

This is a repository copy of *Benchmarking and Inter-Comparison of Sentinel -1 InSAR* velocities and time series.

White Rose Research Online URL for this paper: https://eprints.whiterose.ac.uk/162927/

Version: Accepted Version

Article:

Sadeghi, Z, Wright, TJ orcid.org/0000-0001-8338-5935, Hooper, AJ orcid.org/0000-0003-4244-6652 et al. (4 more authors) (2021) Benchmarking and Inter-Comparison of Sentinel -1 InSAR velocities and time series. Remote Sensing of Environment, 256. 112306. ISSN 0034-4257

https://doi.org/10.1016/j.rse.2021.112306

© 2021 Elsevier Inc. Licensed under the Creative Commons Attribution-NonCommercial-NoDerivatives 4.0 International License (http://creativecommons.org/licenses/by-ncnd/4.0/).

Reuse

This article is distributed under the terms of the Creative Commons Attribution-NonCommercial-NoDerivs (CC BY-NC-ND) licence. This licence only allows you to download this work and share it with others as long as you credit the authors, but you can't change the article in any way or use it commercially. More information and the full terms of the licence here: https://creativecommons.org/licenses/

Takedown

If you consider content in White Rose Research Online to be in breach of UK law, please notify us by emailing eprints@whiterose.ac.uk including the URL of the record and the reason for the withdrawal request.



eprints@whiterose.ac.uk https://eprints.whiterose.ac.uk/ **1** Benchmarking and Inter-Comparison of Sentinel-1 InSAR velocities and time series

2 Z.Sadeghi^{a*}, T.J. Wright^a, A.J.Hooper^a, C. Jordan^b, A. Novellino^b, L. Bateson^b, J.Biggs^c

- ^a COMET, School of Earth and Environment, University of Leeds, Leeds, UK
- ^b British Geological Survey, Environmental Science Centre, Keyworth, Nottingham, UK
- ^c COMET, School of Earth Sciences, University of Bristol, Bristol, UK
- 6 * Corresponding author, email:z.sadeghi@leeds.ac.uk

7 Key Points: InSAR, Comparison, Sentinel-1, Ground Motion

8 Abstract

9 Different InSAR algorithms and methods produce velocities and times series that are not identical, even using the same data for the same area. This inconsistency can cause confusion 10 11 and be a barrier to uptake and widespread use of the data in the commercial sector. With the widespread availability of Sentinel-1 SAR data and a suite of new algorithms in the commercial 12 and academic sectors, it is timely to develop a method for comparison of different results. In 13 14 this study, we focus on developing and testing an independent and robust methodology for assessment of different InSAR processing results. Our proposed method is adapted from the 15 Terrafirma Process Validation project; we compare geocoded line-of-sight velocities and time 16 series, density and coverage, as well as some qualitative metrics. We use Sentinel-1 data from 17 an area in Glasgow (UK) processed using 4 different approaches modified RapidSAR, 18 19 SqueeSAR, GAMMA-IPTA and conventional StaMPS. The main areas of ground motion are 20 detected using all approaches, with the average standard deviation of velocity differences for 21 all inter-comparison pairs in all polygons equal to 1.1 mm/yr. Sentinel-1 InSAR therefore 22 provides comparable results that are independent of processing approaches. However, there are considerable differences in some aspects of the results, in particular in their density and 23 coverage. We discuss the reasons for these differences and suggest a framework for validation 24 25 that could be used in future national or pan-national ground motion services.

26 Abbreviation:

27	Asc	Ascending
28	APS	Atmospheric Phase Screen
29	Des	Descending
30	DePSI	Delft PS-InSAR processing package
31	DP	Deforming Polygon
32	DS	Distributed Scatterers
33	ESA	European Space Agency
34	EU-GMS	European Ground Motion Service
35	FRInGE	Full-Resolution InSAR time series using Generalised Eigenvectors
36	ISBAS	Intermittent Small Baseline Subsets
37	NERC	Natural Environment Research Council
38	PS	Persistent Scatterer
39	PSI	Persistent Scatterer InSAR
40	SB	Small Baseline
41	SHP	Statistically Homogeneous Pixels
42	StaMPS	Stanford Method for Persistent Scatterers
43	RapidSAR	Rapid Time Series InSAR
44	RP	Rural Polygon
45	TOPS	Terrain Observation by Progressive Scan
46	UP	Urban Polygon
47	WAP	Wide Area Processing

48 **1 Introduction**

Interferometric Synthetic Aperture Radar (InSAR) is an Earth Observation technique based on radar satellite imagery that can measure surface deformation with millimetre level precision (Bamler and Hartl 1998; Gabriel et al. 1989; Hanssen 2001). In order to improve the performance in extracting deformation signals from noisy InSAR data, many different InSAR time-series approaches have been developed (Osmanoğlu et al. 2016; Pepe and Calo 2017).

Persistent Scatterer InSAR (PSI) exploits strong, stable scatterers that display coherent scattering behaviour over time to overcome temporal decorrelation, which restricts the use of conventional InSAR (Ferretti et al. 2000, 2001). By a combination of spatial and temporal filtering, the contribution of atmospheric errors can also be reduced significantly. The original PSI algorithms work where there are large number of strong scatterers (often man-made structures) with a deformation behaviour close to the assumed linear velocity model, although more sophisticated versions of the algorithm, capable of dealing with PS affected by non-linear 61 motion, have also been developed. The Stanford Method for Persistent Scatterers (StaMPS) focusses on improving the number of measurement points in rural areas, and on providing an 62 open source algorithm (Hooper et al. 2007). The major difference between StaMPS and the 63 traditional PS approach is that StaMPS uses the spatial correlation of phase for identifying PS 64 pixels and does not use phase triangulation which forms a spatial network connecting all PS 65 pixels (Hooper et al. 2007). All PSI algorithms use a single-master stack of differential 66 67 interferograms to process PS pixels (Hooper et al. 2012). For satellites such as ERS-1/2 and Envisat, with a relatively large orbital tube and hence a large range of perpendicular baselines 68 69 in individual interferograms, only point scatterers remain coherent in a single master stack.

Distributed Scatterers (DSs) can also be used for extracting velocities and times series from 70 InSAR. These contain coherent information when temporal and orbital baselines are relatively 71 short/small but can be incoherent in interferograms with relatively long time intervals and large 72 perpendicular baselines (different viewing geometries). Small baseline (SB) approaches build 73 74 time series by connecting interferograms with small temporal and perpendicular baselines (Berardino et al. 2002; Schmidt and Bürgmann 2003). By combining PSI and SB approaches, 75 hybrid approaches can increase the measurement density (Hooper 2008; Lanari et al. 2004). 76 77 However, there may still be useful interferometric measurements within the stack of SAR data that are excluded from a hybrid PS/SB analysis, particularly in rural areas where pixels may 78 have intermittent coherence. The multi-interferogram method (Biggs et al. 2007) implemented 79 in PiRate (Wang et al. 2009) and ISBAS (Intermittent Small Baseline Subsets) method (Sowter 80 81 et al. 2013) are based on a modification of the SBAS method (Berardino et al. 2002) and exploit 82 intermittent coherence in order to obtain average velocities for a greater number of DS.

However, time series approaches that only use short-timespan, multi-looked interferograms
suffer from potential biases (Ansari et al. 2020).

85 SqueeSAR forms all possible interferograms, selects neighbouring pixels with similar scattering mechanisms, known as statistically homogeneous pixels (SHP), and provides a 86 87 synergistic analysis of PS and DS without the need for significant changes to the traditional 88 PSI processing chain (Ferretti et al. 2011; Fornaro et al. 2015; Monti-Guarnieri and Tebaldini 89 2008). It improves the density, coverage and quality of measurement points with respect to 90 conventional PSI, over non-urban areas at the cost of a large increase in processing time. In 91 contrast, RapidSAR (Rapid Time Series InSAR) was designed to allow fast ingestion of new images and limited computational load (Spaans and Hooper 2016). This method identifies SHP 92 pixels (named siblings) with a more computationally efficient algorithm than SqueeSAR and 93 does not use phase triangulation. RapidSAR enables coherence in newly formed interferograms 94 to be calculated quickly – the results can be used in a modified SBAS approach to produce time 95 96 series and velocities (Spaans and Hooper 2016).

