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The Impact of Ground Heat Capacity on Drinking Water Temperature

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15 ABSTRACT

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Temperature is known to impact physical, chemical and biological processes in Drinking Water Distribution Systems (DWDS), but it is rarely considered or modelled. This research evaluates the impact of considering a finite heat capacity for the ground, which has been assumed infinite in previous DWDS research. The aim of this work is to explore and quantify the region where the difference between considering infinite or finite heat capacity for the ground is significant, i.e. the distance over which water-ground heat transfer interaction is important. A detailed model comparison is carried out for key pipe materials, diameters and hydraulic conditions. Temperature effects are found to exist for up to tens of kilometres (i.e. several hours) into the DWDS. While the differences found were only a few degrees Celsius, this will affect all reaction rates, such as chlorine decay, and is at the start of the DWDS so will impact the entire downstream network. This work highlights the importance of considering temperature in DWDS, and in particular the finite heat capacity of the ground, in ensuring the provision of safe drinking water.

28 INTRODUCTION

Temperature affects all physical, chemical and biological processes occurring within Drinking 29 Water Distribution Systems (DWDS). In general, temperature increase is associated with water 30 quality deterioration. For example, it is well-known that temperature increases the rate at which 31 chlorine decays (Fisher et al. 2012; Monteiro et al. 2017). This is important for the operation of 32 DWDS, as many countries aim to ensure a minimum target level of disinfectant at the customer tap to 33 minimise risks to public health (WHO 2011). Temperature is also known to affect bacterial-fungal 34 communities (Calero et al. 2021) and precipitation reactions like iron and manganese (Mounce 35 et al. 2016), potentially contributing to discolouration. At present, there is hardly any temperature 36 monitoring, so it is not easy to see trends and/or quantify the implications of temperature increase 37 on water quality (Agudelo-Vera et al. 2020). The data that is available shows that there is correlation 38 with water quality and air temperature, e.g. higher discolouration contacts during summer (van 39 Summeren et al. 2015; Cook et al. 2015). This justifies the current need to monitor and better 40 understand temperature dynamics within DWDS, especially in the face of climate change challenges. 41 Hydraulic and water quality models usually consider a constant temperature (Fisher et al. 2012). 42 According to Agudelo-Vera et al. (2020), only Blokker and Pieterse-Quirijns (2013) and Piller and 43 Tavard (2014) present validated models to compute water temperature across DWDS. Both rely 44 on assuming a fixed ground temperature that constitutes the boundary condition for heat transfer 45 to the fluid. This is equivalent to assume an infinite heat capacity for the ground, which remains 46 unaffected by heat exchange to and from the water within the pipes. This simplification enables 47 a decoupling of the problem: (1) calculate the undisturbed ground temperature, and (2) calculate 48 the water temperature based on the previously estimated ground temperature (typically a daily 49

average). Blokker and Pieterse-Quirijns (2013) propose a micrometeorology model to compute the 50 undisturbed ground temperature and then use EPANET-MSX (Shang et al. 2008) to simulate water 51 temperatures. After analysing several Dutch case studies, they conclude that water reaches the 52 ground temperature at a rate that depends on pipe diameter, pipe thickness, pipe material and flow 53 velocity. This simplified heat transfer model enables identification of overall water temperature 54 behaviours: water temperature in transport mains stays similar to that at the inlet, whereas water 55 temperature in distribution mains approaches the undisturbed ground temperature. What happens in 56 between, i.e. how large and significant is the transition zone, has not been specifically addressed. To 57 compute this region, it is important to acknowledge that in reality ground has a finite heat capacity 58 and the temperature of the ground around a pipe is affected by the drinking water temperature 59 (and vice versa), i.e. it is a coupled problem. In winter, input water is usually colder than the 60 ground, so water heats over the pipeline and the surrounding ground loses temperature in the 61 process. In summer, incoming water is likely to be warmer than the ground, so water cools along 62 the pipeline and the surrounding ground heats in the process. Even though the few studies that 63 address temperature modelling at DWDS usually make the infinite ground heat capacity hypothesis 64 (Blokker and Pieterse-Quirijns 2013; Piller and Tavard 2014), it is known that the heat transfer 65 process is a complex phenomenon (Agudelo-Vera et al. 2020). 66

The importance of fluid-ground interaction has been studied for other buried pipe infrastructure 67 systems where the temperature difference between the ground and the fluid is larger and so it is not 68 acceptable to assume infinite heat capacity for the ground. This is the case of sewer systems, where 69 temperature modelling has become important to assess the potential of heat recovery applications. 70 In these systems, it is usual to assume a penetration depth for heat exchange around the sewer. 71 The ground temperature is usually measured, and the depth of this influence area (where heat 72 conduction through the ground takes place) is usually adjusted (Durrenmatt and Wanner 2008; 73 Abdel-Aal et al. 2014). This approach is effectively approximately equivalent to considering a 74 finite heat capacity for the ground within a zone of influence. It is highly dependent on field 75 measurements, so its application is empirical, relying on physically measured temperature data 76

from the ground around pipes. Ground-fluid temperature interaction has also been studied in the 77 context of Ground Source Heat Pumps (GSHP). These systems are specifically designed to exploit 78 the temperature difference between the circulating fluid and the ground for different purposes (Soni 79 et al. 2015). For these applications it is essential to consider the finite heat capacity of the ground 80 to correctly model thermal interaction. Horizontal GSHP extend beneath the ground surface in a 81 similar way to DWDS. Their behaviour can be modelled experimentally, numerically or analytically 82 (Gan 2019). Analytical approaches are especially interesting because they have potential to provide 83 a conceptual framework for systematic assessment. Most analytical approaches compute fluid 84 temperature variation along the pipe according to the Finite Line Source (FLS) model: the pipe 85 behaves as a line that releases or receives heat within a semi-infinite ground domain (Claesson 86 and Dunand 1983; Fontaine et al. 2011). GSHP have a well-established steady analysis. Unsteady 87 analysis involves a temporal convolution that is time-consuming to solve (Lamarche 2017). Several 88 researchers have proposed the use of accelerating schemes to solve this convolution, like the Fast 89 Fourier Algorithm (Marcotte and Pasquier 2008). Such an approach requires knowledge of the heat 90 history beforehand, which is not trivial in systems that run near the ground surface (both horizontal 91 GSHP and DWDS), because they are exposed to weather variations (Lamarche 2017). Lamarche 92 (2019) has recently proposed a non-history scheme to compute the heat transfer and output fluid 93 temperature based on the input fluid temperature evolution over time. This unsteady ground model 94 could be applied at DWDS to assess the impact of water-ground heat transfer interaction at different 95 periods of the year. 96

The main objective of this paper is to assess the importance of ground heat capacity on drinking water temperature. This is possible by fulfilling three specific aims. First, to compare the water temperatures obtained with the usually adopted decoupled model (infinite ground heat capacity) and a more realistic coupled approach (finite ground heat capacity). The importance of making one assumption or the other can be assessed by estimating the transition region. The transition region is here defined as the distance or equivalent residence time required for the water to attain the undisturbed ground temperature. It represents the region where heat interaction is important.

Therefore, the second aim of this work is to quantify the transition region. We aim to derive an 104 explicit expression to compute the transition region by assuming steady flow and ground conditions, 105 but requiring simplification of the less tractable unsteady ground behaviour. Thus, the third aim 106 of this paper is to assess if the analytical expression derived when considering steady ground 107 conditions can approximate transition regions when considering the annual cycle of the ground 108 and input water temperatures. The novelty of this work lies in providing a conceptual framework 109 to better describe the steady and unsteady water-ground heat transfer interaction of DWDS. Our 110 hypothesis is that there is a significant transition region and hence that assuming infinite ground 111 heat capacity is not a good enough approximation for a significant part of the DWDS. 112

113 **METHODOLOGY**

The steady and unsteady approaches presented in this section build on the principles of ground heat transfer. Note that steady flow conditions will be assumed as a first approximation to the complex heat interaction problem.

