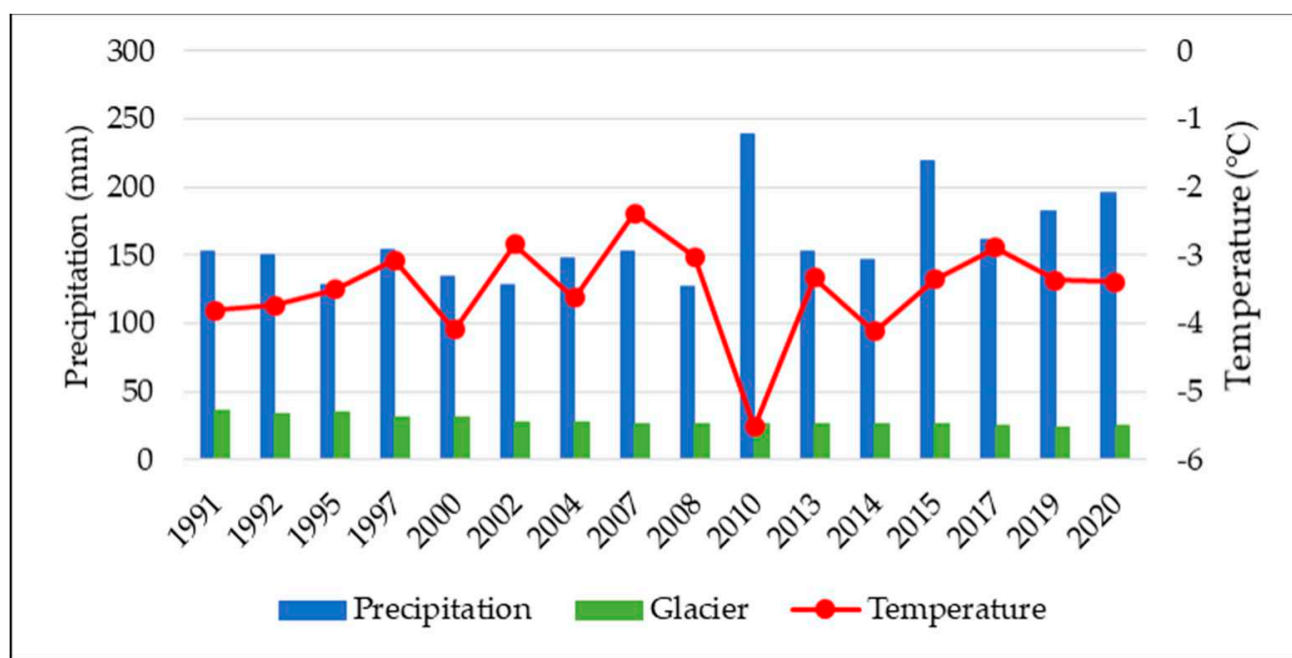


Figure 5. Glacier area change and average climate elements.

3.2. Air Temperature Analysis

The average air temperature has been increasing due to climate change (Figure 6).

