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The global distribution of leaf chlorophyll content

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37

38 **Abstract**

39 Leaf chlorophyll is central to the exchange of carbon, water and energy between the biosphere and the
40 atmosphere, and to the functioning of terrestrial ecosystems. This paper presents the first spatially
41 continuous view of terrestrial leaf chlorophyll content (Chl_{Leaf}) across a global scale. Weekly maps of
42 Chl_{Leaf} were produced from ENVISAT MERIS full resolution (300 m) satellite data with a two-stage
43 physically-based radiative transfer modelling approach. Firstly, leaf-level reflectance was derived from
44 top-of-canopy satellite reflectance observations using 4-Scale and SAIL canopy radiative transfer models

45 for woody and non-woody vegetation, respectively. Secondly, the modelled leaf-level reflectance was
46 used in the PROSPECT leaf-level radiative transfer model to derive Chl_{Leaf} . The Chl_{Leaf} retrieval algorithm
47 was validated with measured Chl_{Leaf} data from 248 sample measurements at 28 field locations, and
48 covering six plant functional types (PFTs). Modelled results show strong relationships with field
49 measurements, particularly for deciduous broadleaf forests ($R^2 = 0.67$; $\text{RMSE} = 9.25 \mu\text{g cm}^{-2}$; $p < 0.001$),
50 croplands ($R^2 = 0.41$; $\text{RMSE} = 13.18 \mu\text{g cm}^{-2}$; $p < 0.001$) and evergreen needleleaf forests ($R^2 = 0.47$; RMSE
51 $= 10.63 \mu\text{g cm}^{-2}$; $p < 0.001$). When the modelled results from all PFTs were considered together, the
52 overall relationship with measured Chl_{Leaf} remained good ($R^2 = 0.47$, $\text{RMSE} = 10.79 \mu\text{g cm}^{-2}$; $p < 0.001$).
53 This result was an improvement on the relationship between measured Chl_{Leaf} and a commonly used
54 chlorophyll-sensitive spectral vegetation index; the MERIS Terrestrial Chlorophyll Index (MTCI; $R^2 = 0.27$,
55 $p < 0.001$). The global maps show large temporal and spatial variability in Chl_{Leaf} , with evergreen broadleaf
56 forests presenting the highest leaf chlorophyll values with global annual median of $54.4 \mu\text{g cm}^{-2}$. Distinct
57 seasonal Chl_{Leaf} phenologies are also visible, particularly in deciduous plant forms, associated with
58 budburst and crop growth, and leaf senescence. It is anticipated that this global Chl_{Leaf} product will make
59 an important step towards the explicit consideration of leaf-level biochemistry in terrestrial water,
60 energy and carbon cycle modelling.

61

62 Key words: Radiative transfer, 4-Scale, SAIL, PROSPECT, leaf biochemistry, MERIS, satellite, remote
63 sensing, leaf physiology, carbon cycle, ecosystem modelling, phenology

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69 1.0 Introduction

70 Chlorophyll molecules facilitate the conversion of absorbed solar radiation into stored chemical energy,
71 and the exchange of matter and energy fluxes between the biosphere and the atmosphere. Our ability
72 to accurately model these fluxes is important to forecasting carbon dynamics, within the context of a
73 changing climate. However, within conventional carbon modelling approaches, the parameterisation of
74 vegetation structure and physiological function over both spatial and temporal domains, with an
75 acceptable level of accuracy remains challenging (Groenendijk et al. 2011, Houborg et al. 2015). Within
76 such modelling approaches, leaf area index (LAI) is a core biophysical parameter used to represent
77 vegetation density, seasonal phenology and the fraction of absorbed PAR by vegetation that is
78 converted to biomass (Bonan et al. 2011). The ecological importance of LAI has led to well-validated
79 datasets of LAI maps at global scales and fine spatial resolution (~1 km) (Baret et al. 2013, Deng et al.
80 2006). However, recent studies have found that a temporal decoupling between vegetation structure
81 and function can occur (Croft, Chen, and Zhang 2014b, Croft, Chen, Froelich, et al. 2015, Walther et al.
82 2016), particularly in deciduous vegetation with a strong seasonal phenology. Chlorophyll molecules
83 comprise an important part of a plant's "photosynthetic apparatus" (Peng et al. 2011), through the
84 harvesting of light and the production of biochemical energy for use within the Calvin-Benson cycle
85 (Porcar-Castell et al. 2014). Leaf chlorophyll content (Chl_{leaf}) therefore represents a plant's physiological
86 status, and is closely related to plant photosynthetic function; demonstrating a strong relationship to
87 plant photosynthetic capacity (V_{cmax}) (Croft et al. 2017). Neglecting to consider chlorophyll phenology
88 within carbon models can lead to an overestimation of the amount of plant carbon uptake at the start
89 and end of a growing season in deciduous vegetation (Croft, Chen, Froelich, et al. 2015, Luo et al. 2018).
90 The incorporation of inter- and intra- annual variations of chlorophyll within ecosystem models has been
91 shown to improve the simulations of carbon and water fluxes (Luo et al. 2018, He et al. 2017).

92

93 Global efforts to map Chl_{Leaf} have been hampered by the complexity in the relationship between
94 satellite-derived canopy reflectance and plant biophysical and biochemical variables. Thus far, satellite
95 remote sensing applications to map Chl_{Leaf} have largely been limited to empirical methods, via the
96 derivation of statistical relationships between spectral vegetation indices (VIs) and leaf or canopy
97 chlorophyll content (Le Maire, François, and Dufrêne 2004, Sims and Gamon 2002). Indices that include
98 'red-edge' wavelengths (690–740 nm) are the most strongly related to Chl_{Leaf} (Croft, Chen, and Zhang
99 2014a, Malenovský et al. 2013), due to the ready saturation of chlorophyll absorption bands when Chl_{Leaf}
100 exceeds $\sim 30 \mu\text{g cm}^{-2}$ (Croft and Chen 2018). Some studies have shown promising results using empirical
101 methods (Datt 1998, Haboudane et al. 2002). However, this is usually achieved at local scales, within
102 closely related species (Blackburn 1998) or for uniform, closed canopies, where the vegetation stand
103 essentially behaves as a 'big leaf' (Gamon et al. 2010), and where contributions from other variables,
104 such as background vegetation and non-photosynthetic elements are low. At the leaf level, variations in
105 internal leaf structure, leaf thickness and water content, differentially affect leaf reflectance (Serrano
106 2008, Croft, Chen, and Zhang 2014a). At the canopy scale, vegetation architecture including LAI, foliage
107 clumping, stand density, non photosynthetic elements and understory vegetation, in addition to the
108 sun-view geometry, affect measured reflectance factors (Demarez and Gastellu-Etchegorry 2000,
109 Verrelst et al. 2010, Malenovský et al. 2008).

110

111 An alternative satellite-based approach for deriving Chl_{Leaf} from top-of-canopy reflectance data is
112 through the use of physically-based radiative transfer models. Radiative transfer models provide a direct
113 physical relationship between canopy reflectance and Chl_{Leaf} because they are underpinned by physical
114 laws that determine the interaction between solar radiation and the vegetation canopy. Leaf-level
115 estimation of foliar chlorophyll is achieved by coupling a canopy model and leaf optical model, in a two-
116 step process (Croft and Chen 2018): firstly to derive leaf level reflectance from canopy reflectance and

117 then to derive leaf pigment content from the modeled leaf reflectance (Zhang et al. 2008, Croft et al.
118 2013, Zarco-Tejada et al. 2004). A number of canopy models have been used for this purpose, ranging
119 from turbid medium models (e.g. SAIL) (Verhoef 1984), hybrid geometric optical and radiative transfer
120 models (e.g. 4-SCALE (Chen and Leblanc 1997), GeoSAIL (Huemmrich 2001), DART (Gastellu-Etcheberry,
121 Martin, and Gascon 2004)) in which the turbid media are constrained into a geometric form (i.e. a leaf,
122 shoot, branch and/or crown), and ray-tracing techniques (FLIGHT) (North 1996). At the leaf level, the
123 most widely used leaf optical model is PROSPECT (Jacquemoud and Baret 1990), which has been
124 extensively applied and validated across a wide range of plant species (Malenovsky et al. 2006), due to
125 the small number of input parameters required and its ease of inversion. Previous research has
126 demonstrated the strength of this physically-based method for a number of different ecosystems (Zarco-
127 Tejada et al. 2004, Zhang et al. 2008, Croft et al. 2013, Demarez and Gastellu-Etcheberry 2000, Houborg
128 and Boegh 2008). However, this approach has yet to be applied at the global scale.

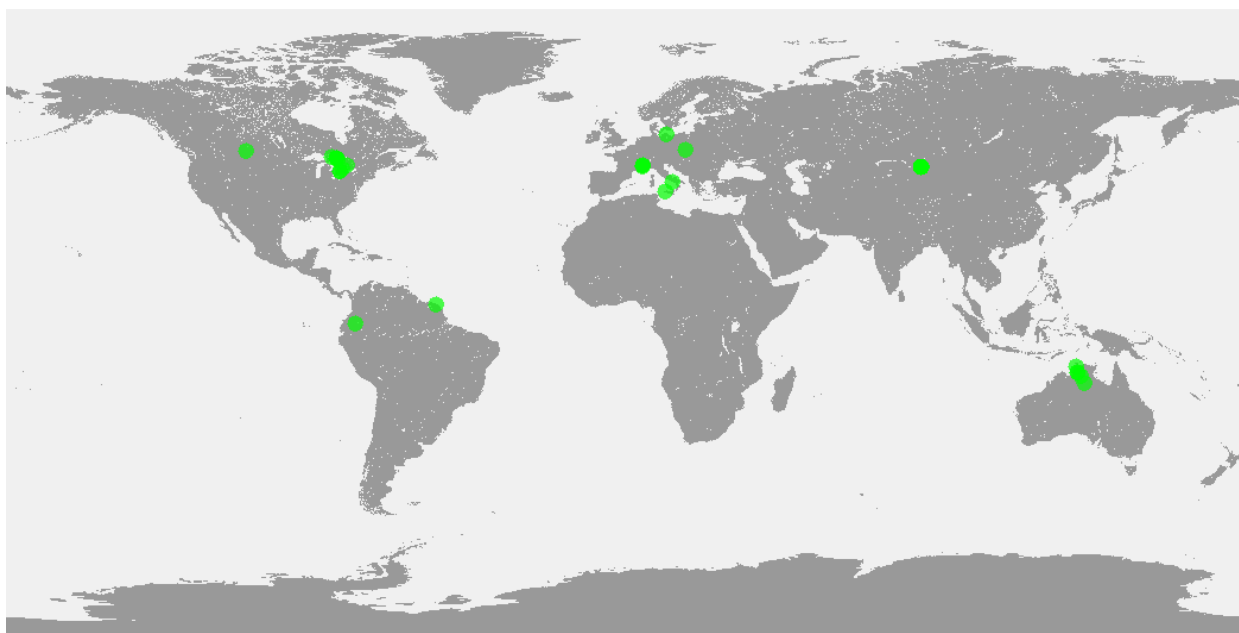
129
130 This paper presents the first global Chl_{Leaf} map from satellite data using physically-based radiative
131 transfer models. Chl_{Leaf} is defined on a leaf-area basis, as chlorophyll content per half the total surface
132 leaf area. Expressing Chl_{Leaf} by leaf area (as opposed to by dry mass) is the closer representation of what
133 is directly measured by a satellite instrument, and is most appropriate for linking Chl_{Leaf} to ecosystem
134 processes, such as carbon and water fluxes in relation to surfaces (Wright et al. 2004). Chl_{Leaf} is modelled
135 from ENVISAT MEdium Resolution Imaging Spectrometer (MERIS) 300 m reflectance data in a two-step
136 modelling approach, using coupled canopy and leaf radiative transfer models. Modelled Chl_{Leaf} results
137 were subsequently validated using measured ground data at a range of different field sites over six
138 different plant functional types (PFTs). Chl_{Leaf} maps are produced at the global scale every seven days for
139 an entire calendar year 2011 in order to provide spatially- and temporally-distributed leaf chlorophyll
140 content, and consequent physiological information for ecosystem modelling and ecological applications.