Using all possible interferograms at full SAR resolution for Sentinel-1 or other wide-swath
SAR missions is challenging due to the large data volume. The Sequential Estimator approach
has therefore been proposed to form interferograms efficiently for long InSAR time series by
processing the data in small batches and forming compressed artificial interferograms from
each (Ansari et al. 2017). Alternatively, FRInGE (Full-Resolution InSAR time series using
Generalised Eigenvectors) generates a full coherence matrix efficiently and selects both PS and
DS pixels at full resolution.

Different InSAR time series methods use different strategies to extract information from SAR
 images. They are also different in terms of dealing with the contributions of various phenomena
 impacting the interferometric phase including long wavelength trends, atmospheric phase

107 screens (APS) and nonlinear deformation. Moreover, different strategies can be applied to remove these terms, e.g. spatial and temporal filter size for removing APS, size of spatial scale 108 to de-trend and/or type of nonlinear deformation model (i.e. periodic, exponential). Therefore, 109 products of different InSAR algorithms are not identical and can be dissimilar in terms of 110 111 quantitative and qualitative metrics. In order to assess the quality of InSAR data for a specific case study, there is a requirement to evaluate the consistency of available InSAR data produced 112 113 by different time series approaches. Up to now, several different methods to compare InSAR products have been presented, all of which have limitations and none of which have used 114 115 Sentinel-1 images, (see Section 2). Therefore, there is a need to present an inter-comparison method which addresses the limitations of previous studies including the application to 116 Sentinel-1 InSAR data. 117

Sentinel-1 is a two-satellite imaging radar constellation, providing global C-band imagery 118 designed to supply the data needs of Europe's Copernicus programme. Sentinel-1A & -1B offer 119 120 a six-day revisit cycle and unprecedented coverage of Europe, with 12-day imagery acquired globally. Sentinel-1 uses the Terrain Observation by Progressive Scan (TOPS) mode, sweeping 121 the beam in the fight direction, and is designed primarily for InSAR applications (De Zan and 122 123 Monti Guarnieri 2006). Raw data acquired by Sentinel-1 are freely available, addressing the limitation of cost and/or lack of data and providing research and commercial opportunities 124 increasingly, Sentinel-1 data are being used to form nationwide/international ground motion 125 maps. All Sentinel-1 imagery is acquired within a narrow orbital tube, maximizing 126 interferometric coherence. To exploit Sentinel-1 data, a European Ground Motion Service (EU-127 128 GMS) is under development, by the European Environment Agency 129 (https://land.copernicus.eu/user-corner/technical-library/european-ground-motion-service), to provide consistent, regular, standardised, harmonised and reliable information on ground 130 motion over Europe and across national borders, with millimetre accuracy (Crosetto et al. 131

2020). The ground motion results will be derived from time series analyses of Sentinel-1 data, 132 most likely using different PS and DS InSAR approaches. Several Copernicus Participating 133 States including Germany, Italy, Norway, Spain, Denmark, and France have already or are in 134 the process of implementing national ground motion services. These services will benefit from 135 EU-GMS by standardising national service components and encouraging the use of 136 deformation data by both public and commercial users. To make the outputs useful for 137 138 operational applications, quality assessment of ground motion maps is a fundamental priority, and an important aspect of quality assessment is data consistency, particularly at borders or 139 140 boundaries, where different methods may have been used. The nationwide/international ground motion map will be likely processed by multiple suppliers, therefore there is a need to assess 141 and ensure consistency of InSAR results. 142

Our main goal in this research is to develop and test a fair and robust methodology to assess 143 144 the similarities and differences between results from different InSAR processing chains, and to 145 recommend a validation strategy for any nationwide/international (e.g. UK/EU) ground motion map. We review the history of InSAR comparison approaches with their characteristics and 146 limitations in Section 2. In Section 3, we describe an approach we have developed. In section 147 4, we use the method to compare 4 processing algorithms for a test area in Glasgow. We present 148 results in Section 5 and discuss the major differences and similarities between the InSAR 149 results in Section 6, providing recommendations for future nationwide/international products 150 151 and validation activities. Finally, we summarise the main conclusions in Section 7.

152 2) Review of previous InSAR comparison and validation approaches

Several previous projects have compared and validated InSAR velocities and time series.
Following the 2003 Fringe meeting, the European Space Agency (ESA) initiated a blind InSAR validation project, PSIC4 (Persistent Scatterer Interferometry Codes Cross Comparison and Certification for long term differential interferometry), to assess the performance of PSI for

157 land deformation monitoring using Envisat and ERS images (Crosetto et al. 2007b; Raucoules et al. 2009). The project analysed results for the same area provided by Altamira Information 158 (Crosetto et al. 2008a), DLR (German Space Agency) (Adam et al. 2005), Gamma Remote 159 160 Sensing (Werner et al. 2003), IREA-CNR (Institute for Electromagnetic Sensing of the Environment National Research Council of Italy) (Berardino et al. 2002), TRE (Tele-161 Rilevamento Europa) (Ferretti et al. 2007), TUDelft (Delft University of Technology) (Kampes 162 163 2005), UPC (Catalonia Polytechnics University) (Mora et al. 2003) and Vexcel (Van der Kooij et al. 2005). Pre-processing of the data prior to inter-comparison comprised applying 164 165 geolocation shifts and spatially referencing each data set to the same reference area. The most relevant indicators used to compare the results were the average deformation rate and the 166 density and distribution of the selected PS points. The PSIC4 test area was a coal mining area 167 168 in the South of France, which was undergoing rapid subsidence and did not include stable features. The reference area was a stable local area outside the mining work. The results showed 169 that for the case under consideration, the main area of subsidence could not, or could only 170 partly, be assessed by most of the InSAR teams due to the low density of PS in the area of 171 interest. Moreover, the standard deviation of velocity differences between the data sets ranged 172 between 0.6 and 1.9 mm/year which can be considered as an estimate of local uncertainties. 173 One of the most important conclusions of PSIC4 concerned the characteristics of the coal 174 175 mining test site in which none of the conditions to measure deformation with millimetric 176 accuracy by PSI was fully realised. The severe characteristics of the PSIC4 test site were nonoptimal for PSI due to i) abrupt nonlinear motion and ii) rapid velocities which were prone to 177 aliasing with the 35-day revisit time of Envisat/ERS. The project recommended future SAR 178 179 missions with more frequent acquisitions in order to improve the ability of PSI to detect rapid velocities (Raucoules et al. 2009). PSIC4 used "blind conditions" with no a priori information 180 about the deformation or the goal of the PSI analysis. The teams used a standard PSI approach 181

182 instead of tailoring the processing to a specific objective, which could partly explain the lack of PS in the mining area. PSIC4 demonstrated that, at that time, PSI performance was highly 183 dependent on the application, and the limitations were real. A wider area inter-comparison, 184 "Provence Inter-Comparison", was later presented using the same data as PSIC4 but covering 185 a larger area, including both deforming and stable areas (Crosetto et al. 2007a). One difference 186 between the Provence Inter-Comparison study and PSIC4, was that it compared the data set in 187 188 the radar coordinate system, to avoid validation issues associated with geocoding errors. The Provence Inter-Comparison showed a greater degree of consistency between the velocity maps 189 190 and the time series from different providers (Crosetto et al. 2007a). It was largely based on data outside the mining area, where the results of the two projects were similar. 191

The Terrafirma project (Capes et al. 2009), part of the EU/ESA Global Monitoring for 192 Environment and Security (GMES) programme, the precursor to Copernicus, also established 193 a PSI process validation approach, known as the Terrafirma Validation Project, which built on 194 the earlier studies. The Terrafirma validation project had two aims: result validation via 195 comparison with ground truth levelling and inter-comparison of the results of different InSAR 196 providers. The inter-comparison methodology initially compared four InSAR data sets from 197 different providers (TRE, Altamira Information, Gamma Remote Sensing, and Fugro NPA) in 198 radar coordinate systems to a reference processing result (GENESIS, DLR PSI processing), 199 which was defined as the "truth" (Adam et al. 2009). Pre-processing steps included checking 200 the global consistency of the data sets and the coregistration in radar space, referencing the data 201 202 to the same reference in the time and space dimensions and removing potential tilts by detrending. Velocities, time series, topographic corrections, detection capabilities and data 203 densities were compared. The project produced a set of global statistics, which concerned large 204 sets of PS pixels and provided information on the global inter-comparison behaviour of 205 different metrics. The average standard deviation of the velocity differences and the mean 206

standard deviation of the time series differences were 0.5-0.7 mm/yr and 1.5-5.6 mm, respectively. These values were used to derive error bars to indicate the quality of the estimate derived by PSI, which was key information for Terrafirma end users. Since deformation rates in the case studies in the Terrafirma Validation Project were moderately low, one should be careful in extending these statistics to areas involving higher deformation rates. Moreover, the results showed remarkable differences in PS density between the providers, which resulted from the use of different criteria during PS selection (Crosetto et al. 2008b).