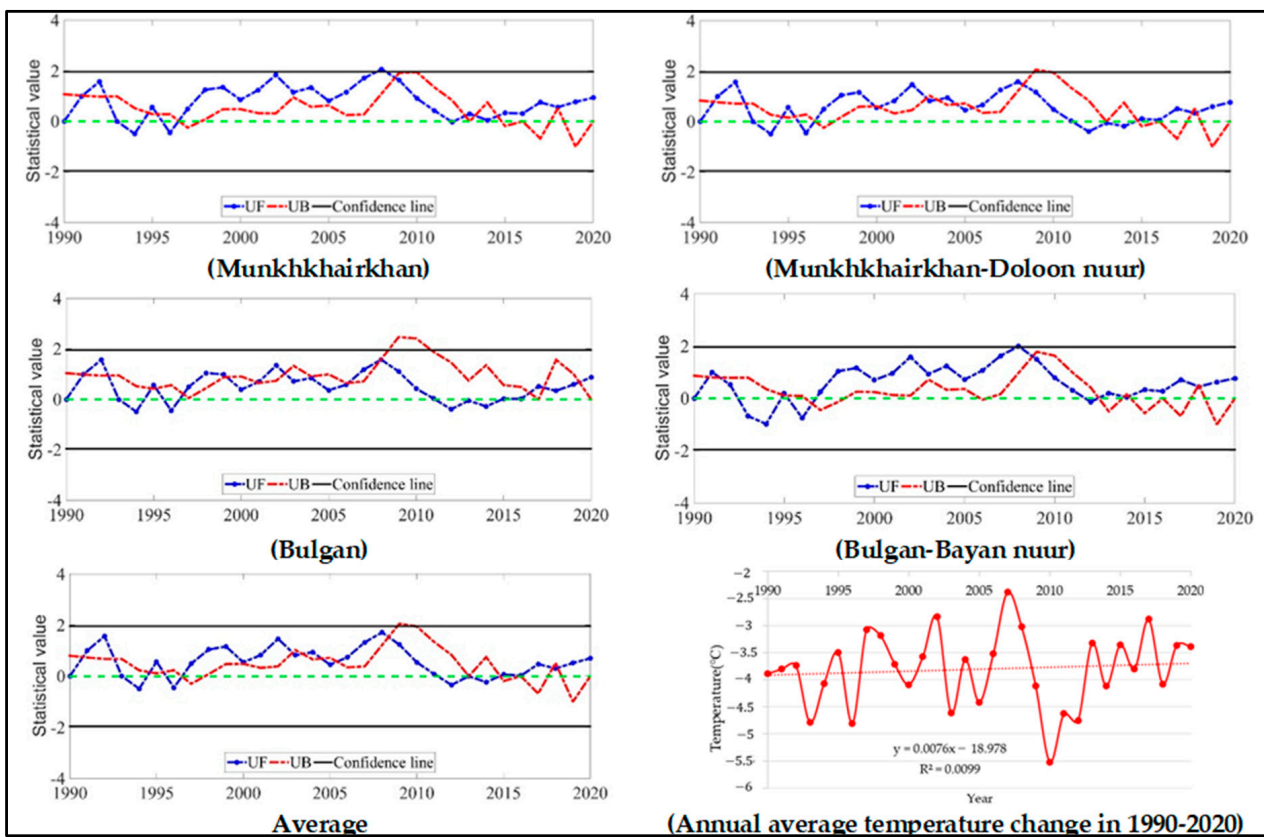


Figure 6. Trends in the annual average temperature near Mt. Munkhkhairkhan, where UF and UB are parameters of this change. UF is the change trend and UB is the reverse order of the trend.

The results of the Mann–Kendall analysis showed that the average air temperature trend rose between 1990 and 2020 ($Z = 0.96$). The result of the Innovative Trend analysis was $\phi = -0.18$, and the Sen’s Slope Estimator Test showed $\beta = 0.01$ (Table 2).

Table 2. The results of the air temperature change (1990–2020) Mann–Kendall test (Z), Innovative Trend Analysis (Φ), and Sen’s Slope Estimator test (β).

S/No.	Name of Stations	Z (MK)	Φ	β
1	Munkhkhairkhan	1.29	−0.18	0.01
2	Doloonuur	1.03	−0.22	0.01
3	Bulgan	1.20	−0.17	0.01
4	Bayannuur	1.03	−0.22	0.01
5	Average	0.96	−0.18	0.01

Although short-term decreases have been observed at times, these decreases have stabilized at a high level and temperatures have increased again. Specifically, in 1990–1995, the air temperature had a steady decreasing trend, and in 1995–2008, it increased sharply and approached +2 °C in statistical terms; in 2008–2012, it gradually decreased to a statistically negative value, but from 2015, an increasing trend has been observed (Figure 6).

3.3. Precipitation Analysis

Precipitation is the main cause of losses in permafrost, and winter precipitation as snow balances the loss of glacier mass caused by summer temperatures in glaciers in cold regions. In Mongolia, the precipitation volume is low in the winter and higher in the summer and autumn.

According to long-term observations, the average annual precipitation in Western Mongolia ranges from 100 mm in the south to 250 mm in the north; about 80% falls from

May to September. According to observations, the amount of precipitation from June to August changes from 50 to 110 mm.

The results of the Mann–Kendall analysis methods showed that precipitation tended to increase from 1990 to 2020 ($Z = 2.81$). The Innovative Trend Analysis showed $\varphi = 1.72$, and the Sen’s Slope Estimator Test result was $\beta = 1.38$ (Table 3, Figure 7).

Table 3. The results for precipitation change Mann–Kendall test (Z), Innovative Trend analysis (Φ), and Sen’s Slope Estimator test (β).

