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Zhuo, L. orcid.org/0000-0002-5719-5342 and Han, D. (2017) Multi-source hydrological soil moisture state estimation using data fusion optimisation. *Hydrology and Earth System Sciences*, 21 (7). pp. 3267-3285. ISSN 1027-5606

<https://doi.org/10.5194/hess-21-3267-2017>

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Multi-source hydrological soil moisture state estimation using data fusion optimisation

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Received: 13 September 2016 – Discussion started: 14 October 2016

Revised: 17 May 2017 – Accepted: 3 June 2017 – Published: 4 July 2017

Abstract. Reliable estimation of hydrological soil moisture state is of critical importance in operational hydrology to improve the flood prediction and hydrological cycle description. Although there have been a number of soil moisture products, they cannot be directly used in hydrological modelling. This paper attempts for the first time to build a soil moisture product directly applicable to hydrology using multiple data sources retrieved from SAC-SMA (soil moisture), MODIS (land surface temperature), and SMOS (multi-angle brightness temperatures in H–V polarisations). The simple yet effective local linear regression model is applied for the data fusion purpose in the Pontiac catchment. Four schemes according to temporal availabilities of the data sources are developed, which are pre-assessed and best selected by using the well-proven feature selection algorithm gamma test. The hydrological accuracy of the produced soil moisture data is evaluated against the Xinanjiang hydrological model's soil moisture deficit simulation. The result shows that a superior performance is obtained from the scheme with the data inputs from all sources ($NSE = 0.912$, $r = 0.960$, $RMSE = 0.007$ m). Additionally, the final daily-available hydrological soil moisture product significantly increases the Nash–Sutcliffe efficiency by almost 50 % in comparison with the two most popular soil moisture products. The proposed method could be easily applied to other catchments and fields with high confidence. The misconception between the hydrological soil moisture state variable and the real-world soil moisture content, and the potential to build a global routine hydrological soil moisture product are discussed.

1 Introduction

Soil moisture is a key element in the hydrological cycle, regulating evapotranspiration, precipitation infiltration, and overland flow (Wanders et al., 2014). For hydrological applications, the antecedent wetness condition of a catchment is among the most significant factors for accurate flow generation processes (Berthet et al., 2009; Matgen et al., 2012a). Norbiato et al. (2008) reported that initial wetness conditions are essential for efficient flash flood alerts. Additionally, an operational system requires reliable hydrological soil moisture state updates to reduce the time-drift problem (Aubert et al., 2003; Berg and Mulroy, 2006; Dumedah and Coulibaly, 2013). However, currently there is no available soil moisture product that can be used directly in hydrology modelling, primarily because soil moisture is difficult to define and there is no single shared meaning in various disciplines (Romano, 2014).

Although there have been many soil moisture measuring projects (e.g., satellite missions such as advanced scatterometer (ASCAT), soil moisture and ocean salinity (SMOS), and Soil Moisture Active Passive (SMAP); ground-based networks such as Soil Climate Analysis Network (SCAN), U.S. Surface Climate Observing Reference Networks, and COsmic-ray Soil Moisture Observing System), they are not sufficiently used in hydrology due to the following reasons: (1) misconception between the hydrological soil moisture state variable and the real-field soil moisture content (Zhuo and Han, 2016a); (2) unawareness of data availability and strength/weakness of different data sources; (3) the existing soil moisture products are mainly evaluated against point-based ground soil moisture observations or airborne retrievals, which have significant spatial mismatch (both horizontally and vertically) to catchment-scales, and are there-

fore less applicable to hydrological modelling (Pierdicca et al., 2013); (4) underutilisation of multiple data sources (e.g., multi-angle raw observations by satellite sensors).

Some studies have attempted to directly utilise the existing soil moisture products (i.e., data from satellites, land surface models, and in situ methods directly) for flood prediction improvement, for example, Brocca et al. (2010) explored that utilising the soil water index from ASCAT sensor could improve runoff prediction mainly if the initial catchment wetness conditions were unknown; Aubert et al. (2003) assimilated in situ soil moisture observations into a simple rainfall–runoff model and acquired better flow prediction performance; Javelle et al. (2010) suggested that estimations of antecedent soil moisture conditions were useful in improving flash flood forecasts at ungauged catchments; contrarily, the Chen et al. (2011) study showed assimilating ground-based soil moisture observations was generally unsuccessful in enhancing flow prediction; Matgen et al. (2012b) revealed that satellite soil moisture products added little or no extra value for hydrological modelling. Clearly those results are rather mixed. Challenges remain in integrating soil moisture estimated outside the hydrological field into hydrological models. We believe if a hydrologically directly applicable soil moisture product could be produced, the aforementioned studies' results would be significantly improved.

Therefore, the aims of this paper are to clarify the aforementioned misconception between the hydrological model's soil moisture state and the real-world soil moisture, assess the data availabilities for direct hydrological soil moisture state estimation, and fuse those available data sources using a hydrologically relevant approach. It is hoped that the final product has a superior hydrological compatibility over the existing soil moisture products. To achieve these aims, the Xinanjiang (XAJ) (Zhao, 1992) operational rainfall–runoff model is used as a target to simulate flow and soil moisture state information (i.e., soil moisture deficit, SMD) for the Pontiac catchment in the central United States (U.S.). The reason for adopting XAJ is explained in the following section. For the purpose of hydrological soil moisture state estimation, it is effective to adopt the data-driven method, which can map multiple data sources into the desired dataset without computational burden. In this study the local linear regression (LLR) model is used. The multiple data sources applied in this study include the SMOS (Kerr et al., 2010) multi-angle brightness temperatures (T_{bs}) with both horizontal (H) and vertical (V) polarisations, the moderate resolution imaging spectroradiometer (MODIS) (Wan, 2008) land surface temperature, and the soil moisture product by SAC-SMA (Xia et al., 2014). The detail explanations of those datasets are covered in the methodology section. A well-proven feature selection algorithm gamma test (GT) (Stefánsson et al., 1997; Zhuo et al., 2016b) is employed to pre-assess the selected data inputs and find the optimal combination of them for soil moisture state calculation. In addition, an M -test (Remesan et al., 2008) is adopted to explore the best size of the training

data. The desired soil moisture product is trained and tested by the XAJ SMD simulation. In total four data-input schemes are developed according to the temporal availability of the selected data inputs, which are then combined to give a daily hydrological soil moisture product.

Compared with previous work, our study contains the following new elements: (i) a hydrologically directly usable soil moisture product is proposed; (ii) the GT and LLR techniques are used for the first time in a data fusion of multiple data sources for hydrological soil moisture state estimation; (iii) the use of multiple data sources is useful, which allows data users to analyse the availability of the different products and compare the relative benefits of them.

2 Material and methods

2.1 Study area

In this study, the Pontiac catchment (1500 km², Fig. 1) is used for the calibration and the validation of the XAJ model. The Pontiac (40.878° N, 88.636° W) lies on the north-flowing Vermilion River, which is a tributary of the Illinois River of the state of Illinois, U.S. The worst flood in this area occurred on 4 December 1982, cresting at 5.84 m above mean sea level (m.s.l.), and the most recent flood occurred on 9 January 2008, cresting at 5.75 m m.s.l.; therefore, this catchment is likely located within a winter-flooding region. Pontiac is covered with moderate canopy (the annual mean normalised difference vegetation index retrieved from the MODIS satellite is around 0.4), and when compared with a densely vegetated catchment, it has more accurate soil moisture estimations from satellites (Al-Bitar et al., 2012). Based on the Köppen–Geiger climate classification, this medium-sized catchment is dominated mainly by hot summer continental climate (Peel et al., 2007). With reference to the University of Maryland Department Global Land Cover Classification, it is used primarily for agriculture purposes (Bartholomé and Belward, 2005; Hansen et al., 2000). The soil mostly consists of Mollisols, which has deep and high organic matter, and the nutrient-enriched surface soil is typically between 60 and 80 cm in depth (Webb et al., 2000). The study period is from January 2010 to December 2011. The reason for using this 2-year period of data is because there have been many data gaps from 2013 to 2017, and the data quality in 2012 was poor. As a result, only the data in 2010–2011 are consistent and of high quality. As pointed out by Liu and Han (2010), “Traditionally, hydrologists use rules of thumb to select a certain period of hydrological data to calibrate the models (i.e., 6-year data)”. However, their study has shown “the information content of the calibration data is more important than the data length; thus, 6-month data may provide more useful information than longer data series”. Therefore, the 2 years of high quality data adopted in the study are better than a longer period of poor quality data.