117 2.1 Ground heat transfer principles

Figure 1 shows a pipe buried in the ground and exchanging heat between the ground and the fluid that is being transported. Assuming that the ground is a homogeneous semi-infinite medium, its temperature distribution T(t, x, y, z) behaves according to the heat conduction equation (Lamarche 2019):

$$\frac{1}{\alpha_{ground}} \frac{\partial T(x, y, z, t)}{\partial t} = \nabla^2 T(x, y, z, t)$$
(1)

123 with

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122

$$T(x, y, 0, t) = T_{surf}(t)$$
⁽²⁾

$$q''(t) = -k_{ground} \frac{\partial T(x, y, z, t)}{\partial n} \Big|_{r=r_p}$$
(3)

¹²⁷ Note that *T* represents the temperature field in the ground and $\alpha_{ground}[m^2/s] = \frac{k_{ground}}{\rho_{ground} \cdot C_{ground}}$ ¹²⁸ is the thermal diffusivity of the ground, which can be obtained by dividing its conductivity ¹²⁹ $k_{ground}[W/m/K]$ by its density $\rho_{ground}[kg/m^3]$ and specific heat capacity $C_{ground}[J/kg/K]$. Eq. (2) represents the boundary condition at the surface $T_{surf}(t)$ and Eq. (3) the heat exchange with the pipe, where $q''[W/m^2]$ is the heat flux per unit area and $r_p[m]$ the pipe radius (radial coordinates).

¹³³Due to the linearity of the heat equation, the solution to the problem is typically computed by ¹³⁴making use of the superposition principle (Claesson and Dunand 1983). According to this principle, ¹³⁵the original problem can be divided in two (see Figure 1): (1) computing the temperature field ¹³⁶associated with the heat extraction/release to the pipe, assuming that there is a zero-temperature ¹³⁷boundary condition at the ground surface, and (2) computing the temperature field associated with ¹³⁸the changing surface temperature, as if there was no pipe:

$$T(x, y, z, t) = {}^{1}T(x, y, z, t) + {}^{2}T(z, t)$$
(4)

Note that problem 2 aims to compute the temperature field of the undisturbed ground at a depth *z*, disregarding the presence of the pipe. Therefore, it can be solved by assuming a surface temperature behaviour $T_{surf}(t)$, which propagates through the ground. Different models can be adopted to simulate the surface temperature. For example, it can be modelled by assuming an annual sinusoidal variation at the ground boundary (Lamarche 2019):

$$T_{surf}(t) = T_0 - A \cdot \cos(\omega \cdot (t - t_{shift}))$$
(5)

Kusuda and Achenbach (1965) proposed an analytical solution for this boundary condition:

¹⁴⁷
$${}^{2}T(z,t) = T_0 - A \cdot \exp\left(-z\sqrt{\frac{\omega}{2 \cdot \alpha_{ground}}}\right) \cdot \cos\left(\omega \cdot (t - t_{shift}) - z\sqrt{\frac{\omega}{2 \cdot \alpha_{ground}}}\right)$$
(6)

¹⁴⁸ Where $T_0[^{\circ}C]$ is the mean ground surface temperature, $A[^{\circ}C]$ is the variation amplitude of temper-¹⁴⁹ ature at the surface, $\omega = 2\pi/8760 h^{-1}$ is the annual frequency, $t_{shift}[h]$ is the time for the coldest ¹⁵⁰ day of the year and $\alpha_{ground}[m^2/h]$ is the thermal diffusivity of the ground. Thermal diffusivity is ¹⁵¹ here expressed in m^2/h to be consistent with annual simulations, which are typically carried out

every hour when analysing ground temperatures (Lamarche 2019). Thermal diffusivity is a critical 152 parameter when computing the undisturbed ground temperature distribution and is here assumed 153 constant over the simulation period. Sand is typically used as backfill for pipe installation. Table 1 154 shows typical ground parameters for an average dry and wet sand (Blokker and Pieterse-Quirijns 155 2013). Figure 2 shows the corresponding temperature distributions over a year according to Eq. 156 (6) at different depths. More complex models exist and could be used to simulate the undisturbed 157 ground temperature, leading to different shapes in Figure 2, but they will not impact the core of 158 the questions that this paper aims to answer. Note that pipes are usually buried at depths between 159 0.6-2.5 m depending on the country (Agudelo-Vera et al. 2020). Blokker and Pieterse-Quirijns 160 (2013) observed that there is a damping effect of air temperature daily variations at a 1 m depth in a 161 case study in The Netherlands. This supports the assumption of only considering annual variations 162 for the ground temperature at usual pipe depths. 163

164 **2.2 Steady ground model**

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Section 2.1 has shown that the ground experiences annual changes over the year. However, a simplified steady state approach can be assumed to start simulating the heat exchange when considering periods of days or weeks. The fluid temperature and heat variation along a horizontal pipe has been studied before for the steady-state case (Claesson and Dunand 1983). Fluid temperature decays exponentially along the pipe (Fontaine et al. 2011), so the water temperature when considering a ground surface temperature equal to 0°C (problem 1 in Figure 1) would be equal to:

$$T_w(x) = T_{in} \cdot \exp\left(\frac{-x}{Q \cdot \rho_w \cdot C_w \cdot R}\right)$$
(7)

¹⁷² Where $T_w(x)[^{\circ}C]$ represents the water temperature at a distance *x* from the inlet, $T_{in}[^{\circ}C] = T_w(0)$ ¹⁷³ represents the water temperature at the inlet, $Q[m^3/s]$ is the water volume flow rate, $\rho_w[kg/m^3]$ is ¹⁷⁴ the fluid density, $C_w[J/kg/K]$ is the water specific heat capacity and $R[m \cdot K/W]$ is the unit length ¹⁷⁵ thermal resistance. Eq. (7) can be referenced to the undisturbed ground temperature (problem 2 in ¹⁷⁶ Figure 1) by applying superposition:

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$$T_w(x) = (T_{in} - T_{ground}) \cdot \exp\left(\frac{-x}{Q \cdot \rho_w \cdot C_w \cdot R}\right) + T_{ground}$$
(8)

Where $T_{ground}[^{\circ}C] = {}^{2}T(z_{p}, \infty)$ represents the undisturbed ground temperature at the installation 178 pipe depth z_p in an infinite time (i.e. steady state). Note that the installation depth is here assumed 179 equal to the difference in elevation between the ground surface and the center line of the pipe. This is 180 a simplification required for the analytical steady approach to be consistent with the unsteady ground 181 model presented in Section 2.3. Figure 2 shows the undisturbed ground temperature evolution at 182 the center line of a 300 mm diameter pipe buried at $z_p = 1$ m, as well as at its top (z = 0.85 m) 183 and bottom (z = 1.15 m). The maximum difference is approximately 1°C, i.e. there are $\pm 0.5^{\circ}C$ 184 deviations with respect to the temperature at the center of the pipe through the year. At the same 185 time, the installation depth refers to the elevation of the center line of the pipe in the middle of the 186 pipeline. 187

Thermal resistance can be computed as a summation of resistances (Çengel and Ghajar 2011),
 which is different depending on whether the heat transfer interaction is considered or not:

$$R = \frac{\ln(\frac{2\cdot z_p}{r_o})}{2\cdot \pi \cdot k_{ground}} + \left(\frac{\ln(\frac{r_o}{r_i})}{2\cdot \pi \cdot k_{pipe}} + \frac{1}{Nu \cdot k_w \cdot \pi}\right) = R_{ground} + R_{pipe}$$
(9a)

$$R = \left(\frac{\ln(\frac{r_o}{r_i})}{2 \cdot \pi \cdot k_{pipe}} + \frac{1}{Nu \cdot k_w \cdot \pi}\right) = R_{pipe}$$
(9b)

Where $r_o[m]$ and $r_i[m]$ represent the inner and outer pipe radius, $k_{pipe}[W/m/K]$ is the pipe conductivity, $k_w[W/m/K]$ is the water conductivity and Nu[-] is the Nusselt number, which must be computed as:

$$Nu = \begin{cases} 3.66 & \text{if } Re < 2300\\ \frac{(f/8) \cdot (Re-1000) \cdot Pr}{1+12.7 \cdot (f/8)^{0.5} \cdot (Pr^{\frac{2}{3}}-1)} & \text{if } Re \ge 2300 \end{cases}$$
(10)

where f[-] represents the friction factor, Re[-] is the Reynolds number and Pr[-] corresponds

¹⁹⁷ to the Prandtl number:

$$Re = \frac{\rho_w \cdot v \cdot 2 \cdot r_i}{\mu_w} \tag{11}$$

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$$Pr = \frac{\mu_w \cdot C_w}{k_w} \tag{12}$$

with v[m/s] being equal to the water velocity and $\mu_w[kg/m/s]$ representing the dynamic viscosity of water.