S/No.	Name of Stations	Z (MK)	Φ	β
1	Munkhkhairkhan	2.01	1.28	0.96
2	Doloonuur	3.21	1.79	1.40
3	Bulgan	3.21	1.79	1.40
4	Bayannuur	3.43	2.09	1.56
5	Average	2.81	1.72	1.38

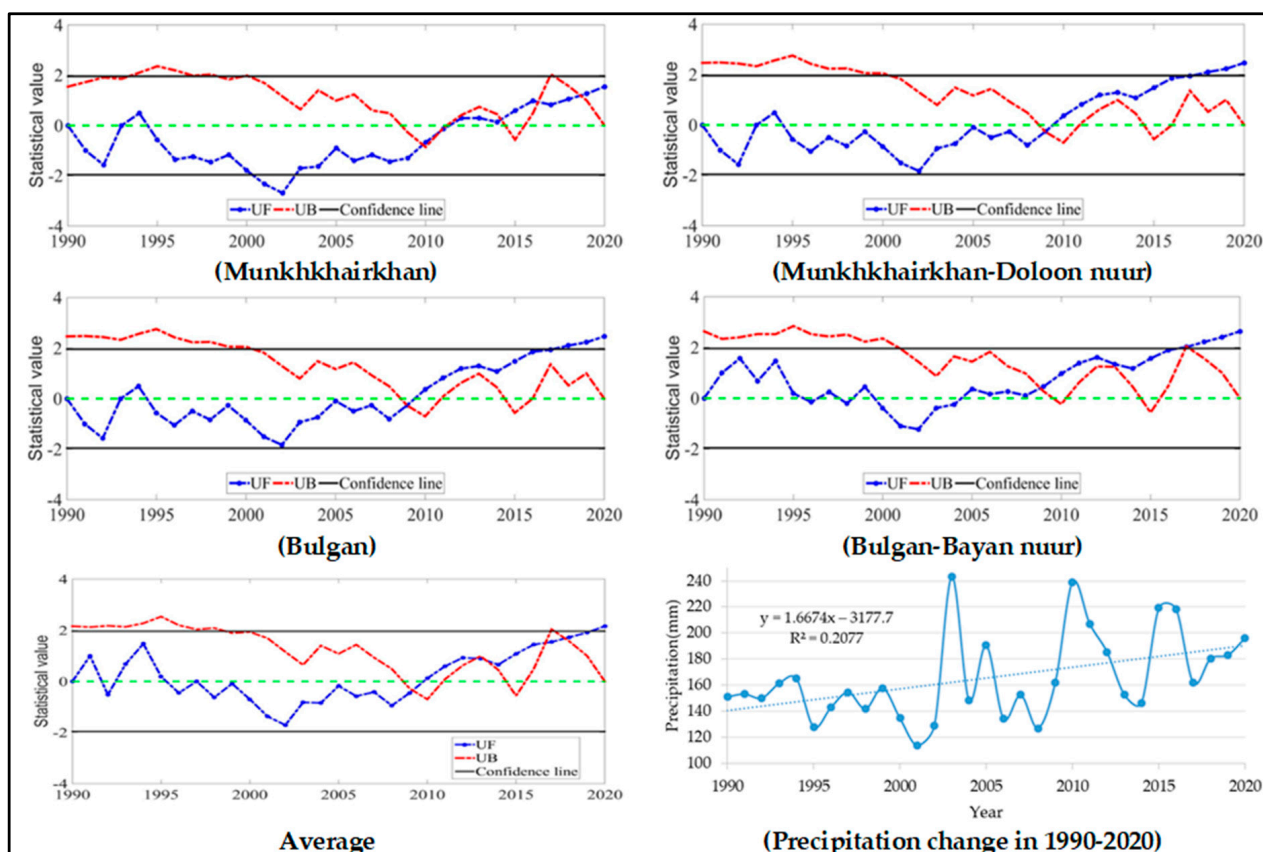


Figure 7. Trends of precipitation near Mt. Munkhkhairkhan.

A timeseries analysis of precipitation showed that precipitation tended to increase over cycles of 2–3 years and then decline. Precipitation trends showed an overall decrease from 1990 to 2003 but an increase over the last 20 years, (Figure 7). In summer, thunderstorms and heavy rains occur.

Precipitation gradually increased between 1990 and 1994 and approached +2 in statistical value, but inbetween 1995 and 2002 it decreased sharply and reached a statistical value of -2 ; it was negative between 2002 and 2009, but between 2009 and 2020 it increased from a 0 to a +2 value—a statistically significant increase over the last 20 years.