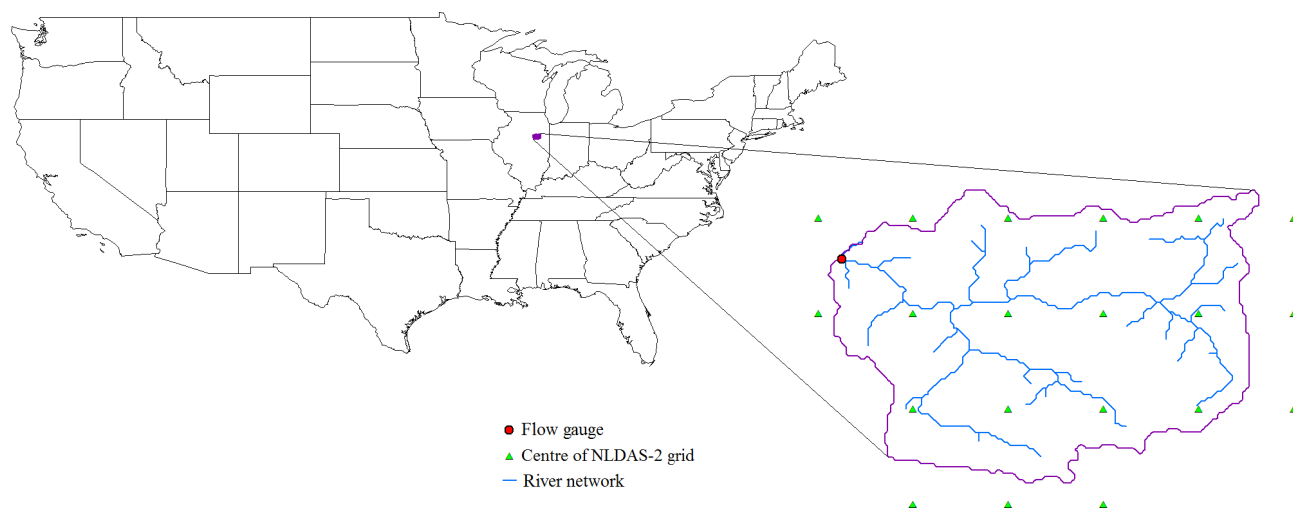


Figure 1. The location and river network of the Pontiac catchment in the U.S., with the flow gauge and NLDAS-2 central grid points (Zhuo et al., 2015a).

The North American Land Data Assimilation System 2 (NLDAS-2) (Mitchell et al., 2004) provides precipitation and potential evapotranspiration information to run the XAJ model. Both data forces are at 0.125° spatial resolution and have been converted to daily temporal resolution. In order to use those distributed forcing into the lumped XAJ model, both forcing have been interpolated with the area-weighted average method instead of the more complicated Kriging approach, because the latter could produce errors if not well controlled (Wanders et al., 2014). The average annual rainfall depth is about 954 mm, and the average annual potential evapotranspiration is approximately 1670 mm. The daily observed flow data are acquired from the U.S. Geological Survey.

2.2 Hydrological model

The XAJ hydrological model is used for the simulation of SMD and river flow at a daily time step. It is a simple lumped rainfall–runoff model with many applications performed in world-wide catchments (Chen et al., 2013; Gan et al., 1997; Shi et al., 2011; Zhao, 1992; Zhao and Liu, 1995; Zhuo et al., 2016a, 2015b). Since XAJ can obtain rather effective flow modelling performances and requires only two meteorological forcing (precipitation and potential evapotranspiration) inputs (Peng et al., 2002), it is used more widely than the more complicated semi-distributed/fully distributed hydrological models for operational applications.

As shown in Fig. 2, the XAJ model has three main components: evapotranspiration, runoff generation, and runoff routing. XAJ consists of soil layers (upper, lower and deep) in its evapotranspiration calculations. Because XAJ adopts the multi-bucket variable-size method in its modelling concept, it has unfixed soil depths, which is more effective than the fixed depths models (Beven, 2012). Other widely used mod-

els such as PDM (Moore, 2007), VIC (Liang et al., 1994), and ARNO (Todini, 1996) also follow this concept.

In XAJ, the three-layer soil moisture state variables are all calculated as SMD, which is an important soil-wetness variable in hydrology. SMD is defined as the amount of water to be added to a soil profile to bring it to the field capacity (Calder et al., 1983; Rushton et al., 2006). In this study, only the surface SMD (i.e., top layer) referring to the vegetation and the very thin topsoil, is utilised as a hydrological soil moisture target. This is because the water held in the top few centimetres of the soil has been widely recognised as a key variable associated with water fluxes (Eltahir, 1998; Entekhabi and Rodriguez-Iturbe, 1994). Moreover, the current satellite technology is only capable of acquiring the Earth information from the outermost layer of the soil. Therefore, as a case study based on the XAJ model, we only focus on the surface soil moisture state investigation here. Future research will focus on the root-zone soil moisture product development by using a similar method proposed in this study.

In this study, a modified version of the XAJ model is adopted, and interested readers are referred to Zhuo and Han (2016b) for more details. All the XAJ's 17 parameters are used during the model calibration, which are shown in Table 1. In this study, the genetic algorithm (Wang, 1991) is used for parameter optimisation. Based on the genetic algorithm result, minor trial and error adjustments to the parameters EX, *B*, WUM, WLM, and WDM are also carried out to obtain the best model performance (Chen and Adams, 2006). The calibration and the validation results (during January 2010–April 2011 and May 2011 to December 2011, respectively) of the XAJ model can be found in Zhuo et al. (2015a). Discussion regarding the river flow and SMD simulation results in this catchment have been published in Zhuo and Han (2016b), with Nash–Sutcliffe efficiency (NSE)

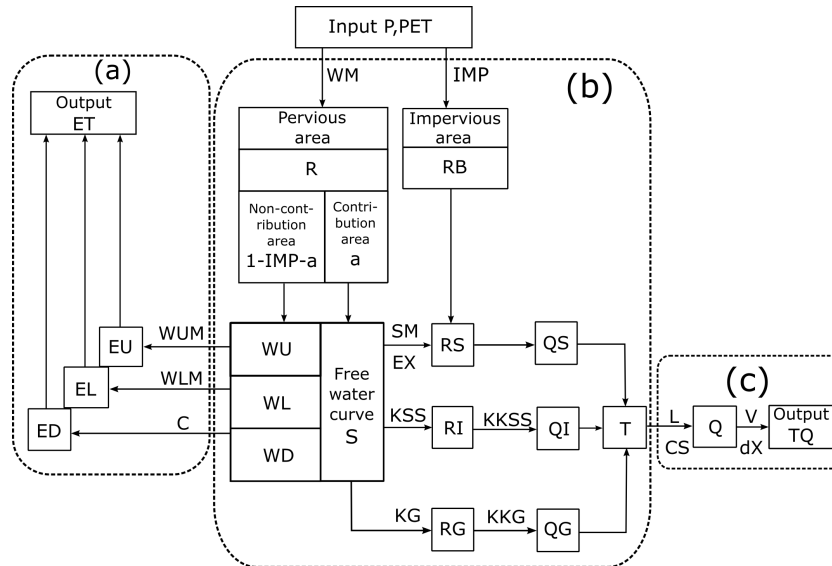


Figure 2. Adopted flowchart of the XAJ model (Zhao, 1992). The model consists of an evapotranspiration component (a), a runoff generating component (b), and a runoff routing component (c). P, PET, and ET are the precipitation, potential evapotranspiration, and the simulated actual evapotranspiration respectively; WU, WL and WD represent the upper, lower, and deep soil layers' areal mean-tension water storage respectively; WM is the areal mean field capacity; EU, EL, and ED stand for the upper, lower, and deep soil layers' evapotranspiration output respectively; S is the areal mean free water storage; a is the portion of the sub-catchment producing runoff; IMP is the factor of impervious area in a catchment; RB is the direct runoff produced from the small portion of impervious area; R is the total runoff generated from the model with surface runoff (RS), interflow (RI), and groundwater runoff (RG) components respectively. These three runoff components are then transferred into QS, QI, and QG and combined as the total sub-catchment inflow (T) to the channel network. The flow outputs Q from each sub-catchment are then routed to the catchment outlet to produce the final flow result (TQ). The rest of the symbols are explained in Table 1.

Table 1. The XAJ model parameters used in the Pontiac catchment.

Symbol	Model parameters	Unit	Range
<i>K</i>	Ratio of evapotranspiration	[-]	0.10–1.20
WUM	The areal mean field capacity of the upper layer	mm	30–50
WLM	The areal mean field capacity of the lower layer	mm	20–150
WDM	The areal mean field capacity of the deep layer	mm	30–400
IMP	Percentage of impervious and saturated areas in the catchment	%	0.00–0.10
<i>B</i>	Exponential parameter with a single parabolic curve, which represents the non-uniformity of the spatial distribution of the soil moisture storage capacity over the catchment	[-]	0.10–0.90
<i>C</i>	Coefficient of the deep layer that depends on the proportion of the catchment area covered by vegetation with deep roots	[-]	0.10–0.70
SM	Areal mean free water capacity, which represents the maximum possible deficit of free water storage	mm	10–50
KG	Outflow coefficient of the free water storage to groundwater relationships	[-]	0.10–0.70
KSS	Outflow coefficient of the free water storage to interflow relationships	[-]	0.10–0.70
EX	Exponent of the free water capacity curve	[-]	1.10–2.00
KKG	Recession constant of the groundwater storage	[-]	0.01–0.99
KKSS	Recession constant of the lower interflow storage	[-]	0.01–0.99
CS	Recession constant in the lag and route method for routing through the channel system with each sub-catchment	[-]	0.10–0.70
<i>L</i>	Lag in time	[-]	0.00–6.00
<i>V</i>	Parameter of the Muskingum method	$m s^{-1}$	0.40–1.20
<i>dX</i>	Parameter of the Muskingum method	[-]	0.00–0.40

Table 2. General data-input properties relevant for this study.

	SMOS- T_{bs}	MODIS-LST	SAC-SMA-SM
Product	brightness temperature	land surface temperature	soil moisture
Unit	Kelvin (K)	Kelvin (K)	$m^3 m^{-3}$
Near-real-time (NRT)	Yes	Yes	Yes
Spatial resolution (km)	35–50	5.6	14
Data time step	~ every 3 days	~ daily	Daily
Data availability for the studying period (days)	217	458	730

obtained larger than 0.80 during both the calibration and validation periods. The results are not repeated here.

2.3 Multiple data sources for hydrological soil moisture state estimation

Data sources from SMOS, MODIS, and SAC-SMA are used (Table 2). All data sources have been converted into catchment-scale datasets by the area-weighted average method. The detail description of each data source is given as follows. The main reason for choosing those three data sources is due to their near-real-time (NRT) availabilities (MODAPS Services, 2015; Rodell, 2016) (SMOS becomes available in NRT recently; ESA Earth Online, 2016), which allows for fast implementation in flood forecasting.

