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1 **Substantial increases in Eastern Amazon and Cerrado biomass**
2 **burning-sourced tropospheric ozone**

3
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19 **Abstract:**

20 The decline in Amazonian deforestation rates and biomass burning (BB) activity (2001-2012) has
21 been shown to reduce air pollutant emissions (e.g. aerosols) and improve regional air quality (AQ).
22 However, in the Cerrado region (savannah grass-lands in north-eastern Brazil) satellite observations
23 reveal increases in fire activity and tropospheric column nitrogen dioxide (an ozone precursor)
24 during the burning season (August-October, 2005-2016), which have partially offset these AQ
25 benefits. Simulations from a 3-D global chemistry transport model (CTM) capture this increase in
26 NO₂ with a surface increase of ~1 ppbv/decade. As there are limited long-term observational

27 tropospheric ozone records, we utilise the well-evaluated CTM to investigate changes in ozone.
28 Here, the CTM suggests that Cerrado region surface ozone is increasing by ~10 ppbv/decade. If left
29 unmitigated, these positive fire-sourced ozone trends will substantially increase the regional health
30 risks and impacts from expected future enhancements in South American BB activity under climate
31 change.

32

33 **1. Introduction**

34 Fire is a widely used method for land clearance in the tropics, allowing rapid conversion of regions of
35 natural vegetation to agricultural land (Cochrane, 2003). In the Amazon, these deforestation
36 practices led to a 15-20% decrease in forest cover between 1976 and 2010 (Hansen et al., 2013;
37 Aragão et al., 2014; Davidson et al., 2012). Vegetation fires are a large source of reactive trace gases
38 and aerosol to the Amazon atmosphere (Hodgson et al., 2018, Sena and Artaxo, 2015; Bela et al.,
39 2015; Ward et al., 1992), especially during the dry season, with substantial impacts on the
40 atmospheric radiation balance (Sena and Artaxo, 2015; Scott et al., 2018; Thornhill et al., 2018;
41 Kolusu et al., 2015) and surface air quality (Johnston et al., 2012; Jacobson et al., 2014; Reddington et
42 al., 2015; Pacifico et al., 2013; Artaxo et al., 2013).

43

44 Over the period 2001 to 2012, deforestation rates in Brazil decreased by ~40% (Hansen et al., 2013)
45 leading to improvements in particulate air quality (Reddington et al., 2015). These decreasing
46 deforestation rates correlated with reductions in satellite-observed aerosol optical depth (AOD) over
47 the Amazon, which were attributed to significant reductions in surface emissions of aerosol from
48 deforestation fires (Reddington et al., 2015). Reductions in surface PM_{2.5} (mass of particulate matter
49 with diameter less than 2.5 µm) concentration, resulting from the reduction in fire emissions, were
50 estimated to prevent 400-1700 premature deaths annually in South America. The reduction in fire
51 activity is particularly evident in the arc of deforestation, around the southern edge of the Amazon
52 Basin, where the majority of Amazon deforestation fires occur (van der Werf et al., 2017). Satellite-

53 observed fire-burned area (FBA) in this region has decreased by 0.25-0.5 %/year between 1998 and
54 2015 (Andela et al., 2017). Despite these reductions in the arc of deforestation, FBA in north-eastern
55 Brazil, has increased by 0.5-1.0% per year, resulting in increased regional pollutant emissions (Andela
56 et al., 2017; Chen et al., 2013). This region represents the transition biome between the eastern
57 Amazon, the Cerrado grasslands and the Caatinga desert vegetation (Santo et al., 2017). However,
58 this region is predominantly covered by savannah grasslands and hereafter we refer to it as the
59 “Cerrado Region” (i.e. the Cerrado biome in Figure 1a & b and the boarding regions of the Caatinga
60 and Amazonia biomes).

61

62 As a major source of ozone precursors (nitrogen oxides (NO_x) and volatile organic compounds
63 (VOCs)), vegetation fires have been observed to lead to enhancement in tropospheric ozone
64 concentrations (Bela et al., 2017; Artaxo et al., 2013; Jaffe and Wigder, 2012; Kirchhoff and Marinho,
65 1994; Kirchhoff et al., 1996). Tropospheric ozone is an important air pollutant, which causes adverse
66 effects on human health (Jerrett et al., 2012; Doherty et al., 2017), crops (Hollaway et al., 2012; Van
67 Dingenen et al., 2009) and natural vegetation (Sitch et al., 2007). Deforestation fires tend to be large
68 in scale (due to extensive above-ground biomass) with a greater amount of smouldering, while
69 savannah fires tend to be smaller and burn at higher temperatures with more flaming combustion
70 (Hodgson et al., 2018; Alencar et al., 2015; Longo et al., 2009). Consequently, these different fire
71 characteristics can result in very different emissions. For instance, emission factors (i.e. mass of trace
72 gas/aerosol emitted per mass of dry matter burnt) for deforestation fires, release larger quantities of
73 primary aerosol emissions (13.0 gkg^{-1}) than savannah fires (8.5 gkg^{-1}), while the converse is the case
74 for NO_x emissions (deforestation emission factor = 2.55 gkg^{-1} and savannah emission factor = 3.9 gkg^{-1})
75 ¹) (Akagi et al., 2011). Therefore, understanding how changes in fire activity impact NO_x and ozone
76 pollution requires knowledge of the underlying vegetation type and emission characteristics.

