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Effect and fate of Ni during aging and thermal-induced phyllomanganate-

to-tectomanganate transformation

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Abstract: Phyllomanganates are ubiquitous in a variety of environments and 1 commonly enriched in transition metal elements, such as Ni. The effect of such foreign 2 3 metal cations on phyllomanganate transformation is widely documented under aqueous conditions together with the induced modification of Ni geochemical behavior. A 4 5 similar knowledge is lacking however on phyllomanganate transformation and on the induced fate of associated metal elements that may occur under dry conditions, that 6 prevail in deserts and arid areas increasingly exposed to severe droughts or wildfires. 7 The present study shows that crystallinity, morphology, Mn oxidation state, and Ni 8 9 binding mechanisms are essentially unaffected when aging hexagonal birnessite (Mn oxidation state ~3.90 and Ni/Mn molar ratios of 0.00 and 0.13) in the dry state at room 10 temperature for up to 8 years. In contrast, heating aged Ni-doped birnessite to 25-200 11 °C results in an increased proportion of edge-sharing Ni-Ni(Mn) pairs with increasing 12 temperature induced by the migration of interlayer Ni to birnessite octahedral layers 13 and/or by an increased sharing of coordination oxygens by interlayer Ni/Mn from 14 15 adjacent layers. Further heating to 400 °C does not change this proportion, with 16 birnessite layer structure being retained. Transformation of Ni-doped birnessite to cryptomelane is complete at 500 °C, while that of Ni-free birnessite is achieved at 400 17 °C, suggesting that Ni doping increases birnessite thermal stability. Birnessite-to-18 19 cryptomelane transformation comes with a strong increase of Mn oxidation state, whereas this parameter remains unchanged in heated birnessite samples. Ni 20 21 incorporation in the cryptomelane framework, reduces its release during reductive acid dissolution by a factor of 396 ± 15 compared to initial birnessite. These results shed 22

light on mineral transformation affecting layered manganates under dry conditions andon the fate of associated transition metal elements.

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Keywords: manganese oxide, mineral transformation, dry aging, thermal treatment,
transition metals

28

29 **1. Introduction**

30 Manganese (oxyhydr)oxide minerals are widespread in geological settings such as 31 soils, sediments, ocean nodules, and desert varnishes (Post, 1999; Lu et al., 2019). Minerals from the birnessite family are the most common naturally occurring Mn 32 oxides and may form through both biogenic and abiotic processes (Villalobos et al., 33 34 2003; Morgan, 2005; Lanson et al., 2008). These phyllomanganates are composed of layers of edge-sharing [Mn(IV)O₆] octahedra with hydrated interlayers (Giovanoli, 35 1980). Existence of Mn vacancies and/or isomorphous substitutions of low valence 36 37 cations such as Mn(III), Ni(II), Co(III), and Fe(III) for Mn(IV) in birnessite layers induce a net negative charge in the layers, which is compensated for by the presence of 38 hydrated cations and/or protons in the interlayer region (Drits et al., 1997; Silvester et 39 al., 1997). Owing to their unique structure (tiny particle sizes, vacancies, and mixed Mn 40 41 valences), birnessites have strong adsorption and/or oxidation capacities for a variety of trace metals, and thus host a large number of metal cations (Peacock, 2009; Simanova 42 43 and Pena, 2015). Natural birnessite minerals are generally enriched in metals such as Co, Ni, and Fe (Taylor et al., 1964; Manceau et al., 2014; Lee et al., 2019). 44

45	Birnessite-like minerals can be precursors of other naturally occurring Mn oxides
46	such as tectomanganates. Natural cryptomelane, which has a 2×2 tunnel structure, is
47	mainly found in supergene oxidation zones of Mn deposits, lateritic weathering profiles,
48	and volcanic ashes, and may form through authigenic precipitation or transformation
49	from birnessite minerals (Lu and Li, 2015). Birnessite can transform to cryptomelane
50	via dissolution-reprecipitation pathways (Jothiramalingam et al., 2006; Zhang et al.,
51	2011, 2016; Sinha et al., 2013) or via topotactic transformation mechanisms (Tu et al.,
52	1994; Ching et al., 2001; Grangeon et al., 2014, 2015). It was proposed that Mn(III)
53	content in the octahedral layers plays a key role in the conversion from
54	phyllomanganates to tectomanganates (Atkins et al., 2014; Grangeon et al., 2014, 2015;
55	Yang et al., 2018). Most studies investigating the phyllomanganate-to-tectomanganate
56	transformation were conducted in solution (Tu et al., 1994; Zhang et al., 2011; Atkins
57	et al., 2014; Yang et al., 2018, 2019; Wu et al., 2019, 2021), consistent with the
58	prevalence of wet conditions in soils and sediments. Dry conditions prevail however in
59	desert soils, that cover more than 20% of Earth's land area (Zhu et al., 2022), and in
60	soils affected by seasonal droughts. As a result of the present climate change, the
61	frequency of extreme weather patterns such as severe droughts and extreme heat events
62	has increased over the past two decades, a long-term tendency according to climate
63	projections (Vicente-Serrano et al., 2020). As the phyllomanganate-to-tectomanganate
64	transformation processes even in the dry state (Atkins et al., 2014; Grangeon et al., 2014
65	2015; Yang et al., 2018), it is important to investigate such mineral transformation and
66	the fate of associated metal cations in dry conditions.

Very few studies have reported dry-state transformation over long periods of time 67 (Grangeon et al., 2014, 2015) and, to our knowledge, no study has investigated dry-68 69 state transformation at elevated temperatures linked for example to the increasing frequency of wildfires (Murphy and Timbal, 2008). During large wildfires, soil 70 temperature may exceed 600 °C (Rein et al., 2008), possibly resulting in the 71 transformation of soil minerals (Johnston et al., 2016; Li et al., 2019). In particular, 72 laboratory experiments showed that birnessite can be thermally transformed to 73 cryptomelane (Chen et al., 1986; Golden et al., 1986). Such transformation may 74 75 significantly affect the fate of elements in the initial minerals. For example, heating As(V)-coprecipitated iron oxyhydroxysulfate mineral schwertmannite at temperatures 76 higher than 400 °C results in the formation of hematite and causes the initially 77 78 structurally incorporated As to be surface-complexed, thus enhancing As mobilization (Johnston et al., 2016). Similar processes have also been reported for transformation of 79 naturally occurring goethite-rich materials upon heating (Perrier et al., 2006; Landers 80 81 et al., 2009). However, no studies have yet investigated the fate of metal elements associated with Mn oxide minerals during their thermal-induced transformation. 82 83 The present study thus investigates the possible phase transformation affecting Ni-

doped hexagonal birnessites during aging in the dry state at room temperature for 5 and 84 gears and subsequent heating over a wide range of temperatures (50-500 °C) and the 85 induced fate of associated Ni. Powder X-ray diffraction, electron microscopies, 87 thermogravimetric analysis, and Ni and Mn K-edge X-ray absorption fine structure 88 spectroscopy are combined with acid dissolution experiments for this purpose. Nickel was chosen as a model metal cation because of being 1) one important trace metal nutrient and contaminant, 2) a redox-insensitive element, and 3) because its binding mechanisms on birnessite-like minerals have been thoroughly studied in the literature. The present results provide insights into the mobility and bioavailability of metal nutrients and contaminants associated with the phase changes of host Mn oxide minerals, and may deepen our understanding of Mn-mediated metal geochemical behaviors in dry soils and during extreme heating events.

96 2. Materials and methods

97 **2.1 Dry aging of Ni-doped birnessites at room temperature**

Hexagonal birnessite, HB, and Ni-doped birnessite, Ni10, from our previous study 98 (Yin et al., 2012) were used here. These samples were synthesized according to the 99 100 protocol of McKenzie (1971). Briefly, this protocol involves the reduction of a boiling KMnO₄ solution (300 mL, 0.667 M) by an HCl solution (45 mL, 6 M) in the 101 absence/presence of NiCl₂ (Ni/Mn molar ratio of 0.10). The pHs of the synthetic 102 103 suspensions after mineral crystallization were ~7.2 and ~5.5, respectively. After washing thoroughly the obtained solids, they were dried at 40 °C for several days, 104 ground carefully to particle sizes below 100 mesh, and reserved for further use. HB and 105 Ni10 samples have specific surface areas (SSAs) of 19 and 71 $m^2 \cdot g^{-1}$, respectively (Yin 106 et al., 2012). The Ni content in Ni10 is 6.1 wt.%, and the K contents in HB and Ni10 107 are 8.4 and 3.8 wt.% respectively (Yin et al., 2012). 108

Aliquots of these birnessite powders were stored in sealed polystyrene tubes and
kept in the dark at room temperature. After aging for 5 and 8 years, some of the samples

were thoroughly characterized. As-obtained, aged samples were correspondingly named as HB Ny and Ni10 Ny (N = 5 or 8; Table 1).