141

142 2.0 Methods

143 2.1 Ground data and validation sites

144 We used measured leaf chlorophyll data from 248 sampling measurements within 28 field locations
145 covering six PFTs (Figure 1 and Table 1) to validate the Chl_{Leaf} retrieval algorithm. These data included 49
146 measurements in deciduous broadleaf forests (DBF), 9 measurements in evergreen broadleaf forests
147 (EBF), 100 measurements in needleleaf forests (ENF), 55 measurements in croplands (CRP), 28
148 measurements in grassland (GRS) and 7 measurements in shrublands (SHR).



149

150 *Figure 1: Locations of the twenty-eight field locations used for leaf chlorophyll content validation.*

151

152 For each field location, the number of discrete sites or dates over which individual ground
153 measurements were taken are reported in Table 1. The number of replicates and spatial sampling varied
154 between the field sites, due to the nature of data collection in individual projects and between
155 researchers, and represents the large number of sources that have come together to produce this work.

157 *Table 1: Details of the ground measurements of leaf chlorophyll content used for validation used in this*
 158 *study, including the sampling location (Lat = latitude, Long = longitude) and sampling year, dominant*
 159 *species, plant functional type (PFT), mean site leaf area index (LAI) value and chlorophyll determination*
 160 *method (Chl. method).*

Site name	Country	Sampling years	Lat/ Long*	PFT	Dominant species	Mean LAI	Mean Chl. ($\mu\text{g cm}^{-2}$)	No. of sites/dates	Chl. method†	Reference
Haliburton	Canada	2004	45.24 -78.54	DBF	Sugar maple	5.5	38.3	8	Lab	Zhang et al. (2007)
Borden Forest	Canada	2013-15	44.32 -79.93	DBF	Red maple	4.2	37.8	13	Lab	Croft et al. (2015)
Bioindicators DBF	Canada	2002	46.84, -81.41	DBF	White birch, poplar	-	34.3	26	Lab	-
Sudbury DBF	Canada	2007	47.16, -81.71	DBF	Trembling aspen	4.3	48.8	2	Lab	Simic et al. (2011)
Amazon FG	French Guyana	2007-08	5.28, -52.92	EBF	-	6.7	62.1	3	Chl. meter	Rowland et al. (2014)
Amazon Ecuador	Ecuador	2012	-0.18, -76.36	EBF	-	-	53.6	6	Field spec.	Arellano et al. (2017)
Sudbury ENF	Canada	2007	47.18, -81.74	ENF	Black spruce		29.0	10	Lab	Simic et al. (2011)
Chapleau	Canada	2012	47.58, -83.01	ENF	Jack pine	3.3	49.3	8	Lab	Croft et al. (2014)
Bioindicators ENF	Canada	2001-02	47.11, -81.82	ENF	Black spruce, Jack pine	-	32.8	60	Lab	-
Sudbury Zhang	Canada	2003-04	47.16, -81.74	ENF	Black spruce	3.2	30.4	18	Lab	Zhang et al. (2008)
Le Casset	France	2008	44.98, 6.48	ENF	European larch	-	60.8	2	Lab	Homolová et al., (2013)
Bily Kriz	Czech Republic	2004, 2006	49.50, 18.54	ENF	Norway spruce	8.2	39.8	2	Lab	Homolová et al. (2013); Malenovský et al. (2013)
Col du Lautaret Terrasses	France	2008	45.04, 6.35	GRS	Mixed alpine sp.	-	38.5	3	Lab	Homolová et al. (2014)
Roche Noir	France	2008	45.06, 6.38	GRS	Mixed alpine sp.	-	39.3	5	Lab	Pottier et al. (2014)
GRS National Park	Canada	2012-14	49.16, -107.56	GRS	Mixed-grass prairie	1.0	35.4	20	Lab	Tong and He (2017)
Stratford Wheat	Canada	2013	43.45, -80.86	CRP	Wheat	2.2	38.7	5	Lab	Dong et al. (2017)

Stratford Corn	Canada	2013	43.46, -80.81	CRP	Maize	2.7	42.4	10	Lab	Dong et al. (2017)
Mosuowanzhen	China	2011	44.19, 85.49	CRP	Cotton	1.87	64.3	15	Lab	Yi et al., (2014)
Sele River Plain	Italy	2009	40.52, 15.00	CRP	Maize, fruit trees	2.3	42.6	29	Chl. Meter •	Vuolo et al. (2012)
Demmin	Germany	2006	53.99, 13.27	CRP	Maize, wheat, Sugarbeet	-	35.4	4	Chl meter	Hajnsek et al. (2006)
Trapani	Sicily	2010	37.64, 12.85	CRP	Olive trees	1.5	39.6	3	Chl. meter	-
Howard Springs	Australia	2009	-12.49, 131.15	SHR	Eucalyptus	1.3	57.7	1	Lab	Amiri (2013)
Daly Uncleared	Australia	2009	-14.16, 131.39	SHR	Eucalyptus	0.8	50.1	1	Lab	Amiri (2013)
Daly Regrowth	Australia	2009	-14.13, 131.38	SHR	Eucalyptus	0.9	45.1	1	Lab	Amiri (2013)
Adelaide River	Australia	2009	-13.08, 131.12	SHR	Eucalyptus	0.7	61.5	1	Lab	Amiri (2013)
Dry Creek	Australia	2009	-15.26, 132.37	SHR	Eucalyptus	0.8	45.9	1	Lab	Amiri (2013)
Sturt Plains Shrubland	Australia	2009	-17.15, 133.35	SHR	Eucalyptus	0.0	58.2	1	Lab	Amiri (2013)
Sturt Plains Woodland	Australia	2009	-17.18, 133.35	SHR	Acacia	0.7	45.1	1	Lab	Amiri (2013)

161 * where multiple sampling sites are present at a given field location, approximate central co-ordinates are given.

162 DBF = Deciduous broadleaf, ENF = Evergreen needleleaf, EBF = Evergreen broadleaf, GRS = Grassland, CRP =

163 Cropland, SHR = Shrubland. †Lab = laboratory extraction; Chl meter = handheld optical meter; Field spec = field

164 spectrometer. • indicates species specific chl. meter calibration equations.

165

166 Methods of chlorophyll measurement presented in Table 1 were either through laboratory analysis

167 techniques (Lab) or via optical methods (Croft and Chen 2018), i.e. a handheld chlorophyll meter (Chl.

168 meter) or a field spectroradiometer (Field spec.). Where species specific calibration equations were

169 used, further details of the calibration equations can be found in the corresponding referenced paper.

170 The field spectrometer-based retrieval (Arrellano et al., 2017) used the PROSPECT radiative transfer

171 model (Jacquemoud et al., 1990) to estimate chlorophyll content. It is recognised that the different

172 methods of chlorophyll determination may lead to a certain degree of variability between validation
173 measurements. The algorithm was validated according to the closest available ENVISAT MERIS satellite
174 date and location to the field sampling date. Due to the relative scarcity of available *in situ* Chl_{Leaf} data,
175 the existing validation data were maximised by using data collected in years shortly preceding or
176 following the MERIS operational time frame (i.e. earlier than 2002 or later than 2011; Table 1). In this
177 case, the correct day of year (DOY) from the closest calendar year within the 2002-2011 time period,
178 was used. The validation results (Section 3) are partitioned to indicate whether the results are from the
179 current year or closest matched year.

180

181 **2.2 Satellite data**

182 **2.2.1 MEduium Resolution Imaging Spectrometer (MERIS) satellite data**

183 ENVISAT MERIS satellite-derived reflectance data was selected as the input remote sensing product for
184 this study, due to the sampling of chlorophyll-sensitive red-edge bands, a short temporal revisit time
185 (every 2-3 days) for global application, medium spatial resolution (300 m) and its high radiometric
186 accuracy (Curran and Steele 2005). We used the full resolution (FR) surface reflectance (SR) product,
187 which is produced as a 7-day temporal synthesis from data collected at the original 2-3 day revisit
188 frequency. The global SR time series are produced by a series of pre-processing steps within the MERIS
189 pre-processing chain, including radiometric, geometric, bidirectional reflectance distribution function
190 (BRDF), pixel identification, and atmospheric correction with aerosol retrieval. The 7-day product
191 normalises reflectance to nadir view, and the solar zenith angle is that of 10h00 local time for the
192 median day of the compositing period. MERIS reflectance data is provided in 13 bands (spectral
193 resolution = ~10 nm) in the visible, red-edge and near infra-red bands, with the atmospheric bands
194 (bands 11 and 15) removed. The MERIS FR surface reflectance global time-series covers the 2003-2012
195 time period.

196

197 **2.2.2 Landcover map**

198 A global land cover map produced from 300 m spatial resolution MERIS data, as part of the European
199 Space Agency Climate Change Initiative (CCI-LC) project, from the 5-year period (2008-2012) was used to
200 define PFTs. The legend is based on the UN Land Cover Classification System (LCCS) with the view to be
201 compatible with the Global Landcover 2000 (GLC2000), GlobCover 2005 and 2009 products.

202

203 **2.2.3 Leaf area index (LAI)**

204 Copernicus Global Land Service GEOV1 LAI product derived from SPOT-VGT satellite (Baret et al., 2013)
205 was used in the algorithm, with a global LAI coverage from 1999 to the present at 10-day temporal
206 intervals and a spatial resolution of 1 km. The GEOV1 LAI product is derived from the CYCLOPES v3.1 and
207 MODIS c5 biophysical products, through a neural-network machine-learning algorithm (Baret et al.
208 2013). The selection of CYCLOPES v3.1 and MODIS c5 products takes advantage of previous
209 development efforts and capitalises on the performances of each product (Camacho et al. 2013). As part
210 of the processing chain, the LAI product is reprojected onto the Plate-Carrée 1/112° grid, temporally
211 smoothed, interpolated at the 10-day frequency, and re-scaled to fit the expected range of variation
212 (Verger et al. 2015). GEOV1 LAI products consider clumping as a weighted contribution of CYCLOPES and
213 MODIS products, where MODIS LAI accounts for clumping at plant and canopy scales (Knyazikhin et al.
214 1998), with the exception of needleleaf forests, for which shoot clumping is not accounted for. In the
215 CYCLOPES algorithm, landscape clumping is represented by considering fractions of mixed pixels (Baret
216 et al. 2013). Recent validation studies indicated that GEOV1 outperformed most existing products in
217 both accuracy and precision (Camacho et al. 2013). Validation results showed good spatial and temporal
218 consistency and accuracy, smooth and stable temporal profiles, good dynamic range with reliable
219 magnitudes for bare areas and dense forests (Camacho et al. 2013). The GEOV1 LAI product was

220 selected for this project, because it met the following criteria: a high temporal frequency (every 10
221 days), acceptable spatial resolution (1 km), a long archive of data covering the complete MERIS mission
222 (2002-2012), and strong product validation.