77

78 It is only recently, in the satellite era, that space-borne observations from various missions have
79 made it possible to study long-term changes in global tropospheric composition. Here, we present
80 the first study using these long-term (2005-2016) records, in conjunction with a chemistry transport
81 model (CTM), to investigate the impact of changing wildfire/biomass burning (BB) activity on
82 tropospheric ozone air quality across the Amazon during the BB season (August-September-October,
83 ASO). We exploit satellite measurements of tropospheric nitrogen dioxide (NO_2) and fire activity, as
84 well as new state-of-the-art lower tropospheric ozone retrievals. Our focus is on detecting and
85 analysing trends in tropospheric/surface ozone and its precursors to investigate the impact on
86 regional AQ. Section 2 introduces the satellite measurements and CTM used in this study, section 3
87 presents our results and the conclusions are discussed in section 4.

88

89 **2. Methods and Data:**

90 ***2.1. Satellite Data***

91 We use satellite measurements of tropospheric column NO_2 (TCNO_2) and sub-column (0-6 km) ozone
92 (SCO_3) from the Ozone Monitoring Instrument (OMI) from 2005-2016. OMI is a nadir viewing
93 instrument on-board NASA's Aura satellite (2004-present) (Boersma et al., 2007). The OMI TCNO_2
94 data was downloaded as Level 2 swath data from the Tropospheric Emissions Monitoring Internet
95 Service (TEMIS, www.temis.nl). The SCO_3 data was provided by the Rutherford Appleton laboratory
96 (RAL), which uses an optimal estimation technique (Miles et al., 2015) to produce a state-of-the-art
97 product with peak vertical sensitivity in the lower troposphere. Aura is polar orbiting with an
98 approximate local overpass time of 13.30 and OMI has nadir-viewing spectral ranges of 270-500 nm
99 (Boersma et al., 2011). All data sets have been quality controlled for geometric cloud fraction less
100 than 0.2, good quality flags and the OMI row anomalies (Braak, 2010) where applicable. Detailed
101 analysis of TCNO_2 retrieval uncertainties is provided by Boersma et al., (2004), while Boussez
102 (2014) quantify the impact of biomass burning aerosol uncertainty on retrieved TCNO_2 . Miles et al.,
103 (2015) provide error/uncertainty analysis on the retrieval scheme used for the RAL OMI SCO_3

104 product. These satellite data sets were mapped onto high-resolution spatial grids of $0.05^\circ \times 0.05^\circ$ for
105 TCNO₂ and $1.0^\circ \times 1.0^\circ$ for SCO₃ (Pope et al., 2018). Global Ozone Monitoring Experiment – 2 (GOME-
106 2) TCNO₂ data was also acquired from TEMIS as a Level 3 monthly mean gridded product for August-
107 September-October (ASO) 2007-2016, to confirm that the Cerrado Region TCNO₂ positive trends
108 (**Figure 2a**) were realistic and not an artefact of the OMI row anomaly.

109

110 We use two different satellite-derived fire activity datasets to investigate regional patterns in fires
111 and their trends over time. These are fire-burned-area (FBA) from the Global Fire Emissions
112 Database (GFED vn4.0 (van der Werf et al., 2017) and fire radiative power (FRP) from the Global Fire
113 Assimilation System (GFAS vn1.2; Kaiser et al., 2012). Both fire products are derived from Moderate
114 Resolution Imaging Spectroradiometer (MODIS) measurements, on-board NASA's Aqua and Terra
115 satellites (Remer et al., 2005) and are used to produce gas-phase and aerosol emissions from fires
116 through application of emissions factors (Wooster et al., 2018). While GFAS only provides FRP and
117 total emissions, GFED provides information on different vegetation types burned, including
118 contributions from deforestation and savannah fires.

119

120 **2.2. TOMCAT Model Setup and Evaluation**

121 The TOMCAT global off-line chemical transport model (CTM) (Chipperfield, 2006) is forced by
122 ECMWF ERA-Interim reanalysis meteorology (Dee et al., 2011) and has a horizontal spatial resolution
123 of $2.8^\circ \times 2.8^\circ$ with 31 vertical levels up to 10 hPa. The model includes detailed tropospheric
124 chemistry, including 229 gas-phase reactions and 82 advected tracers (Monks et al., 2017), and
125 heterogeneous chemistry driven by size-resolved aerosol from the GLOMAP module (Mann et al.,
126 2010). Simulations used here include anthropogenic emissions from the Coupled Model
127 Intercomparison Project Phase 6 (CMIP6) (Hosely et al., 2018) and fire emissions from GFED vn4.0
128 (van der Werf et al., 2017). Biogenic Volatile Organic Compounds (VOCs) emissions are from the
129 Chemistry-Climate Model Initiative (CCMI) (Morgenstern et al., 2017).