2.3.1 SMOS multi-angle brightness temperatures (SMOS- T_{bs})

The SMOS (1.4 GHz, L-band) level-3 T_{bs} data covering the studying period are available from the Centre Aval de Traitement des Données SMOS (CATDS) (Jacquette et al., 2010). The reason for choosing the SMOS satellite is because compared with other satellite techniques (i.e., optical, and thermal infrared), microwave bands (especially with longer wavelength such as L-band, 21 cm) can penetrate deeper into the soil (~5 cm) and have less interruptions from weather conditions (Njoku and Kong, 1977). Additionally, SMOS has a relatively longer period of data record compares with other satellite missions such as SMAP. SMOS retrieves the thermal emission from the Earth in both H and V polarisations with wide ranges of incidence angles from 0 to 60°. The observation depth of SMOS is approximately 5 cm with a spatial resolution of 35–50 km depending on the incident angle and the deviation from the satellite ground track (Kerr et al., 2012, 2010, 2001).

SMOS provides T_{bs} retrievals at all incidence angles averaged in 5° width angle bins, which have been transformed into the ground polarisation reference frame (i.e., H, and V polarisations). Therefore, the number of the SMOS- T_{bs} inputs for the hydrological soil moisture estimation can be as high as 24 (12 angle bins per polarisation), with the centre of the first angle bin at 2.5° in both polarisations (Rodriguez-

Fernandez et al., 2014). As the satellite progresses, any given location on the Earth's surface is scanned a number of times at various incidence angles, depending on the location with respect to the satellite subtrack: the further away, the fewer the angular acquisitions (Kerr et al., 2010). The data availabilities of the SMOS- T_{bs} are illustrated in Fig. 3 (the availabilities for H and V polarisations are the same). It can be seen that the data availabilities among various incidence angles are rather different. In this study the only angle range that gives the most available record of data is from 27.5 to 57.5° (i.e., 7 for H and 7 for V polarisation), which is therefore chosen for the hydrological soil moisture development. This angle range is in line with the angle selection in Rodriguez-Fernandez et al. (2014). In addition the SMOS level-3 soil moisture product from the CATDS (SMOS-SM) is also acquired for a comparison with the estimated soil moisture product. Retrievals that are potentially contaminated with radio-frequency interference have been removed. Readers are referred to Kerr et al. (2012) for a full description of the SMOS-retrieving algorithms, and Njoku and Entekhabi (1996) for good knowledge of how passive microwaves relate to soil moisture variations.

2.3.2 MODIS land surface temperature (MODIS-LST)

The MODIS/Terra (Earth Observing System AM-1 platform) (Wan, 2008) daily MOD11C1-V5 land surface temperature covering the studied period is downloaded from the Land Processes Distributed Active Archive Centre website. MODIS is chosen among other operational optical satellites for its suitable features, mostly, due to its frequent revisiting time and free NRT data availability. It measures 36 spectral bands between 0.405 and 14.385 μm , and acquires data at three spatial resolutions 250, 500, and 1000 m respectively, while the adopted MOD11C1 V5 product incorporates 0.05° (5.6 km) spatial resolution. The benefit of adding land surface temperature information is that previous studies have shown the variations in soil moisture have a strong linkage with land surface temperature (Carlson, 2007; Goward et al., 2002; Mallick et al., 2009). One reason is the changes of land surface temperature are mainly affected by albedo and diurnal heat capacity, and the diurnal heat capacity is mainly controlled by soil moisture (Price, 1980). Wan (2008) com-

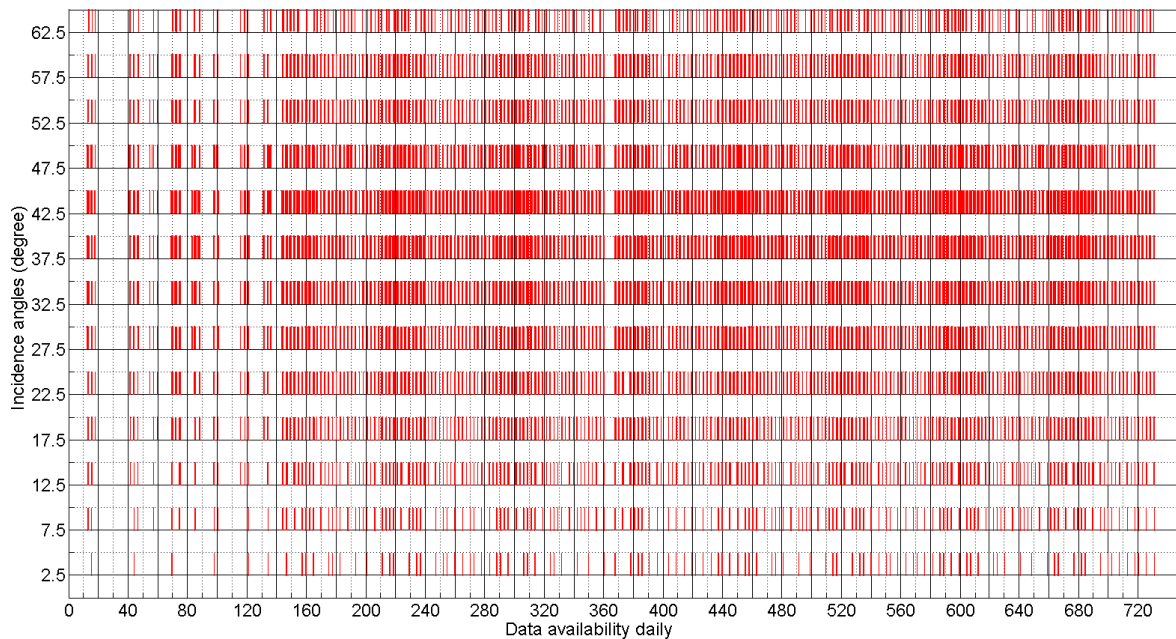


Figure 3. SMOS- T_{bs} data availabilities. It is noted that the available dates for the horizontal and the vertical polarisations are the same; therefore, only one is shown here.

pared MOD11C1-V5 land surface temperatures in 47 clear-sky cases with in situ measurement and revealed that the accuracy was better than 1 K in the range from -10 to 58 °C in about 39 cases. Cloud-contaminated data have been removed by a double-screening method, and its details can be found in Wan et al. (2002).

2.3.3 SAC-SMA soil moisture estimation (SAC-SMA-SM)

The reason for choosing the SAC-SMA land surface modelled soil moisture product is because satellites can often have missing data due to various weather and canopy conditions (e.g., rainfall, frozen weather, and vegetation coverage); therefore, this daily dataset is essential in producing a temporally completed hydrological soil moisture product. In this study, the surface soil moisture (0–10 cm) simulated from the SAC-SMA model is selected. This is because its estimated soil moisture gives a high accuracy against the observational soil moisture and a good correlation with the XAJ SMD (Zhuo et al., 2015b). The daily SAC-SMA-SM is given in a spatial resolution of 0.125° . The dataset can be download from <http://www.emc.ncep.noaa.gov/mmb/nldas/>. Readers are referred to Xia et al. (2012) for a full description of the SAC-SMA data products.

2.3.4 Data availabilities

As shown in Table 2, the availability of the three data sources is rather different. Unlike SMOS and MODIS, SAC-SMA-2 SM is a model-based product that runs in a NRT mode, and

therefore it produces valid data every day during the whole studying period. Whereas the two satellites' data are more exiguous and depend on weather and surface conditions. Compared with MODIS, the SMOS's retrieval is even sparse and the biggest data shortage normally occurs in the winter season where its returned microwave signal is mostly affected by frozen soils (Zhuo et al., 2015a). Based on the data availability analysis, the proposed hydrological soil moisture product is built from four data-input schemes as presented in Table 3. Those four schemes enable us to test and compare the estimated soil moisture state more comprehensively. Since the continuity of a soil moisture product is essential for any operational applications, SAC-SMA-SM is included in all of the schemes.

2.4 Data fusion

2.4.1 Gamma test for feature selection

Before model building, it is important to carry out a feature selection process, because it can simplify the model inputs, shorten training times, and reduce overfitting problems. In this study a proper combination of the incidence angles from the SMOS T_{bs} is vital for the best soil moisture state calculation. For this purpose, a feature selection method called GT is adopted. It has been effectively used in numerous studies for model-input selection (Durrant, 2001; Jaafar and Han, 2011; Noori et al., 2011; Remesan et al., 2008; Tsui et al., 2002; Zhuo et al., 2016b). In addition to the feature selection, GT can also give a useful indication of the underlying model complexity. It is a near-neighbour data analysis routine that

Table 3. Four data-input schemes: scheme 1: SAC-SMA-SM; scheme 2: SAC-SMA-SM and MODIS-LST; scheme 3: SAC-SMA-SM and SMOS- T_{bs} ; scheme 4: SAC-SMA-SM, MODIS-LST, and SMOS- T_{bs} .