223

224 **2.3 Algorithm development**

225 To derive leaf chlorophyll content from remote sensing data at the global scale, a two-step modelling
226 approach was selected. The first step in the algorithm is the modelling of leaf-level reflectance spectra
227 from satellite-derived canopy reflectance data, using physically-based radiative transfer models to
228 account for the influence of canopy architecture, image acquisition conditions and background
229 vegetation on canopy reflectance. The second step was to retrieve leaf chlorophyll content from the
230 modelled leaf reflectance derived in Step 1, using a leaf optical model. A two-step inversion method is
231 favoured over a coupled one-step inversion because the output of each stage can be assessed
232 individually, and may be validated against measured leaf-level reflectance data at field sites (Croft et al.
233 2013, Zhang et al. 2008). This physically-based canopy inversion method has been previously
234 successfully demonstrated using different combinations of canopy and leaf models (Croft et al. 2013,
235 Moorthy, Miller, and Noland 2008, Zarco-Tejada et al. 2004, Kempeneers et al. 2008).

236

237 For the first step, two canopy reflectance models were selected, according to the structural
238 characteristics of the vegetation present. For spatially heterogeneous 'clumped' vegetation types (i.e.
239 deciduous and coniferous trees, shrubs), the 4-Scale geometrical-optical model (Chen and Leblanc
240 1997) was used (see Section 2.3.1). For homogenous canopies that can be treated as one-dimensional
241 (1D) turbid media, such as agricultural crops, we used the SAIL model (Verhoef 1984) (Section 2.3.2). To
242 invert the canopy radiative transfer models, individual LUTs for different PFTs were created, based on
243 variable and fixed input parameters (Table 2 and Table 3). The LUT approach was selected to optimise

244 computational resources and reduce problems associated with appearances of local minima, given
245 sufficient sampling of the variable space (Jacquemoud et al. 2009). Whilst these structural
246 parameterisations are important, their influence on canopy reflectance is mediated by LAI, which is the
247 dominant driver of modelled canopy reflectance (Zhang et al. 2008 and Section 2.4). The availability of
248 accurate and well validated global-scale LAI products facilitated the treatment of LAI as a variable
249 parameter, which was derived from the GEOV1 LAI product (Section 2.2.3) at 10 day intervals and 1 km
250 spatial resolution. In the second step, leaf chlorophyll content was retrieved using a leaf optical model
251 (PROSPECT; Jacquemoud and Baret 1990) (Section 2.3.3) from the modelled leaf reflectance that was
252 derived in step 1.

253

254 **2.3.1 Step 1: Leaf reflectance inversion - 4-Scale model**

255 The 4-Scale model (Chen and Leblanc 1997) was selected for the inversion of leaf level reflectance from
256 satellite-derived images in forested and spatially clumped ecosystems. 4-Scale considers the three-
257 dimensional (3D) spatial distribution of vegetation groups and vegetation crown geometry, in addition
258 to the structural effects of branches and scattering elements. In closed canopies, structural variables
259 such as crown shape and clumping of foliage elements are dominant, whereas in open canopies the
260 effects of background reflectance and shadows prevail. The 4-Scale model simulates the Bidirectional
261 Reflectance Distribution Function (BRDF) based on canopy architecture described at four scales: 1)
262 vegetation grouping, 2) crown geometry, 3) branches, and 4) foliage elements (Chen and Leblanc 2001).
263 A crown is, therefore, represented as a complex medium, where mutual scattering occurs between
264 shoots or leaves. Consequently, sunlit foliage can be viewed on the shaded side and shadowed foliage
265 on the sunlit side. Reflected radiance from different scene components is calculated by first separating
266 sunlit and shaded components through first-order scattering, and then adding multiple scattering from
267 subsequent interactions with vegetation elements or background material (Chen and Leblanc 2001). To

268 model canopy reflectance, the 4-Scale model was run in the forward mode, with variable (LAI, solar
269 zenith angle and view zenith angle) and fixed structural parameters and background reflectance spectra
270 (Table 2), set according to ground measurements and values reported in the literature (Chen and
271 Leblanc 1997, Leblanc et al. 1999). The element clumping index (ΩE) represents vegetation clumping at
272 scales larger than the shoot, and is associated with canopy architecture and structural variables, such as
273 crown size and vegetation density. ΩE is an important parameter for estimating radiation interception
274 and distribution in plant canopies (Chen, Menges, and Leblanc 2005). The non-random spatial
275 distribution of trees is simulated using the Neyman type A distribution to permit clumping and
276 patchiness within a forest stand (Chen and Leblanc 1997).

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292 *Table 2: Fixed and variable parameters used in the 4-Scale model for LUT generation.*
 293 *Broadleaf/needleleaf classes include respective deciduous and evergreen plant functional types.*

	Broadleaf trees	Needleleaf trees	Shrubland	References
LAI (m² m⁻²)	0.1-8	0.1-8	0.1-8	Baret et al., (2013)
Solar zenith angle (°)	20-70	20-70	20-70	MERIS metadata
View zenith angle (°)	0	0	0	MERIS metadata
Relative azimuth angle (°)	0	0	0	MERIS metadata
Stick height (m)	10	10	3	Los et al. (2012)
Crown height (m)	8	10	7	Los et al. (2012)
Crown shape	Spheroid	Cone & cylinder	Spheroid	Chen and Leblanc (1997)
Tree density (trees/ha)	1400	3000	1000	Chen et al. (1997) & Leblanc et al. (1999)
Crown radius (m)	1.25	1.0	1.25	Chopping et al. (2008), Evans et al. (2015) & Thorpe et al. (2010)
Neyman grouping	3	4	3	Leblanc et al. (1999)
Clumping index (ΩE)	0.90	0.80	0.80	Pisek et al. (2011)
Needle to shoot ratio (γE)	1	1.4	1	Pisek et al. (2011)
Foliage element width (m)	0.15	0.1	0.15	Chen and Cihlar (1995) & Leblanc et al. (1999)
Background composition	Green vegetation and soil	Green vegetation and soil	Dry grasses and soil	Observations/measurements

294

295 According to the specified input parameters, the 4-Scale model calculates canopy reflectance (ρ) as a
 296 linear summation of four components:

297

$$298 \quad \rho = \rho_{PT\lambda} F_{PT} + \rho_{ZT\lambda} F_{ZT} + \rho_{PG\lambda} F_{PG} + \rho_{ZG\lambda} F_{ZG} \quad [\text{Eq. 1}]$$

299

300 where the reflectance factors from each scene component are: sunlit vegetation ($\rho_{PT\lambda}$), shaded
 301 vegetation ($\rho_{ZT\lambda}$), sunlit ground ($\rho_{PG\lambda}$) and shaded ground ($\rho_{ZG\lambda}$), and F_{PT} , F_{PG} , F_{PG} and F_{ZT} represent the
 302 probability of viewing each component, respectively (Chen and Leblanc 2001). The 4-Scale model

303 includes a multiple scattering scheme, for 2nd order scattering and above, through computing the
 304 interactions among the four scene components based on view factors from one component to the other
 305 (Chen and LeBlanc, 2001). In this unique way, the geometrical effects propagate to all orders of multiple
 306 scattering. In order to derive leaf reflectance ($\rho_{L\lambda}$) from sunlit crown reflectance (ρ_{PT}), the enhancement
 307 of both sunlit and shaded reflectance due to multiple scattering is accounted for using a multiple
 308 scattering factor (Zhang et al. 2008, M factor; Croft et al. 2013, Simic, Chen, and Noland 2011). The M
 309 factor can be calculated using the output sunlit and shaded scene components from 4-Scale:

310

$$311 \quad M_{\lambda} = \frac{\rho_{\lambda} - \rho_{PG\lambda} F_{PG}}{\rho_{L\lambda} F_{PT}} \quad [\text{Eq. 2}]$$

312

313 Finally, the M factor (Eq. 2) is used to derive leaf reflectance from canopy-level satellite reflectance
 314 ($\rho_{satellite}$), by converting sunlit crown reflectance into sunlit leaf reflectance, and allowing the inclusion of
 315 the less variable shaded components (Eq. 3).

316

$$317 \quad \rho_{L\lambda} = \frac{\rho_{satellite\lambda} - \rho_{PG\lambda} F_{PG}}{M_{\lambda} F_{PT}} \quad [\text{Eq. 3}]$$

318

319

320 **2.3.2 Step 1: Leaf reflectance inversion - SAIL model**

321 The SAIL radiative transfer model (Verhoef 1984) was used to derive leaf reflectance from satellite-
 322 derived images in cropland and grassland ecosystems, where the distribution of foliage approaches
 323 randomness, and information on canopy structural variables such as crown shape and the clumping of
 324 foliage elements is not applicable. Turbid medium models such as SAIL (Verhoef, 1984) assume that the
 325 canopy is composed of homogeneous, horizontal layers of Lambertian scatterers randomly distributed in

326 space. The SAIL model is one of the most popular and well validated models used in agricultural systems
 327 (Jacquemoud et al. 2009, Darvishzadeh et al. 2008, Clevers and Kooistra 2012). SAIL is based on Suit's
 328 model (Suits 1971), which is founded on a set of four differential equations: (1) diffuse incoming flux, (2)
 329 diffuse outgoing flux, (3) direct solar flux, and (4) flux with radiance in the direction of remote sensing
 330 observation. Table 3 details the fixed and variable parameters used within the SAIL model, based on
 331 field observations and literature (Privette et al. 1996, Darvishzadeh et al. 2008) that were used to model
 332 canopy reflectance for the LUT-based inversion.