3.4. Warming and Glacier Melting

The northwest and northeast parts of Mt. Munkhkhairkhan are cooler and have more rainfall compared to the southwest and south-east, which is one of the reasons for the

concentration of glaciers near the northern part of the mountain. In comparison, the annual average temperature and precipitation for the Altai (south-east) are 2.8 °C and 69 mm and for Uyench (south-west) they are 2.4 °C and, 71.8 mm, compared with—in the north-west—Munkhkhairkhan at 0.7 °C and, 110 mm and 1.1 °C, and 137.1 mm in the northwest of the mountain (Table 4 and Figure 8).

Table 4. The long-term average temperature and precipitation of the weather stations closest to Mt. Munkhkhairkhan (1990–2020).

Station	T Average (°C)	T Max (°C)	T Min (°C)	Precipitation (mm)
Bulgan (Baitag)	3.1	39	−49.5	80.7
Munkhkhairkhan	0.7	30.9	−36.3	110.6
Must	0.7	31.9	−38.8	87.6
Altai	2.8	43.6	−42.9	69
Uyench	2.4	39.6	−39.4	78.1
Duut	−1.9	32.7	−41.4	113.9
Bulgan (Duchinjil)	−1.1	32.5	−40.5	137.1
Deluun	−1.9	31.1	−42.7	110.7