	SAC-SMA-SM	MODIS-LST	SMOS- T_{bs}
Scheme 1	x		
Scheme 2	x	x	
Scheme 3	x		x
Scheme 4	x	x	x

determines the minimum mean-squared error (MSE) that can be achieved based on the input–output dataset utilising any continuous non-linear models (Zhuo et al., 2016b). The calculated minimum MSE is referred to as the gamma statistics and denoted as Γ . For detailed calculations about the GT algorithm, interested readers are referred to Koncar (1997), Pi and Peterson (1994), and Stefánsson et al. (1997). Here only the basic knowledge about the GT is shown:

$$\{(\mathbf{x}_i, y_i), \quad 1 \leq i \leq M\}, \tag{1}$$

where the inputs $\mathbf{x}_i \in R^m$ are vectors restricted by a closed bounded set $C \in R^m$, and their corresponding outputs $y_i \in R$ are scalars, M stands for the sample points. The outputs y are determined by the input vectors \mathbf{x} that carry predictively useful messages. The only assumption made is that their latent relationship is from the following function:

$$y = f(\mathbf{x}_1 \dots \mathbf{x}_m) + r, \tag{2}$$

where f is built up as a smooth model with r representing random noise. Without loss of generality, the assumption of r noise distribution is that its mean is always zero, because all the constant bias has been considered within the f model. Additionally, r 's variance ($\text{Var}(r)$) is restricted within a set boundary. The observations' potential model is now defined within the class of smooth functions.

The Γ is related to $N[i, k]$, which represents as the k th ($1 \leq k \leq p$) nearest neighbours of each vector \mathbf{x}_i ($1 \leq i \leq M$), written as $\mathbf{x}_{N[i,k]}$ ($1 \leq k \leq p$), where p is a fixed integer. In order to determine the gamma function from the input vectors, the delta function is used:

$$\delta_M(k) = \frac{1}{M} \sum_{i=1}^M |\mathbf{x}_{N[i,k]} - \mathbf{x}_i|^2 \quad (1 \leq k \leq p), \tag{3}$$

where the function $|\mathbf{x}_{N[i,k]} - \mathbf{x}_i|$ calculates the Euclidean distance. The gamma function for its output values is expressed as in Eq. (4), and the Γ can be determined from Eqs. (3) and (4):

$$\gamma_M(k) = \frac{1}{2M} \sum_{i=1}^M |y_{N[i,k]} - y_i|^2 \quad (1 \leq k \leq p), \tag{4}$$

where $y_{N[i,k]}$ is the corresponding output values for the k th nearest neighbours \mathbf{x}_i ($\mathbf{x}_{N[i,k]}$). To find Γ a least-squared regression line for the p points $(\delta_M(k), \gamma_M(k))$ is built using the following equation:

$$\gamma = A\delta + \Gamma, \tag{5}$$

where Γ can be determined when δ is set as zero. The detailed explanation is

$$\gamma_M(k) \rightarrow \text{Var}(r), \quad \text{when } \delta_M(k) \rightarrow 0. \tag{6}$$

Equation (5) gives us valuable information about the underlying system; not only that the Γ is a useful indicator of the optimal MSE result that any smooth functions can achieve, but also its gradient A provides guidance about the underlying model complexity (i.e., the steeper the gradient the more sophisticated the model should be adopted). In this study, the winGamma™ software is used for GT calculation (Durrant, 2001). The mathematical feasibility of GT has been published in Evans and Jones (2002).

2.4.2 M-test for training data-size selection

A common practice in non-linear modelling is to split the dataset into training and testing parts. However, there is no universal solution on how to divide the datasets (i.e., the proportion of each part) so that the best modelling results could be obtained. Here, an M -test is carried out, where M stands for the training data size. M -test is accomplished by calculating the Γ for increasing the M value (i.e., expanding the training data) and exploring the resultant graph to judge whether the Γ approaches a stable asymptote. Such an approach is straightforward and effective in finding the optimal sizes of training and testing datasets, while avoiding overfitting problems and reducing unsystematic attempts.

2.4.3 Local linear regression

Various data fusion techniques have been developed (Prakash et al., 2012; Srivastava et al., 2013; Wagner et al., 2012); however, their methods require high computational time to run and this, in a real-time flood forecasting framework, could not match the operational needs. Comparatively, the LLR model is a simpler method and requires relatively low computational time. Therefore it is chosen in order to test if a simple method is able to provide effective performance. LLR is a non-parametric regression model that has been applied in Liu et al. (2011), Pinson et al. (2008), Sun et al. (2003), and Zhuo et al. (2016b) for forecasting and smoothing purposes. LLR builds local linear regression based on the nearest points (p_{\max}) of a targeted point, and repeats such a process over the whole training dataset to produce a piecewise linear model. There are many methodologies in selecting the p_{\max} , in this study a method called influence statistics is used (Durrant, 2001; Remesan et al., 2008), which is outlined as below.

Assume there are p_{\max} nearest points, then Eq. (7) can be built:

$$\mathbf{X}m = y, \quad (7)$$

here \mathbf{X} is a $p_{\max} \times d$ matrix, which shows the d -dimensional information of p_{\max} , \mathbf{x}_i are the nearest points confined between 1 and p_{\max} , y is the output vector with p_{\max} dimension, and m is a set of parameters formed in a vector, which plays an important role in mapping the solution from \mathbf{X} to y . Therefore Eq. (7) can be expanded as

$$\begin{pmatrix} \mathbf{x}_{11} & \mathbf{x}_{12} & \mathbf{x}_{13} & \cdots & \mathbf{x}_{1d} \\ \mathbf{x}_{21} & \mathbf{x}_{22} & \mathbf{x}_{23} & \cdots & \mathbf{x}_{2d} \\ \vdots & \vdots & \vdots & \ddots & \vdots \\ \mathbf{x}_{p_{\max}1} & \mathbf{x}_{p_{\max}2} & \mathbf{x}_{p_{\max}3} & \cdots & \mathbf{x}_{p_{\max}d} \end{pmatrix} \begin{pmatrix} m_1 \\ m_2 \\ \vdots \\ m_d \end{pmatrix} = \begin{pmatrix} y_1 \\ y_2 \\ \vdots \\ y_{p_{\max}} \end{pmatrix}. \quad (8)$$

In order to solve the equation, the following two conditions are set: (a) if \mathbf{X} is square and non-singular then Eq. (7) can be simply calculated as $m = \mathbf{X}^{-1}y$; (b) if \mathbf{X} is not square or singular, Eq. (7) needs to be rearranged and m can be obtained by finding the minimum of

$$|\mathbf{X}m - y|^2 \quad (9)$$

with the distinct solution of

$$m = \mathbf{X}^{\#}y, \quad (10)$$

where $\mathbf{X}^{\#}$ is the pseudo-inverse matrix of \mathbf{X} (Penrose, 1955, 1956).

3 Results

In this section, different combinations of input data (Table 3) are adopted to examine their impacts on hydrological soil moisture estimation. XAJ SMD is used as a hydrological soil moisture state benchmark for the training and testing. More discussion about the misconception between the hydrological model's soil moisture state variable and the real-world soil moisture content is covered in Sect. 4. During GT and M -test processes, all data inputs need to be normalised so that their mean is zero and standard deviation is 0.5. This step is necessary in reducing the impacts of numerical difference from various inputs, hence improving the GT efficiency (Remesan et al., 2008). Five statistical indicators are used for the soil moisture estimation analysis: Pearson product moment correlation coefficient (r), MSE, which is the same value as the gamma statistic Γ , standard error (SE), NSE (Nash and Sutcliffe, 1970), and root mean square error (RMSE).

3.1 Scheme 1: SMD estimation using SAC-SMA-SM as input

Although in this scheme, there is no need for data feature selection because only one data input is involved, the GT is still carried out to explore the useful information about the underlying relationship between the XAJ SMD and the SAC-SMA-SM. The calculated gamma statistics are shown in Table 4. The Γ of 0.072 indicates that the optimal MSE achievable using any modelling technique is 0.072; and the small value of SE shows the precision and accuracy of the GT result. Γ is a significant target value in the M -test to find the most suitable training data size. As presented in Fig. 4a, when more training data (i.e., M increases in steps of one) is used the Γ changes dramatically. Eventually at $M = 292$, Γ starts to stabilise around 0.072. The M -test allows us to confidently apply the first 292 datasets to build a model of a given quality, in the sense of predicting with a MSE around the asymptotic level. The corresponding gamma gradient (A) suggests the complexity of the underlying system: the larger the A value is the more complex the system is. For example if A is significantly large, a more complicated model like a support vector machine might be required, but $A = 1.353$ in scheme 1 is small (Remesan et al., 2008); therefore, a LLR model should be able to simulate the system. For LLR modelling, its complexity level is controlled by the p_{\max} parameter. As illustrated in Fig. 5, p_{\max} is identified from a trial and error method. The procedure is to increase the LLR p_{\max} value from 2 to 100 to analyse the variations of their corresponding Γ results. It can be seen from Fig. 5 that the smallest Γ is achieved at $p_{\max} = 4$, which is therefore adopted for the LLR modelling. The training and testing scatter plots for the LLR modelling are shown in Fig. 6a. It is observed that there are some points lying far above the bisector line during the training period signifying higher estimations, whereas some points sit far below the bisector line during the testing period indicating underestimation of the SMD. For the testing results, when XAJ-simulated soil moisture states have already reach the total dryness (i.e., XAJ SMD peaks at around 0.080 m), the predicted soil moisture state is still in the drying process. Figure 7a plots the time series of the estimated and the targeted SMD. The plot shows that the estimated SMD follows the seasonal trend of the soil moisture fluctuations well; therefore, it is wetter during the winter season and exsiccated during the hot summer season. However, it is clear to see that the model is not able to capture the extreme situations very well, especially during the wet season when the XAJ SMD becomes smaller (e.g., between day 300 and day 350).

3.2 Scheme 2: SMD estimation using SAC-SMA-SM and MODIS-LST as inputs

Land surface temperature is the product of the soil temperature multiplied by the emissivity, and the emissivity depends on the dielectric constant of the soil and soil moisture

Table 4. Model statistical performances and modelling information, where Γ is the calculated gamma statistic, which is the minimum MSE that can be achieved from a modelling method; A is the gamma gradient; SE is the Standard error; p_{\max} is the nearest points for LLR modelling; M is the training data size; and SMOS IA is the chosen incidence angle of SMOS- T_b s.