333

334 *Table 3: Fixed and variable parameters used in the SAIL model for LUT generation.*

	Cropland	Grassland	References
Leaf area index	0.1 - 8	0.1 - 8	Baret et al., (2013)
Solar zenith angle (°)	20-70	20-70	MERIS metadata
View zenith angle (°)	0	0	MERIS metadata
Relative azimuth angle (°)	0	0	MERIS metadata
Leaf inclination distribution function	Ellipsoidal	Erectophile	Clevers and Kooistra (2012) & Zou and Möttus (2015)
Hot spot parameter (m m⁻¹)	0.1	0.1	Darvishzadeh et al. (2008) & Vohland and Jarmer (2008)
Soil factor (dimensionless)	1	0.4	Observations/ measurements

335

336

337 The leaf inclination distribution function (LIDF) describes the frequency distribution of leaf inclination
 338 angles in different directions. For agricultural crops, LIDF is ellipsoidal (Campbell, 1986), with the mean
 339 leaf inclination angle set at 40°, in line with values measured within previous studies (Zou and Möttus
 340 2015, Liu, Pattey, and Jégo 2012). For grasslands the LIDF is erectophile (LiDFa = -1, LiDFb = 0) (Clevers
 341 and Kooistra 2012, Sandmeier et al. 1999, He, Guo, and Wilmshurst 2007, Verma et al. 2017, Migliavacca
 342 et al. 2017). The hot spot (Hs) parameter is defined as the ratio between the average leaf size and the
 343 height of the canopy (Darvishzadeh et al. 2008, Verhoef 1985). The soil factor (Sf) accounts for

344 variations in the brightness of the soil background, where a value of 0 denotes a wet soil and 1 a dry soil.
345 A LUT was generated from the fixed and variable parameters shown in Table 3, which contains the ratio
346 of the modelled canopy reflectance from the SAIL model to the SAIL input leaf reflectance as a scaling
347 factor ($S Factor_{\lambda}$). Modelled leaf reflectance ($\rho_{L\lambda}$) is retrieved from satellite canopy reflectance ($\rho_{Satellite\lambda}$)
348 according to Equation 4.

349

$$350 \quad \rho_{L\lambda} = \rho_{Satellite\lambda} * S Factor_{\lambda} \quad [Eq. 4]$$

351

352 **2.3.3 Step 2: Deriving leaf chlorophyll content – PROSPECT model**

353 The leaf radiative transfer model PROSPECT (Jacquemoud and Baret 1990, Feret et al. 2008) is used to
354 derive Chl_{leaf} from the modelled leaf reflectance spectra generated in step one (Section 2.3.1 and 2.3.2).
355 In PROSPECT-5, leaf reflectance and transmittance (400-2500 nm) are defined as a function of six
356 parameters: structure parameter (N), chlorophyll (a+b) concentration (C_{ab}), carotenoid content (Car),
357 brown pigment (C_b), dry matter (C_m) and equivalent water thickness (C_w). Absorption is calculated as the
358 linear summation of the specific absorption coefficients of the biochemical constituents and their
359 respective concentrations. PROSPECT has had widespread validation across a large number of
360 vegetation species and plant functional types (Croft, Chen, Zhang, et al. 2015, Demarez and Gastellu-
361 Etchegorry 2000, Darvishzadeh et al. 2008, Malenovský et al. 2006). Whilst previous studies inverted
362 PROSPECT by iteratively minimising a merit function (Croft et al., 2013), the large volume of data used in
363 this study precluded this option. In order to optimise computational resources, using a vegetation index
364 (VI) based inversion approach. PROSPECT was run in the forward mode to derive chlorophyll-sensitive
365 spectral vegetation indices, according to incrementing chlorophyll content values. The leaf biophysical
366 parameters were varied for each PFT (Table 4), based on measured values and within reported ranges in
367 the literature.

368

369

370 *Table 4: Leaf parameters used in the PROSPECT model, along with the vegetation index (VI) derived from*
 371 *forward-modelled reflectance spectra for each plant functional type. C_{ab} = chlorophyll content; N =*
 372 *structural parameter; Car = carotenoid content; C_b = brown matter; C_m = dry matter content; C_w =*
 373 *equivalent water thickness.*

374

	C_{ab} ($\mu\text{g cm}^{-2}$)	N (unitless)	Car ($\mu\text{g cm}^{-2}$)	C_b	C_m (g cm^{-2})	C_w (g cm^{-2})	Vegetation index	Reference
DBF	0-100	1.2	Chl/7	0	0.005	0.01	MERIS _{SR}	Féret et al. (2011)
EBF	0-100	1.8	Chl/7	0	0.005	0.01	MERIS _{SR}	Arellano et al. (2017)
DNF	0-100	2.8	Chl/7	0	0.05	0.01	MERIS _{ND}	De Santis et al. (2009) & Kötz et al. (2004)
ENF	0-100	2.8	Chl/7	0	0.05	0.01	MERIS _{ND}	De Santis et al. (2009) & Kötz et al. (2004)
SHR	0-100	1.8	Chl/7	0	0.005	0.01	MERIS _{ND}	Enrique et al. (2016) & Sow et al. (2013)
CRP	0-100	1.4	Chl/7	0	0.015	0.01	MERIS _{ND}	Jacquemoud et al. (2000)
GRS	0-100	1.2	Chl/7	0	0.005	0.01	MERIS _{ND}	Darvishzadeh et al. (2008)

375

376 The VIs used in the VI-based PROSPECT inversion are focussed on the red-edge reflectance bands that
 377 are resistant to saturation at high chlorophyll values. Croft et al., (2014a) found strong relationships
 378 between normalised difference (ND) and simple ratio (SR) vegetation indices with chlorophyll content
 379 (Zarco-Tejada et al. 2001, Mutanga and Skidmore 2004), which were modified to MERIS red-edge
 380 spectral bands as follows:

381

$$382 \quad MERIS_{ND} = \frac{R_{754} - R_{709}}{R_{754} + R_{709}} \quad [Eq. 5]$$

383

384

$$MERIS_{SR} = \frac{R_{754}}{R_{709}} \quad [\text{Eq. 6}]$$

385

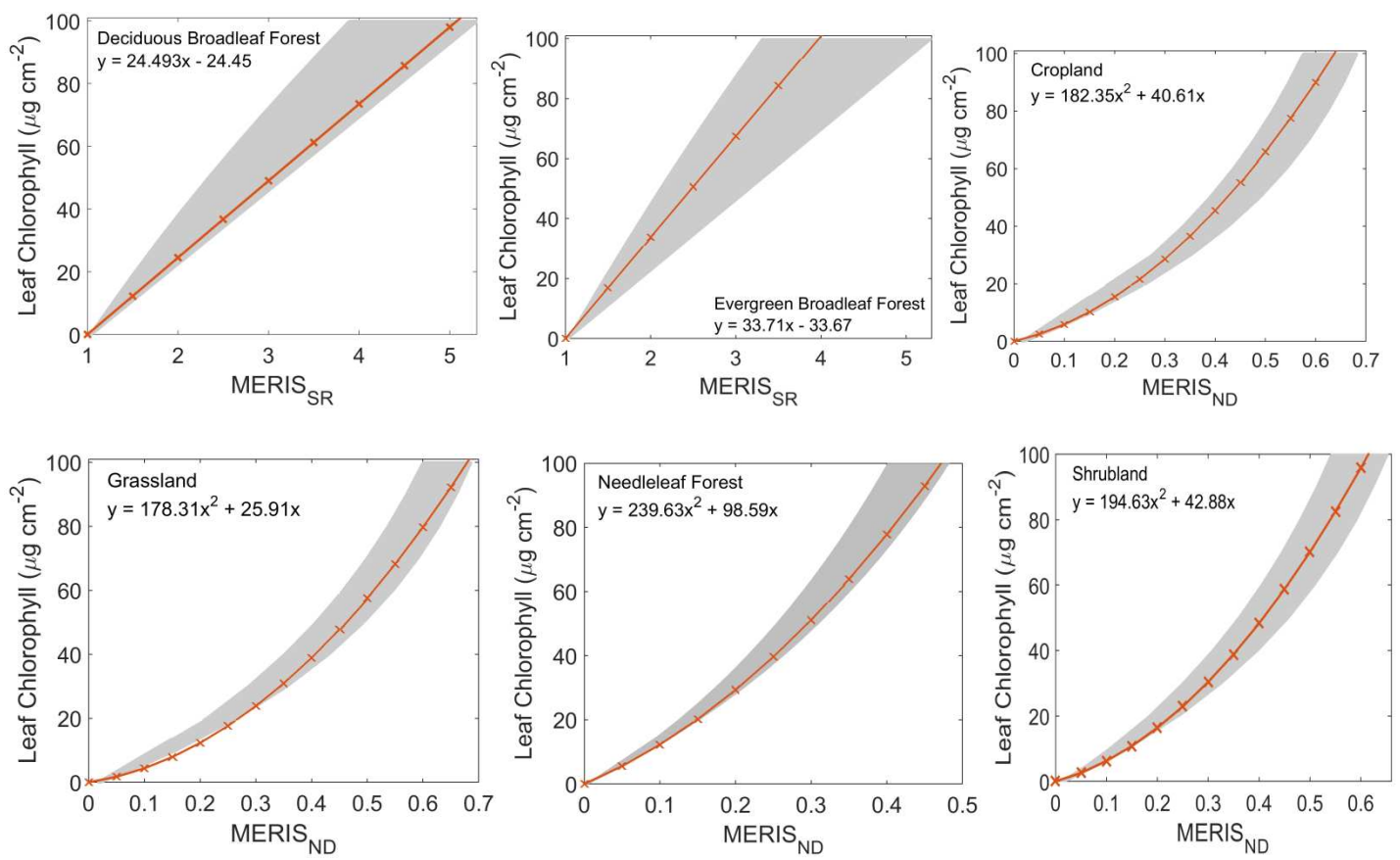
386 where R is reflectance at given wavelength λ in nm (R_λ). Regression equations were then generated

387 between incrementing chlorophyll content and the modelled VI from PROSPECT (Figure 2), which are

388 then applied to the modelled leaf level reflectance derived from canopy model inversion in Step 1

389 (Sections 2.2.1 and 2.2.2) to derive the final Chl_{Leaf} values.

390



391

392

393 *Figure 2: Regression equations for the PROSPECT-based leaf chlorophyll inversion. The shaded grey area*

394 *represents the potential uncertainty of the forward modelled vegetation index values, according to a*

395 *range of measured PROSPECT parameter values from published data. For all plant functional types apart*

396 *from needleleaf forests, the range in input values is two standard deviations either side of the mean from*

397 *the LOPEX dataset. For needleleaf forests uncertainty values from Kötz et al. (2004) were used.*

398

399 The shaded grey area in Figure 2 indicates the degree of uncertainty arising from the VI-inversion step,
400 according to the fixed biophysical parameters that are used to parameterise the PROSPECT model.
401 Measured N , C_m and C_w values from the LOPEX published dataset of leaf reflectance and biophysical
402 parameters (Hosgood et al. 1995) were used for DBF, SHR, GRS, CRP species. For the needleleaf and EBF
403 species, for which large measured datasets are not so widely available, we used parameter uncertainty
404 values from Kötz et al. (2004) and standard deviations from Ferreira et al. (2013) for needleleaf and EBF
405 species, respectively. For DBF, SHR, GRS, CRP and EBF species, the PROSPECT model was run in the
406 forward mode for input parameter values that represented two standard deviations of the mean N , C_m
407 and C_w values within the published datasets (represented by grey shading), for each PFT individually. The
408 orange line represents the relationship between the VI and Chl_{Leaf} using the structural parameters in
409 Table 4. The regression equation shown is the empirical model that is used to convert modelled leaf
410 level reflectance derived in Step 1 (Sections 2.3.1 and 2.3.2) to Chl_{Leaf} .