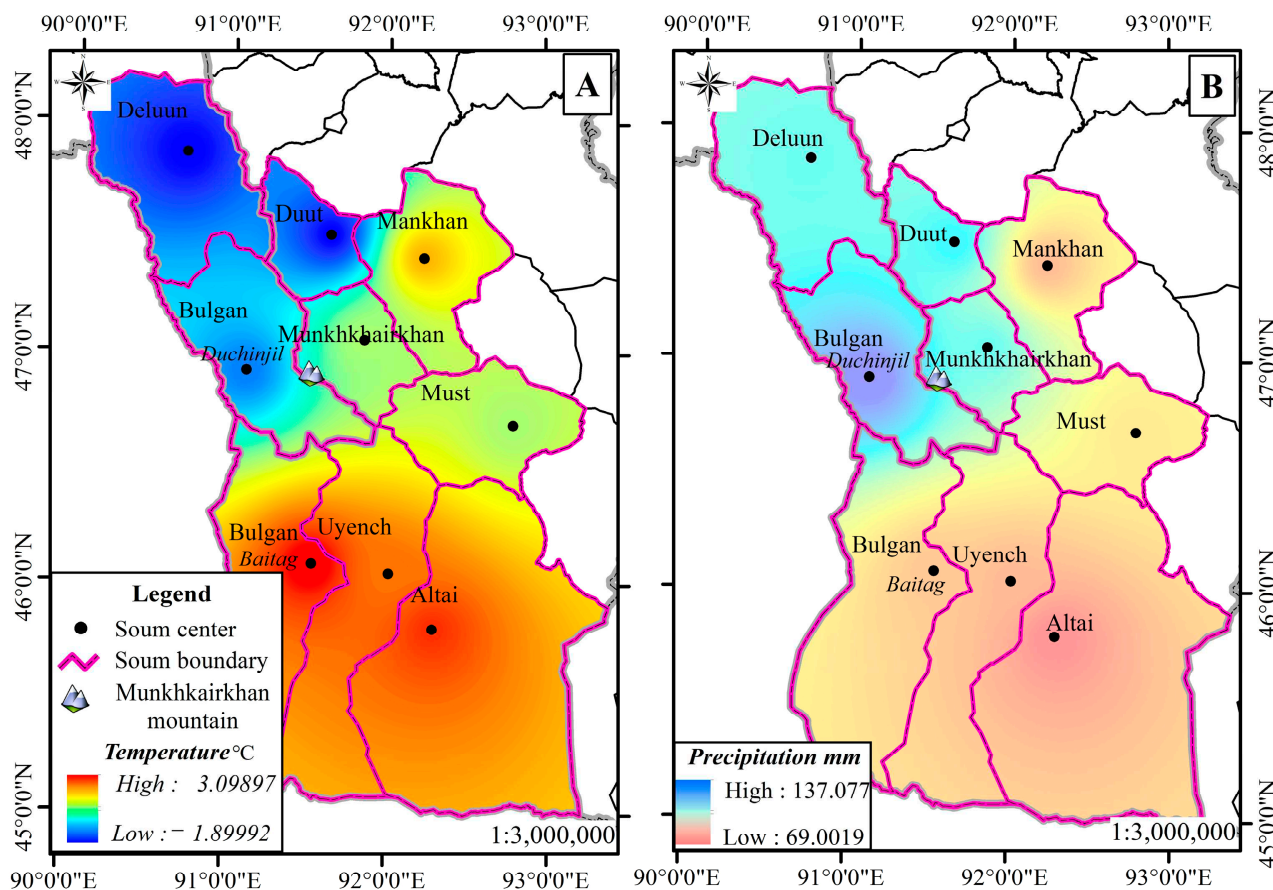


Figure 8. The long-term average temperature (A) and precipitation (B) of the weather stations closest to the mountain.

The geographical position of Mt. Munkhkhairkhan is in a dry climate from the south of the Trans-Altai Gobi, and according to data from the Baitag station is 5 °C higher than other north, northwest, and north-eastern exposures. Precipitation at the Baitag and Altai stations was 68–70 mm lower than at the other stations. Consequently, glaciers in the southern exposure were absent (Table 4 and Figure 8).

When the air temperature was lower than average, and the precipitation was higher than average, the melting of the glaciers was lower (2002, 2015, 2017)—whereas when the temperature was higher than average, and the precipitation was lower, the melting of the glaciers was higher (2004, 2010; Figure 9).

