



# 1 Measurement of OH reactivity by laser flash photolysis coupled with laser-induced 2 fluorescence spectroscopy

3 Daniel Stone<sup>1</sup>, Lisa K. Whalley<sup>1,2</sup>, Trevor Ingham<sup>1,2</sup>, Peter M. Edwards<sup>1,\*</sup>, Daniel R. Cryer<sup>1</sup>, Charlotte A. Brumby<sup>1</sup>,  
4 Paul W. Seakins<sup>1,2</sup> and Dwayne E. Heard<sup>1,2</sup>

5

6 <sup>1</sup> School of Chemistry, University of Leeds, Leeds, LS2 9JT, UK

7 <sup>2</sup> National Centre for Atmospheric Science (NCAS), University of Leeds, Leeds, LS2 9JT, UK

8 \*now at Wolfson Atmospheric Chemistry Laboratories, Department of Chemistry, University of York, Heslington,  
9 York, YO10 5DD

10

## 11 Abstract

12 OH reactivity ( $k'_{\text{OH}}$ ) is the total pseudo-first-order loss rate coefficient describing the removal of OH radicals to all  
13 sinks in the atmosphere, and is the inverse of the chemical lifetime of OH. Measurements of ambient OH reactivity  
14 can be used to discover the extent to which measured OH sinks contribute to the total OH loss rate. Thus, OH  
15 reactivity measurements enable determination of the comprehensiveness of measurements used to predict air  
16 quality and ozone production, and, in conjunction with measurements of OH radical concentrations, to assess our  
17 understanding of OH production rates.

18 In this work, we describe the design and characterisation of an instrument to measure OH reactivity using laser  
19 flash photolysis coupled to laser-induced fluorescence (LFP-LIF) spectroscopy. The LFP-LIF technique produces OH  
20 radicals in isolation, and thus minimises potential interferences in OH reactivity measurements owing to the  
21 reaction of HO<sub>2</sub> with NO which can occur if HO<sub>2</sub> is co-produced with OH in the instrument. Capabilities of the  
22 instrument for ambient OH reactivity measurements are illustrated by data collected during field campaigns in  
23 London, UK, and York, UK.

24 We also present the coupling and characterisation of the LFP-LIF instrument to an atmospheric chamber for  
25 measurements of OH reactivity during simulated experiments, and provide suggestions for future improvements  
26 to OH reactivity LFP-LIF instruments.

27

## 28 1. Introduction

29 OH radicals dominate atmospheric oxidation chemistry, controlling the lifetimes of most primary pollutants and  
30 greenhouse gases emitted into the atmosphere, including methane, CO, volatile organic compounds (VOCs), NO<sub>2</sub> and  
31 SO<sub>2</sub>, whilst also contributing to the production of secondary pollutants such as ozone, sulphuric acid and secondary  
32 organic aerosol (SOA). Appreciation of the factors controlling atmospheric OH radical concentrations is thus  
33 essential to understanding the processing and fate of trace species in the atmosphere, and to our ability to  
34 understand and predict air quality and climate change. Moreover, the short chemical lifetimes of the OH radical and  
35 the closely related HO<sub>2</sub> radical make OH and HO<sub>2</sub> ideal species for testing the chemical mechanisms used in  
36 atmospheric models since their concentrations are controlled by in situ chemistry alone and are not influenced by  
37 transport processes. However, model simulations of OH concentrations require calculation of both OH production  
38 and loss rates, and there is potential for agreement between modelled and observed OH concentrations based on  
39 opposing errors in the production and loss terms.

40 Observations of OH radical concentrations made in conjunction with measurements of OH reactivity ( $k'_{\text{OH}}$ ), the total  
41 loss rate of OH and the inverse of the OH chemical lifetime ( $\tau_{\text{OH}}$ ), thus provide a means to separate the production



42 and loss terms for OH, enabling a more robust test of our understanding of OH radical concentrations and of  
 43 atmospheric oxidation chemistry. In addition, comparison of measured OH reactivity with calculated OH reactivity,  
 44 based on observed concentrations of OH sinks and known rate coefficients for their reactions with OH (Equation 3),  
 45 also provides an indication of the presence and importance of unmeasured OH sinks.