	Γ	A	SE	p_{\max}	M	SMOS IA
Scheme 1	0.072	1.353	0.004	4	292	–
Scheme 2	0.060	0.568	0.002	2	199	–
Scheme 3	0.033	0.152	0.004	7	120	H: 27.5–47.5, 57.5° V: 27.5–42.5, 52.5, 57.5°
Scheme 4	0.029	0.119	0.006	5	62	H: 37.5–57.5° V: 37.5–42.5, 57.5°

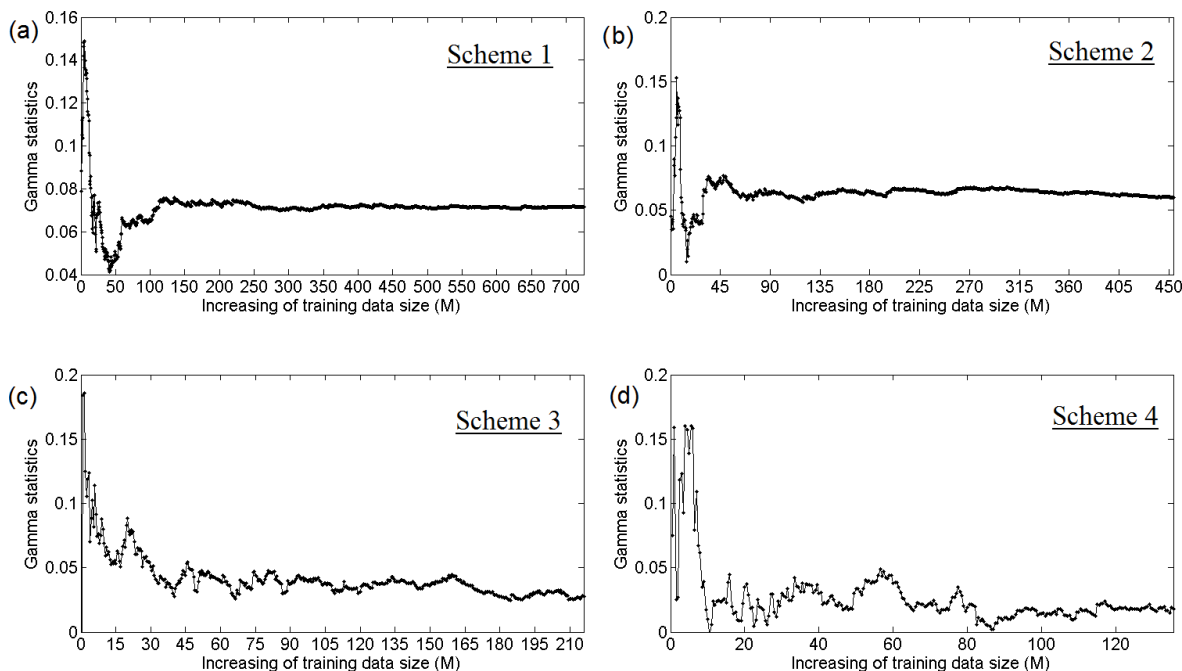


Figure 4. M -test to find the best training data size: (a) scheme 1, (b) scheme 2, (c) scheme 3, and (d) scheme 4.

(Rodriguez-Fernandez et al., 2015). Therefore, the additional MODIS-LST information could potentially improve the soil moisture estimation. The modelling process is the same as in scheme 1. In Table 4, it is clear to observe that by adding the MODIS-LST input, the Γ is improved to 0.060 and its corresponding gradient A is reduced significantly to less than half of scheme 1. Meanwhile the SE value is decreased remarkably as well showing the accuracy of the GT. The M -test in Fig. 4b shows the graph settles to an asymptote around 0.060, which is consistent with the calculated Γ result. Training data size of 199 is chosen here because it gives the lowest Γ value. For the LLR modelling, the best p_{\max} value is found to be 2 from the trial and error result in Fig. 5. The LLR training and testing performances are presented in Fig. 6b. Although the problem of underestimation of extremely dry soil still exists (i.e., the points concentrate at the right end of the training and testing plots), overall the model's prediction ability during both phases is better than that of scheme 1 (i.e., data points

are closer to the 45° line). The improvement can also be seen clearly in the time series plot in Fig. 7b. For example, the big disparities between the estimated and the targeted SMDs around day 300 and day 350 are reduced evidently.

3.3 Scheme 3: SMD estimation using SAC-SMA-SM and SMOS- T_b s as inputs

The multi-angle T_b s retrievals are the main data inputs for SMOS soil moisture calculation; therefore, their inclusion should also add a positive effect to the hydrological soil moisture estimation. As aforementioned, an efficient feature selection of the SMOS incidence angles is important for the best SMD calculation. In this study all the possible combinations from all inputs variables are examined with the Γ result as the statistical indicator. This method is capable of examining every combination (16 383 embeddings in this case) of data inputs to target the optimal combination that gives the

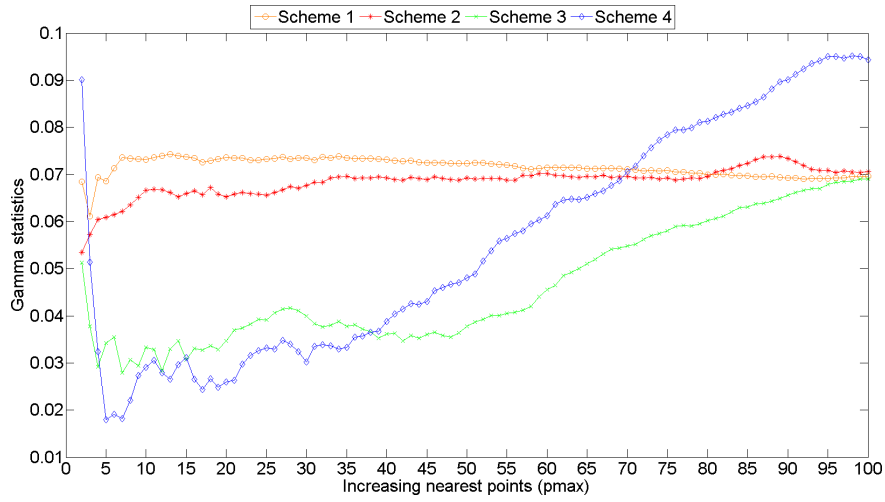


Figure 5. Gamma statistic (Γ) variations for increasing the LLR p_{max} value.

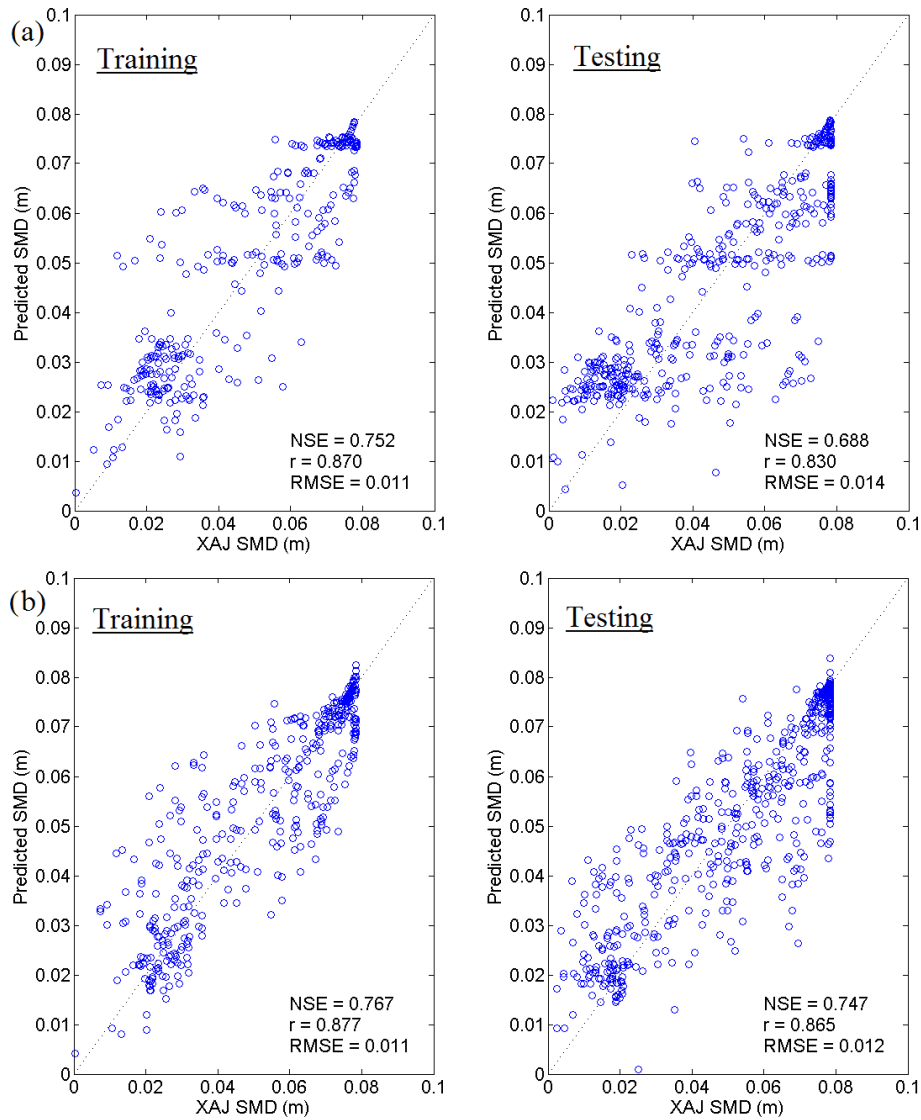


Figure 6. LLR modelling during the training and testing phases for (a) scheme 1 and (b) scheme 2.