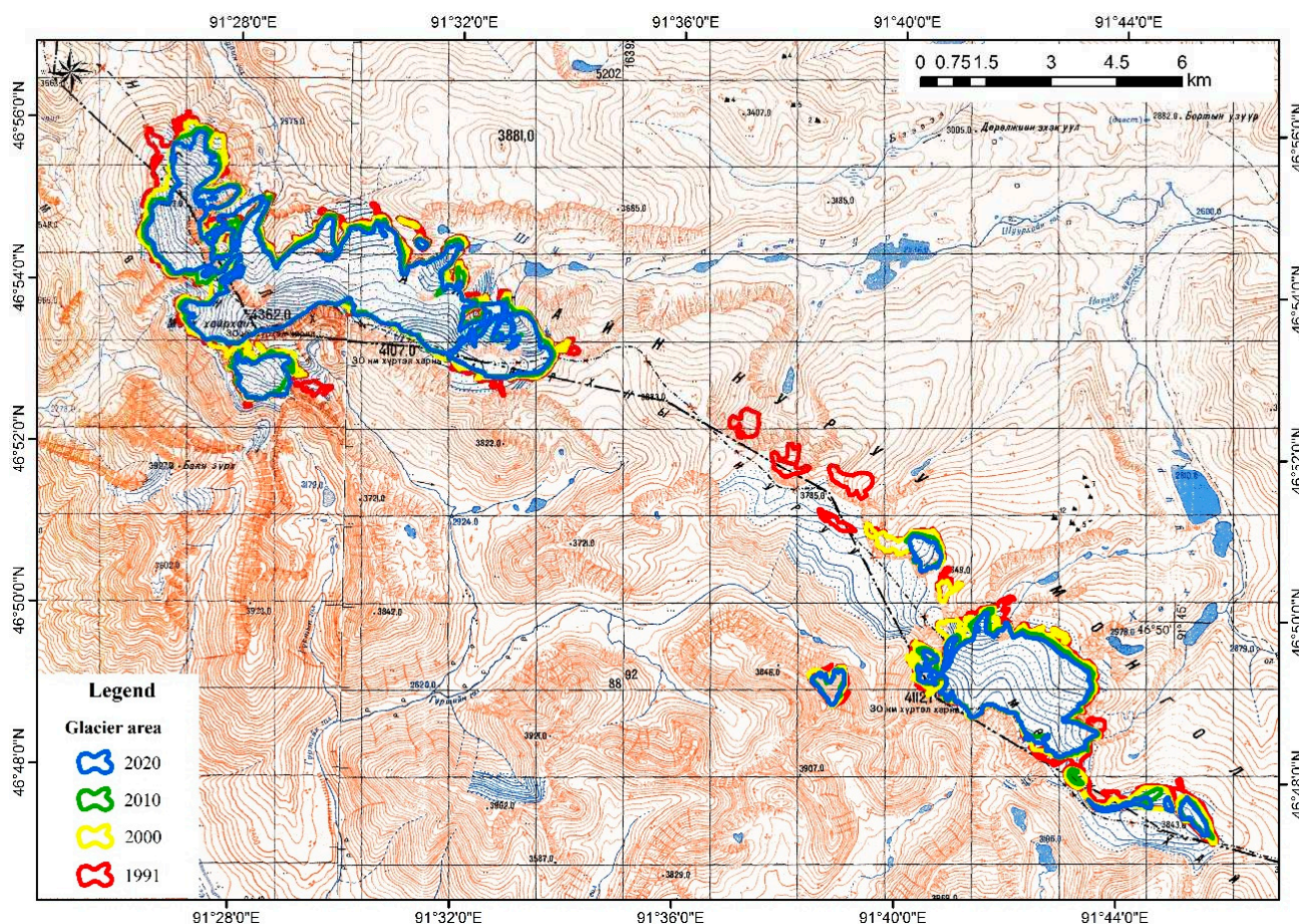


Figure 9. Glacier area Changes over the last 30 years (1991–2020).

Due to the glacial melting, there is a high risk of ecosystem change for the surrounding areas. Near the Musun Buurug Lake, Shuurkhain Basin on Mt. Munkhkhairkhan, permafrost thawing is clearly observable. A thermokarst has formed in the lake area due to the melting of permafrost at the lake’s shore (Figure 10A–D). Due to the retreat of the glacier, the area of the Musun Khanatai Lake located near the glacier of the basin of Doloonuur has increased by 48 percent in the last 30 years (1990–2020).