411

412 The VI-based PROSPECT inversion method was assessed against the conventional PROSPECT inversion
413 method of iteratively minimising a merit function (Feret et al., 2008), for the ground validation sites
414 detailed in Table 1. The relationship between the modelled Chl_{Leaf} from the VI-inversion and modelled
415 Chl_{Leaf} from the merit function inversion was $R^2 = 0.74$ (percentage bias = 20% and RMSE = $11.01 \mu\text{g cm}^{-2}$).
416 Moreover, the VI-inversion increased the relationship between modelled and measured Chl_{Leaf} by 20%
417 ($p < 0.001$), when compared to the merit function Chl_{Leaf} inversion. Using VIs centred on the chlorophyll-
418 sensitive red-edge rather than inverting the model across the full leaf spectra reduces the confounding
419 influence of other biophysical variables, such as leaf structure and carotenoid content. Several studies
420 have found a high redundancy of wavelength channels in vegetation studies (Jacquemoud et al., 1995;
421 Simic and Chen, 2008, Croft et al. 2015). Thenkabail et al. (2004) reported that the data volume can be

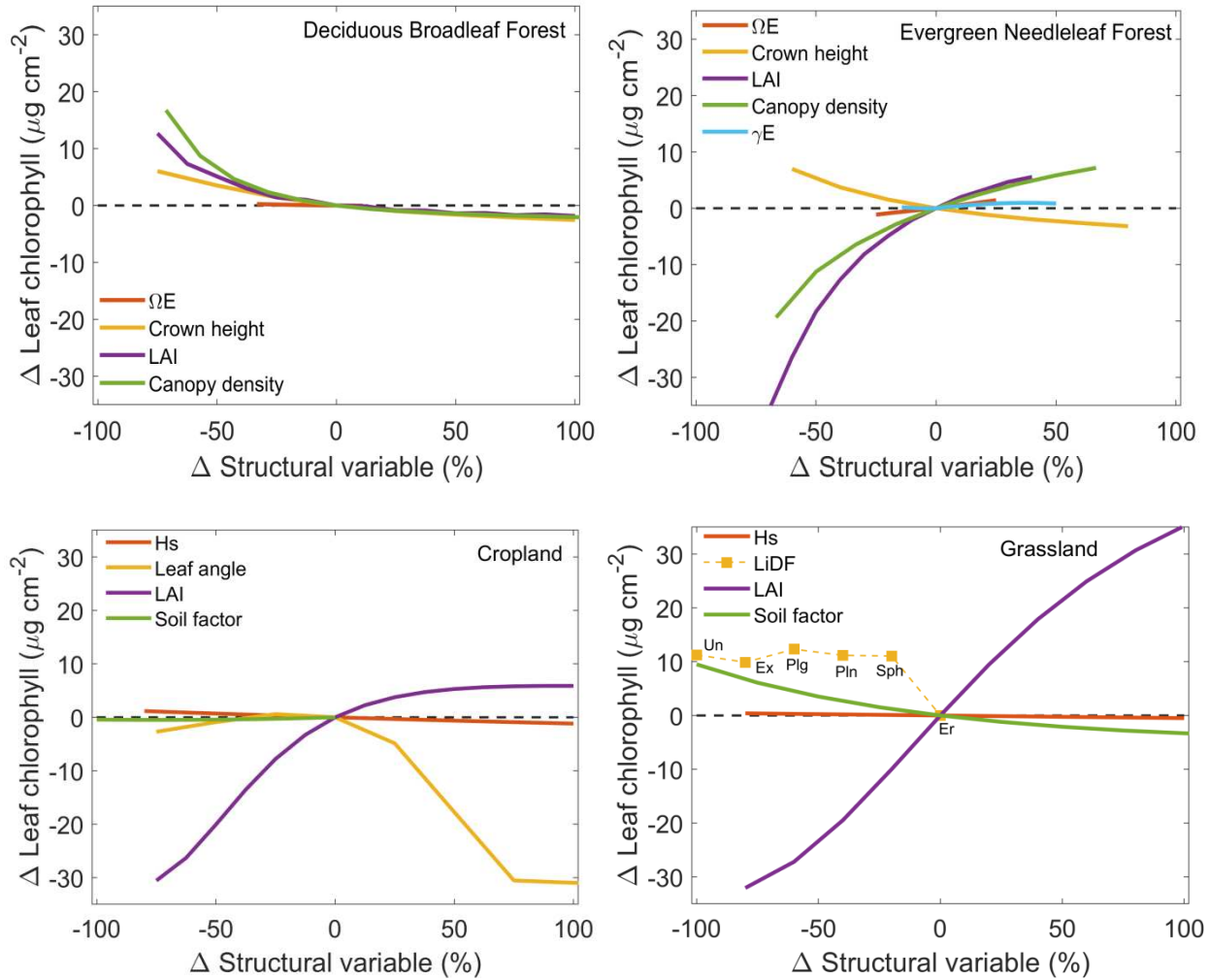
422 reduced by 97% when hyperspectral wavebands are reduced to the first five principal components, and
423 still explain close to 95% of data variability. A reduction in spectral data may therefore permit a better
424 model fit over the more sensitive and relevant wavelength ranges.

425

426 **2.4 Canopy model inversion sensitivity analysis**

427 A sensitivity analysis was undertaken to assess the impact of canopy structural parameters values
428 differing from the fixed value used in the inversion algorithm on the modelled Chl_{Leaf} value. Four of the
429 major PFTs with different structural parameterisations and that employed radiative transfer models
430 (DBF, ENF, CRP, GRS) were tested. All values were held constant according to the values used in the
431 chlorophyll inversion algorithm (Tables 2 and 3), with one parameter incrementally perturbed. The value
432 zero represents the condition under which the structural parameters are consistent with the values used
433 in the main chlorophyll inversion algorithm, as outlined in Tables 1 and 2 (i.e. the reference case). In the
434 case of LAI, which is a variable parameter in the algorithm we used a representative value (LAI = 2.5
435 (GRS), 4.0 (CRP), 5.0 (ENF), 4.0 (DBF)) as found in a global analysis by Asner et al., (2003), for each PFT to
436 avoid bias from an extreme case. Solar and view zenith angles were held constant at 40° and 0° ,
437 respectively. To achieve comparability between structural parameters, results are scaled to a
438 percentage difference from the reference case. The chlorophyll values are shown as an absolute
439 difference from the reference case, in order to provide an estimate of the impact that deviations in the
440 actual structural parameter values from the one used in the modelling algorithm has on modelled leaf
441 chlorophyll content values. The structural parameters vary slightly between the PFTs and two models
442 (PROSAIL and 4-Scale), according to the structural parameters present in the model.

443



444

445

446

447 *Figure 3: Sensitivity of modelled chlorophyll values to variations in structural parameters within the 4-*

448 *Scale and SAIL models, for a) deciduous broadleaf forests, b) evergreen needleleaf forests, c) Croplands,*

449 *d) Grasslands. Zero percentage change (x axis) represents the reference case, where all structural*

450 *parameters are those used in the algorithm inversion in Table 2 and 3. For grassland results, the different*

451 *LiDFs are imposed on the x axis scale with LiDF names indicated, where Un = Uniform, Ex = Extremophile,*

452 *Plg = Plagiophile, Pln = Planophile, Sph = Spherical, Er= Erectophile.*

453

454 Figure 3 reveals different sensitivities of modelled Chl_{Leaf} values in response to changing a specific
455 structural parameter. In all cases variations in LAI presented a dominant effect on the modelled
456 chlorophyll content. For crops a 25% difference in LAI from the reference case (i.e. LAI = 3.0 and 5.0),
457 resulted in a change in modelled chlorophyll by $-7.8 \mu\text{g cm}^{-2}$ and $+3.76 \mu\text{g cm}^{-2}$. The asymmetric
458 response of the modelled Chl_{Leaf} values to LAI deviations represents the non-linear response of canopy
459 reflectance to increasing LAI, which saturates at higher LAI values. In the GRS results, this saturation
460 does not occur because of the erectophile leaf inclination angle distribution, and because the LAI
461 reference case is lower (LAI = 2.5), resulting in smaller maximum values (GRS $+\Delta 100\%$ LAI = 5.0, CRP
462 $+\Delta 100\%$ LAI = 8.0). For ENF trees, using the 4-Scale model, variations in canopy density also presented a
463 high degree of sensitivity in modelled Chl_{Leaf} values at extreme low canopy density values. A 33%
464 difference in density from the reference case (i.e. 2000 stems/ha and 4000 stems/ha), resulted in a
465 change in modelled chlorophyll by $-6.4 \mu\text{g cm}^{-2}$ and $+4.4 \mu\text{g cm}^{-2}$, while a 66% difference resulted in a -
466 $19.4 \mu\text{g cm}^{-2}$ and $+7.17 \mu\text{g cm}^{-2}$ change. For ENF and DBF, deviations in other structural parameters (Ω ,
467 crown height and γ E for ENF species) presented a smaller impact on the modelled Chl_{Leaf} . A 20%
468 difference in crown height, for example, only affected Chl_{Leaf} values by $+1.5 \mu\text{g cm}^{-2}$ and $-1.18 \mu\text{g cm}^{-2}$.
469 For the SAIL model parameters in CRP plants, in addition to LAI, the leaf angle distribution (fixed value =
470 40°) also strongly affected modelled Chl_{Leaf} at larger leaf inclination angles, approaching more
471 erectophile leaf inclination angles ($>60^\circ$). From leaf angle distribution = 10° to 50° the imposed
472 variations in modelled chlorophyll were $-2.7 \mu\text{g cm}^{-2}$ to $-4.9 \mu\text{g cm}^{-2}$. This result was also consistent for
473 the GRS results, where the different LIDFs are noted in Figure 3d. The LIDFs exhibited little change until
474 the leaf inclination angle approached those associated with erectophile canopies, with the plagiophile
475 LIDFs exhibiting the largest difference in chlorophyll values from the modelled erectophile reference
476 case ($+12.3 \mu\text{g cm}^{-2}$). Changes in the soil factor and the hotspot parameter induced very little differences
477 in modelled Chl_{Leaf} . For example, in CRP plants a $+100\%$ change in the hotspot value ($H_s = 0.2$) only

478 resulted in a $-1.2 \mu\text{g cm}^{-2}$ in Chl_{Leaf} , confirming the findings of Vohland and Jarmer (2008). Deviations in
479 Chl_{Leaf} values that are close to zero, with increasing percentage change in the structural parameter
480 indicates a greater tolerance of the model algorithm to differences in the actual structural parameter
481 value at a given site, to the value that is fixed in the model. The inclusion of LAI as a variable parameter
482 mitigates against much of the uncertainty generated in incorrect structural parameters, provided the
483 satellite-derived LAI value is reasonable. Underestimations of the actual ground LAI value by the satellite
484 product may therefore cause substantial uncertainty in modelled Chl_{Leaf} results.