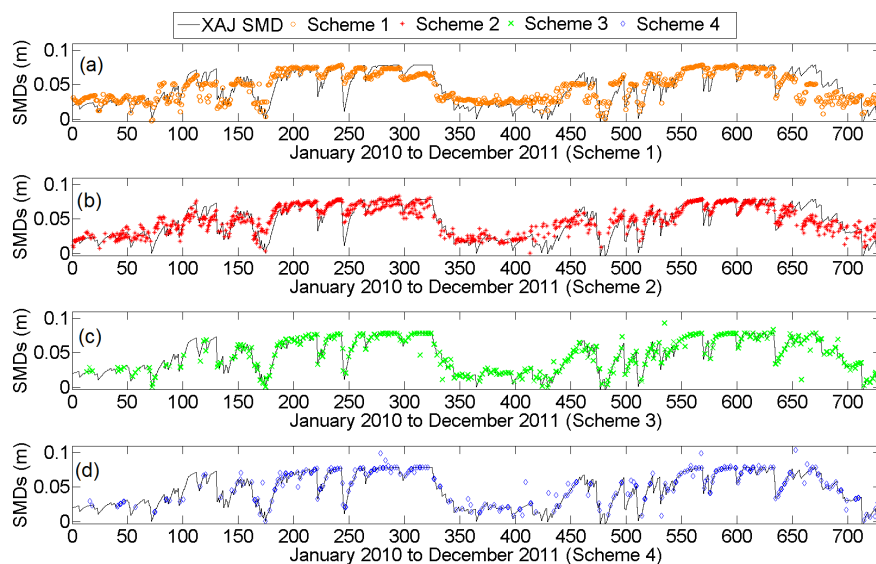


Figure 7. The time series plots of the XAJ SMD and the estimated SMD from the four schemes: (a) scheme 1, (b) scheme 2, (c) scheme 3, and (d) scheme 4.

smallest absolute Γ value. As discussed in Sect. 2.3.4, SAC-SMA-SM is a compulsory data input; therefore, it is not included in the selecting process. The best set of SMOS- T_b s to retrieve soil moisture state is composed of H polarisation at the incidence angles of 27.5–47.5, 57.5°, and V polarisation at the incidence angles of 27.5–42.5, 52.5, 57.5°. This result demonstrates that using a combination of H and V T_b s gives a better soil moisture estimation, which is logically sensible because different polarisations carry distinct information of the Earth's surface. However, some incidence angles could hold common features, which when putting together could result in a negative effect to the LLR modelling, and are therefore not included. The detailed investigation of the possible common features is outside the scope of this paper, which is mainly due to the SMOS working mechanism.

As seen from Table 4, the inclusion of SMOS- T_b s significantly improves the Γ result by 54 %, while the gradient A is reduced greatly by 89 % as compared with scheme 1. The small A value illustrates that the underlying system is more straightforward and easier to model than that of scheme 1. The M -test analysis in Fig. 4c produces an asymptotic convergence from a 120 training data size of Γ value around 0.033. It is interesting to see that the proportion of the required training data is relatively larger than those in schemes 1 and 2. The potential reason could be explained by the larger amount of data inputs in this scheme. For LLR modelling, the p_{\max} that gives the smallest Γ is 7 (Fig. 5). The SMD estimations during the training and the testing are presented in Fig. 8a. It can be seen that the SMD prediction ability of this scheme is remarkably better than the previous ones, as most of the points lie on the bisector line albeit there are still some under- and overestimations. The reason SMOS outperforms MODIS in SMD estimation could

be due to the long wavelength the microwave has; therefore, it presents the top few centimetres of the soil while MODIS LST (thermal infrared) only provides information at the soil surface. The used LLR algorithm has been double checked to filter out the potential of an overfitting problem. The checking processes are performed by muddling the SMD target in the testing datasets as well as altering the input file, and its efficiency stays the same. Hence, it is believed that the LLR model is very useful in calculating SMD from this scheme. Generally the NSE, r and RMSE statistical indicators show a high agreement during both training and testing phases. For the time series plot in Fig. 7c, it is clear to see that most of the estimated points lie closely to the benchmark line. The observed outliers could be partly due to the data shortage in this scheme so that not all the scenarios are covered in the datasets.

3.4 Scheme 4: SMD estimation using SAC-SMA-SM, MODIS-LST, and SMOS- T_b s as inputs

In this scheme, all the three data sources are used to test if the modelling performance can be further improved. Here the full embedding calculation is again carried out to explore the most suitable incidence angles from the SMOS- T_b s. This is because the added MODIS-LST data could carry identical (i.e., redundant) features with some of the SMOS- T_b s datasets. As a result of the full embedding calculation, the best set of SMOS- T_b s is composed of H polarisation at the incidence angles of 37.5–57.5°, and V polarisation at the incidence angles of 37.5–42.5, 57.5°. As seen in Fig. 4d, the total amount of data is significantly reduced due to the shortage of simultaneously available days between the MODIS and the SMOS observations. Interestingly the M -test graph vi-

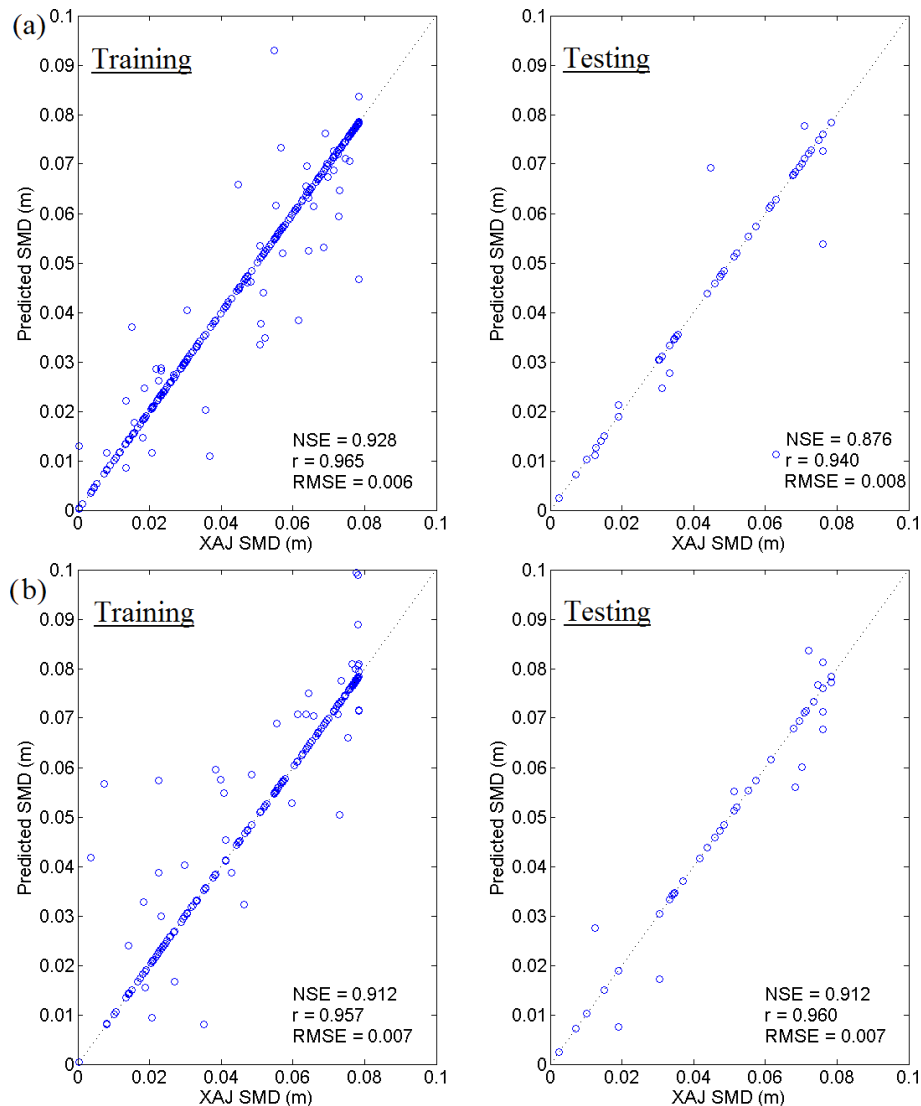


Figure 8. LLR modelling during the training and testing phases for (a) scheme 3 and (b) scheme 4.

brates more significantly than the other three schemes, which could be due to the smaller data size and the larger amount of data inputs in this scheme. Here the training data size is chosen as 62 with Γ obtained at around 0.030. The optimal p_{\max} is identified to be 5 (Fig. 5). The LLR modelling results are shown in Figs. 7d and 8b. It is obvious that this scheme further improves the accuracy of the SMD estimation, especially with the high statistical performances achieved during both training and testing phases. Comparatively, this scheme is more stable for SMD estimation, albeit it requires more data inputs and is only realisable when both the MODIS and the SMOS observations are available.