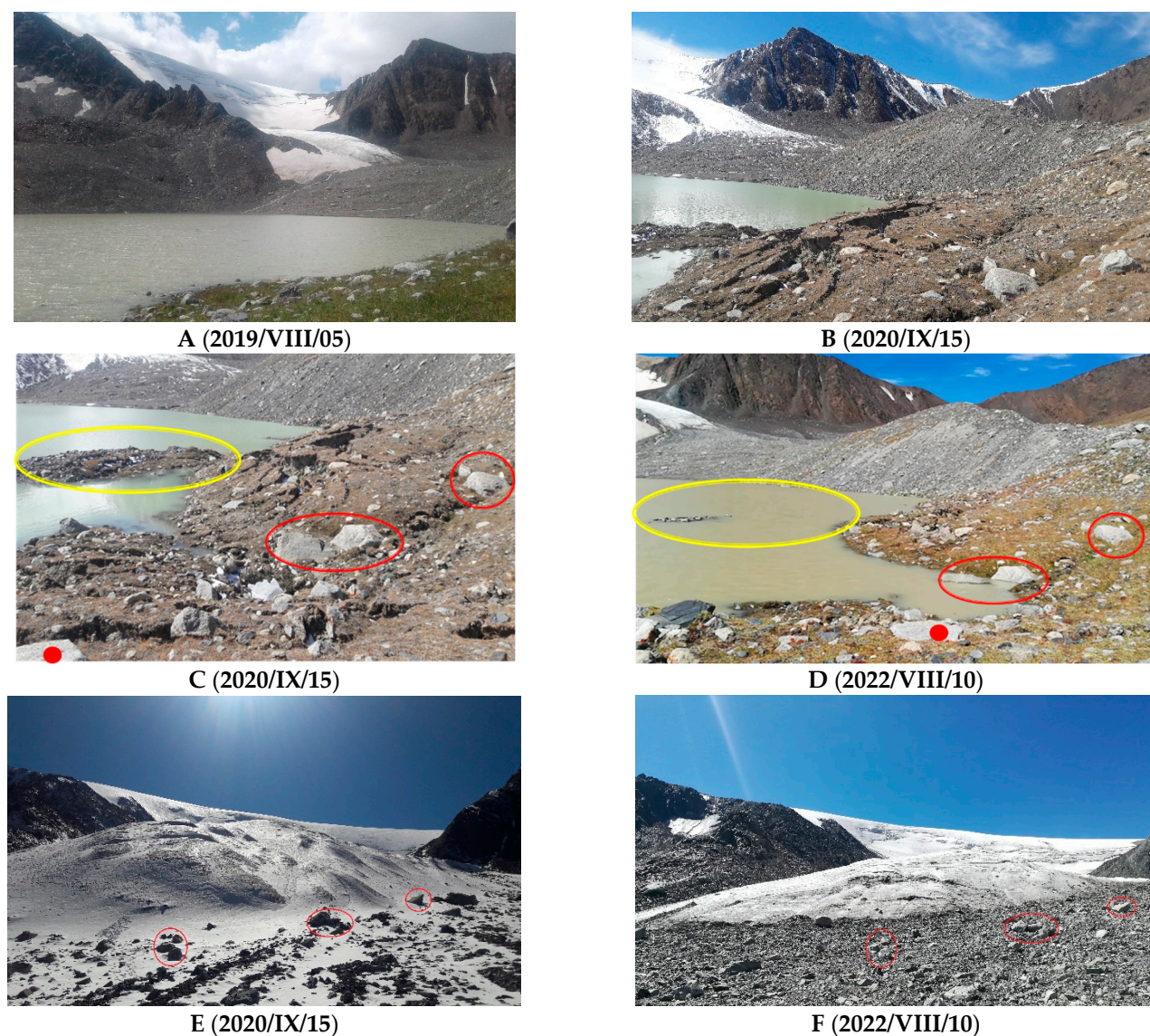


Figure 10. Glacier degradation and thermokarst along the lake shore. Red and yellow rings and red dots denote field observation markers, 2019–2022 (Bayarmaa.M., 2019, 2020, 2022).

Starting from 2019, we marked defining points for measuring the retreat of the glaciers of the Shuurkhai and Doloonuur glaciers and lakes. Observations were made on how ice and snow processes affect the retreat of these glaciers and the morphometry of the lakes. Thus, between 2019 and 2022, the Shuurkhai glaciers had retreated from the front of the tongue of the glaciers by 26.7 m and from the Doloonuur glaciers by 28.8 m. Between 2019 and 2022, Shuurkhai River Valley glacier No.3 retreated by 26.7 m, and Doloonuur River Valley glacier No. 7 by 28.8 m (Figure 10A,B,E,F).

4. Discussion

Important indicators of climate change are air temperature increases and glacial melting. Glacier mass balance is the outcome between temperatures that melt ice and winter precipitation as snow that forms ice. Mass balance is the controlling of both the volume of glaciers and their area. Counter-intuitively, the number of glaciers can increase when mass balance decreases because glacier “branches” can be detached from the main body of a glacier [37].

P. Baast made the first attempt to inventory some glaciers of the Mongolian Altai in 1998. In 2012, D. Otgonbayar, inventoried glaciers of Mt. Sutai, Tsambagarav, and

Munkhkhairkhan. D. Otgonbayar calculated a total of 24 glaciers covering 48.13 km² in 1970, and again in 2008 when 17 glaciers covered 26.57 km² [13,15]. Additionally, Kamp and Pan independently estimated 27.4 km² glacial area in 2011. By the estimation of Ganyushkin, 26.86 km² glaciers were present [17]. Ganyushkin and other researchers studied glacial retreat on the Altai ranges of Mongolia and Russia from the Little Ice Age. They estimated that during the Little Ice Age, 78.67 ± 2.82 km² of Mt. Munkhkhairkhan was covered by glaciers [17].

Pan and other researchers calculated that the glacial area of the Mongolian Altai decreased by 43% between 1990 and 2016, and the maximum glacial retreat occurred during the warmest summer temperatures [14].

Our study demonstrates that Mt. Munkhkhairkhan had a 24.5 km² glacial area as of 2020. The distribution of glaciers across altitudes shows that 97.86% were located above 3400 m a.s.l and, between 1990 and 2020, the glacier area located below 3600 m a.s.l decreased by 70% on average. The total area of glaciers decreased by 32.3% between 1990 and 2020.

The regression analysis showed that changes in air temperature, precipitation, and glacial area were strongly correlated, and an increase in annual precipitation during the warm season in summer and autumn in recent years may be able to reduce the rate of glacial melting although the overall average temperature increased.

As a result of this study, we have determined different periods of climate and glacier area changes from 1990 to 2020. These periods include:

1990–2000: The first decade was extremely warm, dry, and glacier melting was intense (1.2 km²).

2000–2010: Cooler, wetter mid-decade; less glacier melting compared to the previous decade (0.9 km²).

2010–2020: The last decade has been extremely warm, but with high precipitation and low glacier melting (0.3 km²; Figures 6 and 8).