485

486 **2.5 Data smoothing and gap-filling**

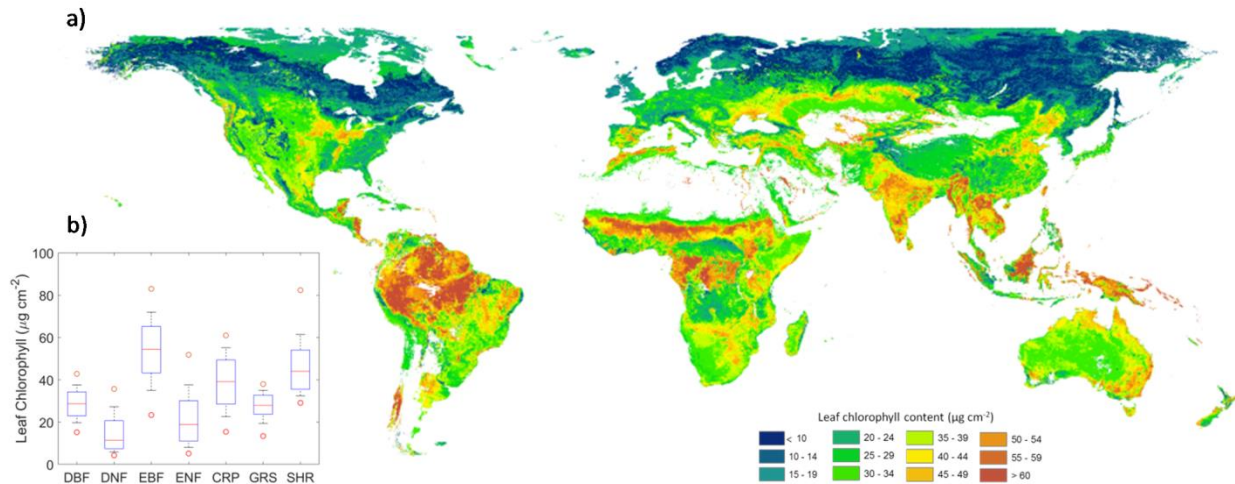
487 The Locally Adjusted by Cubic-Spline Capping (LACC) method (Chen, Deng, and Chen 2006) was
488 developed to produce continuous seasonal trajectories of satellite-derived surface parameters
489 contaminated by clouds and other atmospheric effects. The LACC method previously has been used
490 successfully to smooth time series of clumping index and NDVI (He et al. 2016). Contaminated points can
491 be identified in the time series and replaced with expected values through temporal interpolation
492 between adjacent valid points. For identifying atmospherically contaminated points in a time series, a
493 cubic-spline curve fitting technique with a curvature constraint was developed. Without this curvature
494 constraint, the fitted polynomial curve would also fit the contaminated points, defeating the study
495 purpose. In the LACC algorithm, a maximum global curvature of 0.5 is prescribed in the first step to
496 identify contaminated points. In the second step, the curvature is adjusted locally (i.e. at different dates
497 of the year) according to the local shape of the curve fitted with the global curvature constraint. This
498 adjustment is to enhance the large curvatures generally found at the beginning and the end of a growing
499 season, or double cropping seasons, and to reduce the curvature during a non-growing season and/or
500 around the peak of growing seasons. In this way, smooth trajectories that also follow the rapid growing
501 season variation patterns can be produced. In this study, the LACC algorithm is used to identify

502 contaminated data points and to gap-fill missing data points. However, in tropical areas there are often
503 fewer than the required six data points needed to run the LACC algorithm. In this case, in order to derive
504 temporally continuous data across the year, we removed the highest and lowest data point and linearly
505 interpolated across the year. In northern latitude evergreen needleleaf forests where snow cover
506 prevented acquisition of surface reflectance data from the vegetation canopy, we took the mean
507 seasonal value and extended it across the year. Whilst such temporal gap-filling is likely to introduce
508 some uncertainty in the modelled chlorophyll results, the temporal variations in chlorophyll content for
509 these PFTs is small relative to deciduous forests or croplands. The annual global distribution of leaf
510 chlorophyll content maps that are presented in this paper are for the last full year that MERIS was
511 operational (2011). Missing data due to cloud or MERIS image acquisition limitations (Tum et al. 2016)
512 are gap-filled with corresponding 2010 data from the same DOY within the LACC smoothing algorithm
513 for spatial completeness and visual assessment. The original 2003-2011 global maps are also available
514 for use within the academic community.

515

516 **3.0 Results and Discussion**

517 The annual global distribution of leaf chlorophyll content is presented in Figure 4a, with the geographic
518 trends varying both within and between biomes. Figure 4b depicts the annual range in chlorophyll
519 values for the major PFTs, as an integration of the temporal and spatial variability in chlorophyll content
520 over a year and across the globe.



521

522

523 *Figure 4: a) The spatial distribution of median annual global distribution of leaf chlorophyll content (at 9*
 524 *km resolution), and b) the median and interquartile range of leaf chlorophyll content, along with*
 525 *extreme values given for 9% and 91% (whiskers) and 2% and 98% (circles) of the data range.*

526

527 On an annual basis, the highest median Chl_{Leaf} values are present in evergreen broadleaf forests ($54.4 \mu\text{g cm}^{-2}$
 528 cm^{-2} ; Figure 4b), which is in part due to the high chlorophyll content of the vegetation present, and also
 529 due to the lack of seasonal leaf loss. The lower median annual chlorophyll values (Figure 4b) are
 530 consequently found in deciduous biomes (broadleaf and needleleaf forests, $28.8 \mu\text{g cm}^{-2}$ and $11.4 \mu\text{g cm}^{-2}$
 531 cm^2 , respectively). Ground measurements have previously demonstrated that nitrogen-poor needleleaf
 532 species typically exhibit lower chlorophyll contents by unit area than broadleaf species (Middleton et al.
 533 1997), which is also evident in the median annual results shown in Figure 4 (ENF, $18.9 \mu\text{g cm}^{-2}$).
 534 Environmental constraints may also impact Chl_{Leaf} values, through temperature extremes or water
 535 availability in evergreen PFTs. Temperate grasslands, for example, are often located in regions that
 536 experience extreme annual temperature variability (McGinn 2010), which affects chlorophyll
 537 biosynthesis (Ashraf and Harris 2013). From the global map of median annual Chl_{Leaf} , the footprints of

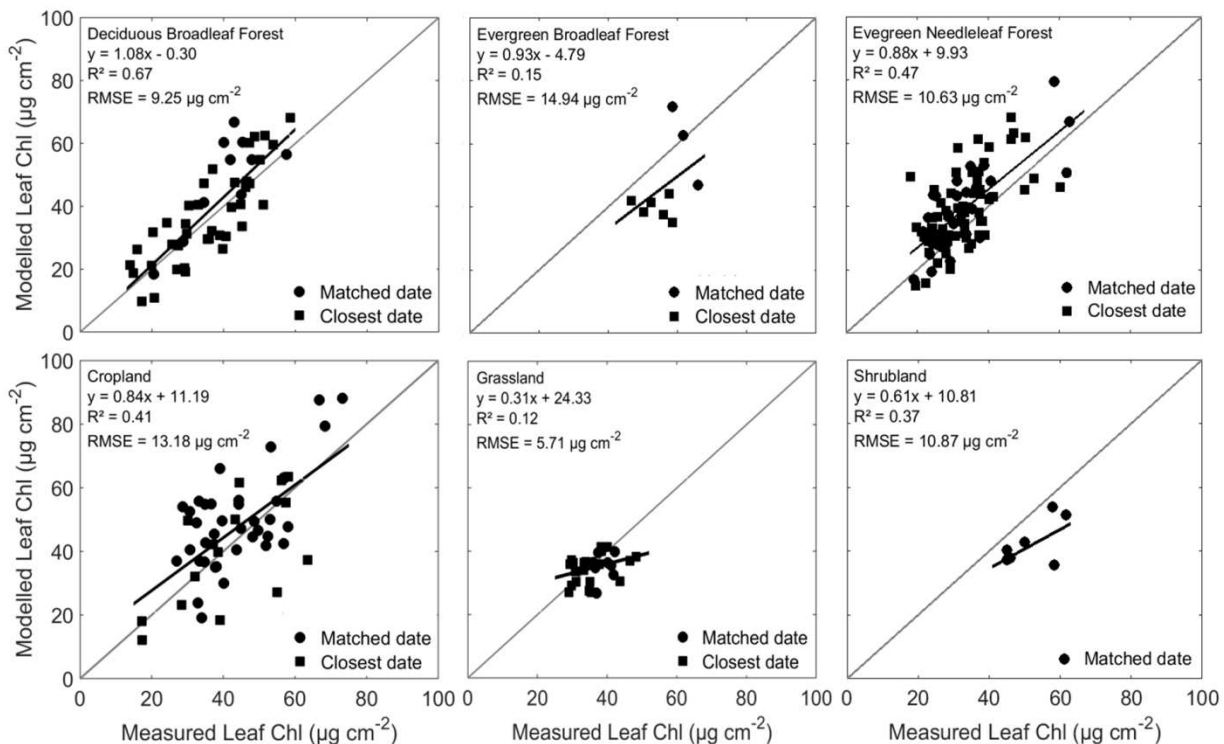
538 some agricultural regions are also clearly detectable, for example the ‘Corn Belt’ in Midwestern USA,
 539 throughout India, the Loess belt in central Europe and in northeastern China, due to their higher annual
 540 peak values, which may be related to fertilization application.

541

542 3.1 Validation of modelled leaf chlorophyll estimates for individual PFTs

543 The leaf chlorophyll algorithm was validated using 248 samples, across a number of different sampling
 544 dates and years. The modelled chlorophyll results are plotted against ground-measured values for all
 545 available PFTs (Figure 5).