Note that Eq. (9a) includes three terms that correspond to heat conduction through the ground, 203 heat conduction through the pipe wall and convection through water, respectively. In other words, 204 the first term is related to the ground $R_{ground}[m \cdot K/W]$ (Claesson and Dunand 1983; Fontaine et al. 205 2011; Lamarche 2019) and the other two (within brackets) refer to the pipe $R_{pipe}[m \cdot K/W]$. All 206 terms should be considered when analysing the water-ground interaction (i.e. Eq. 9a to simulate 207 the ground finite heat capacity), but only the last two terms (R_{pipe}) should be considered when 208 disregarding this interaction (i.e. Eq. 9b to assume an infinite heat capacity for the ground). The 209 implications of assuming Eq. (9a) or (9b) will be discussed in Section 3.1, where Eqs. (8) and (9) 210 will be applied to assess the temperature evolution over a pipeline under steady state conditions. 211

Eqs. (8) and (9) can be rearranged to explicitly compute the transition length $(L_t[m])$. This length is quantified by identifying the position where the absolute error between the undisturbed ground temperature and the water temperature is lower than a specified tolerance, i.e. $|T_w - T_{ground}| \le tol$. This leads to:

/|T

$$L_{t} = Q \cdot \rho_{w} \cdot C_{w} \cdot \left[R_{ground} + R_{pipe}\right] \cdot \ln\left(\frac{|T_{in} - T_{ground}|}{tol}\right) =$$

$$= Q \cdot \rho_{w} \cdot C_{w} \cdot \left[\frac{\ln(\frac{2 \cdot z_{p}}{r_{o}})}{2 \cdot \pi \cdot k_{ground}} + \left(\frac{\ln(\frac{r_{o}}{r_{i}})}{2 \cdot \pi \cdot k_{pipe}} + \frac{1}{Nu \cdot k_{w} \cdot \pi}\right)\right] \cdot \ln\left(\frac{|T_{in} - T_{ground}|}{tol}\right) \quad (13a)$$

$$L_{t} = Q \cdot \rho_{w} \cdot C_{w} \cdot \left[R_{pipe}\right] \cdot \ln\left(\frac{|T_{in} - T_{ground}|}{tol}\right) =$$

$$= Q \cdot \rho_{w} \cdot C_{w} \cdot \left[\left(\frac{\ln(\frac{r_{o}}{r_{i}})}{2 \cdot \pi \cdot k_{pipe}} + \frac{1}{Nu \cdot k_{w} \cdot \pi}\right)\right] \cdot \ln\left(\frac{|T_{in} - T_{ground}|}{tol}\right) \quad (13b)$$

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These transition lengths could be turned into equivalent residence times as L_t/v , where v[m/s] is

the water velocity assuming average constant velocity. Note that Eq. (13a) provides the transition 219 length when considering finite heat capacity for the ground (i.e. coupled model) and Eq. (13b) 220 provides the equivalent assuming infinite heat capacity for the ground (i.e. decoupled model). 221 As explained in the Introduction, assuming infinite heat capacity is the assumption made in the 222 few existing applications for temperature modelling in DWDS and as has been coded to simulate 223 network behaviour taking advantage of the EPANET-MSX functionality. Reality is coupled, and 224 the difference between results obtained with Eqs. (13a) and (13b) will show the implications of 225 simplifying by assuming infinite heat capacity. 226

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2.3 Unsteady ground model

Section 2.1 has highlighted that seasonal changes in air temperature penetrate to typical DWDS pipe burial depths. Input water temperatures are also expected to change over the year according to temperature trends at the source water, water treatment plant and/or service reservoir. These annual changes will impact water temperature dynamics, and so the associated water quality.

The application of the ground model proposed by Lamarche (2019) to DWDS considers that 232 the volume flow rate of the fluid $Q[m^3/s]$, the input water temperature record $T_{in}(t)$ [°C] and the 233 undisturbed ground temperature distribution ${}^{2}T(z,t)[{}^{\circ}C]$ determine the total heat transfer flux at 234 each time q(t)[W]. They all condition the temperature at the end of the pipeline $T_{out}(t)[^{\circ}C]$ and 235 the temperature of the ground at the pipe surroundings $T_p(t)$ [°C]. In this work, the formulation is 236 explained for a prototypical water pipeline with a length L[m] located at a constant depth $z = z_p[m]$ 237 (horizontal pipe). Heat transfer flux cannot be assumed constant along the pipe due to the temporal 238 and spatial variations in temperature gradients, so it is split into $n_s = L/L_i$ segments as proposed 239 by Fontaine et al. (2011), where $L_i[m]$ is the length of each segment *i*. Therefore, the unknowns 240 to be solved at each time step are the input water temperature at each segment except the first one 241 $(T_{in,i}(t) [^{\circ}C]; \forall i = 2, ..., n_s)$ and the temperature of the ground at the surroundings of each pipe 242 segment $(T_{p,i}[^{\circ}\mathbf{C}]; \forall i = 1, \ldots, n_s).$ 243

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The solution depends on how the temperature reduction $\theta_i[-]$ varies along each segment. It is

defined as:

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$$\theta_{i} = \frac{\tilde{T}_{out,i}(t) - {}^{1}T_{p,i}(t)}{\tilde{T}_{in,i}(t) - {}^{1}T_{p,i}(t)}$$
(14)

with $\tilde{T}_{out,i}(t) = T_{out,i}(t) - {}^{2}T(z_{p}, t)$ and $\tilde{T}_{in,i}(t) = T_{in,i}(t) - {}^{2}T(z_{p}, t)$ according to the superposition principle and ${}^{2}T(z_{p}, t)$ given by Eq. (6). At the same time, temperature variation across each segment can be assumed to follow an exponential decay (Lamarche 2019):

$$\theta_i = \exp\left(\frac{-L_i}{Q \cdot \rho_w \cdot C_w \cdot R_{pipe}}\right) \tag{15}$$

Only $R_{pipe}[m \cdot K/W]$ is here considered as thermal resistance because the ground effect is considered in the unsteady ground model formulation. The pipe and the ground domains are later solved jointly through a system of equations (Eqs. 17 to 28). Note that assuming an exponential decay for each segment (Eq. 15) implies considering that water achieves the steady state at each time, neglecting the plug flow along the pipeline. This is a simplification (see Section 4), but it enables computation of an equivalent heat transfer coefficient $X_i[W/m/K]$ for each segment as:

$$X_i = \frac{Q \cdot \rho_w \cdot C_w}{L_i} \cdot (1 - \theta_i) \tag{16}$$

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According to Lamarche (2019), the temperature in the pipe surroundings can be computed by considering the pipe segment interaction along the pipe:

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$${}^{1}T_{p,i}(t) = S_{p,i}(t) + \sum_{j=1}^{n_s} S_{q,ij} \cdot X_j \cdot (\tilde{T}_{in,j}(t) - {}^{1}T_{p,j}(t))$$
 (17)