113 **2.2 Thermal treatment of the 8-year aged samples**

In addition, the 8-year aged samples (HB 8y and Ni10 8y) were heated to 114 115 temperatures (50, 100, 150, 200, 300, 400 and 500 °C), spanning the common range of 116 soil temperatures reported during wildfires (Rein et al., 2008). Typically, ~0.1 g of each sample was placed in a ceramic crucible and heated in a muffle furnace in air. After 117 increasing the furnace temperature to the set value (rate of 10 °C·min⁻¹), the sample 118 119 powder was kept at this temperature for 2 h, consistent with previous studies (Cai et al., 2001). The furnace door was then opened to let the powder cool down naturally to room 120 temperature, and as-obtained solids were stored for further analysis. These samples 121 122 were labelled HB 8y T and Ni10 8y T, where T stands for the temperature used in °C (Table 1). 123

124 **2.3 Solid characterization**

125 Powder X-ray diffraction (XRD) analysis was used to identify structural changes of Mn oxide samples. Data were collected on a Bruker D8 Advance diffractometer. As 126 the XRD profile changes during drying aging at room temperature were expected to be 127 minimal, the diffraction data of the 5- and 8-year aged samples were collected at a step 128 129 size of 0.04 °20 Cu Ka using a counting time of 10 or 40 sec per step. The XRD data of the thermal transformation products were collected using a scanning rate of 10° per 130 131 min with a step size of 0.02 °20. Field emission scanning electron microscopy (FESEM) images of the samples were obtained using a Jeol JSM-6700F microscope after coating 132

the minerals with a gold evaporated film for 30 s. High angle annular dark field (HAADF) scanning transmission electron microscopy (STEM) images of Ni10_8y_500 were obtained by a JEOL JEM-ARM200F microscope equipped with a spherical aberration correction system. Thermogravimetric analysis (TGA) of the Ni-doped birnessites before and after 8-year aging was conducted from ~35 °C to 800 °C on a NETZSCH TG 209 thermal analyser in N₂ or dry air at a flow rate of 20 mL·min⁻¹ with a heating rate of 10 °C·min⁻¹.

140 2.4 X-ray absorption fine structure (XAFS) spectra

The XAFS spectra were measured at room temperature on the 1W1B beamline at the Beijing Synchrotron Radiation Facility (Yin et al., 2012, 2015). Ni K-edge data were obtained in fluorescence mode over the 8157-9063 eV range. Monochromator energy was calibrated using a Ni metal foil ($E_0 = 8333$ eV). Mn K-edge data were collected in transmission mode over the 6363-7311 eV range for most samples (6340-7104 eV for 8-year aged HB and Ni10 samples heated to 50 and 150 °C). A Mn metal foil was used to calibrate the monochromator ($E_0 = 6539$ eV).

Data reduction and analysis were performed using Ifeffit/SIXPack (Ravel and Newville, 2005; Webb, 2005). Parameters used for background removal of Ni K-edge spectra were: $E_0 = 8343$ eV, Rbkg = 1.0 Å and k-weight = 2. After background removal and normalization, spectra were converted into the k³-weighted functions and Fourier transforms (FTs) were calculated over 2.8-12.5 Å⁻¹ (k range) and 1-6 Å (R range), respectively. Mn K-edge spectra were background-subtracted using the following parameters: $E_0 = 6553$ eV, Rbkg = 1.0 Å and k-weight = 2. XAFS spectra were then

155	converted into the k^3 -weighted functions, and FTs were calculated over 3.4-12.1 Å ⁻¹
156	and 1-4 Å (k and R ranges, respectively). Structural parameters (distance, R;
157	coordination number, CN; and Debye-Waller factor, σ^2) for Ni local environments were
158	obtained by fitting the experimental k ³ -weighted EXAFS spectra to the standard
159	EXAFS equation (Kelly et al., 2008) using several single-scattering paths, as described
160	previously (Peña et al., 2010; Yin et al., 2012). Phase and amplitude functions for
161	single-scattering paths were calculated using FEFF7 (Rehr et al., 1992). Clusters of
162	atoms used for the analysis of Ni K-edge EXAFS spectra were based on chalcophanite
163	crystal structure (Post and Appleman, 1988) and substituting Ni for interlayer Zn or
164	layer Mn. An amplitude reduction factor (S_0^2) of 0.848 was adopted for Ni, consistent
165	with our previous study (Yin et al., 2012).

To determine the relative proportions of heterovalent Mn cations (Mn⁴⁺, Mn³⁺ and 166 Mn²⁺) in Mn oxides, Mn K-edge XANES spectra were analyzed using the Combo 167 method (Manceau et al., 2012). In a typical procedure, the 17 standard spectra used to 168 develop the method (Table S1) were first imported. 4-8 references with non-negative 169 fractions were then used for the linear combination fitting of normalized XANES 170 spectra. Mn K-edge EXAFS data were fitted using birnessite (Grangeon et al., 2010) or 171 cryptomelane (Yin et al., 2015) model. An amplitude reduction factor (S_0^2) of 0.973 172 was adopted from our previous study (Yin et al., 2014). 173

In all EXAFS fits, the number of independent variables used was smaller than the
number of independent data points. During Fourier transform and EXAFS data fitting,
a Hanning window was used.

177 **2.5 Acid dissolution experiments**

Consistent with previous studies (Beak et al., 2008; Qin et al., 2017) acid 178 179 dissolution experiments were performed to investigate Ni crystal chemistry, stability, and bioaccessibility in these samples. Briefly, 0.0400 g of solid were dissolved at 37 °C 180 181 in 250 mL of 2 M HCl solution while stirring at a constant rate. These conditions are similar to in vitro measurements of metal bioaccessibility in the gastric phase (Beak et 182 al., 2008). Suspension aliquots (5 mL) were withdrawn at predetermined intervals and 183 filtered immediately through 0.22 µm membrane filters. The concentrations of Mn²⁺ 184 and Ni²⁺ in the filtrate were analyzed using atomic absorption spectrometry (AAS, 185 Varian AAS 240FS). The metal (Me) dissolution ratio was calculated as χ_{Me} = 186 Me(t)/Me(total), where Me(t) is the aqueous metal concentration at time t, and Me(total)187 188 is the final metal concentration after complete dissolution of the solid.

189 **3 Results**

190 **3.1 Changes in mineralogical phases**

Powder XRD patterns of the Ni-doped birnessite samples before and after dry 191 aging at RT for 5 and 8 years (Fig. 1a,b) are typical for turbostratic hexagonal 192 birnessites, and can be indexed using a C-centered two-dimensional unit cell (Drits et 193 al., 1997, 2007). The unsplit (31, 02) bands and the ratios of (11, 20) to (31, 02) d values 194 close to $\sqrt{3}$ both confirm the hexagonal layer symmetry (Drits et al., 2007). XRD 195 profiles of aged samples show only minimal differences with those of fresh samples. 196 For both HB and Ni10 samples, the position of the (31, 20) diffraction band shifts 197 slightly towards lower angle for the 8-year aged sample compared to fresh and 5-year 198

aged samples, indicating a slight increase of the layer lattice parameter b, without
significant distortion of the hexagonal layer symmetry however (Grangeon et al., 2008).
Consistent with the high values of Mn average oxidation state in the initial products
(3.86 and 3.95 for HB and Ni10, respectively – Yin et al., 2012), similarity of the XRD
profiles of fresh and aged samples supports the stability of birnessite layer and
interlayer structures and the stability of these minerals in the dry state at room
temperature.

Upon heating, powder XRD patterns of the 8-year aged samples undergo 206 207 significant changes (Fig. 1c,d). When heated to temperatures ≤ 300 °C, all samples essentially retain their layer structure. After heating to 400 °C, HB 8y is fully 208 transformed to cryptomelane (K(Mn⁴⁺₇Mn³⁺)O₁₆ – ICDD 74-1451), while Ni10 8y 209 210 retains most of its layer structure, being only partly transformed to cryptomelane. After heating to 500 °C for 2 h, Ni10 8y is essentially transformed to cryptomelane, with a 211 minor NiMn(IV)O₃ (ICDD 75-2089) impurity. Further linear combination fitting (LCF) 212 213 analysis of Mn K-edge EXAFS spectra of thermal transformation products of HB 8y and Ni10 8y by using spectra of Ni10 8y 300 and HB 8y 500 as birnessite and 214 cryptomelane standards, respectively, confirms the XRD results. For example, LCF 215 analysis shows that Ni10 8y 400 is composed of $87 \pm 1\%$ birnessite and $13 \pm 1\%$ 216 cryptomelane (Fig. S1 and Table S2). 217

218 **3.2** Thermal stability of 8-year aged birnessites in N₂ or dry air

Eight-year aged birnessites are more stable than their fresh counterparts upon heating in N₂ atmosphere, although weight losses occur almost at the same temperatures

221	(Fig. 2 and Table S3a). The amount of physisorbed water is identical in HB and HB_8y
222	(35-150 °C $- 8.8-9.3\%$), while that in Ni10_8y (35-180 °C $- 11.0 \pm 0.3\%$) is slightly
223	lower than that in Ni10 (12.0 \pm 0.1 %). The amount of structural water (150-700 °C),
224	including hydroxyl groups and interlayer water molecules (Lee et al., 2007; Yin et al.,
225	2012), is significantly lower in aged birnessites (4.8 \pm 0.5 % and 5.8 \pm 0.4 % in HB_8y
226	and Ni10_8y, respectively) compared to corresponding fresh samples (7.5 \pm 0.9 % and
227	7.2 ± 0.2 %).The weight losses assigned to lattice oxygen (700-800 °C) are alike in HB
228	and HB_8y (4.3 ± 1.0 % and 3.8 ± 0.5 %, respectively). Contrastingly, this weight loss
229	is lower in Ni10_8y (4.1 \pm 0.2 %) compared to Ni10 (7.5 \pm 0.5 %). The lattice oxygen
230	weight loss is related to $Mn(IV)$ reduction and formation of phases such as Mn_2O_3 (Lee
231	et al., 2007).