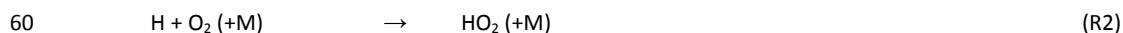
$$46 \quad -d[\text{OH}]/dt = \sum k_x[\text{X}][\text{OH}] \quad (\text{Equation 1})$$

$$47 \quad = k_{\text{OH}}'[\text{OH}] \quad (\text{Equation 2})$$

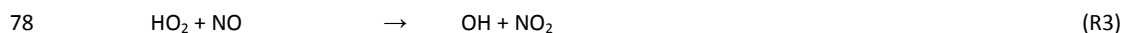
$$48 \quad k_{\text{OH}}' = \sum k_x[\text{X}] \quad (\text{Equation 3})$$

49 where  $k_x$  is the rate coefficient for reaction of OH with species X and  $k'_{\text{OH}}$  is the OH reactivity (the pseudo-first-order  
 50 rate coefficient for reaction of OH with all reaction partners present). Finally, using both  $[\text{OH}]$  and  $k_{\text{OH}}'$  to determine -  
 51  $d[\text{OH}]/dt$  experimentally, it is possible to evaluate the completeness of our knowledge of OH sources, which when  
 52 added together should equal  $+d[\text{OH}]/dt$  if the steady-state budget is closed (Martinez et al., 2003; Whalley et al.,  
 53 2011; Fuchs et al., 2013; Lu et al., 2013).

54 Measurements of OH reactivity in the atmosphere have been made by three different techniques – the flow tube  
 55 technique, the laser flash photolysis technique and the comparative reactivity method, with all three methods  
 56 relying on production of above ambient concentrations of OH radicals and monitoring of the OH decay rate, either  
 57 directly or indirectly. The flow tube method typically generates OH radicals at the tip of a sliding injector by  
 58 photolysis of water vapour (R1) using a mercury vapour lamp, also resulting in production of HO<sub>2</sub> radicals (R2).



61 The OH radical signal is monitored downstream of the injector after mixing with a flow of ambient air in the main  
 62 tube. By changing the position of the sliding injector relative to the point at which OH is detected it is possible to  
 63 vary the mixing time of OH with the ambient air, and thus to determine the total loss rate for OH in the flow tube.  
 64 However, the technique has a number of disadvantages. The time resolution of measurements made by the flow  
 65 tube method is relatively poor, owing to the need to measure OH signals at a number of different injector positions  
 66 to obtain a kinetic profile, during which time the ambient OH reactivity could show significant variability, although  
 67 Mao et al. (2009) overcome this issue for airborne measurements of OH reactivity by reducing the number of time  
 68 points used to determine the OH decay rate. The flow rates of sampled air in the flow tube method are relatively  
 69 high (~ 300-900 slm (standard litres per minute)), with turbulent flow conditions leading to high wall loss rates of OH  
 70 in the flow tube and relatively high uncertainties in determinations of OH reactivity owing to uncertainties in the wall  
 71 loss rates. Knowledge of the flow velocity in the flow tube, requiring direct measurement or knowledge of the flow  
 72 regime, total flow rate and cross-section of the flow tube, is also needed to convert the distance over which OH and  
 73 ambient air are mixed to reaction time, and can lead to uncertainties in the contact time between OH and reactants  
 74 in ambient air. A significant disadvantage of the flow tube method is the generation of equal concentrations of OH  
 75 and HO<sub>2</sub> following photolysis of water vapour at the tip of the sliding injector (reactions R1-R2), leading to the  
 76 potential for production of OH in the flow tube on the timescale of the experiment owing to the reaction of HO<sub>2</sub> with  
 77 ambient NO (R3).

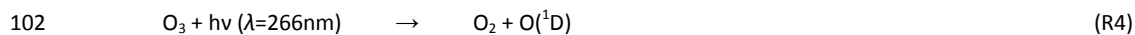


79 The production of OH from R3 reduces the observed decay rate of OH in the flow tube, and measurements of OH  
 80 reactivity using the flow tube method thus also require simultaneous measurements of ambient NO concentrations  
 81 in order to correct for interferences from HO<sub>2</sub> + NO, which can be quite significant. For example, for 75 ppb NO, a  
 82 typical rush hour mixing ratio in Mexico City, a correction factor of ~1.7 was required to account for the production  
 83 of OH from HO<sub>2</sub> + NO within the flow tube (Shirley et al., 2006).