3.5 Produce an unintermitted soil moisture product

The data availability of the four schemes varies. As shown in Fig. 9, scheme 1, which has the poorest soil moisture state

estimation, gives the most data availability, while scheme 4, which has the most accurate soil moisture state estimation, shows the least data availability. In order to produce an unintermitted hydrological soil moisture product, the four schemes need to be combined together to complement each other. The combining method is by selecting the best available soil moisture estimation. For example, if all the schemes have available data at the same time, the best scheme's soil moisture data are chosen (i.e., scheme 4 in this situation), whereas if just one scheme has data on that day, only that scheme's soil moisture data are used. The performances of the four schemes as well as the combined product are summarised in Table 5. Although the combined soil moisture state is obtained with lower statistical performances than that of schemes 3 and 4, it is still hydrologically very accurate especially when compared to the SMOS's official soil mois-

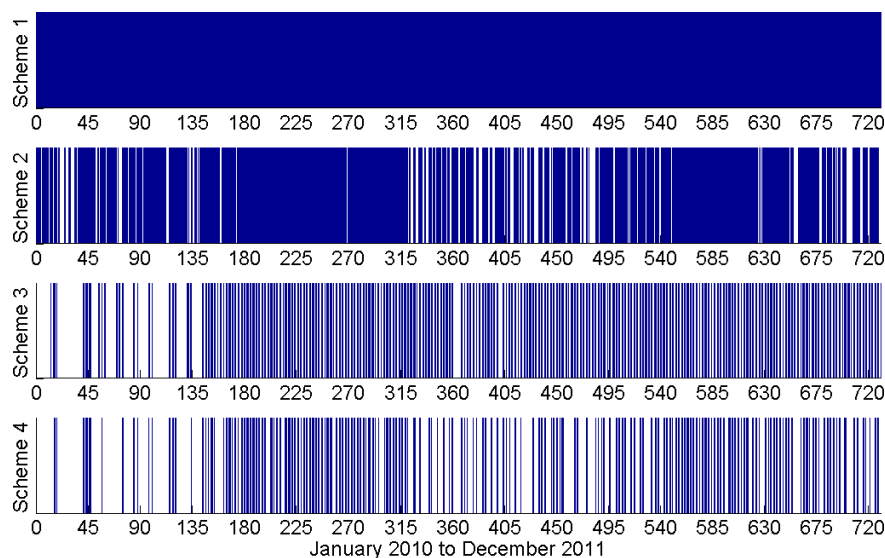


Figure 9. Data availability plots of the four schemes: scheme 1: SAC-SMA-SM input; scheme 2: SAC-SMA-SM and MODIS-LST inputs; scheme 3: SAC-SMA-SM and SMOS- T_b s inputs; scheme 4: SAC-SMA-SM, MODIS-LST, and SMOS- T_b s inputs. The total available days for the four schemes are 730, 458, 217, and 140 respectively.

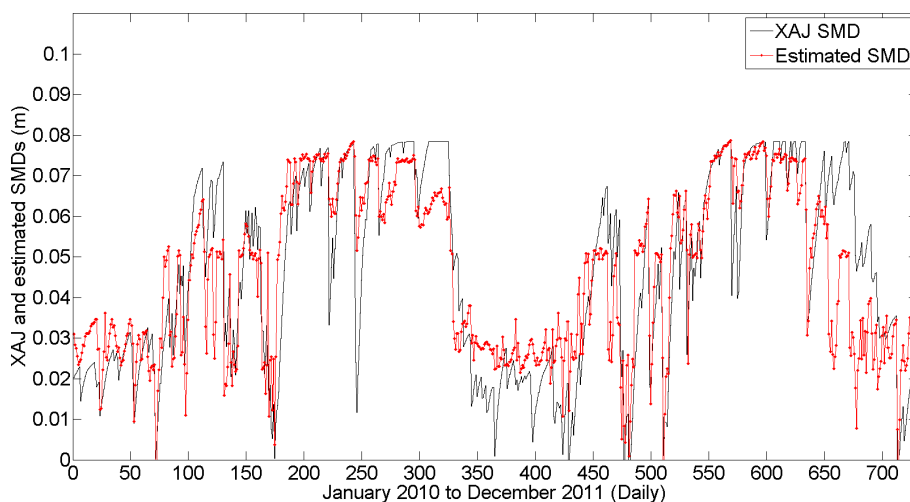


Figure 10. Time series plot of the combined daily hydrological soil moisture state estimations.

ture product (Table 5). The time series of the combined soil moisture state is plotted in Fig. 10. It can be seen that the general trend of the produced soil moisture state follows the targeted data very well. However, it tends to overestimate some of the wet events during the rainy season and significantly underestimate the dryer soil condition in September 2011. Those poor estimations are mostly from schemes 1 and 2 where schemes 3 and 4 are not available. Since more and more microwave satellite observations are becoming obtainable, those new data sources could add extra benefits into the proposed model, and the accuracy of the soil moisture product is expected to be further enhanced.

4 Discussion

4.1 What is a soil moisture state variable?

This study uses the XAJ's SMD simulation as a target because it is directly produced by a hydrological model. However, it is argued that models with different parameters values can generate equally good flow results called the equifinality effect, because they are all calibrated based on the observed flow. For this reason, their soil moisture state variables can be distinct among each other.

In order to investigate this effect in more details, the XAJ model is manipulated by increasing one of its parameters WUM by 30%. By doing so, the XAJ's flow simulation re-

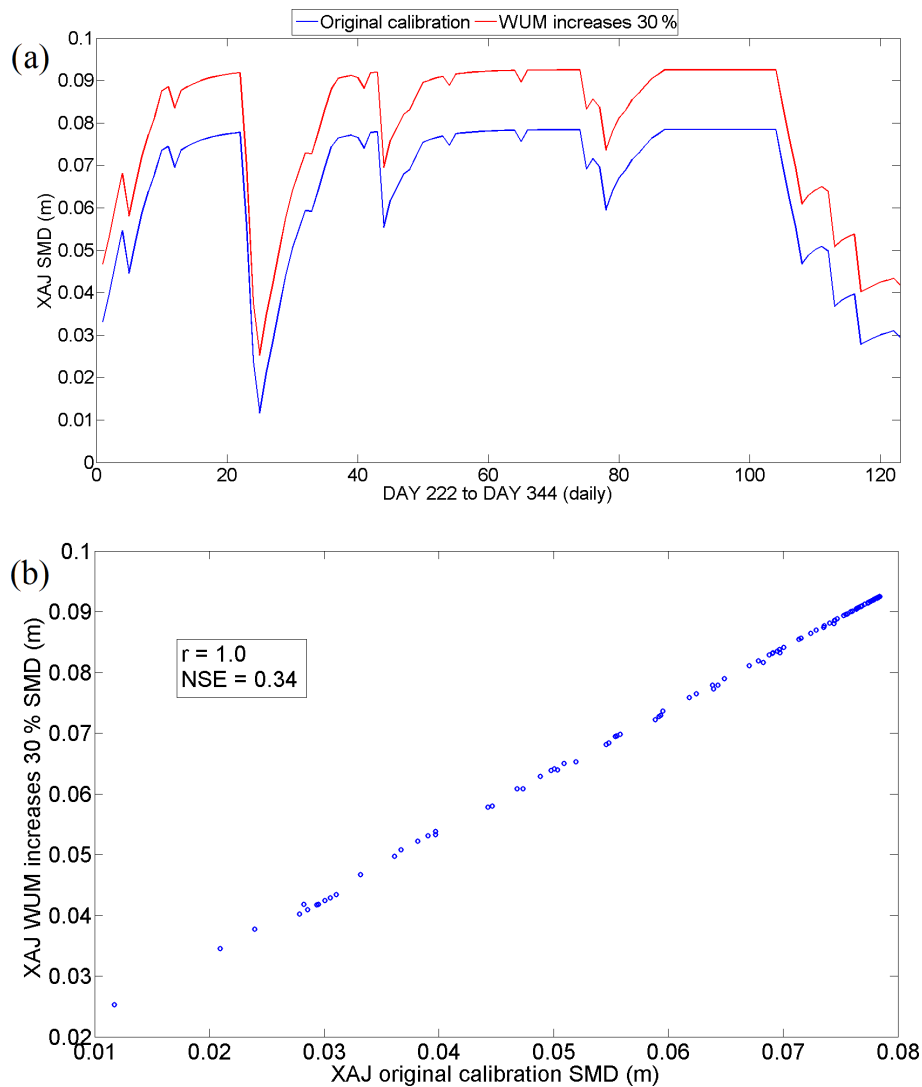


Figure 11. SMD variations from the manipulated XAJ calibration (i.e., the WUM parameter is increased by 30 %) and its original calibration.

Table 5. Summary of SMD estimation performances. It is noted that RMSE is in the unit of metres.

	Training			Testing		
	NSE	r	RMSE	NSE	r	RMSE
Scheme 1	0.752	0.870	0.011	0.688	0.830	0.014
Scheme 2	0.767	0.877	0.011	0.747	0.865	0.012
Scheme 3	0.928	0.965	0.006	0.876	0.940	0.008
Scheme 4	0.912	0.957	0.007	0.912	0.960	0.007
Combined	–	–	–	0.790	0.889	0.011
SMOS-SM	–	–	–	0.420	0.650	0.017

mains as effective as its original form (the same NSE values), but its soil moisture state changes significantly from its original values. For a better visualisation, an enlarged plot of the SMD simulations between day 222 and day 344 is presented. As seen from Fig. 11a although the soil moisture state

variables from two equally good calibrations have a wide range of value differences ($NSE = 0.34$), they both follow the same pattern: when it rains they become wet by the similar amount; when there is a dry period they all move into a dryer state in a similar rate to the actual evapotranspiration. Therefore, they appear as in parallel movements and the latter plot (Fig. 11b) shows a very strong linear correlation ($r = 1.0$) between them. It is important to note that the selection of the dry period (i.e., high SMD values) is because it is the most critical period of time for the need of accurate soil moisture values for hydrological modelling. This is because during the real-time flood forecasting, after a long period of dryness, the accumulation of error in the hydrological models can become larger and larger with time. With accurate soil moisture information, the error could be corrected.