232 The profiles for the release of both physisorbed and structural water molecules from the 8-year aged samples upon heating in air are similar to those in N₂ (Fig. 2b and 233 Table S3b). With increasing temperature, a weight increase is observed however for 234 both HB 8y (460-800 °C - $0.4 \pm 0.6\%$) and Ni10 8y (520-660 °C - $1.0 \pm 0.3\%$). These 235 weight gains are likely due to the oxidation of low valence $Mn^{2+/3+}$ by O₂ (Tang et al., 236 2014). Weight loss occurs again from 660 °C for Ni10 8y, owing to the Mn reduction 237 and lattice O loss during thermal decomposition to phases such as Mn₂O₃ (Tang et al., 238 2014). These results are consistent with previous studies that indicate oxidation of 239 $Mn^{2+/3+}$ at intermediate temperatures in air and reductive decomposition of MnO_2 at 240 high temperatures both in air and N₂ (Lee et al., 2007; Tang et al., 2014). 241

242 **3.3 Morphology changes**

243	After aging for 8 years, birnessite morphology, 3D hierarchical microspheres
244	composed of 2D disk-like plates, changes little, with the thickness of platy crystals
245	remaining also unchanged (Fig. 3). Heating HB_8y up to temperatures \leq 300 °C slightly
246	decreases microsphere size however, and even more so plate thickness (Figs. S2-4).
247	HB_8y_400 particles appear as prisms, consistent with the transformation to
248	cryptomelane. This morphology is different from commonly reported needles (Sinha et
249	al., 2013; Grangeon et al., 2015; Zhang et al., 2016), but similar to that of cryptomelane
250	obtained from thermal transformation of triclinic birnessite (Chen et al., 1986). After
251	heating HB_8y further to 500 °C, cryptomelane prismatic crystals grow in all three
252	dimensions compared to HB_8y_400. Similar to the corresponding fresh crystals,
253	Ni10_8y crystals are smaller than HB_8y ones (Fig. 3, Figs. S2-4). Heating Ni10_8y to
254	temperatures \leq 400 °C, modifies crystal size and plate thickness only marginally.
255	Although powder XRD analysis shows a minor amount of cryptomelane in
256	Ni10_8y_400, no prismatic crystals were observed, whereas Ni10_8y_500 consists
257	mainly of prismatic crystals (Fig. 3g). Both the HAADF-STEM image and the
258	corresponding Fast Fourier transform (FFT) pattern obtained on a prismatic
259	Ni10_8y_500 crystal confirm the nature of these crystals exhibiting cryptomelane [200]
260	facets with a d spacing of 4.98 Å (Fig. 3h). Atomic image (Fig. 3i) shows the 2×2
261	tunnels of the cryptomelane framework. In this image, the uniform dark stripes
262	represent the K atoms and the coordinated water molecules rather than Ni in the
263	interlayer regions. This indicates that Ni is most likely structurally incorporated in the
264	cryptomelane framework rather than adsorbed in tunnels and/or to crystal surface.

265 **3.4 Changes in Mn Average Oxidation State (AOS)**

Manganese K-edge XANES analysis using the Combo method (Manceau et al., 266 2012) shows that upon dry aging, the proportion of low valence Mn ($Mn^{3+/2+}$) increases 267 only slightly at the expense of Mn⁴⁺ (Fig. 4 and Table S1). Mn AOS values calculated 268 269 for HB 8y and Ni10 8y (3.81 and 3.90, respectively) are similar to those of HB and Ni10 (Yin et al., 2012), within the accuracy (± 0.04) of the method (Manceau et al., 270 2012). As birnessite Mn AOS is a key parameter for their reactivity towards various 271 redox reactions (Wang et al., 2018), this similarity suggests that aged birnessites can 272 effectively mediate the geochemical behavior of various redox-sensitive contaminants 273 in the environment. 274

Heating of HB 8y and Ni10 8y modifies the relative proportions of various Mn 275 276 species depending on temperature. Below 300 °C, the proportions of heterovalent Mn cations in HB 8y and Ni10 8y are essentially constant whereas increasing temperature 277 to 400 °C for HB 8y and 500 °C for Ni10 8y induces the oxidation of low-valence Mn 278 cations to Mn⁴⁺ (Fig. 4). Consistently, Mn AOS values calculated for HB 8y heated to 279 50-300 °C range 3.81-3.84, and increase to 3.91-3.93 after heating to temperatures \geq 280 400 °C. Similarly, Mn AOS values calculated for Ni10 8y heated to 50-150 °C remain 281 constant at 3.89-3.91. When heating Ni10 8y further to 200-400 °C, Mn AOS values 282 decrease to 3.82-3.85, most likely owing to Mn⁴⁺ reduction. Over this temperature 283 range, the weight loss due to hydroxyl groups probably results also in the loss of oxygen 284 anions and thus in the reduction of high valence Mn cations. Contrasting evolutions of 285 Mn AOS observed for HB 8y and Ni10 8y over this temperature range are probably 286

ascribed to the larger loss of hydroxyl groups in Ni10 8y compared to HB 8y, possibly 287 because of the lower crystallinity of the former (Fig. 1a,b and Fig. 2b). At 500 °C, the 288 289 content of low valence Mn in Ni10 8y 500 is greatly decreased, and its Mn AOS is sharply increased to 3.98. The increase in Mn AOS at 400-500 °C is most likely due to 290 291 the oxidation of low valence Mn by O₂ in the air, consistent with TGA results (Tang et al., 2014). In the latter case, weight gain related to absorption of O_2 by these samples 292 occurs at 460-520 °C however, most likely owing to the steady temperature increase 293 rate (Fig. 2b). Such an increase of Mn AOS after the birnessite-to-cryptomelane 294 transformation is consistent with previous reports (Cai et al., 2001; Birkner and 295 Navrotsky, 2017). 296

297 **3.5 Changes in local atomic environments of Mn**

298 Consistent with the results of powder XRD analysis, dry aging at room temperature changes little the birnessite layer structure. Similarly, upon heating to 299 temperatures ≤ 300 °C, the k³chi(k) functions and corresponding FTs (light gray bands 300 in Fig. 5a,c) remain essentially unchanged, confirming the birnessite structure of heated 301 HB_8y samples. After heating to 400 °C, the shape of the oscillation at 7.2-9.8 Å⁻¹ 302 (k³chi(k) function) changes to match that of cryptomelane (Yin et al., 2015) and the 303 magnitude of the third peak (R + Δ R ~ 3.1 Å) in the corresponding FT is greatly 304 increased, confirming the transformation of birnessite to cryptomelane. The oscillation 305 profile and FT of Ni10 8y (light gray bands in Fig. 5b,d) indicate the same 306 mineralogical evolution after heating to 500 °C. 307

308 Fittings of the Mn K-edge EXAFS spectra for heated birnessite samples based on

309	a birnessite structure model, and on a cryptomelane structure model for HB_8y_400,
310	HB_8y_500, and Ni10_8y_500, provide good agreements (Fig. 5 and Table S4). All
311	Mn-O distances in the [MnO ₆] octahedra of birnessite and cryptomelane are similar at
312	1.904 ± 0.001 Å (n = 17) and 1.895 ± 0.001 Å (n = 3), respectively. Edge-sharing Mn-
313	Mn distances in birnessite derived from HB (2.881 ± 0.002 Å; n = 8) are slightly longer
314	than those derived from Ni10 (2.876 \pm 0.003 Å; n = 9) at an α significance level of 0.05
315	(t-test). These longer distances are likely related to the higher content of layer Mn(III)
316	in birnessite derived from HB compared to those derived from Ni10 (Fig. 4 and Table
317	S1). Edge-sharing Mn-Mn distances in cryptomelane (2.885 \pm 0.003 Å; n = 3) are
318	statistically independent of the precursor and are slightly larger than those in birnessite
319	layers $(2.878 \pm 0.004 \text{ Å}; n = 17)$.

Significant changes are observed for the distances between interlayer Mn(Ni) to 320 the nearest layer Mn(Ni) ($Me^{E}-Me^{TC}$; Me = Mn, Ni) (Fig. 6a and Table S4). HB 8y, 321 Ni10 8y and layered (birnessite) samples derived from them all exhibit similar Me^E-322 Me^{TC} distances (3.48-3.52 Å – Fig. 6b), consistent with values reported for hexagonal 323 birnessites (Grangeon et al., 2010). This distance is significantly decreased to 3.44-3.45 324 Å for cryptomelane (Fig. 6c), consistent with previous studies (Lee et al., 2007; Yin et 325 al., 2015). Shortening of Me^E-Me^{TC} distances in cryptomelane compared to 326 phyllomanganates is consistent with the higher proportion of Mn⁴⁺ cations in the former. 327

328 **3.6 Evolution of Ni binding mechanisms in the mineral**

Nickel K-edge k³chi(k) oscillations and corresponding FTs of Ni10 and Ni10_8y
are similar, indicating similar Ni local environments (Fig. 7). In the FTs, three main

331	peaks located at R+ Δ R ~ 1.6 Å, ~2.4 Å and ~3.0 Å mainly correspond to the first Ni-O
332	pair in [NiO ₆] octahedron, edge-sharing Ni-Ni(Mn) pair (Ni ^E) and triple corner-sharing
333	pairs between interlayer Ni(Mn) and layer Ni(Mn) (Ni ^{TC} – Peacock, 2009; Peña et al.,
334	2010; Yin et al., 2012). Intensity ratios of these three peaks are constant indicative of
335	similar proportions of Ni ^E and Ni ^{TC} (Fig. 7 and Table 2). In Ni10_8y, $74 \pm 4\%$ of the Ni
336	content is located above/below vacant layer sites, which is statistically equivalent to
337	Ni10 (Yin et al., 2012).