130

131 TOMCAT surface/tropospheric ozone was evaluated against surface observations from Manaus
132 (60.2°W, 2.6°S) in the Amazon (2010-2011). During the BB season (ASO), the model successfully
133 captures peak observed surface ozone concentrations of 13-15 ppbv. However, between January
134 and May there is systematic positive bias of ~5 ppbv (although this is within the observational
135 variability). Comparisons with aircraft observations from the the South AMerican Biomass Burning
136 Analysis (SAMBBA) campaign (Darbyshire et al., 2019) (September-October, 2012) show the model
137 successfully reproduces the boundary layer vertical ozone profile between 30-45 ppbv, with a slight
138 positive bias of 2-3 ppbv. These observations and comparison are also consistent with ozonesondes
139 from Natal (2007-2008). More details on the model evaluation are located in the **SM 2**.

140

141 Overall, the TOMCAT model successfully captures the Amazon ozone seasonality and absolute
142 concentrations in the lower troposphere, giving some confidence in the ozone simulations used to
143 investigate long-term changes in surface ozone in the following analysis.

144

145 **3. Results**

146 ***3.1. Satellite fire emission signals and trend detection***

147 The period of peak fire activity over the Amazon occurs during ASO (Hodgson et al., 2018; van der
148 Werf et al., 2017; Kaiser et al., 2012), when FBA (**Figure 1a**) and FRP (**Figure 1c**) reach over 10% and
149 100 mW/m², respectively. In contrast, during the non-BB season (February-March-April, FMA), there
150 are minimal fire signals, peaking at <1% (**Figure 1b**) and 10 mW/m² (**Figure 1d**). Substantial fire-
151 related TCNO₂ signals (2.5-4.0×10¹⁵ molecules/cm², **Figure 1e**) are present across the Amazon during
152 ASO, while in FMA TCNO₂ concentrations are less than 1.0 ×10¹⁵ molecules/cm² (**Figure 1f**). During
153 ASO peak TCNO₂ is located over the Cerrado region where savannah-type fires (both natural and
154 anthropogenic) burn at high temperatures (Hodgson et al., 2018) releasing large quantities of NO_x
155 (Akagi et al., 2011). There will also be contributions from deforestation and agricultural fires (e.g. de

156 Araújo et al., 2019). In comparison, TCNO₂ hotspots over large cities (e.g. São Paulo and Rio De
157 Janeiro) show limited seasonality, remaining above 5.0×10^{15} molecules/cm² year-round. The SCO₃
158 ASO signal is more homogeneous, consistent with ozone being a secondary NO_x-induced pollutant
159 formed downwind of source regions. A clear SCO₃ enhancement (22-25 Dobson units, DU) is present
160 during ASO over the Amazon compared with FMA (12-18 DU). SCO₃ over the South Atlantic is also
161 elevated (over 22 DU) in ASO as a result of ozone-enriched outflow from the southern African
162 biomass burning (Moxim and Levy, 2000).

163

164 A long-term trend analysis (**Figure 2**) shows significant increases in both fire activity and TCNO₂
165 across the Cerrado Region. Trends are calculated using a linear least-squares fit over the ASO
166 composition average for each year between 2005 and 2016, where we remove extreme
167 drought/anomalous fire years (e.g. 2005, 2007, 2010 and 2012; hereafter defined as ED-AF). Though
168 the 2015 intense positive El Niño-Southern Oscillation (ENSO) event substantially enhanced drought
169 conditions in the Central Amazon (Anderson et al., 2018), it had a limited impact on fire activity over
170 our primary region of interest (black box in Figure 1a – see **SM 1**). Hence, 2015 was not defined as
171 ED-AF (see **SM 1**). Here, we follow the approach of Reddington et al., (2015) and remove ED-AF years
172 from our analysis as they represent different emission regimes when compared to the normal state
173 (Aragão et al., 2014). For aerosol emissions, Reddington et al., (2015) found that ED-AF years were
174 1.5-2.8 times greater than normal years, while for NO_x emissions we find a factor increase of 14-15.
175 This avoids skewing of the long-term trends due to larger than usual fire activity and tropospheric
176 pollutant loading in those years. Black polygon-outlined regions in **Figure 2** show significant trends at
177 and above the 90% confidence level ($>90\%CL, trend/\sigma_{trend} > 1.645$). Artificial background TCNO₂
178 trends (i.e. OMI row anomaly; Braak, 2010) have been removed from the OMI time series (**SM 1**).
179 Where significant TCNO₂ trends exist, there are insignificant trends in the stratospheric slant NO₂
180 column and in the tropospheric air mass factor.

181

182 Significant positive trends in FBA (**Figure 2a**), FRP (**Figure 2b**) and TCNO₂ (**Figure 2c**) range between
183 0.3-0.7 %/year, 2-5 mW/m²/year and 0.1-0.15 ×10¹⁵ molecules/cm²/year, respectively, potentially
184 driven by changes in Cerrado savannah-type fire activity. Given that all three data sets have spatially
185 consistent trends, it provides us with confidence that the detected TCNO₂ trends are related to fire
186 activity. In the arc of deforestation region, significant (>90% CL) negative trends are also found in all
187 three datasets, consistent with previous studies (Reddington et al., 2012), which identified decreases
188 in deforestation fires and satellite-observed AOD between 2001 and 2012. Though the positive FBA
189 and TCNO₂ signals are spatially less extensive than those of the FRP and AOD negative trends.