214 As the Terrafirma PSI certification process was intended for local (20km×20km) PSI analysis of deformation, the Wide Area Processing (WAP) Terrafirma project later expanded this 215 methodology to validate PSI processing over a significantly greater area (one or more scenes 216 of 100km×100km) than that considered in the initial Terrafirma PSI certification (Adam et al. 217 2013; Brcic et al. 2014). The major differences between the processing chains in the wide area 218 relate to atmospheric compensation and trend removal. Both steps were applied for TRE 219 products, none of them implemented for DLR products and Altamira only removed the long 220 wavelength trend. The results showed that the standard deviations of the deformation velocity 221 differences for coherent pixels were below 1 mm/yr in most of the inter-comparison cases. This 222 was one requirement of Terrafirma PSI certification. It also concluded that the most significant 223 factors affecting compliance with this requirement were: (a) possible long wavelength trends 224 225 affecting the interferograms (resulting from spurious atmospheric components and orbital fringes); (b) systematic phase components associated with the master scene used for the PS 226 227 analysis, and (c) possible phase unwrapping errors, which were strongly dependent on the deformation signal and the presence of any data gaps in the interferograms. 228

Previous validation approaches have several limitations. Firstly, to be useful in real-world applications, InSAR data must be geocoded. By only comparing results from different methods/providers in the radar coordinate system, an important step of InSAR processing is

232 excluded. Secondly, specifying a reference InSAR product as the "truth", as was done in the Terrafirma Validation Project, can also lead to an unfair comparison, as it excludes the 233 possibility that the reference data set also has errors. Thirdly, validation projects to date have 234 235 used data from Envisat and ERS; the improved spatial and temporal coverage of Sentinel-1 data, and its narrow orbital tube, opens up several new opportunities for InSAR processing, 236 which were not feasible previously. Finally, previous approaches were only applied to 237 238 validating PSI methods; a comparison method that can consider both PS and DS is now required. 239

Several recent studies have compared individual data sets or methods. A comparative study based on the results from DePSI (Delft PS-InSAR processing package) and StaMPS (Stanford Method for Persistent Scatterers), was applied using two data sets from ERS and Envisat and concluded that these methods are complementary (Sousa et al. 2011). The time-series InSAR results generated using ERS data with both a PS method and a SBAS algorithm were compared quantitatively and the calculated discrepancy was found to be consistent with those estimated by the PSIC4 study (Shanker et al. 2011).

In another study, the capability of three InSAR time-series techniques, PSI, SBAS and SqueeSAR, for evaluating landslide deformation, was investigated using TerraSAR-X images (Mirzaee et al. 2017). The estimated average velocity maps and coherence maps produced by the methods were compared and it was concluded that SqueeSAR was more efficient for evaluating landslide kinematics in the rural case study.

Finally, the performance of ISBAS and RapidSAR were compared using Sentinel-1 images to monitor shale-gas operations in Lancashire, outlined as part of the Environmental Baseline Monitoring programme conducted by the British Geological Survey (BGS). The results showed agreement between the approaches to estimate average annual velocity in the study area (Jordan et al. 2019). With the Copernicus European Ground Motion Service now being commissioned, it is timely to formalise requirements for comparison of InSAR results. In this research, we present a methodology for inter-comparison of geocoded InSAR products using Sentinel-1 images. We test this methodology with the InSAR products resulted from different InSAR time series algorithms over a case study where multiple InSAR data sets are available.

262 **3 Methods**

In this section, we introduce our new inter-comparison method. The outline of the proposed 263 approach is shown in Figure 1. We base the approach on the Terrafirma Validation Project, but 264 265 tackle its limitations as follows: 1) As end-users require geocoded InSAR data, we compare all the data sets in geographic rather than radar coordinates. This allows us to consider any 266 potential geocoding errors that can impact on the final product, especially in areas with very 267 local deformation. 2) Because no InSAR processing chain produces perfect, noise-free results, 268 we avoid assuming that any reference InSAR processing is the "truth". 3) We define several 269 polygons with different land cover types and stability. This allows us to assess how the 270 agreement differs between InSAR data with different signals and/or different ground 271 272 conditions. 4) We do not limit the time series processing to PSI algorithms and are open to any 273 other methodologies e.g. both PS and DS InSAR processing. 5) We work with Sentinel-1 274 imagery. Our approach can be split into pre-processing and inter-comparison stages. These are described in more detail below. 275

276 3.1 Pre-processing

277 Before comparing data sets, some pre-processing steps are required:

(i) We assess the consistency of geocoded data sets from different InSAR methods. As
the coordinate system of the points selected by different InSAR methods might be
different, we convert all the InSAR data to an identical geographic coordinate system.

Any geocoding errors are critical when the deforming area is very small and should be noted. Adjustments can be made if necessary to ensure the data are comparable. This pre-processing step was applied in the PSIC4 project. We assume that any translation of coordinates is constant for the whole data set; and assesse them by overlaying the data on an accurate base map and considering some control points.

- (ii) We select pairs of InSAR data sets for comparison, with each data set processed using
 a different method. For the comparison to be valid, both data sets in a pair must use
 data from the same ascending or descending Sentinel-1 pass. Therefore, the inputs of
 this step are individual InSAR data sets from different methods and the outputs are
 different pairs of data sets.
- (iii) The time range of InSAR data for each comparison pair may be different. To ensure
 consistency as much as possible, we re-estimate the velocity using a common time
 range for each comparison pair, by fitting linear velocities to the time series for each
 pixel using only data from the common time range. The re-estimated velocities for
 the InSAR data sets forming each InSAR pair are the outputs of this step.
- (iv) We identify the common dates in the time series for each comparison pair and set thefirst common date as a reference time as follows:
- 298 $d(t)_{Tem-new} = d(t)_{Tem-old} d(t_0)_{Tem-old}$ (1)

where for each selected point in an inter-comparison pair, $d(t)_{Tem-new}$ is the rereferenced deformation time series in temporal space (output), $d(t)_{Tem-old}$ is the original deformation time series (input) and $d(t_0)_{Tem-old}$ is the deformation of the first common date for the corresponding pair.

303 (v) We can optionally apply an identical low pass filter to each time series data set
304 (input), in this case using a triangular filter covering 5 epochs. This helps to remove

the effect of random noise in a similar manner from both time series. The output of
this step is filtered time series of InSAR data for each pair. Ideally, we work with
unfiltered time series before applying this filter, but this may not always be possible.
In this study, only one of our data sets was filtered, and we used the same temporal
filter for the other unfiltered data sets.

(vi) We re-reference the deformation rate and deformation time series of all comparison
pairs to an identical local reference area, which is outside the deforming areas and
contains coherent pixels:

313
$$d(t)_{new} = d(t)_{Tem-new} - d_r ref(t)_{Tem-new}$$
(2)

$$V_{new} = V_{old} - V_{ref_{old}}$$

where for each selected point in an inter-comparison pair, $d(t)_{new}$ is the re-315 referenced deformation time series in spatial and temporal spaces (output), 316 $d_ref(t)_{rem-new}$ is temporally re-referenced deformation time series for the 317 reference area, V_{new} is the re-referenced deformation velocity in spatial space 318 (output), V_{old} is the original deformation velocity and V_ref_{old} is the deformation 319 velocity of the reference area. Unlike the Terrafirma Validation Project, we do not 320 have access to the coherence of selected points for all data sets. Therefore, we apply 321 a noise analysis algorithm to identify high-quality pixels in the reference area 322 323 (Hooper et al. 2007; Sadeghi et al. 2018). First, selected pixels are connected to form a network using Delaunay triangulation. Then, for each arc connecting two 324 pixels, a weighted average phase is calculated from the entire time series, and 325 removed from the original phase of the arc, which is then low pass filtered in time. 326 The resulting phase, with the weighted average phase added back in, provides an 327 estimate for the smooth underlying signal. Phase noise is estimated by subtracting 328

- the smooth phase from the original phase of the arc. Finally, the phase noise of each measurement pixel is obtained from the phase noise of its corresponding arcs. The pixels with a noise level less than a threshold for all data sets are selected in the reference area.
- (vii) We define an identical geographic grid with 40 meters spacing in both easting and
 northing and for each of the InSAR data sets calculate the mean value of any
 measurement points located inside each grid cell. The outputs of this step are the
 time series and velocities for each defined grid cell. We specify no-data for grid cells
 which contain no measurement points.
- (viii) We define polygons covering areas with different scattering and deformation
 characteristics so that the algorithms can be tested in different conditions. In the case
 of our test site in Glasgow, we define urban, rural and deforming polygons (see
 section 4 and Figure 2). Selected grid cells inside each defined polygon for all InSAR
 data are the outputs of this step.