The k³-weighted Ni K-edge EXAFS functions of heated Ni10_8y samples exhibit significant changes (Fig. 8A). Upon heating, oscillations at ~ 7.6 Å⁻¹ and ~ 8.2 Å⁻¹ are gradually combined into one oscillation and a shoulder at ~ 6.6 Å⁻¹ increases in intensity. The corresponding FTs (Fig. 8B) indicate that these evolutions are related to the reversing of Ni^E and Ni^{TC} peak relative intensities. Quantitative fitting indicates Ni^E proportions of 27 ± 4 %, 34 ± 5 %, 37 ± 5 %, 48 ± 5 %, 48 ± 7 %, and 51 ± 10 % for Ni10 8y samples heated to 50, 100, 150, 200, 300, and 400 °C, respectively (Table 2).

345 **3.7 Increased Ni stability during acid dissolution**

Acid dissolution experiments were conducted to detect possible modifications of Ni availability in heat-treated Mn oxides (Fig. 9). For Ni10_8y and Ni10_8y_200, Ni²⁺ and Mn²⁺ release rates in solution are large over the first few hours. With time, the metal release rates are gradually reduced till complete dissolution, because of the lower pool. In both samples, Ni²⁺ and Mn²⁺ releases are concurrent but Ni²⁺ is released slightly faster than Mn²⁺, consistent with the predominance of Ni^{TC}. Compared to Ni10_8y, the releases of Ni²⁺ and Mn²⁺ from Ni10_8y_200 are slower. ~99% of both Ni²⁺ and Mn²⁺ are released after 5 and 7 h for Ni10_8y and Ni10_8y_200, respectively, possibly as a result of a minor crystal growth. Ni10_8y_500 dissolution is much slower as 50 days are needed to achieve complete dissolution. Transformation of the initial birnessite to cryptomelane is the most likely hypothesis for such a dramatic decrease of dissolution rate. In addition, Ni²⁺ and Mn²⁺ releases are not concurrent any more, Mn²⁺ being released to solution much faster than Ni²⁺.

Nickel and Mn release kinetic curves can be fitted using the Kabai equation, 359 $\gamma_{Me}=1-\exp(-(kt)^{\alpha})$, where k and α are coefficients related to the dissolution rate constant 360 and the mineral structure, respectively (Fig. 9 - Kabai, 1973; Kaur et al., 2010). 361 Parameters k for both Ni²⁺ and Mn²⁺ dissolution are steadily decreasing with increasing 362 temperature. Ni and Mn release rates measured for Ni10 8y 200 are $67 \pm 4\%$ and $71 \pm$ 363 364 3% those from Ni10_8y, respectively. Ni and Mn release rates measured forNi10_8y_500 are dramatically reduced to $0.25 \pm 0.01\%$ and $1.51 \pm 0.06\%$ of those 365 from Ni10_8y, respectively. 366

The χ_{Ni} - χ_{Mn} curves obtained for Ni10_8y and Ni10_8y_200 follow the 1:1 line with a slight convexity however (Fig. 9d), and are essentially similar to that of the fresh Ni-doped birnessite (Qin et al., 2017). This suggests a uniform Ni distribution in these birnessite samples (Alvarez et al., 2006, 2007), consistent with elemental mapping (Fig. S5). Contrastingly, the χ_{Ni} - χ_{Mn} curve obtained for Ni10_8y_500 differs significantly from the previous two exhibiting a strong concavity, indicative of Ni location in the inner parts of cryptomelane framework (Alvarez et al., 2006, 2007).

374 **4** Discussion

4.1 Stability of (Ni-doped) birnessite during aging in dry state

When aged in the dry state at room temperature for up to 5 and 8 years, the 376 377 structure and properties of investigated Ni-free and Ni-doped hexagonal birnessites are almost unchanged, preserving their hexagonal layer symmetry, morphology and Mn 378 379 AOS. This suggests that layered Mn oxides such as birnessite may not as unstable as previously thought (Grangeon et al., 2014, 2015). Increased birnessite stability 380 observed in the present study is most likely related to their low content of Mn³⁺ as 381 previously discussed for the conversion of δ -MnO₂, a poorly crystalline analogue of 382 birnessite, to cryptomelane under similar experimental conditions (Grangeon et al., 383 2014, 2015), possibly reinforced by larger crystal sizes. Layer Mn(III) in 384 phyllomanganates plays indeed a key role in the phyllomanganate-to-tectomanganate 385 386 transformation (Cui et al., 2009; Atkins et al., 2014; Grangeon et al., 2014; Yang et al., 2018). Mn(III) contents of HB and Ni10 range 0.05-0.06 (Yin et al., 2012), much lower 387 than those (0.30) of the reacting δ -MnO₂ of Grangeon et al. (2014). Increased crystal 388 389 size of the present birnessites compared to δ -MnO₂ used by Grangeon et al. (2014) likely reinforces phyllomanganate stability. Crystals with small particle sizes generally 390 have high total free energy indeed and are thus prone to phase transformation (Luo et 391 al., 2008; Navrotsky et al., 2008). Theoretically, stability of hexagonal turbostratic 392 393 birnessite should thus be more stable than δ -MnO₂, owing to its larger crystallite sizes. Powder XRD patterns of our Ni-doped birnessites have strong (00l) reflections (Fig. 394 1a,b) that are absent in those of δ -MnO₂, suggesting a larger number of layers stacked 395 coherently along the **c***-axis. Further, the physical sizes of the present birnessite plates 396

(diameters of 100-300 nm) are also much larger than those of the vernadite samples
(mean diameter of ~10 nm – Yin et al., 2012; Grangeon et al., 2015).

399 Subject to a similar (micro-)structure, birnessite may thus be stable in arid environments such as upland soils and rock surfaces in deserts and represent a long-400 401 term sink for trace metal elements. Dark coatings commonly found on the upper surfaces of rocks in arid and semiarid regions, such as Gobi deserts in China and 402 Mongolia and in karst areas, are identified to be mainly composed of birnessite (Lu et 403 al., 2019), consistent with birnessite stability observed in our aging experiments. In the 404 present study, birnessites were aged in the dark with limited O₂ in sealed tubes. On the 405 other hand, rock surfaces exposed to both sun light and air, and birnessite from 406 varnishes and coatings may undergo additional evolution keeping their layer structure 407 408 anyhow. Field and laboratory studies have evidenced the reductive dissolution of birnesssite under photo illumination of aqueous suspensions however (Sunda and 409 Huntsman, 1994; Kim et al., 2012; Liu et al., 2021), consistent with theoretical 410 calculations (Sherman, 2005; Kwon et al., 2009). Further investigation is needed to 411 assess birnessite stability of birnessites under dry light condition and underlying 412 mechanisms at play. 413

414 **4.2 Enhanced thermal stability of birnessite by Ni doping**

The present results clearly suggest that Ni doping hampers the thermal-induced transformation of birnessite to cryptomelane, although thermodynamically favored. This mineral transformation may be affected by the presence of foreign cations into the birnessite structure however (Birkner and Navrotsky, 2017). It was for example

419	reported that Cu ²⁺ adsorption impedes Mn ²⁺ -induced transformation of vernadite to
420	other Mn oxides at room temperature, ascribing the increased stability to the decrease
421	of mineral surface energy induced by cation adsorption (Yang et al., 2019). Ni ²⁺
422	adsorption on and affinity for birnessite surfaces are similar to those of Cu ²⁺ (Kwon et
423	al., 2013; Manceau and Steinmann, 2021) and decrease of Ni-doped birnessite surface
424	energy likely contributes to its stability upon thermal treatment. Mn(III) content may
425	also play an important role for birnessite thermal transformation as reported in aqueous
426	(Cui et al., 2009; Atkins et al., 2014; Yang et al., 2018) and dry conditions (Grangeon
427	et al., 2014, 2015). Consistently, Ni doping decreases birnessite Mn(III) content (0.06
428	and 0.10 in Ni10_8y and HB_8y, respectively – Table S1) and thus increases its thermal
429	stability. Contrastingly, metal doping can significantly decrease birnessite particle size
430	(Yin et al., 2012, 2013), thus increasing its total free energy, and easing its mineralogical
431	transformation (Luo et al., 2008; Navrotsky et al., 2008). After Ni doping, birnessite
432	crystallinity was greatly reduced. The temperature for the transformation of our Ni-
433	doped birnessite to cryptomelane (500 °C) is lower than those reported for the
434	transformation of well crystallized samples (Chen et al., 1986; Golden et al., 1986; Cai
435	et al., 2001). In addition to its lower Mn(III) content, the increased stability of
436	Ni10_8y compared to HB_8y may be related to the overall decrease in the mineral total
437	free energy induced by Ni ²⁺ adsorption on the vacancies and/or edges. Foreign cation
438	adsorption thus appears to increase birnessite stability and could be a key factor for
439	phyllomanganate stability in natural environments (Yang et al., 2019).