²⁶¹ Where $S_{p,i}(t)$ and $S_{q,ij}$ illustrate the pipe segment interaction as a result of heat conduction through ²⁶² the ground (see Eqs. 31 to 36 below). As the temperature at the beginning of each segment must ²⁶³ be equal to the temperature at the end of the previous segment (n_s in series pipes), Eq. (14) can be ²⁶⁴ rewritten as an additional condition:

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$$\tilde{T}_{in,i+1}(t) = \tilde{T}_{out,i}(t) = \theta_i \cdot \tilde{T}_{in,i}(t) + (1 - \theta_i) \cdot {}^1T_{p,i}(t)$$
(18)

Eqs. (17) and (18) can be rearranged as a system of equations where the coefficient matrix A 266 remains constant and the independent term $\mathbf{B}(t)$ and unknown vector $\mathbf{T}(t)$ change over time: 267

$$\mathbf{A} \times \mathbf{T}(t) = \mathbf{B}(t) \tag{19}$$

The coefficient matrix can be built as: 269

$$\mathbf{A} = \begin{bmatrix} \mathbf{U}\mathbf{L}(n_s \times n_s) & \mathbf{U}\mathbf{R}(n_s \times n_a) \\ \mathbf{L}\mathbf{L}(n_a \times n_s) & \mathbf{L}\mathbf{R}(n_a \times n_a) \end{bmatrix}$$
(20)

With $n_a = n_s - 1$ and: 271

$$\mathbf{UL} = \begin{bmatrix} 1 + X_1 \cdot S_{q,11} & X_2 \cdot S_{q,12} & \cdots \\ X_1 \cdot S_{q,21} & 1 + X_2 \cdot S_{q,22} & \cdots \\ \vdots & \vdots & \ddots \end{bmatrix}$$
(21)

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$$\mathbf{UR} = \begin{bmatrix} -X_2 \cdot S_{q,12} & -X_3 \cdot S_{q,13} & \cdots \\ -X_2 \cdot S_{q,22} & -X_3 \cdot S_{q,23} & \cdots \\ \vdots & \vdots & \ddots \end{bmatrix}$$
(22)

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$$\mathbf{LL} = \begin{bmatrix} -(1-\theta_1) & 0 & \cdots \\ 0 & -(1-\theta_2) & \cdots \\ \vdots & \vdots & \ddots \end{bmatrix}$$
(23)

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 $\mathbf{LR} = \begin{bmatrix} 1 & 0 & \cdots \\ -\theta_2 & 1 & \cdots \\ \vdots & \vdots & \ddots \end{bmatrix}$ (24)

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The independent term can be estimated as: 279

$$\mathbf{B}(t) = \begin{bmatrix} \mathbf{B}\mathbf{U}(t)(n_s \times 1) \\ \mathbf{B}\mathbf{L}(t)(n_a \times 1) \end{bmatrix}$$
(25)

281 With:

$$\mathbf{BU}(t) = \begin{bmatrix} S_{p,1}(t) + X_1 \cdot S_{q,11} \cdot \tilde{T}_{in,1}(t) \\ S_{p,2}(t) + X_1 \cdot S_{q,21} \cdot \tilde{T}_{in,1}(t) \\ \vdots \end{bmatrix}$$
(26)

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$$\mathbf{BL}(t) = \begin{bmatrix} \tilde{T}_{in,1}(t) \cdot \theta_1 \\ 0 \\ \vdots \end{bmatrix}$$
(27)

²⁸⁵ The vector of unknowns would be:

$$\mathbf{T}(t) = \begin{bmatrix} {}^{1}\mathbf{T}_{\mathbf{p}}(t)(n_{s} \times 1) \\ \tilde{\mathbf{T}}_{\mathbf{in}}(t)(n_{a} \times 1) \end{bmatrix}$$
(28)

Note that if the pipe has a different depth at the beginning and the end, the difference should be
included in Eq. (27) as suggested by Lamarche (2019).

Once the system of equations in Eq. (19) is solved at a time t, ${}^{1}T_{p,i}(t)$ and $\tilde{T}_{in,i}(t)$ are known. The temperature at the end of each segment $\tilde{T}_{out,i}(t)$ can be directly computed according to Eq. (18). The heat flux per unit length at the next time step $q'_{i}(t + \Delta t)[W/m]$ can be computed as:

$$q'_{i}(t + \Delta t) = X_{i} \cdot \left(\tilde{T}_{in,i}(t) - {}^{1}T_{p,i}(t)\right)$$
(29)

so the total heat load $q(t + \Delta t)[W]$ can be obtained as:

$$q(t + \Delta t) = \sum_{i=1}^{n_s} q'_i(t + \Delta t) \cdot L_i$$
(30)

The process must then be repeated at the next time step. Solving these equations for the whole year would provide the annual water temperature evolution along the pipe. These values could be used to compute the transition length at any time.

Note that the only terms that have not been explained yet are $S_{p,i}(t)$ and $S_{q,ij}$. According to

Lamarche (2017) and Lamarche (2019), $S_{q,ij}$ can be computed as:

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$$S_{q,ij} = \frac{1}{k_{ground}} \sum_{n=1}^{n_d} \left(1 - \exp(-d_n^2 \cdot \Delta \tilde{t}) \right) \cdot u_{ij}(d_n) \cdot \Delta d_n \tag{31}$$

Where d_n is a dummy variable ($n_d = 450$ in this work), $\Delta \tilde{t} = \frac{\alpha_{ground} \cdot \Delta t}{r_o^2}$ is a dimensionless time interval and $u_{ij}(d)$ involves the inverse Laplace transform of the so-called g-function:

$$u_{ij}(d) = -\frac{d}{\pi} \cdot \mathbf{L}^{-1}(g_{ij}(t_k))$$
(32)

The g-function is characteristic of the system. It was tabulated by Eskilson (1987) for different borehole (here equivalent to pipe) configurations, although several authors have worked on deriving analytical formulations (Zeng et al. 2003; Lamarche and Beauchamp 2007). Note that the g-function has been here expressed as a function of t_k because the Gavesh algorithm (Stehfest 1970; Villinger 1985) used to compute the associated Laplace transform needs to evaluate the function at some unknown times:

$$g_{ij}(t_k) = \frac{1}{2} \int_{\frac{1}{2\sqrt{\alpha_{ground} \cdot t_k}}}^{\infty} \frac{\exp(-r^2 \cdot s^2) - \exp(-r_{imag}^2 \cdot s^2)}{s^2} \cdot \{ierf[(\Delta x + L_j) \cdot s] - ierf[(\Delta x + L_j - L_i) \cdot s]\} \cdot ds$$

$$(33)$$

Where Δx refers to the difference between the initial coordinates of segment *i* and segment *j* ($\Delta x = x_{0,i} - x_{0,j}$), $r = r_o$, $r_{imag} = 2 \cdot z_p$, and the *ier f* function is:

³¹³
$$ierf(a) = a \cdot erf(a) - \frac{1}{\sqrt{\pi}} \cdot (1 - \exp(-a^2))$$
 (34)

On the other hand,
$$S_{p,i}(t)$$
 can be calculated as:

$$S_{p,i}(t) = \frac{1}{k_{ground}} \sum_{n=1}^{n_d} \exp(-d_n^2 \cdot \Delta \tilde{t}) \cdot F_i(\tilde{t}, d_n) \cdot \Delta d_n$$
(35)

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³¹⁶ Where $F_i(0, d_n) = 0$ and:

$$F_i(\tilde{t} + \Delta \tilde{t}, d_n) = \exp(-d_n^2 \cdot \Delta \tilde{t}) \cdot F_i(\tilde{t}, d_n) + \sum_{j=1}^{n_s} q'_j(\tilde{t} + \Delta \tilde{t}) \cdot (1 - \exp(-d_n^2 \cdot \Delta \tilde{t})) \cdot u_{ij}(d_n)$$
(36)