190

191 As the significant fire activity and TCNO₂ positive trends over the Cerrado Region are spatially
192 sporadic and disjointed, regional trends (**Figure 2d**) in FBA and TCNO₂ were determined from the
193 black-outlined region in **Figure 2a**. The variability (standard deviation) in the regional FBA and TCNO₂
194 typically ranges between 150-190% and 40-60% of the mean values, respectively, in non ED-AF
195 years. Here, there are significant regional trends (99% CL) in both FBA and TCNO₂. The TCNO₂
196 regional trend from the Global Ozone Monitoring Experiment – 2 (GOME-2) satellite instrument also
197 increases significantly over time (2007-2016), supporting the OMI TCNO₂ results. By sub-sampling
198 OMI TCNO₂ under pixels with FBA >1.0%, concentrations are larger by approximately 0.5 ×10¹⁵
199 molecules/cm², highlighting a similar significant positive trend (red dashed line). This strongly
200 suggests that increasing regional TCNO₂ concentrations are being driven by increased fire activity as
201 TCNO₂ is larger in fire-classified pixels with a consistent significant positive trend.

202

203 Analysis of trends in GFED NO_x emissions, which are based on FBA, for ASO indicates that increases
204 in GFED-defined savannah-type fires are likely responsible for the increases in TCNO₂ over the
205 Cerrado Region. **Figure 3a** shows the GFED NO_x emissions trends for all fire types for 2005-2016 with
206 ED-AF years removed, and shows significant positive trends (0.05-0.15 g/m²/year) over the Cerrado
207 Region, consistent with the FBA, FRP and TCNO₂. When split into fire types, agricultural fires (**Figure**

208 **3d)** show negligible trends, while deforestation fires (Figure 3c) give spatially incoherent trends
209 across the Amazon (partial spatial agreement with savannah fires over portions of the Cerrado
210 Region e.g. around 50°W, 12°S suggesting a range of different vegetation fire-types driving the
211 positive trends seen in **Figure 2** and **3**). However, the significant positive NO_x emission trends (0.05-
212 0.1 g/m²/year) from savannah fires (**Figure 3b**) are closely related to the magnitude and spatial
213 pattern of the trends in **Figure 3a**. Though it should be noted that the GFED fire-type classifications
214 can be subject to uncertainties (van der Werf et al., 2017). Missed or false fire detections in will also
215 introduce further emission uncertainties (e.g. particulate emissions; Reddington et al., (2016)).

216

217 **3.2. Model simulated NO₂ and ozone**

218 TOMCAT model simulations (2005-2016, **Figure 4**), using annual varying GFED vn4.0 fire emissions,
219 are used to investigate the impact of increased fire activity (predominantly savannah-type fires) on
220 air quality over the Cerrado region, and the wider surrounding Amazon area, given the limited
221 observational spatial and temporal coverage. Neither the surface or aircraft data provide long-term
222 records and the ozonesondes are from Natal (a coastal city). The OMI SCO₃ observations provide
223 valuable seasonal information over the wider Amazon, but are not sensitive enough to detect finer
224 scale changes (e.g. for the Cerrado region) in ozone with time. However, evaluation of TOMCAT
225 tropospheric ozone using these observations (surface sites, aircraft data and ozonesondes; **see SM**
226 **2**) show that the model is able to capture the surface ozone seasonal cycle and reproduce the lower
227 tropospheric vertical profile.

228

229 Over the 2005-2016 period, fire-sourced ozone contributes approximately an extra 10-15 ppbv to
230 Amazon surface ozone concentrations during the BB season. **Figure 4a** shows the difference
231 between the TOMCAT “fire-on” and “fire-off” simulations, where fires have substantially contributed
232 to the surface ozone budget in ASO, consistent with the ozone seasonality in the observations.
233 During ASO, the 2005-2016 average surface concentrations range between 40-50 ppbv in the “fire-

234 on” simulation, while decreasing to under 20 ppbv when fire sources are switched off (“fire-off”, see
235 **SM 3**). The red and blue dashed regions in **Figure 4d** represent the Eastern Amazon and Wider
236 Amazon, respectively, and highlight regions where we have detected long-term fire-sourced ozone
237 enhancements. During ASO, the domain average fire-source ozone contribution ranges between 10-
238 15 ppbv and 5-11 ppbv for the Eastern Amazon and Wider Amazon regions (**Figure 4b**). This implies
239 that approximately 50-60% of the ASO surface ozone is from fire sources (see **SM 3**). The ASO
240 Eastern Amazon ozone contributions tends to be larger as more NO_x is emitted from the
241 predominantly savannah-type fire regimes (i.e. **Figure 3b**), when compared with the Wider Amazon
242 region.