4.3 Changes in Ni binding environments during thermal treatment

441	Nickel K-edge EXAFS fitting of Ni10_8y samples after heating to 25-400 °C
442	evidences changes in Ni crystal chemistry. Increasing temperature from 25 °C to 200
443	°C, increases linearly the proportion of edge-sharing Ni-Ni(Mn) pairs ($R^2 = 0.9502$, n
444	= 5, α = 0.05; Fig. 10A), suggesting migration of interlayer Ni cations to the octahedral
445	layer. Such rearrangement of cation configuration was previously observed during Ni ²⁺
446	adsorption on hexagonal birnessites when increasing reaction pH (Peacock, 2009; Peña
447	et al., 2010; Sorensen et al., 2020) and confirmed by density functional theory
448	calculations (Peña et al., 2010; Kwon et al., 2013). In the present study, heating can
449	provide the energy to overcome the energy barrier for this transformation, increasing
450	temperature then favoring the transformation. After heating to 200 °C, the proportion
451	of total Ni incorporated in birnessite octahedral layers, increases from 0.26 ± 0.04 in
452	Ni10_8y to 0.48 \pm 0.05 in Ni10_8y_200 (1-f value – Table 2). Increasing the
453	temperature above 200 °C does not promote further Ni migration into birnessite
454	octahedral layers however, possibly because remaining "interlayer" Ni cations sit on
455	layer edge sites by forming double corner-sharing complexes, which possess a similar
456	Ni-Mn distance of ~3.5 Å (Manceau et al., 2007).

Alternatively, the increased proportion of edge sharing Ni-Ni(Mn) pairs upon heating may result from the loss of the interlayer water molecules and subsequent rearrangement of adjacent layers. d-spacing of Ni10_8y first peak upon heating steadily decreases from birnessite (001) to cryptomelane (110) peak (Fig. 10B). Loss of interlayer water molecules decreases birnessite layer-to-layer distance (Fig. 6b), possibly allowing triple corner-sharing Ni and Mn atoms from adjacent layers to share

463	their coordination O atoms owing to sample turbostratism (systematic random stacking
464	of adjacent layers along the c^* -axis) and creating local tunnel-like domains (Fig. 6c)
465	(Grangeon et al., 2014). TGA data obtained for Ni10_8y in air consistently indicates a
466	significant loss of water molecules over the 35-200 °C range (Fig. 2b), rearrangement
467	of adjacent layers being also favored by increasing temperature. Both effects allow
468	more coordination O atoms to be shared by interlayer Ni/Mn from adjacent layers.
469	Departure of interlayer water molecules over the 35-200 °C range is consistent with
470	previous reports indicating the departure of nonstructural water molecules (sorbed to
471	mineral surfaces) at temperatures \leq 80-100 °C, and of interlayer water molecules from
472	80-100 °C to 205-250 °C (Villalobos et al., 2003; Zhu et al., 2012; Grangeon et al.,
473	2014). Though TGA analysis indicates further loss of hydroxyl groups and interlayer
474	water molecules during heating at 180-520 °C (Fig. 3b and Table S3b), the proportions
475	of edge-sharing Ni-Ni(Mn) pairs are almost stable over this temperature range, the
476	sample keeping its birnessite structure (Fig. 10A and Fig. 1d). Only when the heating
477	temperature is increased to 500 °C, Ni10_8y is totally transformed to cryptomelane (Fig
478	1d).

479 4.4 Mechanisms for enhanced Ni stability during acid dissolution

Along birnessite transformation to tectomanganates such as todorokite and nsutite incorporated foreign metal cations are released to solution to a large extent (Tu et al., 1994; Zhang et al., 2011; Atkins et al., 2014, 2016; Yang et al., 2018, 2019; Wu et al., 2019, 2021). These studies were conducted in aqueous solutions however, and no study is available on the fate of incorporated foreign metal cations during mineral

transformation occurring in the dry state. The present study shows that long-term aging 485 in dry state does not significantly modify Ni binding mechanisms to birnessite and thus 486 its availability. On the contrary, thermal treatment of aged Ni-doped birnessite 487 significantly reduces Ni availability along reductive acid dissolution of the mineral 488 matrix. Although heated solids retain their layer structure at temperatures < 500 °C, 489 Ni²⁺ release rate is gradually reduced with increasing temperature (Fig. 9). After heat-490 treatment to temperatures \geq 500 °C, birnessite is transformed to cryptomelane and Ni²⁺ 491 release rate is decreased by 99.75 \pm 0.01% compared to the initial aged birnessite. 492 493 Dissolution of Mn oxides in HCl solutions involves protonation, complexation, and reductive dissolution, with reductive dissolution by Cl⁻ being most efficient (Qin et al., 494 2017). Under similar conditions (H⁺ and Cl⁻ concentrations, temperature, stirring rate), 495 496 mineral structure, microstructure, and crystal chemistry may further influence dissolution rate. Although poorly crystalline minerals usually dissolve more readily 497 than their highly crystalline counterparts (Liu et al., 2019), Ni²⁺ release rates are 498 499 reduced for Ni10_8y_200 compared to Ni10_8y despite powder XRD and SEM analysis being indicative of Ni10 8y higher crystallinity and larger particle size. Ni 500 binding to the birnessite structure is thus likely responsible for the lower dissolution 501 rates of Ni10_8y_200. At temperatures ≤ 400 °C, increased proportion of edge-sharing 502 Ni-Ni(Mn) pairs (Fig. 10A), that may result both from Ni migration from interlayer to 503 the layer and from interlayer Ni cations sharing their coordination O atoms with Ni(Mn) 504 from adjacent layers, most likely accounts for this reduced availability. Increasing the 505 temperature further to 500 °C induces the transformation of Ni10_8y into cryptomelane, 506

that exhibits larger particle size and improved crystallinity compared to its Ni10_8y birnessite precursor (Figs. 1 and 3). Furthermore, Ni is thoroughly inserted in cryptomelane framework, as evidenced by HAADF-STEM analysis (Fig. 3i), thus providing a dramatically higher stability to this foreign cation. Though not directly relevant to acid leaching or weathering in natural environments, results from acid dissolution experiments might shed light on similar but much less harsh leaching processes in the environment.

514 **5 Conclusions**

515 Crystallinity, morphology, Mn AOS and Ni binding mechanisms are essentially unaffected when aging Ni-doped hexagonal birnessite in the dry state at room 516 temperature for up to 8 years. When heated, the proportion of edge-sharing Ni-Ni(Mn) 517 518 pairs in aged birenssites increases with temperature increasing from 25-200 °C before reaching a steady state when increasing the temperature further from 200-400 °C. Over 519 this whole temperature range, birnessite layer structure is retained and Mn AOS remains 520 constant. At higher temperatures (400-500 °C), Ni-doped birnessites transform to 521 cryptomelane, Ni doping increasing slightly birnessite stability and delaying the 522 523 mineral transformation. At 500 °C, Ni-doped birnessite with an initial Ni/Mn molar ratio of 0.13 fully transforms to cryptomelane, with a minor NiMnO₃ impurity. Along 524 this mineral transformation, Mn AOS is sharply increased, Mn cations being essentially 525 Mn⁴⁺ in the cryptomelane framework, and corner-sharing Mn-Mn(Ni) bond length is 526 decreased. The increasing proportion of edge-sharing Ni-Ni(Mn) pairs in heated 527 birnessite decreases Ni release to solution during reductive acid dissolution, whereas 528

529	incorporation of Ni in the cryptomelane framework essentially inhibits this release.
530	These results provide additional insights in the stability of birnessite-like minerals
531	under dry conditions that will most likely become more frequent owing to climate
532	change, and in the geochemical behavior of associated trace metal elements.
533	
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540	References
541 542 543	Alvarez M., Rueda E. H. and Sileo E. E. (2006) Structural characterization and chemical reactivity of synthetic Mn-goethites and hematites. <i>Chem. Geol.</i> 231, 288-299.
544 545	Alvarez M., Rueda E. H. and Sileo E. E. (2007) Simultaneous incorporation of Mn and Al in the goethite structure. <i>Geochim. Cosmochim. Acta</i> 71 , 1009-1020.
546 547 548	Atkins A. L., Shaw S. and Peacock C. L. (2014) Nucleation and growth of todorokite from birnessite: Implications for trace-metal cycling in marine sediments. <i>Geochim Cosmochim Acta</i> 144, 109-125
549 550	Atkins A. L., Shaw S. and Peacock C. L. (2016) Release of Ni from birnessite during transformation of birnessite to todorokite: Implications for Ni cycling in marine
551 552 553	 Beak D. G., Basta N. T., Scheckel K. G. and Traina S. J. (2008) Linking solid phase speciation of Pb sequestered to birnessite to oral Pb bioaccessibility:
554 555	implications for soil remediation. <i>Environ. Sci. Technol.</i> 42 , 779-785. Birkner N. and Navrotsky A. (2017) Thermodynamics of manganese oxides: Sodium,

Cai J., Liu J., Willis W. S. and Suib S. L. (2001) Framework doping of iron in tunnel
structure cryptomelane. *Chem. Mater.* 13, 2413-2422.

114, E1046-E1053.