318 **RESULTS**

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In this section, both the steady and the unsteady ground models will be applied to a case study. 319 It is important to highlight that both models have been presented assuming that steady flow (i.e. 320 constant volume flow rate) takes place within the pipe. This implies that the formulations here 321 proposed can only be applied where the flow can be approximated to steady state (i.e. assuming 322 approximation to average daily flow), while the ground surface and the input water temperatures 323 may experience changes according to the annual cycle. This simplification is reasonable when 324 seasonal effects are of interest and daily flows and/or temperature variations are not significant. In 325 this work, an average pipe located early in the DWDS will be analysed, as the temperature gradient 326 and so the effect of temperature exchange is greatest along the first kilometres (with how many 327 kilometres, or how much time, effects are important for a key unknown to be elucidated here). 328 Rather than trying to consider the unnecessary complexity of a real network, and in order to prevent 329 results from being arguably specific to the assumed layout, a single long pipe is considered to 330 enable generic estimation of the transition region in terms of distance and/or time. 331

This pipe is assumed to have a constant velocity v = 0.5 m/s and an internal radius $r_i = 0.15$ 332 m (volume flow rate $Q = 0.0353m^3/s$), which can be considered representative of the usual pipes 333 located early in a DWDS. The pipe is considered to be installed at a constant $z_p = 1$ m, which 334 is consistent with typical installation depths (Agudelo-Vera et al. 2020). As previously explained, 335 such a depth already implies that daily temperature oscillations are not perceived in the ground 336 (Blokker and Pieterse-Quirijns 2013). The pipe is considered to be extremely long, with L = 125337 km, so that the simulated residence times are also high (69.4 h, almost 3 days). Note that the 338 average residence time in water systems at distribution level is 24 hours (Husband et al. 2008), and 339 even greater residence times might be obtained at some points (Machell and Boxall 2012; Machell 340

and Boxall 2014). Therefore, even though the case study is a large diameter pipe with steady flow,
it will provide residence times that are representative of what could happen within the distribution
level (see Section 4 for discussion). This length is associated with significant pressure drops along
the pipeline. In real systems, different diameters and velocities exist along the pipeline, leading to
more realistic pressure distributions.

Different materials will be considered, including unlined Cast Iron (CI), Asbestos Cement / Concrete (AC/C), PolyEthylene (PE) and PolyVinil Chloride (PVC). Table 2 summarises the prototypical characteristics assumed for these materials according to previous literature references (Blokker and Pieterse-Quirijns 2013; Blokker et al. 2014) and industry standards (Canal de Isabel II 2021). Pipe thickness $t_{pipe}[m]$ is here computed based on the assumed Standard Dimensional Ratio (*SDR*), which is considered characteristic of each material:

$$t_{pipe} = \frac{2 \cdot r_i}{SDR - 2} \tag{37}$$

In reality, SDR may change with the diameter and/or the age of the pipe (due to manufacturing evolution), but since the aim of this work is to analyse general behaviours and trends, assuming a constant value is considered sufficient.

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In order to compare the results obtained between the steady and unsteady ground models, the 356 steady conditions will simulate the conditions of the most unfavourable day of the year. Since 357 temperature increase is associated with an acceleration of the various processes that degrade water 358 quality, the worst summer scenario (i.e. maximum temperature gradient during summer) will be 359 analysed here. The parameters that condition the ground behaviour over the year assuming a sand 360 backfill have already been presented (see Table 1 and Figure 2). A wet sand will be assumed in 361 this work to simulate the greatest possible amplitude of the ground temperature over the year and 362 greatest heat transfer from the water, so shortest transition zones. This worst case will be important 363 to test if the steady solution can approximate the unsteady solution in Section 3.2. The maximum 364 daily average of the undisturbed ground temperature at $z_p = 1$ m considering a wet sand is 17.5°C. 365

The incoming fluid temperature is also a required input. In this work, an annual variation is also 366 assumed for the incoming water. The same amplitude of the ground at z = 0 m (i.e. ground surface) 367 is assumed, but the input water temperature is considered to be as lagged as the temperature of 368 the ground at $z_p = 1$ m (considering a wet sand) to account for the time that it takes the water 369 to adapt to atmospheric changes (see Figure 4 later on). In reality, each case will have specific 370 conditions depending on the volume of the service reservoir, its relative position to the ground, 371 etc., but this assumption is reasonable for this case study, where the aim is to assess general trends 372 and behaviours. The maximum daily average of the input water is therefore 20.0° C. These values 373 are consistent with the expected reality in a pipeline system during summer: water is warmer than 374 the ground, so it will cool over the pipeline. 375

376

3.1 Steady ground model

This subsection compares the water temperatures computed when assuming a ground finite heat capacity (i.e. coupled model) as opposed to the traditional ground infinite heat capacity hypothesis (i.e. decoupled model). Transition regions are also computed thanks to Eqs. (13a) and (13b) for different materials, diameters and hydraulic conditions. This analysis will answer the first two specific aims of this work. Note that results obtained with the infinite heat capacity hypothesis are equivalent to those computed for DWDS by making use of EPANET-MSX software (or any other 1-D numerical model suited for temperature analysis).

Figure 3 shows the water temperature evolution along the pipe according to Eq. (8). The 384 water temperature has been computed by assuming a finite ground heat capacity model (thermal 385 resistance as in Eq. 9a) and an infinite ground heat capacity model (thermal resistance as in Eq. 386 9b) for a $r_i = 0.15$ m pipe (300 mm internal diameter) made of CI (CI300), AC/C (AC/C300), PE 387 (PE300) and PVC (PVC300). Figure 3a shows that it takes over 55 km (approximately 31 h) for 388 the water to attain the undisturbed ground temperature according to the finite ground heat capacity 389 (i.e. coupled) model in a CI pipe. This transition length reduces to virtually zero when assuming 390 an infinite ground heat capacity (i.e. decoupled) model. A similar underestimation is observed in 391 Figures 3b, 3c and 3d for the rest of materials. For example, 83 km / 46 h (finite) and 26 km / 14 392

393 394 h (infinite) are needed in the case of a PVC pipe. These results suggest that assuming infinite heat capacity is not an acceptable simplification for appreciable distances/times into the DWDS.

The transition length can be explicitly computed by making use of Eq. (13a) (finite ground heat 395 capacity) or (13b) (infinite ground heat capacity) and a specified tolerance value ($tol = 0.1^{\circ}$ C). 396 Table 3 shows the importance of the terms involved in the computation of the transition length for 397 the four materials under the two hypotheses. This table shows that R_{ground} is greater in order of 398 magnitude than R_{pipe} . This explains the significant differences perceived in Figure 3: a decoupled 399 approach neglects the most important thermal resistance. This table also shows that the water 400 convection term is almost negligible regardless of the material. This means that the transition 401 length depends almost linearly on water velocity (i.e. it is proportional to the flow rate) and 402 quadratically on pipe diameter. 403

Table 4 shows the computed transition lengths for different pipe materials, diameters and water 404 velocities. Note that some transition lengths exceed the maximum length plotted in Figure 3 (125 405 km), due to the explicit nature of Eqs. (13a) and (13b). Transition region values for v = 0.5m/s406 and pipe diameter 300 mm coincide with those in Table 3. Table 4 shows that CI and PVC have 407 extreme behaviours, leading to minimum and maximum transition lengths, respectively, for a given 408 pipe diameter and water velocity. Moreover, it pinpoints that the transition length/residence time 409 increases with pipe diameter, whereas the equivalent residence time is approximately the same no 410 matter the water velocity. A 0.1 m/s water velocity has been adopted to highlight this point. This 411 means that for a specific pipe (installed at a given depth, surrounded by a specific ground, and with a 412 predefined material and pipe diameter), the residence time and the temperature gradient determine 413 the distance required for heat equilibrium to be achieved. Depending on the water velocity, this is 414 associated with a shorter or longer transition distance. 415

It is important to highlight that the previously computed transition lengths correspond to the steady conditions during the day of summer associated with the maximum gradient. This scenario is associated with maximum differences between the input water and the undisturbed ground temperatures, leading to maximum transition lengths. In spring and autumn, the temperature gradient between the water and the ground is minimum and influence lengths will reduce (see Eqs. 13). The winter scenario is opposite to the summer day here analysed. Note that the infinite ground heat capacity model underestimates water temperature in summer, which is unfavourable when assessing the impacts of climate change in water quality. Winter would be associated with mirror images for Figures 3a to 3d, i.e. temperature would be overestimated with the infinite heat capacity hypothesis during winter.