243
244 Trends in model surface ASO mean NO₂ show significant (>90%CL) increases (>0.1 ppbv/year) over
245 the Cerrado region and decreases (-0.03 to -0.01 ppbv/year) in the arc of deforestation (**Figure 4a**).
246 As expected, the short NO₂ lifetime means that these trends are highly correlated with those in
247 GFED v4.0 NO_x fire emissions (**Figure 3a**). The significant increase in Cerrado fire-related NO_x, in the
248 presence of VOC concentrations (both biogenic emissions from vegetation and emissions from fires),
249 yields significant increases in simulated ozone (over 1.0 ppbv/year, **Figure 4b**). However, insignificant
250 decreases (-0.2 to 0.0 ppbv/year) and increases (0.1-0.3 ppbv/year) occur over the arc of
251 deforestation and western Amazon, respectively. Model-simulated eastern Brazil surface NO₂ and
252 ozone (same region in **Figure 2a, SM 3**) also show significant regional positive trends (>90% CL).

253

254 **4. Discussion and Conclusions**

255 Overall, our results show that there has been a significant increase in fire activity in the Cerrado
256 region (savannah grasslands) of North-eastern Brazil, between 2005 and 2016, yielding substantial
257 increases in fire-related tropospheric pollutants (i.e. NO₂ and ozone). This is important as it
258 highlights different behaviour compared with the “arc of deforestation” region (the south-western
259 Amazon reaching to the north-eastern flank), where a decline in deforestation fires has yielded

260 lower particulate matter emissions and reduced the corresponding health risks (Reddington et al.,
261 2015). As there are limited observations of surface/tropospheric ozone in the Amazon, well-
262 evaluated model simulations offer a method to quantify long-term changes in surface ozone
263 concentrations. In the Cerrado region, TOMCAT simulations suggest considerable long-term
264 increases in fire-sourced surface ozone (~1.0 ppbv/year equating to ~10-12 ppbv regionally over the
265 study period), predominantly from increased burning of savannah grasslands in the Cerrado Region,
266 but also with contributions from deforestation fires.

267 Should these fire activities continue to intensify, as is currently seen in Brazil (TerraBrasilis, 2019;
268 NASA, 2019), the health risks and socioeconomic impacts associated with fire-sourced ozone
269 pollution may increase substantially. This will be confounded by the high probability of more
270 frequent and intense drought BB from future climate change (Page et al., 2017) and land-use change
271 (Fonseca et al., 2019), as well as expected increases in the South American population (United
272 Nations, 2017). The enhanced ozone will also likely further damage vegetation and reduce
273 photosynthesis (Sitch et al., 2007; Pacifico et al., 2015) leading to reductions in crop yields (Hollaway
274 et al., 2012). These effects combined could have substantial impacts on natural vegetation,
275 agriculture, and public health, with potential degradation in ecosystem services and economic
276 losses. Therefore, targeted policies on controlling Cerrado biomass burning would be beneficial to
277 generate local and regional air quality improvements with associated public and ecological health
278 benefits.

279

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289 Rutherford Appleton Laboratory (RAL) OMI SCO₃ data. We also acknowledge the use of fire activity
290 data from the Global Fire Emissions Database (GFED), which were obtained from
291 <https://www.globalfiredata.org/index.html>. FRP data comes from the ECMWF CAMS Global Fire
292 Assimilation System, which can be found at <http://apps.ecmwf.int/datasets/data/cams-gfas/>.
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303 files.