546



547
 548

549 *Figure 5: Validation of modelled leaf chlorophyll with measured ground data for individual plant*
 550 *functional types. Data are distinguished between temporally matched sampling and satellite overpass*
 551 *dates, and for ground data collected outside the MERIS lifespan (2002-2012), the closest satellite date*
 552 *was selected.*

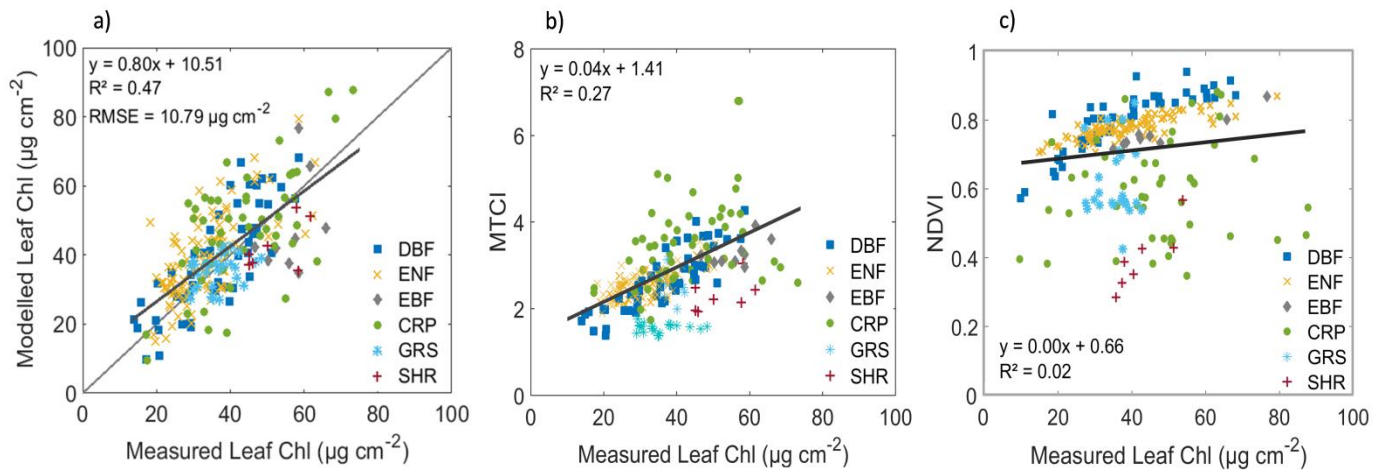
553
554 The strongest performances are seen for DBF ($R^2 = 0.67$; RMSE = $9.25 \mu\text{g cm}^{-2}$, relative RMSE = 25.4%)
555 followed by ENF ($R^2 = 0.47$; RMSE = $10.63 \mu\text{g cm}^{-2}$, relative RMSE = 32.62%) and CRP ($R^2 = 0.41$; RMSE =
556 $13.18 \mu\text{g cm}^{-2}$, relative RMSE = 30.8%). Nonetheless the other three PFTS also showed similar levels of
557 uncertainty, where relative RMSE values were 26.5%, 15.7% and 20.8% for EBF, GRS and SHR,
558 respectively. The statistics reported in Figure 5 are for combined temporally matched sampling and
559 satellite overpass dates, and for ground data collected outside the MERIS lifespan (2002-2012), where
560 the closest satellite date was selected (See Section 2.1). For comparison, the regression results using
561 only the matched dates inside the MERIS operational window are: DBF: $R^2 = 0.63$, $p < 0.001$; ENF: $R^2 =$
562 0.63 , $p < 0.001$; EBF: insufficient data; and CRP: $R^2 = 0.43$, $p < 0.001$. The shrub data used in Figure 5 only
563 contained matched dates. The regression results for all dates combined and only matched dates are
564 comparable, with the largest difference arising for EBF, which has fewer data and a smaller dynamic
565 range.

566

567 **3.2 Overall validation and comparison with vegetation indices**

568 The transferability of an algorithm across spatial and temporal scales is essential for modelling Chl_{Leaf} , or
569 any ecological variable, at the global scale. In comparison to empirical approaches (Gitelson, Gritz, and
570 Merzlyak 2003, Roberts, Roth, and Perroy 2016, Peng et al. 2017), the nature of physically-based
571 retrieval methods can account for relationships between canopy reflectance and Chl_{Leaf} across different
572 species and measurement acquisition conditions. Regressions between measured and modelled Chl_{Leaf}
573 for all PFTs combined are shown in Figure 6, alongside results for two popular vegetation indices with
574 the same measured Chl_{Leaf} data (the chlorophyll-sensitive MERIS Terrestrial Chlorophyll Index (MTCI =
575 $754 \text{ nm} - 709 \text{ nm} / 709 \text{ nm} - 681 \text{ nm}$) and the Normalised Difference Vegetation Index (NDVI = $865 \text{ nm} - 664$
576 $\text{ nm} / 865 \text{ nm} + 664 \text{ nm}$)).

577



578

579

580 *Figure 6: Relationships between measured leaf chlorophyll content and a) physically-based modelled*
581 *chlorophyll content; b) MERIS Terrestrial Chlorophyll Index (MTCl) values and c) Normalised Difference*
582 *Vegetation Index (NDVI) values.*

583

584 Figure 6a demonstrates the suitability of the Chl_{Leaf} algorithm for application across multiple PFTs, with a
585 strong, linear relationship for all PFTs combined ($R^2 = 0.47$; $p < 0.001$). By contrast, the MTCl regression is
586 weaker ($R^2 = 0.27$; $p < 0.001$), with some separation in MTCl values according to PFT. Cropland and DBF,
587 for example, exhibit higher values than grassland and shrubland, for the same Chl_{Leaf} . This stratification is
588 likely to be due to the strong influence of LAI on MTCl values. Nonetheless, MTCl does present a relative
589 improvement over NDVI results ($R^2 = 0.02$; n/s), due to the inclusion of chlorophyll-sensitive red-edge
590 bands within the VI. The biomass-sensitive NDVI values also exhibit a separation according to PFT, due
591 to differences in canopy structure and background, with cropland and grasslands also presenting
592 particular variability in values within the same PFT. This result points to the importance of accounting for
593 variations in canopy structure when deriving leaf-level chlorophyll results, and has implications for

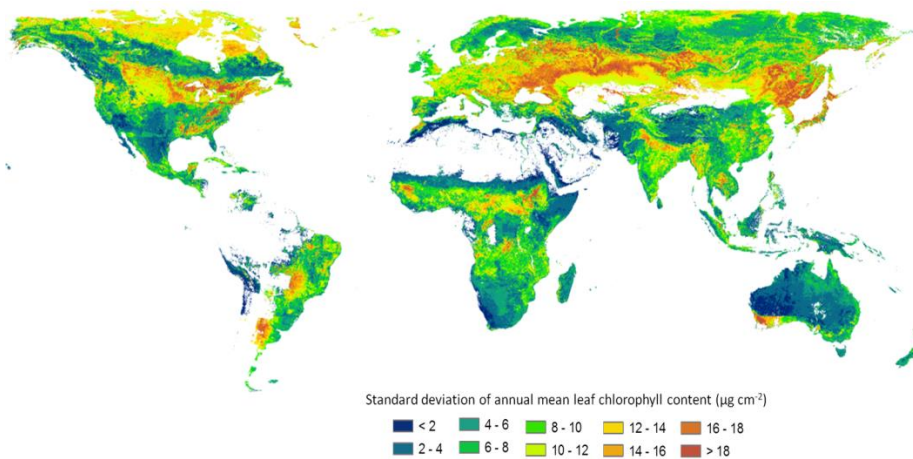
594 applying VIs over large spatial extents to infer information on physiological processes or plant
 595 productivity.

596

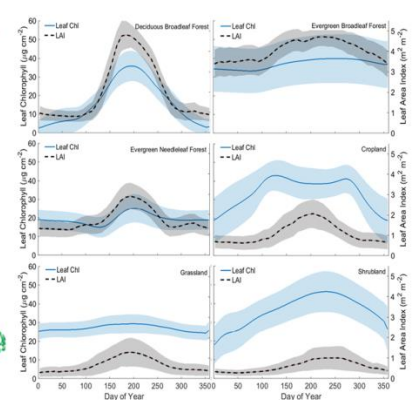
597 3.4 Leaf chlorophyll phenology and temporal trends

598 The annual variability in leaf chlorophyll content was evaluated through the standard deviation of mean
 599 annual Chl_{Leaf} , (Figure 7a). The seasonal phenology of mean Chl_{Leaf} and mean LAI across the northern
 600 hemisphere are shown for different biomes in Figure 7b.

601 a)



b)



602

603 *Figure 7: Global 2011 maps of a) seasonal variability in leaf chlorophyll (one standard deviation of one*
 604 *year time series), and b) mean seasonal phenologies of modelled chlorophyll content and LAI across the*
 605 *northern hemisphere for six different plant functional types during 2011. Shaded area indicates one*
 606 *standard deviation.*

607

608 Global regions that contain large areas of deciduous forests and croplands, which are dominated by a
 609 strong seasonal phenology, present the largest temporal variation and dynamic range in Chl_{Leaf} values
 610 (Figure 7a). In Figure 7b, DBF displays an expected phenology associated with budburst and chlorophyll
 611 biosynthesis in spring (from circa DOY 120) and chlorophyll breakdown during leaf senescence.

612 Temporal variations in cropland Chl_{Leaf} show a high variability globally, which is associated with different
613 species of crops planted, planting regimes (i.e. single or double), fertiliser application and the level of
614 irrigation. ENF temporal Chl_{Leaf} trajectories are relatively consistent across the year, although increasing
615 Chl_{Leaf} within new needles in spring is detectible, along with some chlorophyll breakdown within winter
616 months. Missing local data (predominately in the Amazon) within Figure 7a is due to too few original
617 data points to generate standard deviation values. Importantly, Figure 7b demonstrates the temporal
618 divergence of LAI and Chl_{Leaf} across the growing season, and a clear decoupling of vegetation structure
619 and physiological function (Croft, Chen, and Zhang 2014b, Walther et al. 2016).

620

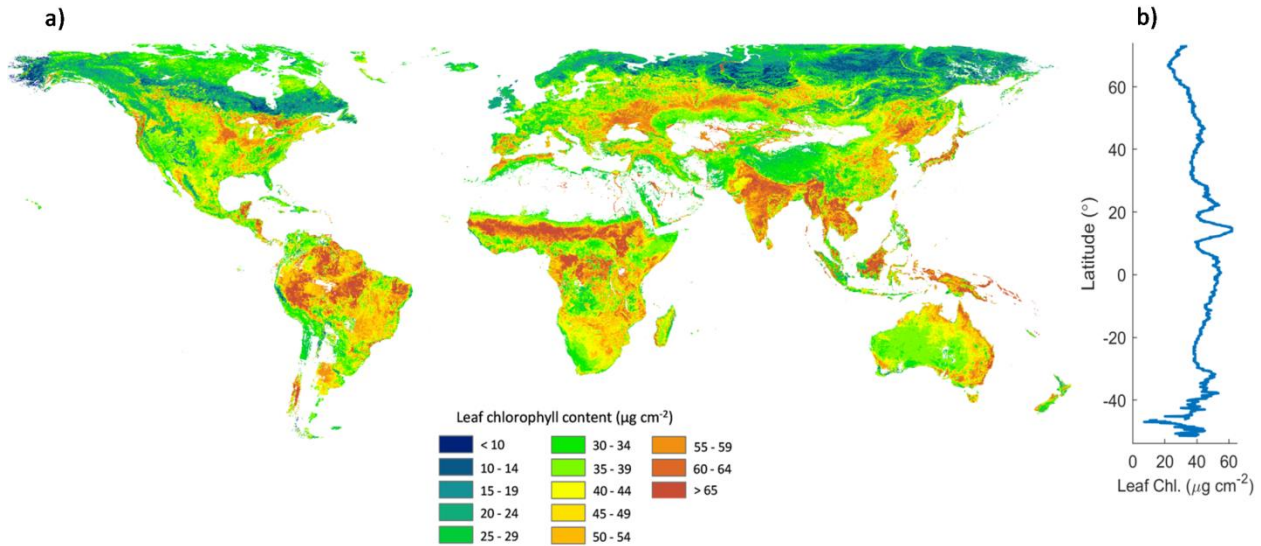
621 **3.3 Spatial and biome-dependent trends**

622 Annual maximum chlorophyll maps offer an opportunity to examine spatial differences in Chl_{Leaf} without
623 the integration of a temporal bias from seasonal change. The geographic variability in the abundance of
624 leaf chlorophyll depends in part on local environmental drivers, and partly on plant resource allocation.