426

3.2 Unsteady ground model

This subsection intends to explore if the steady ground model equations previously applied can 427 be used to approximate the unsteady interaction of the ground over the year (third aim). This implies 428 testing if annual variations are sufficiently slow, so that water temperatures can be approximated 429 with a pseudosteady approach (i.e. steady state every hour) as an alternative to the complex 430 unsteady ground model presented in Section 2.3. Only the finite heat capacity of the ground will 431 be considered for this purpose, because it has already been shown that decoupled models do not 432 sufficiently represent reality for significant distances/times. Simulations are carried out with the 433 original case study ($r_i = 0.15$ m, L = 125 km, v = 0.5 m/s) and only the two materials identified 434 as extreme in Section 3.1 (CI and PVC) will be here assessed. Regarding spatial discretization, 435 $L_i = 500 \text{ m} (n_s = 250)$ will be assumed to start with, although its sensitivity will be tested later on. 436 Figure 4 shows the output water temperature evolution for CI300 and PVC300 pipes at a distance 437 of 25 km from the inlet. This distance has been selected as an arbitrary position within the transition 438 region, so that the difference between materials is noticeable (i.e. temperature equilibrium has not 439 been reached). Figure 4a shows that the water temperature 25 km far from the inlet is close to 440 the ground temperature for CI. This temperature is slightly higher (i.e. closer to the input water 441 temperature) in the PVC pipe (Figure 4b). This result makes intuitive sense due to the higher 442

conductivity of CI, which speeds the heat transfer process and is associated with temperatures
closer to the undisturbed ground temperature at the analysed position (i.e. shorter transition
length). In both cases, results obtained with the unsteady and pseudosteady approximations are
almost identical. They show that, as expected, water heats through the pipeline in winter and cools

along the pipe in summer.

In order to compare the transition regions computed with the unsteady and pseudosteady models, 448 the temperature evolution over the pipe is represented at the worst hour of summer (i.e. the summer 449 day associated with the greatest temperature gradient) in Figure 5a for both CI300 and PVC300 450 pipes. These figures show that there is a slight difference between the temperature computed with 451 the unsteady and pseudosteady models. This is negligible in comparison with other uncertainties 452 in DWDS modelling and water quality reactions in general (Machell and Boxall 2012). Table 5 453 shows the transition lengths computed for the CI and PVC pipes with the unsteady ($L_i = 500 \text{ m}$) 454 and pseudosteady (Eq. 13a) ground models. Pseudosteady results are equal to those obtained with 455 the steady equation (see Table 4), and they are both longer than those obtained with the unsteady 456 simulation. However, the steady/pseudosteady model provides a reasonable approximation to the 457 unsteady model results. In order to illustrate that this difference is not a consequence of the selected 458 spatial discretization, the unsteady model has also been run with $L_i = 250$ m. Table 5 shows that 459 transition lengths are still slightly overestimated (<10%) with the steady/pseudosteady approach. 460 Note that implementing the unsteady ground model in an Intel Core i7-6700 CPU 3.40 GHz 32 461 GB RAM desktop computer (using Matlab R2021a) takes 2018 s (CI300) and 1931 s (PVC300) 462 for $L_i = 500$ m. These times go up to 7505 s (CI300) and 7332 s (PVC300) when considering 463 $L_i = 250$ m, whereas the computational cost of the steady/pseudosteady approach is negligible. 464

Results show that assumed annual changes are sufficiently slow, and the steady/pseudosteady 465 approach can be applied instead of the unsteady ground model to roughly approximate water 466 temperatures and transition lengths. It could be argued that this conclusion is only valid for this 467 particular case study. In order to test its sensitivity, the amplitude of the undisturbed ground and 468 input water temperatures is doubled ($A = 20^{\circ}$ C). This is an exaggeration of reality (it would lead to 469 ground and water temperature values below 0° C), but it can be used to check if this simplification 470 works even when seasonal changes are extreme. Figure 5b shows the water temperature evolution 471 along the CI300 and PVC300 pipes. Like before, the distribution obtained with the unsteady 472 $(L_i = 500 \text{ m})$ and steady/pseudosteady model is reasonably close, and so are the transition lengths 473

474 (see Table 5). Note that this scenario implies doubling the temperature gradient, so the transition
 475 length increases less than twice (natural algorithm of the gradient) with respect to the original
 476 values.

Temperature analysis is case specific, but adopting a steady/pseudosteady approach to analyse how water temperatures and associated transition lengths vary over the year seems to be a reasonable approximation. These analytical expressions could constitute a useful tool to identify the areas of the network where complex heat transfer phenomena take place.

481 DISCUSSION

Results show that the effects of water-ground interaction are important many kilometres/hours 482 into a pipeline system. This means that those areas of the network associated with residence 483 times below a threshold value (which depends on pipe characteristics) are subjected to interaction 484 impacting the water temperature. Residence times may vary widely for different network layouts. 485 Machell and Boxall (2014) published a statistical analysis of the water age at two networks. They 486 identify 9.44 h and 19.86 h as the average water age in these systems, but 5.28 h and 2.68 h as the 487 mode of the mean age in these networks. Mode values are near or below the 5-7 h threshold that 488 corresponds to the transition region of CI and PVC (extreme materials) 100 mm pipes according 489 to a coupled model (see Table 4). Note that this threshold would rise if larger diameters were 490 present. This means that a significant number of pipes are likely to be affected by this interaction, 491 so assuming infinite ground capacity may be a poor simplification for significant parts of a DWDS. 492 The transition region has further implications on water quality analysis. For example, the bulk 493 decay coefficient (k_b) , which partly explains chlorine decay at DWDS, varies with temperature 494

⁴⁹⁵ according to the Arrhenius formula. Wall reactions are also likely to be temperature dependent, but ⁴⁹⁶ given the far greater uncertainty of these coefficients, temperature effects remain unknown. Figure ⁴⁹⁷ 6 shows that a small change in temperature may lead to significant variations in k_b , depending on ⁴⁹⁸ the activation coefficient (E/R). Chlorine bulk decay behaves exponentially with k_b coefficient, ⁴⁹⁹ so temperature should be evaluated and conveniently considered when analysing chlorine decay ⁵⁰⁰ (Díaz and González 2022). There is usually little data available on which to base k_b values,

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and there is typically poor control about whether the assumption of constant temperature over the 501 simulation period provides sufficient accuracy for chlorine simulation. This is particularly true as 502 simulation periods increase for more complex networks with longer residence times. Other water 503 quality parameters and processes will also be affected, from corrosion rates to biological growth. 504 Considering planktonic bacteria which are commonly expected to follow exponential growth trends, 505 meaning increasing numbers towards the extremity of networks, these temperature simplifications 506 may not be significant. But if we consider that >95% of organics including bacteria are in biofilms 507 that are fixed to the pipe walls throughout DWDS (Douterelo et al. 2018), these temperature effects 508 will be important. This is amplified when considering that accelerated growth in summer (due 509 to increased temperature - Calero et al. 2021) or die off (due to the reverse) in winter will seed 510 the entire network. The same applies to reaction precipitations (like iron and manganese), which 511 determine discolouration (Mounce et al. 2016). In other words, at the far extremes of networks either 512 modelling approach will predict that water temperature has approximated ground temperature, so 513 which model is used is not significant for temperature driven effects at such locations. What is 514 important is the temperature driven effects over the complete route to such locations, which will 515 determine the quality of water arriving and then changing further at these locations. This modelling 516 study shows that ignoring ground heat capacity effects could be significant for this. 517