Figure.1) a flowchart showing our proposed inter-comparison methodology.

344 3.2 Inter-comparison of data sets

- After pre-processing, we compare InSAR results in terms of several metrics, including theestimated velocities, time series, density and coverage as follows:
- 347 (i) We calculate the differences between the deformation velocities for the common grid 348 pixels of each pair and estimate their mean (μ_{dV}) and standard deviation (σ_{dV}) . We 349 also calculate the correlation coefficient for estimated velocities (ρ_V) of all common 350 grid pixels.
- (ii) In order to extract statistics from the differences between time-series, (a) in the first 351 step, we compute the differences between the time series for each common grid pixel 352 of each pair and then extract their mean and standard deviation; (b) we calculate the 353 mean of the parameters computed in the previous step for all common grid pixels of 354 a given pair , mean of mean of time-series differences $\mu\mu_{dD}$ and mean of standard 355 deviation of time-series differences $\mu\sigma_{dD}$. The mean values show any potential bias 356 between the estimated deformation velocities/time series of each pair. Standard 357 deviation values provide information on how the deformation velocity/time series 358 differences are distributed. We also calculate the correlation coefficient for the 359 360 estimated deformation time series (ρ_D) of each common grid pixel. This is a useful tool for measuring the degree of similarity of the deformation histories of the 361 analysed time series. 362
- 363 (iii) In order to compare the density and coverage of measurement pixels, we resample 364 the InSAR data onto an identical 100 m × 100 m grid. The number of selected pixels 365 in each cell gives the selected pixel density (*D*); we calculate the average density for 366 each of the polygons with different scattering/deformation characteristics. We also 367 calculate the coverage (*C*) of measurement pixels, which we define as the percentage 368 of 100 m × 100 m grid pixels containing at least 1 InSAR measurement. We note that

in order to make a fair comparison in terms of density and coverage, the noise
analysis described in section 3.1-vi should be applied before comparison and noisy
pixels should be removed from each data set using the same threshold for the phase
noise standard deviation, in this case 1 rad in our case.

4 Data and Case Study

374 We use results from the Clyde Gateway of the Glasgow City Region to test our methods (Figure 2). This is an area of particular interest to the Natural Environment Research Council (NERC) 375 as it is the BGS geothermal energy research field test site of the UKGEOS project 376 377 (https://www.ukgeos.ac.uk/about/project-details). The Glasgow site will help characterise whether water from abandoned mine workings can be used to generate a sustainable and 378 efficient source of energy. Changes in underground water levels, pressure and temperature 379 380 caused by mine water for geothermal energy production activities can lead to surface subsidence/uplift (Heimlich et al. 2015). Therefore, monitoring is required to assess surface-381 level impacts of geothermal abstraction and re-injection research activities (Bateson and 382 Novellino 2019). Although the area is largely urban, it also includes more rural areas, such as 383 the Woodland park within the Cuningar loop. 384

We have access to several Sentinel-1 InSAR data products for this area, processed using four 385 386 different approaches. Results for two of these approaches were provided by commercial companies: SatSense, using a modified RapidSAR algorithm (https://www.satsense.com) 387 (Spaans and Hooper 2016) and TRE-ALTAMIRA, using the SqueeSAR algorithm 388 (https://site.tre-altamira.com) (Ferretti et al. 2011). The results for the other two approaches 389 come from our own processing using GAMMA-IPTA, processed using conventional PSI at 390 BGS (https://www.gamma-rs.ch) and the PS-only option of StaMPS (Hooper et al. 2007). 391 392 Analysis and interpretation of some of the GAMMA-IPTA results can be found in Bateson and

393 Novellino (2019). We used the results of the four different approaches to test our InSAR intercomparison activity. In all we have 5 data sets, 3 in an ascending geometry and 2 in descending. 394 Hereafter, we anonymise the algorithms and label them A-D, in no particular order. We have 395 396 ascending and descending InSAR results for algorithm A, which we used for inter-comparison independently. For algorithm B, we have data for descending geometry only and for algorithm 397 C and algorithm D we have data for ascending geometry only. Therefore, we formed 4 inter-398 399 comparison pairs: A-B (descending), A-C (ascending), A-D (ascending) and C-D (ascending). Table 1 compares the key characteristics of the data sets: the longest time span and the largest 400 401 number of available images are related to A (descending) and B (descending) which used similar Sentinel-1 data sets, while C includes the shortest time range and smallest number of 402 403 Sentinel-1 scenes. B and C used PSI algorithms which select only PS pixels and form single-404 master interferograms, but A and D took advantage of identifying both PS and DS pixels and 405 made a multiple-master interferometric network. Apart from C, spatial de-trending was applied during processing to all of the InSAR data to remove any potential long wavelength trends. 406 407 The different strategies applied by the InSAR algorithms for removing the effects of unwanted elements such as long wavelength trend and APS might have an impact on the level of 408 409 agreement between the algorithms for example by introducing a bias in the average of deformation velocity differences. 410

In algorithm A, the level of noise for a point is assessed by calculating the difference between the smoothed time series and the APS filtered time series. After referencing to the average of its neighbours, this helps to give an idea of which points are inherently noisier since atmospheric effects have been reduced. The phase noise is estimated for algorithm B using method described in section 3.1-vi. For algorithm C, the point quality is measured by calculating the standard deviation of the misfit to a regression through the time series. The standard deviation of the phase misfit depends on the quality of the reference point, the target 418 point and the pre-defined model (usually linear). For algorithm D, a linear model is fitted to 419 the time series for each measurement point before compensating for possible atmospheric 420 components and the standard deviation of the residual phases is calculated to estimate the 421 uncertainty of the average velocity.

Table.1) Key characteristics of the InSAR products: A-descending (Fig 3a), A-ascending (Fig 3b), B-descending (Fig 3c), C-ascending (Fig 3d), D-ascending (Fig 3e).

	Geometry	Time range	Number of Scenes	Measurement Points	Interferogram Network	Trend removal
А	Ascending (Fig 3c)	2015-05-23 2018-12-27	168	PS and DS	Multiple- Master	Yes
	Descending (Fig 3b)	2015-05-01 2019-02-27	175			
В	Descending (Fig 3d)	2015-05-01 2019-02-27	175	PS	Single-Master	Yes
С	Ascending (Fig 3e)	2015-08-15 2017-06-11	35	PS	Single-Master	No
D	Ascending (Fig 3f)	2015-03-12 2017-11-26	107	PS and DS	Multiple- Master	Yes

422 We define three polygons for the test region that broadly cover "deforming" (0.2 km^2) ,

423 "urban" (0.9 km²) and "rural" areas (1.2 km²) (Figure 2). The area west of Cuningar Loop

424 (deforming polygon) was a site developed for the Commonwealth Games Athletes Village

425 and suffers from small 6 mm/yr rates of linear subsidence due to loading of the superficial

426 deposits (Bateson and Novellino 2019).



Figure.2) Location of our case study in Glasgow, the yellow outlined polygons are defined as areas including urban, rural and deforming features.



e)

Figure.3) Estimated LOS deformation velocity in the case study by a) algorithm A using descending Sentinel-1 images, b) algorithm A using ascending Sentinel-1 images, c) algorithm B using descending Sentinel-1 images, d) algorithm C using ascending Sentinel-1 images and e) algorithm D using ascending Sentinel-1 images. The yellow outlined polygons are defined in a) as areas including urban, rural and deforming features. The black outlined ovals show localised subsidence signals.