625

626

627



628

629

630 *Figure 8: a) Maximum annual leaf chlorophyll content (at 9 km resolution) and b) mean values of the*
 631 *maximum along latitudinal bands.*

632

633 Higher annual maximum chlorophyll values ($> 55 \mu\text{g cm}^{-2}$) are associated with tropical forests and
 634 croplands (Figure 8a). Mid-range values ($25\text{-}55 \mu\text{g cm}^{-2}$) are typical of grasslands, temperate broadleaf
 635 forests and arctic tundra vegetation, and lower values ($<25 \mu\text{g cm}^{-2}$) are typical of boreal forests.
 636 Considerable intra-biome chlorophyll variability also exists, for example within croplands higher annual
 637 values are present in India, Eastern China, Western Europe and the great Lakes in USA/Canada, in
 638 contrast to lower values found in Central Europe and the North American mid-west. Reasons for
 639 chlorophyll variability over space are in part due to the abundance of elements needed for chlorophyll
 640 synthesis (nitrogen (N), phosphorous (P), Magnesium (Mg)) (Li et al. 2018), which may vary due to the
 641 nutrient availability status of the soil, the addition of fertilisers in managed landscapes and as a result of
 642 atmospheric N deposition. At the biome-scale, the temperature-dependency of chlorophyll synthesis
 643 reactions, where the optimum temperature of [3,8-DV]-Pchlide α 8-vinyl reductase (DVR) enzyme
 644 activity in chlorophyll synthesis is 30°C (Nagata, Tanaka, and Tanaka 2007), may exert a biotic control on

645 the accumulation of chlorophyll in temperature-limiting ecosystems, such as the arctic. Additionally,
646 plant water availability and drought stress affect the synthesis of chlorophyll, and may also prompt its
647 accelerated breakdown (Ashraf and Harris 2013). As plants allocate resources to optimise physiological
648 processes, it may be expected that solar irradiance conditions, and the leaf economics spectrum (Wright
649 et al. 2004) affect the partitioning of nitrogen, both between and within structural and photosynthetic
650 leaf fractions (Croft et al. 2017, Niinemets and Tenhunen 1997).

651

652 **4.0 Sources of uncertainty in modelled leaf chlorophyll content**

653 The validation results (Figure 5 and 6a) indicate a good algorithm performance and levels of uncertainty
654 comparable to more local scale studies, with reported RMSE values in the literature ranging from around
655 8 to 15 $\mu\text{g cm}^{-2}$ (Houborg et al. 2015). The uncertainties values from all of the PFTs used in this study
656 were within this range; from GRS: 5.71 $\mu\text{g cm}^{-2}$ through to EBF: 14.94 $\mu\text{g cm}^{-2}$, and the overall
657 uncertainty of all PFTs considered together was 10.81 $\mu\text{g cm}^{-2}$. The overall uncertainty of the global
658 Chl_{Leaf} product is therefore considered within reasonable limits of current state of the art approaches to
659 chlorophyll modelling. The sources of uncertainty and future ways to minimise this uncertainty are
660 discussed further below.

661

662 We highlight the main sources of uncertainty that can affect the accuracy of the modelled leaf
663 chlorophyll values within the global maps as: 1) the data quality of the MERIS surface reflectance, 2) the
664 accuracy of input LAI data, 3) the use of fixed structural parameters in canopy models, 4) the spectral
665 contributions of understory or background reflectance to satellite-derived canopy reflectance, and 5)
666 the validation sites used within the study.

667

668 1) Uncertainty in the MERIS surface reflectance product can arise from sources such as sensor
669 calibration issues, cloud contamination, atmospheric correction errors (Garrigues et al. 2008). Due to
670 the nature of the inversion process and the 'ill-posed problem' (Combal, Baret, and Weiss 2002), even
671 small changes in satellite-derived reflectance can lead to large variations in modelled chlorophyll values
672 (Garrigues et al. 2008).

673

674 2) The chlorophyll retrieval algorithm is strongly reliant on the LAI as an input parameter, where small
675 errors in LAI values, particularly at low LAI values can lead to large errors in modelled chlorophyll. This
676 introduces a greater uncertainty into the chlorophyll estimates in areas with little vegetation cover, for
677 example in shrubland and deciduous plant forms at the start and end of the growing season. Errors in
678 LAI at high LAI values, for example due to problems concerning reflectance saturation are less of a
679 concern because the algorithm is less responsive to LAI > 4, or when the canopy behaves as a 'big leaf'
680 and background and branch contributions are minimal. Clumping is considered in the GEOV1 LAI
681 product that we selected (Baret et al., 2013), at both the landscape scale and the canopy scale through
682 the integration of CYCLOPES v3.1 and MODIS c5 biophysical products. Errors in LAI values that arise from
683 a lack of proper consideration of clumping effects mainly affect needleleaf canopies, which can be
684 largely underestimated (Chen, Menges, and Leblanc 2005). The most significant problem, however, for
685 the modelling of Chl_{Leaf} is a lack of accurate LAI during winter months in the northern hemisphere. The
686 presence of snow on the understory and on the actual needleleaf shoots prevents the retrieval of LAI
687 during large periods of time in the winter months. Any satellite-derived LAI data that is retrieved during
688 snow-dominated winter months is highly uncertain, leading to potential problems with the smoothing
689 algorithm. Further uncertainty arises from the LAI data product's coarser spatial resolution (1000 m),
690 compared to the input MERIS surface reflectance product (300 m). The greatest uncertainty arising from
691 this source will be in patchy or spatially variable landscapes, such as croplands or where the LAI values

692 are low (i.e. LAI = <3). This uncertainty will have less of a bearing on more spatially homogenous
693 landscapes (i.e. broadleaf forests) and in vegetation with higher LAI values.

694

695 3) The chlorophyll retrieval algorithm considers the major PFTs separately, through the application of
696 individual LUTs based on measured and reported ground data. However, the structural values used to
697 create the LUTs represent a generalised approximation of the structural parameters, and it is recognised
698 that considerable variation can exist within a given biome. An improved spatial representation of the
699 intra-PFT variability of structural parameters, such as canopy height, stem density and leaf angle
700 distribution, would be beneficial, particularly in needleleaf forests, for example, where spatial variability
701 is driven by the dominant species composition. However, a lack of available spatially-continuous data at
702 fine spatial resolutions currently prevents the explicit parameterisation of these structural variables in
703 the canopy models. Whilst these structural parameterisations are important, their variation over the
704 variable range that may be found within a PFT is likely to be small relative to variations between PFTs,
705 and relative to LAI. The sensitivity analysis presented in Section 2.4 indicates that LAI is the dominant
706 driver of modelled Chl_{Leaf} uncertainty, with other parameters such as crown height, ΩE , γE , Hs parameter
707 the background soil factor having negligible influence. This finding also confirms the work of Zhang et al.
708 (2008), who examined the sensitivity of the 4-Scale model input parameters to the model output,
709 finding that LAI is the dominant driver of modelled canopy reflectance.

710

711 4) The temporal and spatial variability of background contributions are a source of uncertainty in the
712 modelled Chl_{Leaf} values, to a greater or lesser degree depending on the overall canopy coverage. Areas
713 with lower LAI values are more susceptible to errors concerning the erroneous parameterisation of
714 background material in the model. Although substantial progress has been made in retrieving
715 understory reflectance from multi-angular reflectance images (Canisius and Chen 2007, Pisek and Chen

716 2009), the often coarse spatial resolution and lack of global coverage for all vegetation types of these
717 products (Jiao et al. 2014, Liu et al. 2017) precludes their use.

718

719 5) An additional source of uncertainty could come from the validation ground data itself. The relatively
720 small number of measured chlorophyll values sampled *in situ* makes widespread validation challenging,
721 particularly for EBF, SHR and DNF PFTs. The validation data used in this study is through the generous
722 and collaborative efforts of independent researchers, and without an established network collecting
723 regular Chl_{leaf} data or automated measurement systems, direct product validation is time and resource
724 intensive (Garrigues et al. 2008). Consequently, no existing validation data sets are completely
725 representative of all of the global and seasonal variability of vegetation [Baret et al., 2006]. This study
726 used 28 distinct sites to validate the Chl_{leaf} product, which is within the same range of other global
727 biophysical product validations. In the calibration and product development of MODIS LAI products,
728 Yang et al. (2006) selected 25 validation sites, and He et al (2012) used 38 validation sites for their 500 m
729 global clumping index map. Initial Soil Moisture Active Passive (SMAP) soil moisture products were
730 validated using 34 validation sites (Colliander et al. 2017). To an extent, the reported uncertainties on
731 the modelled leaf chlorophyll content values will also be a function of the coverage of geographic and
732 vegetation species contained within the validation sites, and the inherent variability that exists within
733 plant functional types. Future work can focus on evaluating the accuracy of the chlorophyll product in
734 species and management regimes that are underrepresented in the model development and validation.
735 It is likely that sites where the structural properties deviate largest from the values used in the model
736 parameterisation (Table 2) may suffer the largest degree of uncertainty.

737

738

739

740 5.0 Conclusion

741 This research represents the first global view of terrestrial Chl_{Leaf} distribution. Weekly maps of Chl_{Leaf} are
742 produced at the global scale following a two-step physically-based modelling approach. The accuracy of
743 the Chl_{Leaf} product is ultimately dependent on the representation of radiative transfer processes within
744 the canopy and leaf optical models, the structural parameterisation of the radiative transfer models and
745 the accuracy of land cover and leaf area index data. Modelled results show good relationships with
746 measured ground data, in particular for deciduous broadleaf forests ($R^2 = 0.67$; $\text{RMSE} = 9.25 \mu\text{g cm}^{-2}$;
747 $p < 0.001$), croplands ($R^2 = 0.41$; $\text{RMSE} = 13.18 \mu\text{g cm}^{-2}$; $p < 0.001$) and evergreen needleleaf forests ($R^2 =$
748 0.47 ; $\text{RMSE} = 10.63 \mu\text{g cm}^{-2}$; $p < 0.001$). On an annual basis, evergreen broadleaf forests presented the
749 highest median leaf chlorophyll values ($54.4 \mu\text{g cm}^{-2}$). The global values show large temporal and spatial
750 variability expected in Chl_{Leaf} . It is expected that ESA Sentinel-2 series will be used to continue the Chl_{Leaf}
751 time series, due to the presence of red-edge sampling bands and widespread spatial coverage and fine
752 temporal resolution. It is anticipated that this global leaf chlorophyll product will make a significant step
753 towards improving global and regional ecosystem models associated with carbon cycle modelling,
754 through the explicit consideration of foliage biochemistry, and reduce the uncertainty associated with
755 leaf physiology.

756

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762 project and Demin data was collected as part of the AGRISAR 2006 ESA field campaign.

763

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