84 In the comparative reactivity technique, a reactive molecule not usually present in air, typically pyrrole, is entrained  
 85 in a gas flow and the rate of its decay owing to reaction with artificially high concentrations of OH is measured in  
 86 'zero' air and ambient air by proton transfer mass spectrometry (PTR-MS), or gas chromatography with mass  
 87 spectrometry (GC-MS) or photoionisation detection (GC-PID). Comparison of the rates of decay of the molecule in  
 88 'zero' air and ambient air enables determination of the competition between the reaction of OH with the known  
 89 concentration of the reactive molecule and the reaction of OH with sinks in ambient air, thus enabling measurement  
 90 of the ambient OH reactivity. Absolute measurement of the physical loss rate of OH is not required for the  
 91 technique, and the limit of detection of comparative reactivity instruments is determined by the sensitivity to  
 92 changes in the signal corresponding to the concentration of the reactive species. However, OH radicals are typically  
 93 produced in comparative reactivity instruments through reactions R1-R2, in a similar manner to that used in flow  
 94 tube instruments and thus also producing high concentrations of HO<sub>2</sub>. Interferences resulting from OH production  
 95 from HO<sub>2</sub> + NO are thus also potentially problematic for comparative reactivity instruments and knowledge of HO<sub>2</sub>  
 96 and NO concentrations are required to correct for any interferences. In addition, the amount of OH produced is  
 97 dependent on humidity, and it is essential to ensure constant humidity between measurements made in 'zero' air  
 98 and those made in ambient air, with significant corrections often necessary to account for any differences (Michoud  
 99 et al., 2015).

100 The laser flash photolysis technique produces OH in isolation via laser photolysis of O<sub>3</sub>, typically at a wavelength of  
 101 266 nm, followed by reaction of O(<sup>1</sup>D) with ambient H<sub>2</sub>O (reactions R4-R5):



104 The production of OH without initial co-production of HO<sub>2</sub> minimises potential interferences from HO<sub>2</sub> + NO and  
 105 renders the flash photolysis technique more suitable to high NO<sub>x</sub> (NO<sub>x</sub> = NO + NO<sub>2</sub>) environments. The laser flash  
 106 photolysis method also has the advantage that the production of OH radicals is uniform throughout the reaction cell,  
 107 minimising the risk of poor mixing which is potentially problematic for the flow tube and comparative reactivity  
 108 techniques. Flow rates of sampled air are typically lower for the laser flash photolysis instruments (~ 12-20 slm)  
 109 (Sadanaga et al., 2004a) than for those using flow tubes (~ 300-900 slm) (Kovacs and Brune, 2001; Kovacs et al.,  
 110 2003; Ingham et al., 2009; Hansen et al., 2014), and the resulting laminar flow of gas reduces contact of the gas with  
 111 the walls of the instrument, thus reducing the physical loss rate of OH and associated uncertainties. Although  
 112 averaging of data is often required to improve signal-to-noise, a significant advantage of the laser flash photolysis  
 113 technique is the ability to measure ambient OH reactivity in 'real-time' through time-resolved measurements of the  
 114 OH decay following photolysis. The technique has the potential for significantly enhanced time resolution, both in  
 115 terms of the number of time points obtained during the decay of OH, and the averaging time over which the data  
 116 are reported, compared to the flow tube or comparative reactivity methods.

117 The first atmospheric measurements of total OH reactivity were made at an urban background site in Nashville, TN,  
 118 USA, in summer 1999 using the flow tube technique (Kovacs and Brune, 2001; Kovacs et al., 2003). Calculations of  
 119 OH reactivity, using VOC measurements co-located with the reactivity measurements, underestimated the total  
 120 observed reactivity by ~30 % on average owing to unmeasured or unknown VOCs and VOC oxidation products  
 121 (Kovacs et al., 2003). Subsequent measurements at an urban background site in New York, NY, USA, were, on  
 122 average, within 10 % of the calculated reactivity in summer 2001 (Ren et al., 2003), but were underestimated by 30-  
 123 40 % during morning and evening rush hours in winter (Ren et al., 2006a). Significant underestimation of the  
 124 measured OH reactivity in the morning rush hour was also reported for observations in the Mexico City Metropolitan  
 125 Area (MCMA), with the observed reactivity reaching 120 s<sup>-1</sup> (Shirley et al., 2006). High OH reactivity has also been  
 126 observed in Paris during the MEGAPOLI campaign in 2010, with k<sub>OH</sub> reaching 130 s<sup>-1</sup> for continental air masses and  
 127 calculations underestimating the reactivity by up to 75 % (Dolgorouky et al., 2012). Reactivity measurements in  
 128 Tokyo were underestimated in summer, spring and autumn, but reproduced to within 5 % in winter, with the  
 129 reactivity correlating well with NO<sub>x</sub> throughout the year (Sadanaga et al., 2004b; Yoshino et al., 2006; Chatani et al.,