428

430 **5 Results**

In this section, we present the results of our inter-comparison methodology using the available
data sets. Four InSAR comparison pairs can be made: A-B (descending), A-C (ascending), AD (ascending) and C-D (ascending). Before showing the inter-comparison results, we show the
averaged velocities on the defined grid and the common dates for each inter-comparison pair
in Figure 4 and Figure 5, respectively. As can be seen in Figure 4, the reference area is local,
and therefore the estimated inter-comparison statistics represent the local uncertainty.



c)

d)



Figure.4) Average LOS deformation velocity in the inter-comparison grid for the case study by a) algorithm A using descending Sentinel-1 images, b) algorithm A using ascending Sentinel-1 images, c) algorithm B using descending Sentinel-1 images, d) algorithm C using ascending Sentinel-1 images and e) algorithm D using ascending Sentinel-1 images. The yellow outlined polygons are defined as areas including urban, rural and deforming features. The magenta outlined rectangular shows location of reference area. The Purple outlined square in a) shows an area to estimate variograms in figure 8.





437

5.1 Inter-comparison of velocity:

The velocity differences are calculated for the common grid pixels of each pair in the urban,rural and deforming polygons. The mean and standard deviation of deformation velocity

- 440 differences and correlation coefficients of the estimated velocities are extracted and reported
- 441 in Figure 6.



b)

Figure 6) a) An error bar plot showing the mean of velocity differences at the centre of the error bars and the standard deviation of velocity differences as length of the error bars (±1sigma) b) a bar chart showing correlation coefficients of velocities for all common grid pixels between InSAR products in urban, rural and deforming polygons.

442 The mean differences are 1.0 mm/yr at most (C-D, urban polygon), and most are less than 0.1 mm/yr, confirming that there are not any significant biases in estimated velocities. The mean 443 velocity differences associated with comparison pairs A-B and A-D are closer to zero than 444 445 those for A-C and C-D. The standard deviation of the velocity differences can be used to assess the level of agreement between the InSAR algorithms, but does not mathematically represent 446 the uncertainty. The standard deviation related to A-B is well under 1 mm/yr and indicates a 447 good overall agreement. The standard deviations of the other InSAR algorithm pairs are all 448 better than 2 mm/yr and the average of the standard deviations is highest in the rural polygon 449 and lowest in deforming polygon for all pairs. The average of the standard deviations of 450

velocity differences for all polygons and all pairs is 1.1 mm/yr; we use this value to showconfidence bounds for measurements in the scatterplots in Figure 7.

453 All InSAR comparison pairs show low correlation where there is little deformation (in the urban and rural polygons), but have higher correlation in the deforming polygon. Correlation 454 coefficients (ρ_V) are in the range 0.5 to 0.7 showing a good agreement between the different 455 methods, where there is significant deformation. We illustrate the agreement by creating 456 scatterplots of the estimated velocities in the deforming polygon (Figure 7). The correlation is 457 clearest in the comparison between A and B, where both data sets have relatively high density, 458 but a good correlation is also seen in the deforming area for the other data sets, confirming that 459 460 all algorithms are detecting similar deformation signature in this region.



Figure.7) Scatterplot of the estimated LOS velocities by a) A-B, b) A-C, c) A-D, d) C-D for deforming polygon. The color associated with each grid cell in the scatterplot shows the number of measurements points. The blue line shows the y=x axis which is an ideal location for points in a scatterplot and the dashed black lines are showing edge of the confidence bounds (1 sigma) assuming the standard deviation of velocity differences equals 1.1 mm/yr.

In Figure 8, we also show variograms of the estimated velocity differences for all common grid cells of each inter-comparison pair inside the purple outlined square in Figure 4-a. This helps assess the spatial variability in the difference between the estimated velocities by the different methods. The variograms show that the noise level does not increase significantly with the spatial separation of the points on the 1-2 km length scales that we have analysed in this study. However, we would expect the noise level to increase with distance for longer length scales (Emardson et al. 2003).



Figure.8) Experimental variogram (γ) of velocity differences at different separation distances for common grid cells in the purple outlined square in Figure 4-a) between a) A and B (descending), b) A and C(ascending), c) A and D(ascending), d) C and D (ascending).

468 5.2 Inter-comparison of time-series:

As described in the Section 3, we calculate the differences between deformation time series in each comparison pair at each common grid pixels, and then extract mean and standard deviation of these differences at each pixel. Finally, we estimate the mean of the means and the mean of the standard deviations using all of the common grid pixels in each of the polygons (Figure 9).

The mean of mean values $(\mu\mu_{dD})$ is under ±2 mm for all pairs, indicating that there are noticeable systematic effects between the time series pairs. The mean of standard deviations $(\mu\sigma_{dD})$ ranges between 1 mm for the A-B pair and 2 mm for the A-D pair. Time series statistics associated with the A-D pair shows the poorest agreement with respect to the others for all polygons, except the urban polygon where mean of the mean is slightly lower than A-B pair.



Figure.9) An error bar plot showing the mean of mean of time series differences at the centre of the error bars and the mean of standard deviation of time series differences as length of the error bars (±1sigma) for all common grid pixels between InSAR products. The unit for the vertical axis is mm.

We also calculated the correlation coefficient for the estimated deformation time series of each common grid pixel and plotted the percentage of common grid pixels with a correlation coefficient above 0.7 in Figure 10. This figure confirms that the percentage above 0.7 is over 50% for all comparison pairs in the deforming polygon, which means more than half the common grid pixels in this polygon show a high level of similarity between the patterns of estimated deformation. The most similar pattern of deformation time series for all polygons is related to the A-B inter-comparison pair.



Figure.10) A bar chart showing the percentage of common grid pixels where the correlation coefficient between

deformation time series is above 0.7 for urban, rural and deforming polygons.

Comparison pair in descending geometry for point "1":





Comparison pairs in ascending geometry for point "1":



Figure.11) a) Estimated LOS velocity by A(ascending) inside the deforming polygon on our grid used for comparison; data from measurement grid pixel "1" is shown in b-j; b) deformation time series plot of A(descending) and B (descending) using the first common date as a reference in time; c) scatterplot of A(descending) and B(descending; d) deformation time series plot of A(ascending), C(ascending) and D(ascending), the temporal reference of time series is the original reference selected by the InSAR algorithms; e) deformation time series plot of A(ascending) and C(ascending) in the common temporal interval using the first common date as a reference in time; f) scatterplot of A(ascending) and C(ascending); g) deformation time series plot of A (ascending) and D (ascending) in the common temporal interval using the first common date as a reference in time; f) scatterplot of A (ascending) and D (ascending) and D (ascending); i) deformation time series plot of C (ascending) and D(ascending) in the common date as a reference in time; f) scatterplot of C (ascending) and D(ascending) in the common date as a reference in time; j) scatterplot of C (ascending) and D (ascending) in the common date as a reference in time; j) scatterplot of C (ascending) and D (ascending) in the common date as a reference in time; j) scatterplot of C (ascending) and D (ascending) in the common date as a reference in time; j) scatterplot of C (ascending) and D (ascending) and D (ascending).

For illustration purposes, we also compare the time series for one typical subsiding grid cell in 486 487 the deforming polygon (Figure 11-a)) for all pairs. We plot all the time series on the same time axis, although the common temporal interval between the algorithms differs for each InSAR 488 pair. The time histories from the different algorithms and viewing geometries compare well 489 when they are observing the same time periods. Figure 11-b) and 11-c) show the deformation 490 time series plot and deformation time series scatterplot for A and B which used the same 491 492 descending Sentinel-1 data sets. There is an excellent correlation coefficient between the deformation patterns estimated by the two algorithms and no significant bias between the 493 estimated deformation time series can be seen. The reference of the deformation time series is 494 495 the first common date. The deformation time series with the original reference in time selected 496 by the InSAR algorithms are plotted in Figure 11-d) for A, C and D using descending Sentinel-1 images. Then the deformation time series for A-C, A-D and C-D (each comparison pair in 497 498 ascending geometry) are plotted in the common temporal interval using the deformation of the first common date as a reference in time in Figure 11-e), 11-g) and 11-i), respectively. The 499 500 corresponding scatterplots of the estimated deformation time series for A-C, A-D and C-D pair are shown in the Figure 11-f), 11-h) and 11-j), respectively and good agreement between the 501 502 InSAR algorithms in detecting the deforming signal can be seen.