557

560 Chen C. C., Golden D. C. and Dixon J. B. (1986) Transformation of synthetic birnessite

561	to cryptomelane: An electron microscopic study. Clays Clay Miner. 34, 565-571.
562	Ching S., Driscoll P. F., Kieltyka K. S., Marvel M. R. and Suib S. L. (2001) Synthesis
563	of a new hollandite-type manganese oxide with framework and interstitial
564	Cr(III). Chem. Commun., 2486-2487.
565	Cui H., Qiu G., Feng X., Tan W. and Liu F. (2009) Birnessites with different average
566	manganese oxidation states synthesized, characterized, and transformed to
567	todorokite at atmospheric pressure. Clays Clay Miner. 57, 715-724.
568	Drits V. A., Lanson B. and Gaillot A. C. (2007) Birnessite polytype systematics and
569	identification by powder X-ray diffraction. Am. Mineral. 92, 771-788.
570	Drits V. A., Silvester E., Gorshkov A. I. and Manceau A. (1997) Structure of synthetic
571	monoclinic Na-rich birnessite and hexagonal birnessite: I. Results from X-ray
572	diffraction and selected-area electron diffraction. Am. Mineral. 82, 946–961.
573	Giovanoli R. (1980) Vernadite is random-stacked birnessite. Mineral. Deposita 15, 251-
574	253.
575	Golden D., Dixon J. and Chen C. (1986) Ion exchange, thermal transformations, and
576	oxidizing properties of birnessite. Clays Clay Miner. 34, 511-520.
577	Grangeon S., Fernandez-Martinez A., Warmont F., Gloter A., Marty N., Poulain A. and
578	Lanson B. (2015) Cryptomelane formation from nanocrystalline vernadite
579	precursor: a high energy X-ray scattering and transmission electron microscopy
580	perspective on reaction mechanisms. Geochem. Trans. 16, 12.
581	Grangeon S., Lanson B. and Lanson M. (2014) Solid-state transformation of
582	nanocrystalline phyllomanganate into tectomanganate: influence of initial layer
583	and interlayer structure. Acta. Crystallogr. B 70, 828-838.
584	Grangeon S., Lanson B., Lanson M. and Manceau A. (2008) Crystal structure of Ni-
585	sorbed synthetic vernadite: a powder X-ray diffraction study. Mineral. Mag. 72,
586	1279-1291.
587	Grangeon S., Lanson B., Miyata N., Tani Y. and Manceau A. (2010) Structure of
588	nanocrystalline phyllomanganates produced by freshwater fungi. Am. Mineral.
589	95 , 1608-1616.
590	Johnston S. G., Burton E. D. and Moon E. M. (2016) Arsenic Mobilization Is Enhanced
591	by Thermal Transformation of Schwertmannite. Environ. Sci. Technol. 50,
592	8010-8019.
593	Jothiramalingam R., Viswanathan B. and Varadarajan T. K. (2006) Synthesis and
594	structural characterization of copper incorporated manganese oxide OMS-2
595	materials synthesized via potassium birnessite. Mater. Chem. Phys. 100, 257-
596	261.
597	Kabai J. (1973) Determination of specific activation energies of metal oxides and metal
598	oxide hydrates by measurement of the rate of dissolution. Acta Chim. Acad. Sci.
599	Hung 78 , 57-73.
600	Kaur N., Singh B. and Kennedy B. J. (2010) Dissolution of Cr, Zn, Cd, and Pb single-
601	and multi-metal-substituted goethite: relationship to structural, morphological,
602	and dehydroxylation properties. Clays Clay Miner. 58, 415-430.
603	Kelly S. D., Hesterberg D. and Ravel B. (2008) Analysis of soils and minerals using X-
604	ray absorption spectroscopy. In Methods of Soil Analysis, Part 5-Mineralogical

605	Methods (eds. A. L. Ulrey and R. L. Drees). Soil Science Society of America,
606	Wisconsin. pp. 387-463.
607	Kim K., Yoon H. I. and Choi W. (2012) Enhanced dissolution of manganese oxide in
608	ice compared to aqueous phase under illuminated and dark conditions. Environ.
609	Sci. Technol. 46, 13160-13166.
610	Kwon K. D., Refson K. and Sposito G. (2009) On the role of Mn(IV) vacancies in the
611	photoreductive dissolution of hexagonal birnessite. Geochim. Cosmochim. Acta
612	73 , 4142-4150.
613	Kwon K. D., Refson K. and Sposito G. (2013) Understanding the trends in transition
614	metal sorption by vacancy sites in birnessite. Geochim. Cosmochim. Acta 101,
615	222-232.
616	Landers M., Gilkes R. J. and Wells M. A. (2009) Rapid dehydroxylation of
617	nickeliferous goethite in lateritic nickel ore: X-ray diffraction and TEM
618	investigation. Clays Clay Miner. 57, 751-770.
619	Lanson B., Marcus M. A., Fakra S., Panfili F., Geoffroy N. and Manceau A. (2008)
620	Formation of Zn–Ca phyllomanganate nanoparticles in grass roots. Geochim.
621	Cosmochim. Acta 72 , 2478-2490.
622	Lee S., Xu H., Xu W. and Sun X. (2019) The structure and crystal chemistry of vernadite
623	in ferromanganese crusts. Acta. Crystallogr. B 75, 591-598.
624	Lee S. H., Kim T. W., Park D. H., Choy JH., Hwang SJ., Jiang N., Park SE. and
625	Lee YH. (2007) Single-step synthesis, characterization, and application of
626	nanostructured K _x Mn _{1-y} Co _y O _{2-δ} with controllable chemical compositions and
627	crystal structures. Chem. Mater. 19, 5010-5017.
628	Li M., Liu H., Chen T., Wei L., Wang C., Hu W. and Wang H. (2019) The transformation
629	of α -(Al, Fe)OOH in natural fire: Effect of Al substitution amount on fixation
630	of phosphate. Chem. Geol. 524, 368-382.
631	Liu L., Wang X., Zhu M., Ma J., Zhang J., Tan W., Feng X., Yin H. and Liu F. (2019)
632	The speciation of Cd in Cd-Fe coprecipitates: Does Cd substitute for Fe in
633	goethite structure? ACS Earth Space Chem. 3, 2225-2236.
634	Liu Y., Li Y., Li Y., Chen N., Ding H., Jin S., Wang C. and Lu A. (2021) Photo-
635	stimulated anoxic reduction of birnessite (δ -MnO ₂) by citrate and its fine
636	structural responses: Insights on a proton-promoted photoelectron transfer
637	process. Chem. Geol. 561, 120029.
638	Lu A. and Li Y. (2015) Reactivity of natural Mn oxide cryptomelane. In Advances in
639	the Environmental Biogeochemistry of Manganese Oxides (eds. X. Feng, W. Li,
640	M. Zhu, and D. L. Sparks). American Chemical Society, Washington. pp. 89-
641	106.
642	Lu A., Li Y., Ding H., Xu X., Li Y., Ren G., Liang J., Liu Y., Hong H., Chen N., Chu S.,
643	Liu F., Li Y., Wang H., Ding C., Wang C., Lai Y., Liu J., Dick J., Liu K. and
644	Hochella M. F., Jr. (2019) Photoelectric conversion on Earth's surface via
645	widespread Fe- and Mn-mineral coatings. Proc. Natl. Acad. Sci. USA 116, 9741-
646	9746.
647	Luo W., Hu W. and Xiao S. (2008) Size effect on the thermodynamic properties of silver
648	nanoparticles. J. Phys. Chem. C 112, 2359-2369.

649	Manceau A., Lanson M. and Geoffroy N. (2007) Natural speciation of Ni, Zn, Ba, and
650	As in ferromanganese coatings on quartz using X-ray fluorescence, absorption,
651	and diffraction. Geochim. Cosmochim. Acta 71, 95-128.
652	Manceau A., Lanson M. and Takahashi Y. (2014) Mineralogy and crystal chemistry of
653	Mn, Fe, Co, Ni, and Cu in a deep-sea Pacific polymetallic nodule. Am. Mineral.
654	99 , 2068-2083.
655	Manceau A., Marcus M. A. and Grangeon S. (2012) Determination of Mn valence states
656	in mixed-valent manganates by XANES spectroscopy. Am. Mineral. 97, 816-
657	827.
658	Manceau A. and Steinmann S. N. (2021) Nature of high- and low-affinity metal surface
659	sites on birnessite nanosheets. ACS Earth Space Chem. 5, 66-76.
660	McKenzie R. M. (1971) The synthesis of birnessite, cryptomelane, and some other
661	oxides and hydroxides of manganese. Mineral. Mag. 38, 493-503.
662	Morgan J. J. (2005) Kinetics of reaction between O2 and Mn(II) species in aqueous
663	solutions. Geochim. Cosmochim. Acta 69, 35-48.
664	Murphy B. F. and Timbal B. (2008) A review of recent climate variability and climate
665	change in southeastern Australia. Int. J. Climatol. 28, 859-879.
666	Navrotsky A., Mazeina L. and Majzlan J. (2008) Size-Driven Structural and
667	Thermodynamic Complexity in Iron Oxides. Science 319, 1635-1638.
668	Peacock C. L. (2009) Physiochemical controls on the crystal-chemistry of Ni in
669	birnessite: genetic implications for ferromanganese precipitates. Geochim.
670	Cosmochim. Acta 73 , 3568-3578.
671	Peña J., Kwon K. D., Refson K., Bargar J. R. and Sposito G. (2010) Mechanisms of
672	nickel sorption by a bacteriogenic birnessite. Geochim. Cosmochim. Acta 74,
673	3076-3089.
674	Perrier N., Gilkes R. J. and Colin F. (2006) Heating Fe oxide-rich soils increases the
675	dissolution rate of metals. Clays Clay Miner. 54, 165-175.
676	Post J. E. (1999) Manganese oxide minerals: Crystal structures and economic and
677	environmental significance. Proc. Natl. Acad. Sci. USA 96, 3447-3454.
678	Post J. E. and Appleman D. E. (1988) Chalcophanite, ZnMn ₃ O ₇ .3H ₂ O: New crystal-
679	structure determinations. Am. Mineral. 73, 1401-1404.
680	Qin Z., Xiang Q., Liu F., Xiong J., Koopal L. K., Zheng L., Ginder-Vogel M., Wang M.,
681	Feng X., Tan W. and Yin H. (2017) Local structure of Cu ²⁺ in Cu-doped
682	hexagonal turbostratic birnessite and Cu ²⁺ stability under acid treatment. <i>Chem.</i>
683	<i>Geol.</i> 466 , 512-523.
684	Ravel B. and Newville M. (2005) ATHENA, ARTEMIS, HEPHAESTUS: data analysis
685	for X-ray absorption spectroscopy using IFEFFIT. J. Synchrotron Radiat. 12,
686	537-541.
687	Rehr J. J., Zabinsky S. I. and Albers R. C. (1992) High-order multiple scattering
688	calculations of X-ray-absorption fine structure. Phys. Rev. Lett. 69, 3397-3400.
689	Rein G., Cleaver N., Ashton C., Pironi P. and Torero J. L. (2008) The severity of
690	smouldering peat fires and damage to the forest soil. Catena 74, 304-309.
691	Sherman D. M. (2005) Electronic structures of iron(III) and manganese(IV)
692	(hydr)oxide minerals: Thermodynamics of photochemical reductive dissolution