One of the strengths of this work is that it presents an analytical explicit equation to compute 518 the transition length/time. This is possible because the propagation of transport effects (i.e. plug 519 flow along the pipeline) has been disregarded. This means that the thermal balance within the pipe 520 is not considered in the steady equation nor the unsteady model, which accounts for the temporal 521 storage of heat in the ground but not in the water. Assessing the effect of water propagation and 522 analysing the added resolution/need of considering daily patterns or even more random stochastic 523 behaviours (Blokker et al. 2011) is a subject for further research. Addressing this issue is needed 524 to extend the ground finite heat capacity analysis throughout DWDS. 525

⁵²⁶ This work shows that ground characteristics and pipe materials determine the size of the tran-⁵²⁷ sition region. An average wet sand has been considered in this analysis to maximise seasonal

variations, which is important to validate the steady/pseudosteady approach as a reasonable ap-528 proximation to the unsteady ground model. If pipe backfill was dry (low conductivity), the thermal 529 resistance would increase, and so would the associated transition length/time according to Eq. 530 (13a) (see Figure S1 in supplemental data). Ground moisture content varies in reality depending 531 on weather conditions evolution, ground surface characteristics (e.g. grass vs paved surfaces), 532 street flushing, possible leaks, etc. Similarly, a pipe depth of 1 m (sufficient to assume that the 533 undisturbed ground does not experience daily variations) has been here considered. Results do not 534 vary as much when increasing the installation depth to 2.5 m (see Figure S2 in supplemental data). 535 Pipe characteristics, pipe diameter and conductivity have shown to have a significant effect in the 536 thermal behaviour due to their relative importance in Eqs. (9a) and (13a). 537

Water temperature behaviour is case specific, and we currently do not really know how much 538 detail/complexity should be included in a model to simulate realistic temperature behaviours, 539 because there has been insufficient data collected. Water temperatures are typically only measured 540 and recorded at the exit of water treatment works and/or (very limited amount of data) at consumer 541 taps (Agudelo-Vera et al. 2020). Limited data makes validation only possible at these points. 542 Hence why prior research in DWDS has not considered the transition zone. While this paper 543 shows mathematically that the transition zone is significant, and this finding is consistent with the 544 modelling approaches typically adopted for sewer systems (finite heat capacity over an influence 545 area) and GSHP (finite heat capacity), no data exists to verify this. Temperature measurements 546 from physical experiments are needed to further understand temperature dynamics and for model 547 validation. These should cover different spatial and temporal scales under realistic conditions, 548 this could be via suitably complex and scale laboratory conditions initially but is likely to require 549 measurements from DWDS operations. Only then will it be possible to assess to what level of 550 complex models are really needed. The data required includes temperature monitoring but also 551 ground surveys and site inspections (e.g., pipe installation details, weather studies). Accurate 552 hydraulic travel times will also be necessary, requiring more than pressure data for headlosses 553 (Boxall et al. 2004). What is clear is that temperature discussions are especially pressing in the 554

face of climate change. Global warming is expected to increase average temperatures, but also to
 intensify the frequency of extreme phenomena (e.g. heat waves). Their effect in DWDS water
 quality should also be assessed (Pick et al. 2021).

558 CONCLUSIONS

This paper assesses the effect of considering finite ground heat capacity when modelling water 559 temperatures in DWDS. This improvement with respect to the few available previous implemen-560 tations in water supply systems shows that there is a significant transition region between the 561 temperature at the inlet (conditioned by water treatment works and/or service reservoirs) and that 562 at consumer taps (mostly conditioned by the undisturbed ground temperatures). Results for an 563 average pipe early in the system show that the transition region expands at least 5-6 hours in terms 564 of residence time, meaning that the complex water-ground heat interaction process is of importance 565 for a number of pipes within the DWDS. This shows that the traditional assumption of considering 566 an infinite heat capacity for the ground is not enough over this transition area, providing a poor 567 representation of reality. 568

An analytical explicit equation is here provided to quantify the transition length/time under 569 steady flow and ground conditions. It shows that, for a pipe with a predefined diameter and 570 material, installed at a specific depth within a conductive ground environment and an input water 571 - undisturbed ground temperature gradient, the transition length is mainly characterized by the 572 residence time, and water velocity determines the distance over which this transition to equilibrium 573 takes place. This expression also seems to be a good first approximation to the results obtained 574 with complex unsteady ground models, which are time-consuming to run. This analysis must 575 be improved to include the thermal balance within the pipeline and so characterise temperature 576 behaviour for unsteady flows, which are out of the scope of this paper. 577

This work builds on the temperature modelling strategy typically adopted at DWDS. Since temperature drives all reactions and processes from chlorine decay to corrosion and biofilm growth, characterizing this variable is important to ensure the supply of safe clean water at the tap.

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DATA AVAILABILITY STATEMENT

Models and code that support the findings of this study are available from the authors upon reasonable request.

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	$ \begin{array}{c} T_0 \\ (^{\circ}C) \end{array} $	$\begin{array}{c} A \\ (^{\circ}C) \end{array}$	t_{shift} (h)	k_{ground} (W/m/K)	C_{ground} (J/kg/K)	$ ho_{ground} (kg/m^3)$	$lpha_{ground}\ (m^2/h)$
Dry sand	10	10	0	0.95	900	1600	0.0024
Wet sand	10	10	0	3.35	1500	1900	0.0042

TABLE 1. Surface temperature and ground parameters for an average sand: dry sand vs wet sand

	CI	AC/C	PE	PVC
Standard Dimensional Ratio SDR (-)	15	26.5	17	38
Pipe roughness ϵ_{pipe} (mm)	0.2	3	0.03	0.06
Pipe wall conductivity $k_{pipe}(W/m/K)$	60	0.43	0.5	0.16

TABLE 2. Prototypical pipe material characteristics

			Rground	R_p	ipe		
		$Q \cdot \rho_w \cdot C_w$	$\frac{\ln(\frac{2\cdot z_P}{r_O})}{2\cdot \pi \cdot k_{ground}}$	$\frac{\ln(\frac{r_o}{r_i})}{2 \cdot \pi \cdot k_{pipe}}$	$\frac{1}{Nu \cdot k_w \cdot \pi}$	$\ln\left(\frac{ T_{in}-T_{ground} }{tol}\right)$	L_t
		(W/K)	$(m \cdot K/W)$	$(m \cdot K/W)$	$(m \cdot K/W)$	(-)	(km and h)
CI300	Finite		0.1163	3.7959e-4	5.6611e-4		56.1 km (31.1 h)
C1500	Infinite		-				0.5 km (0.3 h)
AC/C300	Finite		0.1193	0.0290	3.7088e-4		71.1 km (39.5 h)
AC/C300	Infinite	1.4809e5	-			3.2	14.1 km (7.8 h)
PE300	Finite	1.400965	0.1171	0.0398	6.3262e-4	5.2	75.4 km (41.9 h)
PE300	Infinite		-	0.0398	0.52020-4		19.4 km (10.8 h)
DVC200	Finite		0.1205	0.0520	6 1 6 0 1 4		83.6 km (46.5 h)
PVC300	Infinite		-	0.0538	6.1681e-4		26.0 km (14.5 h)

TABLE 3. Terms involved in the transition region computation (distance and equivalent residence time) for CI300, AC/C300, PE300 and PVC300 under finite and infinite ground heat capacity hypotheses

TABLE 4. Transition region (distance and equivalent residence time) for different pipe materials, diameters and water velocities