This pattern is close to the X and XI stages of glacier melting in the Altai Dry region. According to Ganyushkin, stage X (1995–2010) was “Extremely warm and dry, intensive melting of glaciers”, and stage XI was “Cool and humid, so melting of glaciers is low” [38]. The melting of glaciers has many important implications for the landscape and for local people. Glacial melt water increases the water level of glacial lakes, the percolation of lake water into the soil, and the thawing of permafrost. Additionally, due to this warming, the thermokarst processes are activated near the shores of glacial lakes, disturbing ecosystems and ecosystem processes [21]. Thawing permafrost carries the risk of lake dams collapsing in the future.

The long-term trends in temperature and precipitation were negatively related. While the average temperature increased from 1990 to 2003, the volume of precipitation declined. When air temperature decreased, annual precipitation amounts increased (Figure 8).

Multiple regression results showed a strong correlation between Temperature, Precipitation, and Glaciers (Multiple R = 0.69, R² = 0.48). Rutherford indicated that Temperature ($t = -2.332$, $p = 0.036$) and Precipitation ($t = -3.212$, $p = 0.007$) were significant predictors in the model. Air temperature and precipitation explained 48 percent of the changes in glacier area, and R = 0.69, which is a strong correlation.

R square (R²) equalled 0.48; this means that the predictors (temperature, and precipitation) explained 48.8% of the variance in Y (Glacier). The coefficient of multiple correlations (R) equalled 0.69; this means that there was a strong correlation between the predicted data (\hat{y}) and the observed data (y). In order to better determine the relationship between glacier climates, we selected 16 samples and used the average values for each period (calculated with 1–3 year intervals of glacier area changes, and 1–3 year average values of air temperature and total precipitation for that period) to calculate correlations. Overall regression: right-tailed, $F(2,13) = 6.20$, p -value = 0.01. Since p -value < α (0.05), we rejected the H₀. All the independent variables (X_i) were significant. The Y-intercept (b): two-tailed, $T = 3.97$, p -value = 0.001. “b” was significantly different from zero.

Potential glacial outburst floods and changes in glacier run-off are likely to affect local people who camp in the mountains during herding activities and local communities that depend on glacial melt water for drinking. At a global scale, glaciers in remote, arid regions such as Mongolia are relatively under-studied, and yet are probably at the greatest risk of being lost during global warming. Therefore, it is necessary to conduct continuous monitoring studies of glaciers and glacial outburst lakes.

5. Conclusions

In this research, we determined the changes in the modern glaciated area of Mt. Munkhkhairkhan, located in the southern part of the Mongolian Altai Mountains, and the temperature and precipitation mode of 1990–2020. We also determined the impact of glacial melt on the highlands.

The Mann–Kendall (MK) test, Innovative Trend Analysis Method (ITAM), and Sen's Slope Estimator Test were used for precipitation and temperature analyses of Mt. Munkhkhairkhan's glaciers, and the remote sensing method was used for estimating glacial area.

Changes in temperature and precipitation over the last 30 years were studied. The average temperature increased by 0.5 °C overall, but there were decreases over a short period of time after which it continued to grow. Precipitation levels tended to decrease between 1990 and 2000, but in the last two decades, they have slightly increased.

According to the observations of the field research, rainfall in recent years has been increasing in the form of storms. Additionally, extreme weather changes (excessive heat, sudden cold, snowfall) have been observed in the summer, while the winter season has been warmer than average for many years.

As of 2020, there was a 24.5 km² total area of glacier at Mt. Munkhkhairkhan—most of it in the north part and above an altitude of 3400 m a.s.l. The area decreased by 13.8% between 1991 and 2000, by 16.02% between 2000 and 2010, and by 6.48% between 2010 and 2020. From 1991–2020, it decreased by 32.3% or 11.7 km². The results show that the glacial area melted by an average rate of 1.07% per year.

Glacier melting was intense in 1991–2010, when the air temperature increased and there was more precipitation. The temperature rate is expected to rise, but since 2000, due to increases in annual precipitation—mainly as snow—the rate of glacial melting has decreased.

Glaciers are very important for maintaining the balance of ecosystems in arid and semi-arid regions and are very sensitive to climate change, so it is important to continue glacier monitoring research. The glaciers and snow area in the study area have changed due to climate warming and precipitation changes and are located in arid and semi-arid regions of Central Asia. This study of Mt. Munkhkhairkhan shows that climate change significantly impacts glaciers and snow.

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