693	in aquatic environments. Geochim. Cosmochim. Acta 69, 3249-3255.
694	Silvester E. J., Manceau A. and Drits V. A. (1997) The structure of monoclinic Na-rich
695	birnessite and hexagonal birnessite. Part 2. Results from chemical studies and
696	EXAFS spectroscopy American Mineralogist, 82, 962–978.
697	Simanova A. A. and Pena J. (2015) Time-resolved investigation of cobalt oxidation by
698	Mn(III)-rich delta-MnO ₂ using quick X-ray absorption spectroscopy. <i>Environ</i> .
699	Sci. Technol. 49, 10867-10876.
700	Sinha A. K., Pradhan M. and Pal T. (2013) Morphological evolution of two-dimensional
701	MnO ₂ nanosheets and their shape transformation to one-dimensional ultralong
702	MnO ₂ nanowires for robust catalytic activity. J. Phys. Chem. C 117, 23976-
703	23986.
704	Sorensen J. V., Gueguen B., Stewart B. D., Peña J., Rouxel O. and Toner B. M. (2020)
705	Large nickel isotope fractionation caused by surface complexation reactions
706	with hexagonal birnessite. Chem. Geol. 537, 119481.
707	Sunda W. G. and Huntsman S. A. (1994) Photoreduction of manganese oxides in
708	seawater. Mar. Chem. 46, 133-152.
709	Tang Q., Jiang L., Liu J., Wang S. and Sun G. (2014) Effect of surface manganese
710	valence of manganese oxides on the activity of the oxygen reduction reaction in
711	alkaline media. ACS Catal. 4, 457-463.
712	Taylor R., McKenzie R. and Norrish K. (1964) The mineralogy and chemistry of
713	manganese in some Australian soils. Soil Res. 2, 235-248.
714	Tu S., Racz G. J. and Goh T. B. (1994) Transformations of Synthetic Birnessite as
715	Affected by pH and Manganese Concentration. Clays Clay Miner. 42, 321-330.
716	Vicente-Serrano S. M., Quiring S. M., Peña-Gallardo M., Yuan S. and Domínguez-
717	Castro F. (2020) A review of environmental droughts: Increased risk under
718	global warming? Earth-Sci. Rev. 201, 102953.
719	Villalobos M., Toner B., Bargar J. and Sposito G. (2003) Characterization of the
720	manganese oxide produced by pseudomonas putida strain MnB1. Geochim.
721	Cosmochim. Acta 67, 2649-2662.
722	Wang Y., Benkaddour S., Marafatto F. F. and Pena J. (2018) Diffusion- and pH-
723	dependent reactivity of layer-type MnO2: Reactions at particle edges versus
724	vacancy sites. Environ. Sci. Technol. 52, 3476-3485.
725	Webb S. M. (2005) SIXPACK: a graphical user interface for XAS analysis using
726	IFEFFIT Physica Scripta T115, 1011–1014.
727	Wu Z., Lanson B., Feng X., Yin H., Tan W., He F. and Liu F. (2021) Transformation of
728	the phyllomanganate vernadite to tectomanganates with small tunnel sizes:
729	Favorable geochemical conditions and fate of associated Co. Geochim.
730	Cosmochim. Acta 295 , 224-236.
731	Wu Z., Peacock C. L., Lanson B., Yin H., Zheng L., Chen Z., Tan W., Qiu G., Liu F.
732	and Feng X. (2019) Transformation of Co-containing birnessite to todorokite:
733	Effect of Co on the transformation and implications for Co mobility. Geochim.
734	<i>Cosmochim. Acta</i> 246 , 21-40.
735	Yang P., Lee S., Post J. E., Xu H., Wang Q., Xu W. and Zhu M. (2018) Trivalent
736	manganese on vacancies triggers rapid transformation of layered to tunneled

737	manganese oxides (TMOs): Implications for occurrence of TMOs in low-
738	temperature environment. Geochim. Cosmochim. Acta 240, 173-190.
739	Yang P., Post J. E., Wang Q., Xu W., Geiss R., McCurdy P. R. and Zhu M. (2019) Metal
740	adsorption controls stability of layered manganese oxides. Environ. Sci. Technol.
741	53 , 7453-7462.
742	Yin H., Dai X. X., Zhu M. Q., Li F. H., Feng X. H. and Liu F. (2015) Fe-doped
743	cryptomelane synthesized by refluxing at atmosphere: Structure, properties and
744	photocatalytic degradation of phenol. J. Hazard. Mater. 296, 221-229.
745	Yin H., Li H., Wang Y., Ginder-Vogel M., Qiu G. H., Feng X. H., Zheng L. R. and Liu
746	F. (2014) Effects of Co and Ni co-doping on the structure and reactivity of
747	hexagonal birnessite. Chem. Geol. 381, 10-20.
748	Yin H., Tan W. F., Zheng L. R., Cui H. J., Qiu G. H., Liu F. and Feng X. H. (2012)
749	Characterization of Ni-rich hexagonal birnessite and its geochemical effects on
750	aqueous Pb ²⁺ /Zn ²⁺ and As(III). <i>Geochim. Cosmochim. Acta</i> 93 , 47-62.
751	Zhang Q., Cheng X. D., Qiu G. H., Liu F. and Feng X. H. (2016) Size-controlled
752	synthesis and formation mechanism of manganese oxide OMS-2 nanowires
753	under reflux conditions with KMnO4 and inorganic acids. Solid State Sci. 55,
754	152-158.
755	Zhang Q., Xiao Z., Feng X., Tan W., Qiu G. and Liu F. (2011) a-MnO ₂ nanowires
756	transformed from precursor δ -MnO ₂ by refluxing under ambient pressure: The
757	key role of pH and growth mechanism. Mater. Chem. Phys. 125, 678-685.
758	Zhu M., Farrow C. L., Post J. E., Livi K. J. T., Billinge S. J. L., Ginder-Vogel M. and
759	Sparks D. L. (2012) Structural study of biotic and abiotic poorly-crystalline
760	manganese oxides using atomic pair distribution function analysis. Geochim.
761	<i>Cosmochim. Acta</i> 81 , 39-55.
762	Zhu Y., Luo Q. and Liu Y. (2022) Deserts and Desertification. IntechOpen, London.
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	Sampla	Dry aging (year)		Thermal treatment of the 8 year aged sample at different temperatures (°C)							
	Sample	5	8	50	100	150	200	300	400	500	
	HB	HB_5y	HB_8y	HB_8y_50	HB_8y_100	HB_8y_150	HB_8y_200	HB_8y_300	HB_8y_400	HB 8y 500	
	Ni10	Ni10_5y	Ni10_8y	Ni10_8y_50	Ni10_8y_100	Ni10_8y_150	Ni10_8y_200	Ni10_8y_300	Ni10_8y_400	Ni10_8y_500	
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 Table 1 Nomenclature of hexagonal turbostratic birnessites and their derivatives.