		CI		AC/C		PE		PVC	
Velocity (m/s)	Diameter (mm)	Finite	Infinite	Finite	Infinite	Finite	Infinite	Finite	Infinite
	100	1.9 km (5.2 h)	0.1 km (0.2 h)	2.2 km (6.1 h)	0.4 km (1.0 h)	2.3 km (6.4 h)	0.5 km (1.4 h)	2.5 km (6.9 h)	0.6 km (1.8 h)
0.1	300	11.4 km (31.7 h)	0.3 km (0.8 h)	14.4 km (39.9 h)	3.0 km (8.2 h)	15.3 km (42.4 h)	4.1 km (11.3 h)	16.9 km (47.0 h)	5.4 km (15.0 h)
	600	32.6 km (90.4 h)	0.7 km (1.9 h)	44.6 km (123.8 h)	11.5 km (32.0 h)	48.0 km (133.3 h)	15.8 km (43.9 h)	54.6 km (151.7 h)	21.1 km (58.7 h)
	100	9.0 km (5.0 h)	0.1 km (0.1 h)	10.7 km (5.9 h)	1.6 km (0.9 h)	11.2 km (6.2 h)	2.2 km (1.2 h)	12.1 km (6.7 h)	2.9 km (1.6 h)
0.5	300	56.1 km (31.1 h)	0.5 km (0.3 h)	71.1 km (39.5 h)	14.1 km (7.8 h)	75.4 km (41.9 h)	19.4 km (10.8 h)	83.6 km (46.5 h)	26.0 km (14.5 h)
	600	160.7 km (89.3 h)	1.3 km (0.7 h)	221.2 km (122.9 h)	56.0 km (31.1 h)	237.9 km (132.2 h)	76.9 km (42.7 h)	271.0 km (150.6 h)	103.5 km (57.5 h)

TABLE 5. Transition region (distance and equivalent residence time) at the worst case summer day according to different finite heat capacity ground models: original vs double temperature amplitude

Temperature amplitude	Ground model	CI300	PVC300
	Unsteady ($L_i = 500 \text{ m}$)	53.5 km (29.7 h)	81.0 km (45.0 h)
Original ($A = 10^{\circ}$ C)	Unsteady ($L_i = 250 \text{ m}$)	53.00 km (29.4 h)	80.75 km (44.9 h)
	Pseudosteady (Eq. 13a)	56.1 km (31.1 h)	83.6 km (46.5 h)
	Unsteady ($L_i = 500 \text{ m}$)	65.0 km (36.1 h)	98.0 km (54.4 h)
Double ($A = 20^{\circ}$ C)	Unsteady ($L_i = 250 \text{ m}$)	64.25 km (35.7 h)	97.75 km (54.3 h)
	Pseudosteady (Eq. 13a)	68.1 km (37.8 h)	101.6 km (56.4 h)

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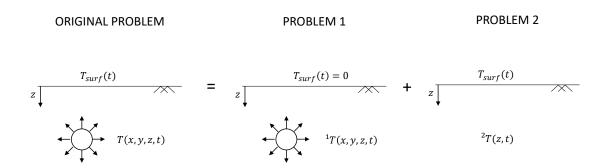


Fig. 1. Superposition principle for heat transfer analysis in a buried pipe

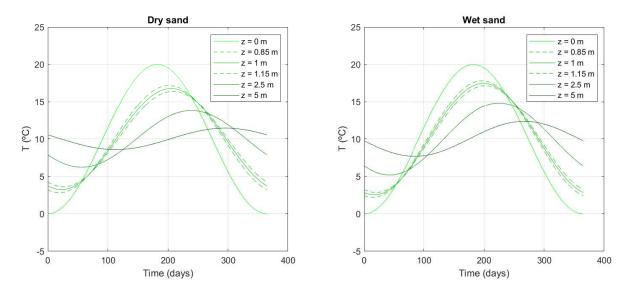


Fig. 2. Undisturbed ground temperature evolution at different depths for an average sand: a) dry sand and b) wet sand

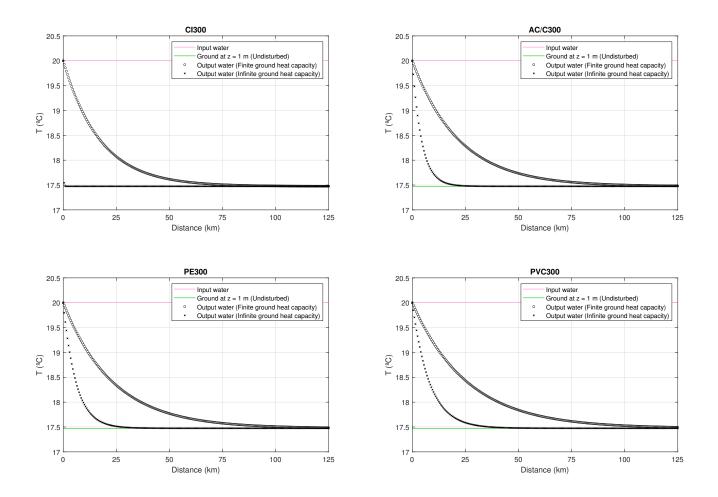


Fig. 3. Temperature evolution along the pipe under steady ground conditions (input water 20.0°C, ground 17.5°C) for wet sand and $z_p = 1$ m: a) CI300, b) AC/C300, c) PE300 and d) PVC300

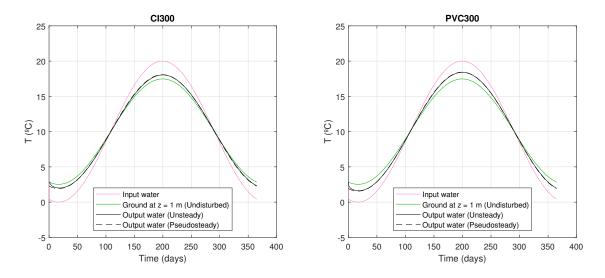


Fig. 4. Annual output water temperature evolution 25 km far from the pipe inlet: a) CI300 and b) PVC300

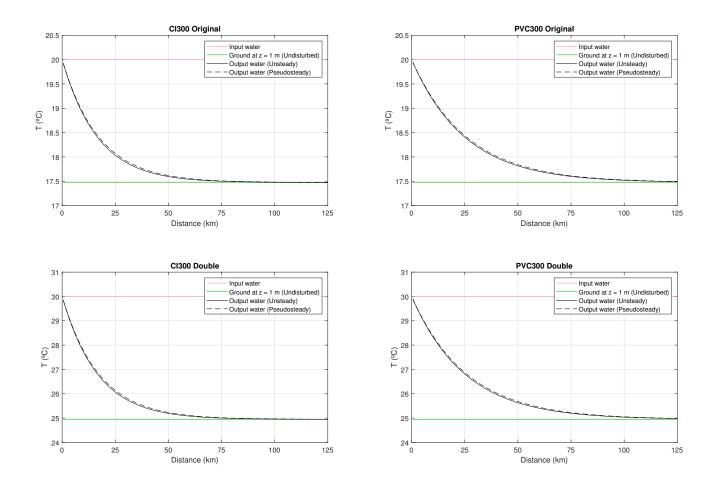


Fig. 5. Temperature evolution along the CI300 and PVC300 pipes under unsteady ground conditions (input water 20.0°C, ground 17.5°C): a) original temperature amplitude and b) double temperature amplitude. Note that the y axis range doubles for b)

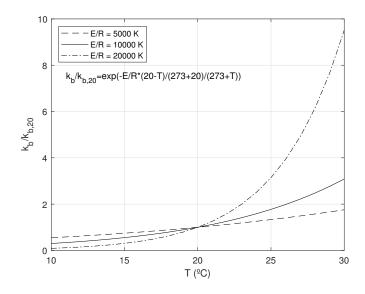


Fig. 6. Chlorine bulk decay coefficient variation with temperature