Although the absolute values of the models' soil moisture state variables are not quite meaningful and comparable, their

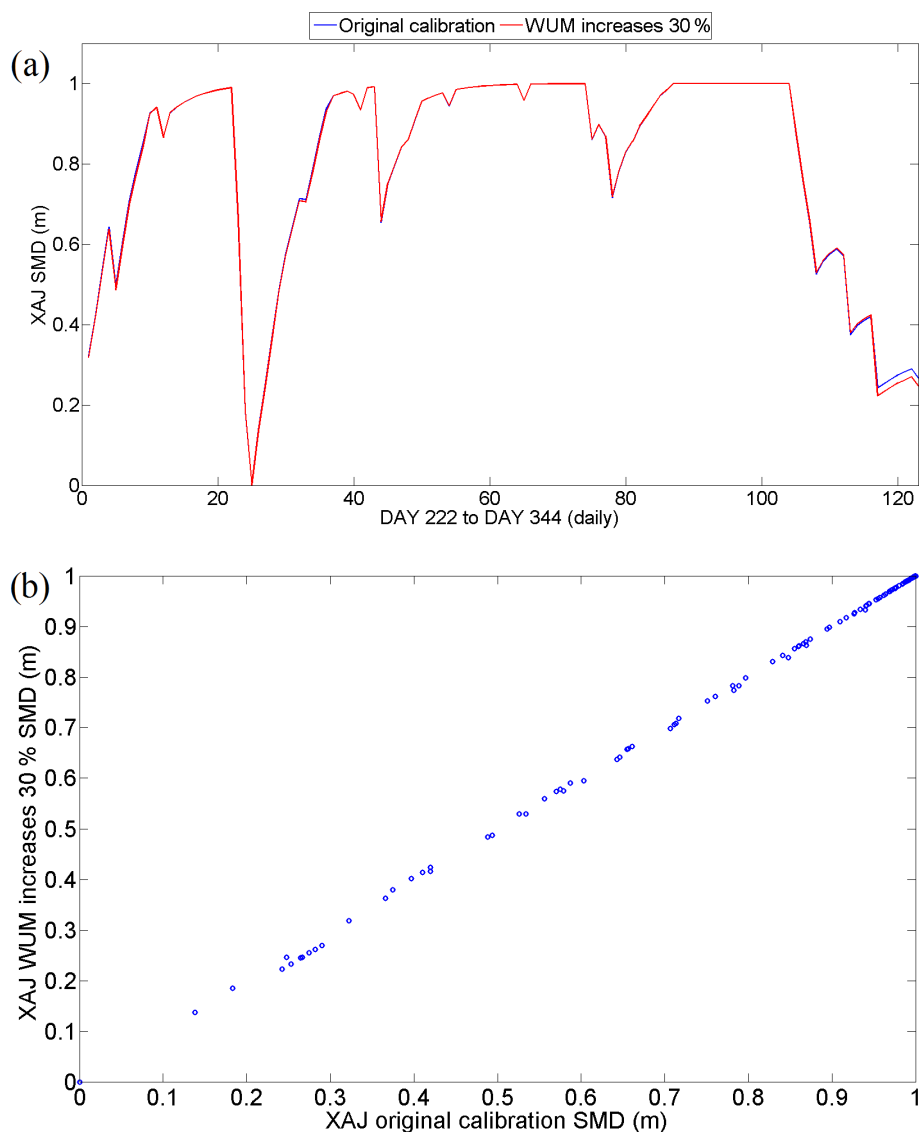


Figure 12. Normalised SMD variations from the manipulated XAJ calibration (i.e., the WUM parameter is increased by 30%) and its original calibration.

variations are the true reflection of the soil moisture fluctuations in the real world. This clarification is a very important concept, because there has been a wide spread of misunderstanding about the hydrological model's soil moisture state and its connection with the real-world soil moisture.

4.2 Soil moisture state normalisation

One deficiency of this study is that the generated soil moisture state is based on a hydrological model's SMD simulation, and therefore it is model parameter dependent. It is desirable to produce a soil moisture indicator that is independent from model parameters and dimensionless with variables between 0 and 1. Normalised hydrological soil moisture state (NHSMS) indicators are produced as presented in

Fig. 12 (corresponding to the SMD simulations shown in Fig. 11). The normalisation method is obtained by adopting the following equation:

$$\text{NHSMS} = \frac{\text{SMD} - \min(\text{SMD})}{\max(\text{SMD}) - \min(\text{SMD})}. \quad (11)$$

Such an approach is very effective as demonstrated by the almost identical SMD curves between the two XAJ simulations. In the future it is planned to use the same process on other hydrological models to test if the normalised soil moisture indicators are not only model parameter independent but also model structure independent. Since all hydrological models are driven by the same physics laws on the conservation of mass, their normalised soil moisture indicators should respond in a similar way (soil becomes wetter when

it rains and drier when there is no rain). If this is true, a new soil moisture product based on NHSMS could be generated as a routine product by the operational organisations such as NASA and ESA. Such a soil moisture product will also be very useful to the meteorological and hydro-meteorological fields in their land surface modelling because the current land surface models suffer from poor performance in their runoff estimations. As aforementioned, all current soil moisture products such as those from ESA and NASA are not optimised for different application fields. Our study gives an example of simulating the soil moisture data targeted to serve the hydrological community. It is possible other products serving farmers in agriculture, ecologists in the environment, and geotechnical engineers in construction could be produced using the proposed method.

4.3 Application of the produced soil moisture data

Another area needing further work is the hydrological application of the produced data. Generally, effective hydrological application of soil moisture data needs three pre-conditions: (1) a good soil moisture data relevant to hydrology, (2) a hydrological model compatible with such data, and (3) an effective data assimilation scheme. This paper tackles the first point, and the other two points would need further research because there are significant knowledge gaps in them. If all the three points are solved, such a data has a huge potential in operational hydrological modelling. For example, initialisation of the model could be shortened, which reduces the need for model warm-up. This is important during real-time flood forecasting when there is not enough data to warm up the model for an imminent flood event. Such a warm-up period could be very long, as demonstrated by the study in Ceola et al. (2015). In addition the XAJ SMD data used here is based on the calibration of the observed rainfall and flow so that the targeted SMD is interpolated between observations and there is a minimum time drift. In real-time flood forecasting, the errors in precipitation and evapotranspiration could accumulate, which cause time-drift problems. Therefore, a soil moisture product such as the one produced in this study (i.e., based on minimal time-drift SMD) could help one avoid such a problem. The proposed soil moisture data are also valuable for the validation of land surface models, especially useful for their runoff simulations. Due to the limit of time and resources, this study has not tackled all the issues, but has laid a good foundation for their future research.

4.4 XAJ model under frozen conditions

The Pontiac catchment is characterised by soil-freezing events in winter seasons. During freezing events, soil moisture transfer fundamentally differs from the unfrozen conditions (e.g., Gelfan, 2006). Although the XAJ model has been successfully applied in simulating flows in frozen soil conditions (e.g., see Zhou et al., 2008), as well as in this case

study, the lumped XAJ model does not explicitly consider soil freezing; thus, SMD simulations can be inaccurate for winter seasons and further research is needed to investigate this issue further.

5 Conclusions

A hydrological soil moisture product is produced for the Pontiac catchment using the GT and the LLR modelling techniques based on four data-input schemes. Three data sources are considered including the soil moisture product from the SAC-SMA model, the land surface temperature retrieved by the MODIS satellite, and the multi-angle brightness temperatures acquired from the SMOS satellite. The four data-input schemes are built from the four combinations of the data sources. The generated soil moisture product (uninterrupted with no missing data) for a period of 2 years (2010–2011) is compared with the XAJ hydrological model's SMD simulation to test its hydrological accuracy. It is concluded that the GT and the LLR modelling techniques together with the chosen data inputs can be used with high confidence to estimate an uninterrupted hydrological soil moisture product, and the proposed method could be easily applied to other catchments and fields.

In this study it has been found that different data sources have their own unique information contents, so that they can complement each other using data fusion technique. Their synergy can be best achieved to produce an enhanced soil moisture product. In data fusion an important principle is MRmr (maximum relevance minimum redundancy). The soil moisture state in this study is generated from a large number of data inputs, and their selection is carried out by the GT, which is one of the methods in MRmr. This is the first time that the GT is used in a data fusion of satellite multiple T_b s scans, land surface temperature and external soil moisture information for producing a hydrological soil moisture product. Future studies should explore other MRmr methods in addition to GT, to compare if they are more effective input selection methods. As to the data fusion regression model, LLR is chosen in this study because it is easily applied and very effective. However, it is possible there may exist other better models. We encourage the community to apply the proposed methodology using other regression models.

Data availability. The NLDAS-2 datasets used in this article can be obtained through the https://ldas.gsfc.nasa.gov/nldas/NLDAS2forcing_download.php (NASA, LDAS, 2017) website, the SMOS level-3 brightness temperatures and soil moisture are from the CATDS at <http://www.catds.fr/Products/Products-access> (CATDS, 2017), and the MODIS level-3 land surface temperature can be obtained from the LP DAAC website at https://lpdaac.usgs.gov/dataset_discovery/modis/modis_products_table/mod11c1 (LP DAAC, 2014).

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. This study is supported by Resilient Economy and Society by Integrated SysTems modelling (RESIST), Newton Fund via Natural Environment Research Council (NERC) and Economic and Social Research Council (ESRC) (NE/N012143/1). We acknowledge the U.S. Geological Survey for making available daily streamflow records (<http://waterdata.usgs.gov/nwis/rt>).

Edited by: Alexander Gelfan

Reviewed by: two anonymous referees

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