304

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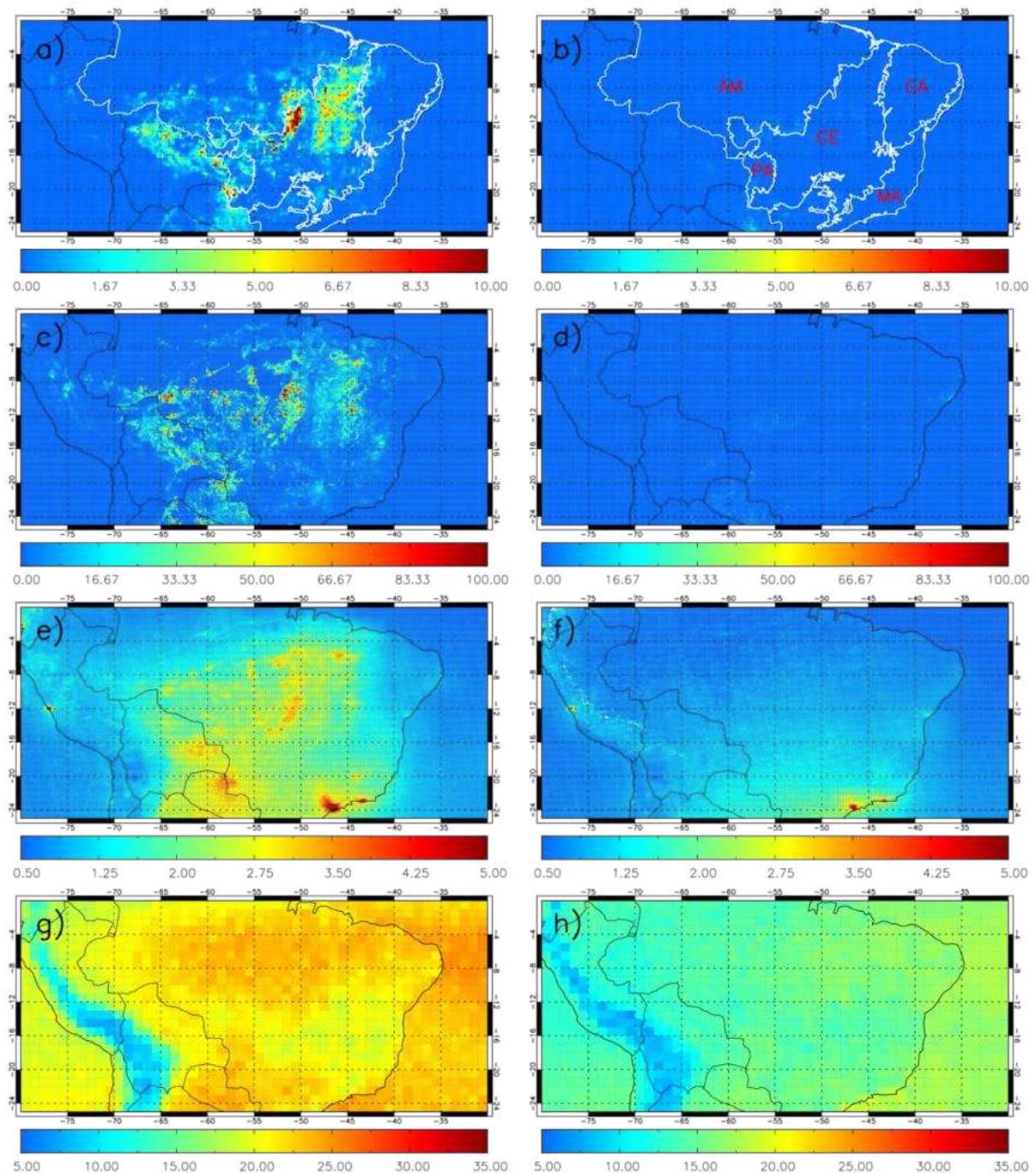
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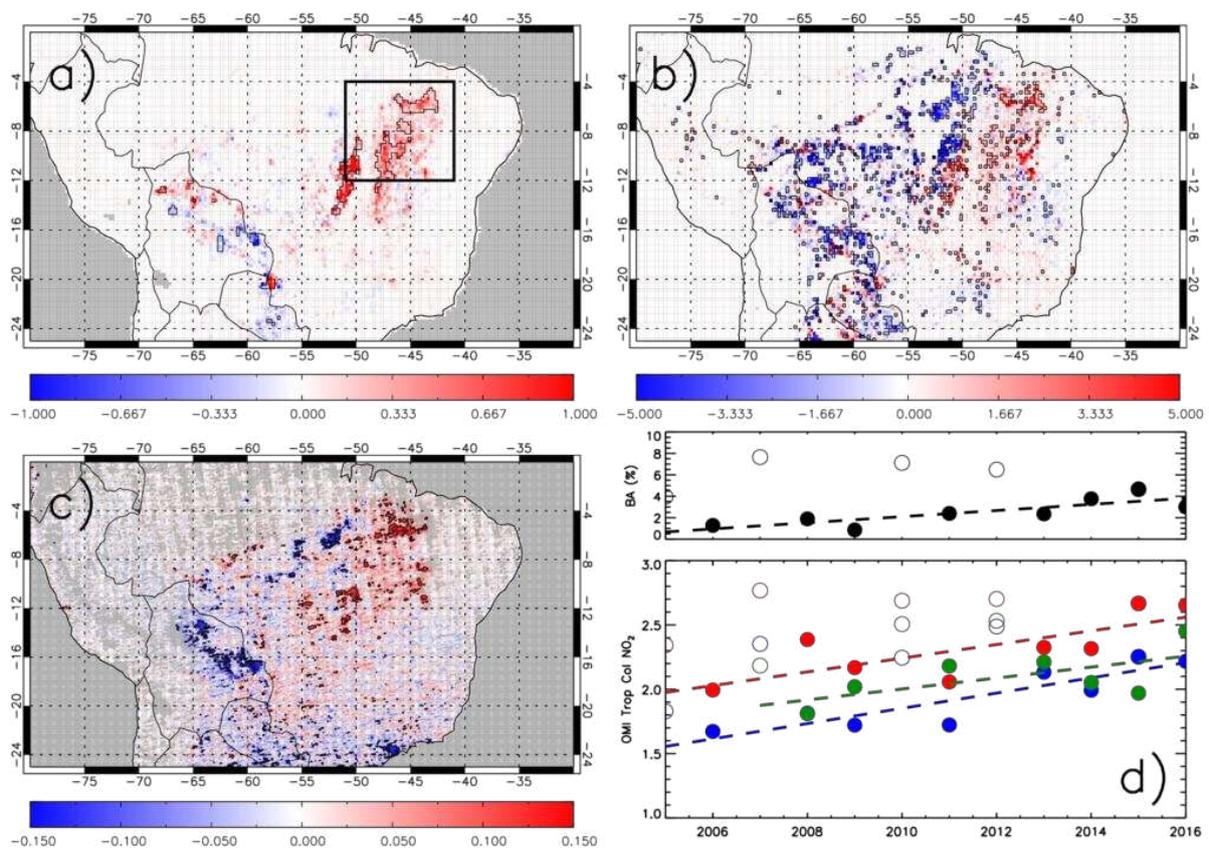
549 **Figures:**