Shells	Parameters	Ni10_8y	Ni10_8y_50	Ni10_8y_100	Ni10_8y_150	Ni10_8y_200	Ni10_8y_300	Ni10_8y_400
Ni-O ₁ ^a	R (Å)	2.040(0.006)	2.039(0.006)	2.038(0.007)	2.034(0.007)	2.035(0.007)	2.013(0.007)	2.015(0.009)
	$\sigma^2(\text{\AA})$	0.0051(0.0004)	0.0050(0.0004)	0.0051(0.0005)	0.0052(0.0005)	0.0050(0.0005)	0.0061(0.0005)	0.0065(0.0006)
Ni-Mn _{1,E}	CN	6×(1-f) ^b	6×(1-f)	6×(1-f)	6×(1-f)	6×(1-f)	6×(1-f)	6×(1-f)
	R (Å)	2.860(0.010)	2.846(0.010)	2.849(0.010)	2.845(0.009)	2.863(0.009)	2.849(0.009)	2.865(0.013)
	$\sigma^2(\text{\AA})$	σ^2 (Ni-Mn _{1,TC})	σ^2 (Ni-Mn _{1,TC})	σ^2 (Ni-Mn _{1,TC})	σ^2 (Ni-Mn _{1,TC})			
Ni-O ₂ ^a	R (Å)	3.340(0.065)	3.326(0.046)	3.317(0.039)	3.337(0.059)	3.357(0.103)	3.299(0.046)	3.293(0.052)
	$\sigma^2(\text{\AA})$	0.0165(0.0086)	0.0103(0.0052)	0.0062(0.0035)	0.0114(0.0067)	0.0217(0.0143)	0.0105(0.0058)	0.0096(0.0062)
Ni-Mn _{1,TC}	CN	6×f	6×f	6×f	6×f	6×f	6×f	6×f
	R (Å)	3.471(0.008)	3.484(0.009)	3.487(0.013)	3.481(0.011)	3.464(0.012)	3.457(0.016)	3.453(0.023)
	$\sigma^2(\text{\AA})$	0.0057(0.0005)	0.0050(0.0006)	0.0046(0.0008)	0.0049(0.0007)	0.0063(0.0006)	0.0072(0.0009)	0.0084(0.0013)
Ni-O ₃	CN	9×f+12×(1-f)	9×f+12×(1-f)	9×f+12×(1-f)	9×f+12×(1-f)	9×f+12×(1-f)	9×f+12×(1-f)	9×f+12×(1-f)
	R (Å)	4.449(0.025)	4.457(0.025)	4.462(0.023)	4.459(0.026)	4.450(0.060)	4.395(0.030)	4.426(0.051)
	$\sigma^2(\text{\AA})$	0.0075(0.0027)	0.0072(0.0026)	0.0054(0.0022)	0.0063(0.0026)	0.0146(0.0082)	0.0094(0.0035)	0.0132(0.0065)
Ni-Mn _{2,E}	CN	6×(1-f)	6×(1-f)	6×(1-f)	6×(1-f)	6×(1-f)	6×(1-f)	6×(1-f)
	R (Å)	4.976(0.039)	4.992(0.044)	4.988(0.046)	4.985(0.040)	5.001(0.028)	4.966(0.023)	4.962(0.042)
	$\sigma^2(\text{\AA})$	σ^2 (Ni-Mn _{2,TC})	σ^2 (Ni-Mn _{2,TC})	σ^2 (Ni-Mn _{2,TC})	σ ² (Ni-Mn _{2,TC})	σ^2 (Ni-Mn _{2,TC})	σ^2 (Ni-Mn _{2,TC})	σ^2 (Ni-Mn _{2,TC})
Ni-Mn _{2,TC}	CN	6×f	6×f	6×f	6×f	6×f	6×f	6×f
	R (Å)	5.472(0.020)	5.437(0.023)	5.448(0.032)	5.475(0.032)	5.545(0.022)	5.493(0.027)	5.457(0.057)
	$\sigma^2(\text{\AA})$	0.0060(0.0018)	0.0065(0.0022)	0.0073(0.0030)	0.0068(0.0027)	0.0035(0.0011)	0.0058(0.0017)	0.0092(0.0039)
	f	0.735(0.036)	0.728(0.039)	0.656(0.050)	0.628(0.049)	0.519(0.045)	0.516(0.065)	0.489(0.104)
E0(eV)		-2.5(1.1)	-0.2(1.1)	-2.5(1.2)	0.1(1.2)	-1.8(1.2)	-3.5(1.1)	-3.3(1.5)
Chi Sq		183.54	289.48	256.49	447.79	360.36	249.14	276.41
R factor		0.0229	0.0215	0.0285	0.0291	0.0355	0.0304	0.0409

Table 2 Ni K-edge EXAFS structural parameters for thermal transformation products of the 8-year aged Ni10 sample.

786 ^aThe coordination numbers of these Ni-O shells are 6. ^bf is the fraction of Ni adsorbed on vacancies as triple corner-sharing complexes (TC) of the total Ni in samples.



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Fig. 1 Powder XRD patterns of pristine and aged turbostratic hexagonal birnessite 789 790 samples HB (a) and Ni10 (b) for 5 and 8 years at 25 °C, and of thermal transformation products derived from 8-year aged samples HB 8y (c) and Ni10 8y (d). In panels (a)-791 (b), the high-angle region (30-80 $^{\circ}$ 2 θ) was scaled by a factor 4 (HB) or 2 (Ni10), and 792 793 in each series the pristine sample was overlaid to emphasize the evolution of the XRD profiles upon aging. In panels (c)-(d), calculated XRD patterns for cryptomelane (ICDD 794 74-1451; dashed light gray lines) and NiMn(IV)O₃ (ICDD 75-2089; solid light gray 795 796 lines) were overlaid to identify the transformation products at high temperatures.



Fig. 2 TGA curves of Ni-free and Ni-doped birnessites before and after aging in dry state at room temperature for 8 years in N_2 atmosphere (a), and those of 8-year aged samples in dry air atmosphere (b). The data for the fresh Ni-doped samples (dashed lines) are from Yin et al. (Yin et al., 2012). For all the samples, the TGA curves are averaged from 2-3 measurements, standard errors being shown as light gray shading.

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Fig. 3 High magnification FESEM images of HB_8y (a), HB_8y_300 (b), HB_8y_400 (c), HB_8y_500 (d), Ni10_8y (e), and Ni10_8y_400 (f) with the low magnification images in the inset of each panel. TEM image of Ni10_8y_500 (g) with an independent crystal in the inset with an area labeled using a blue bracket for the HAADF-STEM image with FFT in the inset (h) and the atomic resolution image (i).



Fig. 4 Proportions of Mn species (Mn⁴⁺, Mn³⁺, and Mn²⁺) in Ni-free HB (a) and Nidoped Ni10 (b) birnessite samples after aging at room temperature for 5 and 8 years
and in samples derived from heat treatment of the 8-year aged samples at different
temperatures, and corresponding variations of the average Mn oxidation state. Data for
fresh HB and Ni10 are from Yin et al. (2012).



Fig. 5 Mn K-edge k³-weighted EXAFS spectra (a-b) and corresponding Fourier 833 transforms (c-d) of aged and heat-treated HB (left) and Ni10 (right) birnessite samples. 834 Samples are colored-coded: fresh samples (red), 5-year aged (green), 8-year aged (blue), 835 836 HB 8y and Ni10 8y samples heated to 50 °C (dark yellow), 100 °C (navy), 150 °C (purple), 200 °C (olive), 300 °C (orange), 400 °C (violet), and 500 °C (light magenta). 837 Solid and dashed lines are experimental data and their best fits, respectively. All spectra 838 839 were fitted using a birnessite model derived from Grangeon et al. (2010), except for HB 8y 400, HB 8y 500, and Ni10 8y 500 spectra that were fitted using a 840 cryptomelane model derived from Yin et al. (2015). Light grey bands indicate the main 841 842 changes in these spectra.



Fig. 6 Evolution as a function of temperature of the EXAFS-derived Mn-Mn(Ni)
corner-sharing distances in HB and Ni10 (a), and illustrations of these distances (blue
bold solid lines) in the birnessite (b) and cryptomelane (c) structures.



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Fig. 7 Nickel K-edge k³-weighted functions (a) and corresponding Fourier transforms (b) of Ni-doped birnessite samples before (light gray lines) and after (orange lines) dry aging at room temperature for 8 years. Fourier transforms (b) were normalized to the first Ni-O shell ($R + \Delta R \sim 1.6$ Å).





Fig. 8 Nickel K-edge k³-weighted EXAFS functions (A) and corresponding Fourier
transforms (FTs, B) of heat-treated 8-year aged Ni-doped birnessite (Ni10_8y).
Experimental data and their best fits are shown as blue and red lines, respectively.
Ni10_8y (a), Ni10_8y_50 (b), Ni10_8y_100 (c), Ni10_8y_150 (d), Ni10_8y_200 (e),
Ni10_8y_300 (f), and Ni10_8y_400 (g).



Fig. 9 Releases of Mn and Ni as a function of time along reductive acid dissolution (2 M HCl solution at 37 °C) of Ni10 8y (a), Ni10 8y 200 (b) and Ni10 8y 500 (c), and corresponding %Ni dissolved (χ_{Ni}) versus %Mn dissolved (χ_{Mn}) curves (d). χ_{Mn} and χ_{Ni} are normalized to the final concentration after complete dissolution. In panels a, b and c, the solid lines are the best fit to the dissolution curve using Kabai equation (Kabai, 1973). Fitting parameters are also listed in tables inserted in the panels. In panel d, the solid black line is the 1:1 line.



883 Fig. 10 Evolution of the proportion of edge-sharing Ni-Ni(Mn) pairs as a function of

- temperature for heat-treated Ni10_8y (A), and shift of the first peak position for heat-
- treated Ni10_8y (B): Ni10_8y (a), Ni10_8y_50 (b), Ni10_8y_100 (c), Ni10_8y_150
- 886 (d), Ni10_8y_200 (e), Ni10_8y_300 (f), Ni10_8y_400 (g), and Ni10_8y_500 (h).