503

5.3 Inter-comparison of density and coverage:

504 One major difference between the InSAR products is the density of pixels. We compare these 505 for all InSAR algorithms in Figure.12. All InSAR algorithms provide the highest and lowest 506 density in urban and rural areas, respectively. We plot density maps for the different InSAR 507 algorithms in Figure.13. The results confirm that A is the most successful InSAR algorithm in

- terms of density of pixels for both ascending and descending geometries, and D identifies the
- 509 lowest density.



Figure.12) A bar chart showing the average density of measurement points by the InSAR algorithms for urban, rural



polygon and deforming polygons.

Figure.13) Density of InSAR measurements (Number of measurement points in 1 km^2) by a) A (descending), b) A (ascending), c) B, d) C and e) D.

We also compare the coverage for each InSAR comparison pair in Figure 14, defined as the percentage of 100×100 m grid pixels containing at least one measurement. The coverage of different InSAR algorithms in the deforming polygons is very similar. Although D provided the lowest density for all polygons, it offers the highest coverage in rural and deforming polygons. Indeed, it was able to select pixels in some locations other InSAR algorithms were not, including inside the Cuningar loop (Figure 3).



Figure.14) A bar chart showing the coverage of measurement points (the percentage of 100×100 m grid pixels containing at least one measurement) by the InSAR algorithms for urban, rural polygon and deforming polygons.

516 **6 Discussion and Recommendations**

In this section, we discuss the major similarities and differences between InSAR results. We 517 then recommend some requirements for a national/international ground motion map. The 518 519 results are quite similar despite the very different algorithms used. The major similarity between all the InSAR data sets is that they all detect similar deformation signals in the 520 deforming polygon. Moreover, all of the methods provide a good density of observations in the 521 522 urban polygon, as bright scatterers are selected appropriately by all the methods. In addition to the motion in the deforming polygon, a number of other features of deformation are seen in all 523 524 data sets. For example, all data methods show localised subsidence (up to 10 mm/yr) on the M74 motorway gantry highlighted with black outlined ovals in Figure 3, which is likely related 525 to instability in the embankment supporting the motorway at this location (Bateson and 526 Novellino 2019). There is particularly good agreement between the velocity and time series 527 products, and density/coverage of measurement points, for algorithm A and algorithm B, even 528 though the methods are dissimilar. The better agreement of the A-B pair with respect to the 529 other pairs is not due to the geometry of the Sentinel-1 data (descending). 530

However, the results are not completely identical. One of the most striking differences between
different InSAR methods is density and coverage of selected pixels. The ability to recover
measurements at a high pixel density and with wide coverage is one of the most important

534 requirements for monitoring many different sources of deformation in different conditions. This can be critical where the deforming signal is very local and occurs in non-urban areas that 535 lack man-made structures. The main reason for the difference in point density is the different 536 537 methodologies used for processing and the criteria that are used for selecting the pixels (Table 1 and Figure 12). In general, those methods that take advantage of both PS and DS, and benefit 538 from making all possible interferograms (e.g. A and D) are more successful at extracting the 539 540 maximum information (density and/or coverage) from the SAR stack. Note, however, that due to the short baseline of the Sentinel-1 interferograms, some DS pixels can remain coherent in 541 542 a single-master interferogram network and would be identified as PS pixels in PSI processing methods such as StaMPS, where phase correlation rather than amplitude is used to identify PS 543 (Hooper et al. 2007). Fully connected networks including interferograms with long temporal 544 545 baselines may suffer from fewer selected pixels compared to the networks with only short 546 baselines. Moreover, the temporal range of processed SAR images has an impact on the density of measurement points. Because scattering behaviour might vary over time, the probability of 547 finding PS pixels with consistent scattering behaviour over longer periods of time reduces. 548 However, in this case, the algorithms with the longest time series have the highest density so 549 550 the time period cannot explain the differences we see. In addition, other factors that can have a major impact on the density of measurements, include the temporal sampling of signal, the 551 552 temporal range of processed data, the configuration of the interferometric network, whether 553 oversampling of the original images is applied, the approach for side-lobe cancelation, and the specific thresholds chosen for an acceptable signal-to-noise ratio (SNR). 554

The sampling rate for Sentinel-1 is such that the single-look pixel spacing is finer than the resolution, and some scatterers can result in more than one PS pixel. In such cases the InSAR algorithm should ensure that any extraneous PSs from a single scatterer are pruned. It should also be noted that time series methods selecting DS pixels can introduce data redundancy when 559 spatial filtering is used to select SHPs. Therefore, given a homogeneous area, the selected points may show identical scattering behaviours. In this case, InSAR algorithms should provide 560 end-users the resolution of the DS pixels and inform them that those measurements do not 561 correspond to that specific point. Techniques that use pixels with intermittent coherence (Biggs 562 et al. 2007; Cigna and Sowter 2017; Sowter et al. 2013) can be more successful in terms of 563 spatial coverage and density, particularly in non-urban areas; however, these methods tend to 564 565 use a high multi-looking factor to improve coherence, and hence the density of observations is often lower. 566

567 Although a high density of measurement points is a desired outcome for an InSAR product, striking a balance between the quality and density of selected pixels is challenging. A higher 568 density can be obtained by not rejecting pixels with higher noise values. The interferometric 569 processing strategy (e.g. the use of a single master or multiple master images) and the 570 methodology of time series filtering/smoothing also have an impact on the level of noise in the 571 572 final results. Decisions may need to be made on a case-by-case basis, depending on the 573 application and the expected magnitude of the deformation signals. Using methodologies that provide both high density and high quality observations of deformation is a key priority for any 574 575 national/international ground motion map.

There might be some systematic effects in difference maps, which are mainly due to different approaches to dealing with long wavelength trends and atmospheric phase screens (APS). In order to evaluate impact of de-trending the products before inter-comparison, we also carried out a comparison with de-trended products from algorithm C. We repeated the calculation of statistics for the velocity and time series inter-comparison described in the section 5-1 and 5-2 for pairs A-C and C-D. The results show that, in this case, de-trending has a negligible impact on the consistency of products from C with those A/D (compare supplementary Figures 1-4 with Figures 6,7,9 and 10). This is likely due to the small size of our polygons (maximum 1.2 km²) which are not significantly affected by the long wavelength trend. It would be appropriate
to de-trend data in inter-comparison activities for large case studies.

586 Different geocoded coordinates for the common selected pixels is another discrepancy between 587 the InSAR products. Overlaying InSAR data on an accurate base map or ortho-rectified aerial 588 photograph and/or using corner reflectors can be solutions for correcting the geocoding shifts. 589 Although a linearly varying shift would probably provide more accurate geocoding corrections, 590 in general, a constant shift is assumed for inter-comparison purposes (Raucoules et al. 2009). 591 Geocoding error correction improves the agreement between different data sets significantly.

InSAR products are different in terms of some qualitative indicators. Spatial resolution is one of the most relevant metrics and ranges from the original high resolution sampling of Sentinel-1 (14.1 m and 2.3 m in azimuth and slant range direction, respectively) to lower resolutions that depend on the associated multi-look factors selected during processing. Some methods can provide both high and low resolutions to be used for different applications (e.g. rural and urban environments); this may be useful for national/international ground motion services.

We also note that ground motion maps are dynamic products, with velocities changing over time. Consequently, the frequency of update and latency period (time delay between acquisition and the update) should be defined by the InSAR algorithms that deliver the product. An appropriate update and latency period should be defined as part of any commissioning process. Some applications, such as hazard monitoring, benefit from rapid updates.

Any future national or international ground motion service using Sentinel-1 InSAR will need to instigate a validation process to ensure data meet minimum standards and are consistent across borders. We propose that this is done in an open and transparent fashion. A single test region, or network of test sites for various applications, should be identified that includes a

range of deformation and land cover types, and bidders should submit their analyses for this 607 region as part of any commissioning process. The results should be open and accessible via an 608 online repository so that InSAR algorithms benefit from understanding how their analyses 609 610 differ from others and so that all can improve their offerings. One of our major challenges in this research, was that different Sentinel-1 images (ascending vs descending; different dates) 611 were processed by the InSAR algorithms, limiting our ability to conduct a fair comparison 612 613 between all approaches. We suggest that in the future a comparison exercise should be repeated periodically and that in each case the time period, acquisition dates and acquisition geometry 614 615 should be explicitly specified.