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551 **Figure 1:** Mean satellite fire activity and tropospheric composition (2005-2016) for August-
 552 September-October (ASO; panels a, c, e and g) and February-March-April (FMA; panels b, d, f and h).
 553 Panels a) and b) show GFED Fire-Burned Area (FBA, %), panels c) and d) show GFAS Fire Radiative
 554 Power (FRP, mW/m²), panels e) and f) show OMI tropospheric column NO₂ (TCNO₂ - 10¹⁵
 555 molecules/cm²) and panels g) and h) show OMI sub-column (0-6 km) ozone (Dobson units – DU). The
 556 white-outlined regions in Figure 1a & b are the Brazilian biomes (Global Forest Watch, 2019) with

557 their corresponding labels in red (AM=Amazonia, CA=Caatinga, CE=Cerrado, MA=Mata Atlantica and
 558 PA=Pantanal).

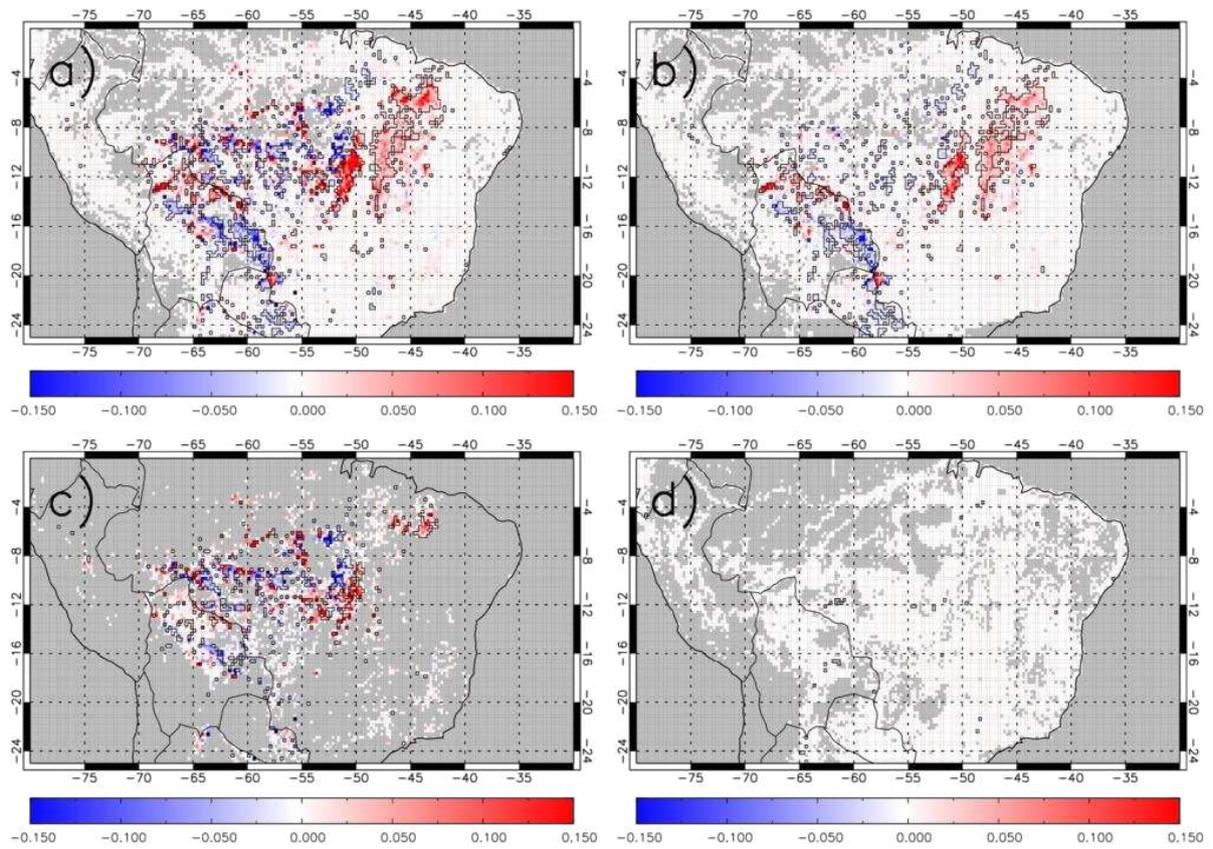


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560 **Figure 2:** Linear trends in the burning-season (ASO) average for each year, between 2005 and 2016,
 561 for a) GFED FBA area fraction (%/year), b) GFAS FRP ($\text{mW}/\text{m}^2/\text{year}$) and c) OMI TCNO_2 (10^{15}
 562 $\text{molecules}/\text{cm}^2/\text{year}$). Black polygon-outlined regions highlight significant trends at and above the
 563 90% confidence level (>90%CL). Panel d) shows significant (>90% CL) regional trends (black box – 2a)
 564 in FBA (%/year), OMI TCNO_2 (10^{15} $\text{molecules}/\text{cm}^2/\text{year}$), GOME-2 TCNO_2 (10^{15} $\text{molecules}/\text{cm}^2/\text{year}$)
 565 and OMI TCNO_2 sampled under FBA area pixels of >1% (10^{15} $\text{molecules}/\text{cm}^2/\text{year}$) represented by
 566 the black, blue, green and red dashed lines, respectively. Open circles represent extreme
 567 drought/anomalous fire (ED-AF) years, and are not included in trend analysis.

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571 **Figure 3:** Trends in GFED NO_x (g/m²/year) ASO total emissions between 2005 and 2016 for a) all fires,

572 b) savannah fires, c) deforestation fires and d) agricultural fires. Black polygon-outlined regions

573 indicate significant trends above the 90% CL.

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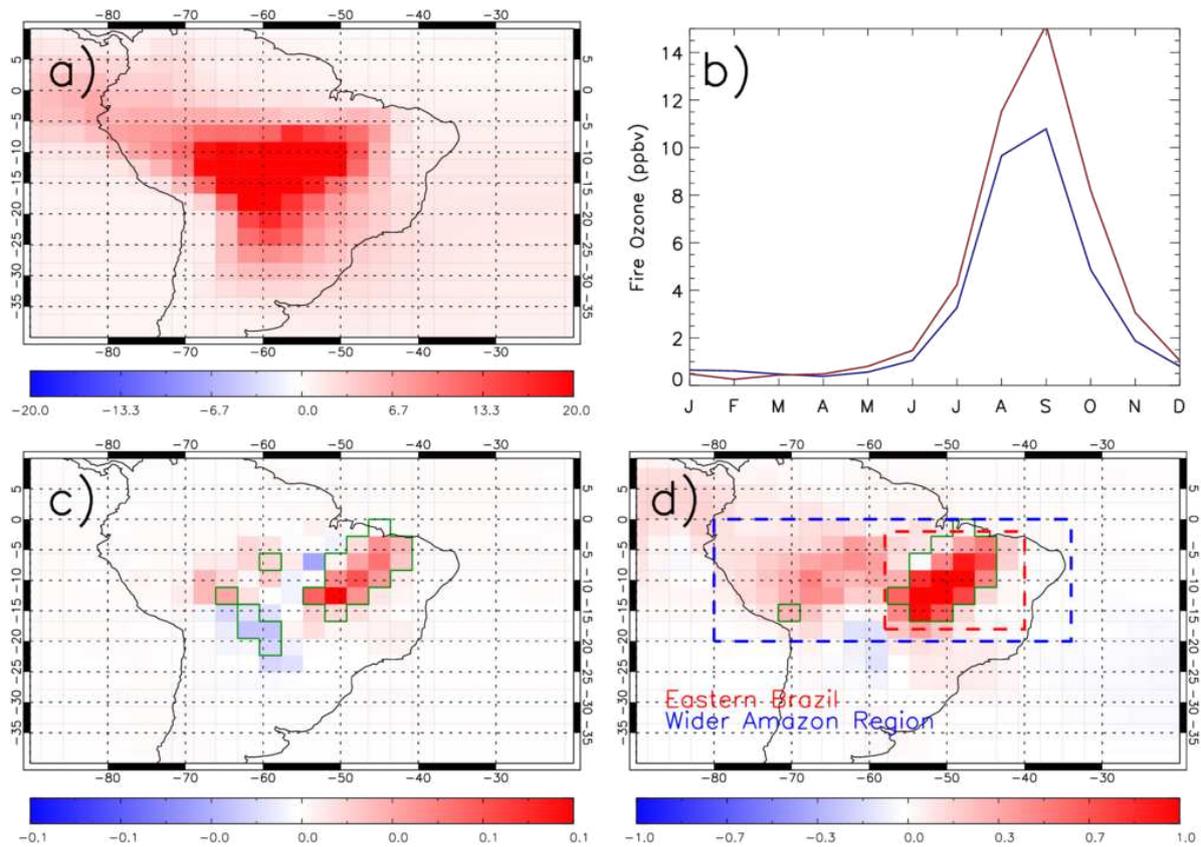
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581 **Figure 4:** a) Average (ASO, 2005-2016) contribution of fire-sourced ozone to Amazon surface ozone

582 concentrations (ppbv). b) Seasonal cycle in surface fire-sourced ozone (ppbv) in the Eastern Brazil

583 and Wider Amazon regions (red and blue dashed regions in panel d). Trends in TOMCAT model

584 surface c) NO₂ (ppbv/year) and d) ozone (ppbv/year) for the ASO average between 2005 and 2016.

585 Green polygon-outlined areas show regions of significant trends above the 90% CL.