In our analysis, the InSAR algorithms produced results over Glasgow to establish a baseline 616 prior to the geothermal exploitation. The deforming areas in Glasgow were very local, and the 617 scattering conditions in the deforming areas were not ideal. In addition, the "truth", estimated 618 through an independent measurement method, was unknown and therefore validation of the 619 620 InSAR products using external measurements was impossible. Test sites should be carefully selected and should cover a range of different deformation types. Independent data should be 621 collected, for example, from dense permanent networks of GNSS and levelling measurements. 622 623 Corner reflectors may be useful for testing geolocation and for providing measurement points with high SNR, and it may be appropriate to process data from a very-high-resolution satellite 624 system such as TerraSAR-X for additional validation of Sentinel-1 results. 625

626 7 Conclusions

In this research, we present an InSAR inter-comparison method, which 1) builds on the Terrafirma Validation Project and 2) addresses the limitations of previously proposed approaches up to now. We tested our method using 5 InSAR time series products including conventional PSI and advanced joint PS and DS InSAR, applied to Sentinel-1 images. We 631 selected an inter-comparison site in Glasgow, for which we had access to multiple InSAR data, and defined three polygons covering urban, rural, and deforming features. It is clear from our 632 results that different InSAR methods detect the same general deformation features, but they are 633 not identical in terms of different metrics. We propose different indicators, which are divided 634 into quantitative metrics e.g. density and coverage of measurement points and qualitative 635 metrics e.g. spatial resolution. Based on our comparison results, we suggest some 636 recommendations, which might be useful for any future nationwide/international InSAR 637 product and validation activities. 638

639 Acknowledgments

640 Benchmarking and Inter-Comparison of Sentinel-1 InSAR velocities and time series is part of

641 Digital Environment project funded by NERC NE/S016104/1. COMET is the NERC Centre

642 for the Observation and Modelling of Earthquakes, Volcanoes and Tectonics, a partnership

between UK Universities and the British Geological Survey. Sentinel-1 data were obtained via

the Copernicus Program of ESA. RapidSAR data were processed by SatSense Ltd. SqueeSAR

data were processed by TRE-Altamira Ltd. GAMMA-IPTA data were processed by British

646 Geological Survey (BGS). The authors are grateful to Dr. A. Ferretti and the anonymous

647 reviewers for comments that improved the manuscript.

648 **References**

Adam, N., Kampes, B., & Eineder, M. (2005). Development of a scientific permanent scatterer system:
Modifications for mixed ERS/ENVISAT time series. In, *Envisat & ERS Symposium, ESA*

651

Adam, N., Parizzi, A., Eineder, M., & Crosetto, M. (2009). Practical persistent scatterer processing
validation in the course of the Terrafirma project. *Journal of Applied Geophysics*, 69, 59-65,
doi:10.1016/j.jappgeo.2009.07.002.

655

Adam, N., Gonzalez, F.R., Parizzi, A., & Brcic, R. (2013). Wide area persistent scatterer

657 interferometry: current developments, algorithms and examples. In, IEEE International Geoscience
 658 and Remote Sensing Symposium-IGARSS.

- Ansari, H., De Zan, F., Bamler, R.J.I.T.o.G., & Sensing, R. (2017). Sequential estimator: Toward
 efficient InSAR time series analysis. *IEEE Transactions on Geoscience Remote Sensing*, 55, 5637-5652,
 doi: 10.1109/TGRS.2017.2711037.
- Ansari, H., De Zan, F., Parizzi, A. (2020). Study of Systematic Bias in Measuring Surface Deformation
 With SAR Interferometry, IEEE Transactions on Geoscience and Remote Sensing, doi:
 10.1109/TGRS.2020.3003421.
- Bamler, R., & Hartl, P. (1998). Synthetic aperture radar interferometry. *Inverse Problems*, *14*, R1-R54,
 doi:10.1088/0266-5611/14/4/001.
- 668
- 669 Bateson, L., & Novellino, A. (2019). Open Report: Glasgow Geothermal Energy Research Field Site -
- 670 Ground motion survey report In: British Geological Survey,
- 671 Available in:http://nora.nerc.ac.uk/id/eprint/524555/1/OR18054.pdf
- 672
- Berardino, P., Fornaro, G., Lanari, R., & Sansosti, E. (2002). A new algorithm for surface deformation
 monitoring based on small baseline differential SAR interferograms. *IEEE Transactions on Geoscience and Remote Sensing*, 40, 2375-2383, doi:10.1109/TGRS.2002.803792.
- Biggs, J., Wright, T., Lu, Z., & Parsons, B. (2007). Multi-interferogram method for measuring
 interseismic deformation: Denali Fault, Alaska. *Geophysical Journal International*, *170*, 1165-1179,
 doi:10.1111/j.1365-246X.2007.03415.x
- Brcic, R., Parizzi, A., Rodriguez Gonzalez, F., & Duro, J. (2014). Technical Report: WP1500:WAP
 Comparison Plan. European Space Agency. ESRIN/Contract No. 4000109669/13/I-AM
- 683

- Capes, R., Marsh, S., Bateson, L., Novali, F., & Cooksley, G. (2009). Terrafirma User Guide: A guide
 to the use and understanding of Persistent Scatterer Interferometry in the detection and monitoring of
 terrain-motion. European Space Agency.
- 687 Available in: https://www.researchgate.net/publication/310799583_Terrafirma_User_Guide 688
- Cigna, F., & Sowter, A. (2017). The relationship between intermittent coherence and precision of
 ISBAS InSAR ground motion velocities: ERS-1/2 case studies in the UK. *Remote Sensing of Environment*, 202, 177-198, doi:10.1016/j.rse.2017.05.016.
- 692 Crosetto, M., Agudo, M., Capes, R., & Marsh, S. (2007a). GMES Terrafirma: Validation of PSI for
 693 users: Results of the Provence inter-comparison. In, *Proceedings of the Envisat Symposium, ESA* (pp.
 694 23-27).
 695
- 696 Crosetto, M., Agudo, M., Raucoules, D., Bourgine, B., de Michele, M., Le Cozannet, G., Bremmer, C.,
 697 Veldkamp, J., Tragheim, D., & Bateson, L. (2007b). Validation of Persistent Scatterers Interferometry
 698 over a mining test site: results of the PSIC4 project. In, *Envisat Symposium,ESA* (pp. 23-27).
 699
- Crosetto, M., Biescas, E., Duro, J., Closa, J., & Arnaud, A. (2008a). Generation of advanced ERS and
 Envisat interferometric SAR products using the stable point network technique. *Photogrammetric Engineering & Remote Sensing*, *74*, 443-450, doi: 10.14358/PERS.74.4.443.
- 703
 704 Crosetto, M., Monserrat, O., Bremmer, C., Hanssen, R., Capes, R., & Marsh, S. (2008b). Ground motion
 705 monitoring using SAR interferometry: Quality assessment. *European Geologist*, 26, 12-15
- 706
- Crosetto, M., Solari, L., Mróz, M., Balasis-Levinsen, J., Casagli, N., Frei, M., Oyen, A., Moldestad,
 D.A., Bateson, L., Guerrieri, L., & Comerci, V. (2020). The Evolution of Wide-Area DInSAR: From
 Regional and National Services to the European Ground Motion Service. *Remote Sensing*, 12, 2043,
 doi:10.3390/rs12122043.
- 711

- De Zan, F., & Monti Guarnieri, A. (2006). TOPSAR: Terrain Observation by Progressive Scans. *IEEE Transactions on Geoscience and Remote Sensing*, 44, 2352-2360, doi:10.1109/TGRS.2006.873853.
- 714

Emardson, TR., Simons, M., & Webb, FH. (2003). Neutral atmospheric delay in interferometric
synthetic aperture radar applications: Statistical description and mitigation. *Journal of Geophysical*

- 717 Research: Solid Earth, 108, doi: 10.1029/2002JB001781.
- Ferretti, A., Fumagalli, A., Novali, F., Prati, C., Rocca, F., & Rucci, A. (2011). A New Algorithm for
 Processing Interferometric Data-Stacks: SqueeSAR. *IEEE Transactions on Geoscience and Remote Sensing*, 49, 3460-3470, doi:10.1109/TGRS.2011.2124465.
- 722
- Ferretti, A., Prati, C., & Rocca, F. (2000). Nonlinear subsidence rate estimation using permanent
 scatterers in differential SAR interferometry. *IEEE Transactions on Geoscience and Remote Sensing*,
 38, 2202-2212, doi:10.1109/36.868878.
- Ferretti, A., Prati, C., & Rocca, F. (2001). Permanent scatterers in SAR interferometry. *IEEE Transactions on Geoscience and Remote Sensing*, *39*, 8-20, doi:10.1109/36.898661.
- Ferretti, A., Savio, G., Barzaghi, R., Borghi, A., Musazzi, S., Novali, F., Prati, C., & Rocca, F. (2007).
 Submillimeter accuracy of InSAR time series: Experimental validation. *IEEE Transactions on Geoscience and Remote Sensing*, 45, 1142-1153, doi:10.1109/TGRS.2007.894440.
- Fornaro, G., Verde, S., Reale, D., & Pauciullo, A. (2015). CAESAR: An Approach Based on Covariance
 Matrix Decomposition to Improve Multibaseline-Multitemporal Interferometric SAR Processing. *IEEE Transactions on Geoscience and Remote Sensing*, *53*, 2050-2065, doi:10.1109/TGRS.2014.2352853.
- 737
- Gabriel, A.K., Goldstein, R.M., & Zebker, H.A. (1989). Mapping small elevation changes over large
 areas: Differential radar interferometry. *Journal of Geophysical Research: Solid Earth*, *94*, 9183-9191,
 doi:10.1029/JB094iB07p09183.
- Hanssen, R.F. (2001). *Radar interferometry: Data interpretation and error analysis*. Springer
 Netherlands.
- 744
- Heimlich, C., Gourmelen, N., Masson, F., Schmittbuhl, J., Kim, S., & Azzola, J. (2015). Uplift around
 the geothermal power plant of Landau (Germany) as observed by InSAR monitoring. *Geothermal Energy 3,2,*doi:10.1186/s40517-014-0024-y.
- Hooper, A. (2008). A multi-temporal InSAR method incorporating both persistent scatterer and small
 baseline approaches. *Geophysical Research Letters*, *35*, doi:10.1029/2008GL034654.
- Hooper, A., Bekaert, D., Spaans, K., & Arikan, M. (2012). Recent advances in SAR interferometry time
 series analysis for measuring crustal deformation. *Tectonophysics*, *514 517*, 1-13,
 doi:10.1016/j.tecto.2011.10.013.
- Hooper, A., Segall, P., & Zebker, H. (2007). Persistent scatterer interferometric synthetic aperture radar
 for crustal deformation analysis, with application to Volcán Alcedo, Galápagos. *Journal of Geophysical Research: Solid Earth*, *112*, *B7*, doi:10.1029/2006JB004763.
- 759
- Jordan, C., Bateson, L., & Novellino, A. (2019). Environmental baseline monitoring for shale-gas
 development: Insights for monitoring ground motion using InSAR analysis. *Science of the total Environment*, 696, 134075, doi:10.1016/j.scitotenv.2019.134075.
- 763
- Kampes, B. (2005). Deformation parameter estimation using permanent scatterer interferometry. In:Delft University of Technology, Delft, The Netherlands.
- 766

- 767 Lanari, R., Mora, O., Manunta, M., Mallorqui, J.J., Berardino, P., & Sansosti, E. (2004). A small-768 baseline approach for investigating deformations on full-resolution differential SAR interferograms. Geoscience Remote 1377-1386, 769 IEEE **Transactions** on and Sensing, 42, 770 doi:10.1109/TGRS.2004.828196.
- Mirzaee, S., Motagh, M., Akbari, B., Wetzel, H.U., & Roessner, S. (2017). Evaluating three insar timeseries methods to assess creep motion, case study: Masouleh landslide in north Iran. *ISPRS Annals of the Photogrammetry, Remote Sensing and Spatial Information Sciences, IV-1-W1*, 223-228,
 doi:10.5194/isprs-annals-IV-1-W1-223-2017.
- Monti-Guarnieri, A., & Tebaldini, S. (2008). On the Exploitation of Target Statistics for SAR
 Interferometry Applications. IEEE Transactions on Geoscience and Remote Sensing, 46, 3436-3443,
 doi:10.1109/TGRS.2008.2001756.
- 780

776

771

- Mora, O., Mallorqui, J.J., & Broquetas, A. (2003). Linear and nonlinear terrain deformation maps from
 a reduced set of interferometric SAR images. *IEEE Transactions on Geoscience and Remote Sensing*, *41*, 2243-2253, doi:10.1109/TGRS.2003.814657.
- Osmanoğlu, B., Sunar, F., Wdowinski, S., & Cabral-Cano, E. (2016). Time series analysis of InSAR
 data: Methods and trends. *ISPRS Journal of Photogrammetry and Remote Sensing*, *115*, 90-102,
 doi:10.1016/j.isprsjprs.2015.10.003.
- Pepe, A., & Calo, F. (2017). A review of interferometric synthetic aperture RADAR (InSAR) multitrack approaches for the retrieval of Earth's surface displacements. *Applied Sciences*, 7, 12, 1264, doi:10.3390/app7121264
- Raucoules, D., Bourgine, B., de Michele, M., Le Cozannet, G., Closset, L., Bremmer, C., Veldkamp,
 H., Tragheim, D., Bateson, L., Crosetto, M., Agudo, M., & Engdahl, M. (2009). Validation and
 intercomparison of Persistent Scatterers Interferometry: PSIC4 project results. *Journal of Applied Geophysics*, 68, 3, 335-347, doi:10.1016/j.jappgeo.2009.02.003.
- Sadeghi, Z., Valadan Zoej, M., Hooper, A., & Lopez Sanchez, J. (2018). A New Polarimetric
 Persistent Scatterer Interferometry Method Using Temporal Coherence Optimization. *IEEE Transactions on Geoscience and Remote Sensing*, 56, 6547-6555, doi:10.1109/TGRS.2018.2840423.
- Schmidt, D.A., & Bürgmann, R. (2003). Time-dependent land uplift and subsidence in the Santa Clara
 valley, California, from a large interferometric synthetic aperture radar data set. *Journal of Geophysical Research: Solid Earth, 108, B9*, doi:10.1029/2002JB002267.
- Shanker, P., Casu, F., Zebker, H.A., & Lanari, R. (2011). Comparison of Persistent Scatterers and Small
 Baseline Time-Series InSAR Results: A Case Study of the San Francisco Bay Area. *IEEE Geoscience and Remote Sensing Letters*, 8, 592-596, doi:10.1109/LGRS.2010.2095829.
- Sousa, J.J., Hooper, A.J., Hanssen, R.F., Bastos, L.C., & Ruiz, A.M. (2011). Persistent Scatterer InSAR:
 A comparison of methodologies based on a model of temporal deformation vs. spatial correlation
 selection criteria. *Remote Sensing of Environment*, *115*, 2652-2663, doi:10.1016/j.rse.2011.05.021.
- 813

805

797

- Sowter, A., Bateson, L., Strange, P., Ambrose, K., & Syafiudin, M.F. (2013). DInSAR estimation of
 land motion using intermittent coherence with application to the South Derbyshire and Leicestershire
 coalfields. *Remote Sensing Letters*, *4*, 979-987, doi:10.1080/2150704X.2013.823673.
- 817

<sup>Spaans, K., & Hooper, A. (2016). InSAR processing for volcano monitoring and other near-real time
applications.</sup> *Journal of Geophysical Research: Solid Earth*, 121, 2947-2960,
doi:10.1002/2015JB012752.

- Van der Kooij, M., Hughes, W., Sato, S., & Poncos, V. (2005). Coherent target monitoring at high
 spatial density: examples of validation results. In, *Fringe*, *ESA*.
- 824
- Wang, H., Wright, T.J., & Biggs, J. (2009). Interseismic slip rate of the northwestern Xianshuihe fault
 from InSAR data. *Geophysical Research Letters*, 36, doi: 10.1029/2008GL036560.
- 827
- 828 Werner, C., Wegmuller, U., Strozzi, T., & Wiesmann, A. (2003). Interferometric point target analysis
- 829 for deformation mapping. In, *EEE International Geoscience and Remote Sensing Symposium-IGARSS*
- 830

.

831