130 2009; Yoshino et al., 2012). Aircraft measurements of OH reactivity have also shown that reactivity tends to  
131 decrease with altitude, with discrepancies between observed and calculated reactivity most pronounced at altitudes  
132 up to 2 km and tending towards agreement at altitudes above 4 km (Mao et al., 2009).

133 Flow tube measurements at an urban site in Houston, US, during the TEXAQS2000 and TRAMP2006 campaigns (Mao  
134 et al., 2010) and at a forested site at Whiteface Mountain, NY, USA (Ren et al., 2006b), were well reproduced by  
135 model calculations. However, measurements made at a coastal site in Norfolk, UK, typically characterised by  
136 relatively 'clean' air were significantly underestimated and attributed to the presence of numerous high molecular  
137 mass VOCs at low concentrations which were not included in the VOC measurement suite (Lee et al., 2009; Ingham  
138 et al., 2009). The presence of unmeasured VOCs was also indicated for the PROPHET 2000 campaign at a forested  
139 site in Michigan, USA, during which the measured OH reactivity was underestimated by ~50 % on average, with the  
140 'missing' OH reactivity exhibiting a strong temperature dependence potentially resulting from temperature-  
141 dependent emissions of unmeasured biogenic VOCs (Di Carlo et al., 2004). Uncertainties in emissions and chemistry  
142 of biogenic VOCs, particularly in the oxidation chemistry of isoprene and its oxidation products, have also been  
143 responsible for underpredictions of observed OH reactivity in forested regions in Suriname (Sinha et al., 2008) and  
144 Borneo (Whalley et al., 2011; Edwards et al., 2013). Model calculations of OH reactivity in Borneo underestimated  
145 the observed diurnal mean reactivity by 30 %, and indicated that uncertainties in the chemistry and deposition rates  
146 of secondary oxidation products could potentially explain the observed reactivity without the need for additional  
147 primary VOC emissions, and that at least 50 % of the carbon-containing compounds which react with OH were not  
148 measured (Edwards et al., 2013). Biogenic VOCs also dominated the daytime OH reactivity in the Pearl River Delta  
149 region, China, with isoprene and its oxidation products comprising ~40 % of the total OH reactivity in the afternoon  
150 and observed reactivity underestimated by ~50 % when calculated from measured OH sinks but reproduced by  
151 model calculations considering the contributions from secondary oxidation products (Lou et al., 2010). However,  
152 observations of OH reactivity in a forested region in Colorado, USA, during the BEACHON-SRM08 campaign were  
153 underestimated by model calculations (~ 30 %), with the dominant VOCs found to be 2-methyl-3-buten-2-ol (MBO)  
154 and monoterpenes (Nakashima et al., 2014).

155 Using a branch enclosure technique, Kim et al. (2011) demonstrated that underestimations of observed OH reactivity  
156 at the PROPHET field site, Michigan, USA, during the 2009 CABINEX campaign were related to oxidation products of  
157 known and measured biogenic VOCs, rather than to unknown or unmeasured primary VOC emissions. Model  
158 calculations were able to reproduce the CABINEX OH reactivity observations below the forest canopy, but  
159 discrepancies were apparent above the canopy, indicating the presence of unmeasured trace gases above the forest  
160 canopy (Hansen et al., 2014). Model calculations and experiments using the comparative reactivity method at a  
161 forested site in Finland also observed differences between OH reactivity measured in the forest canopy and above  
162 the canopy (Mogensen et al., 2011; Nolscher et al., 2012). While the in-canopy reactivity was typically higher than  
163 the above canopy reactivity, transport of wildfire plumes to the site significantly increased the above canopy  
164 reactivity, increasing it above the in-canopy level and increasing the 'missing' reactivity above the canopy from 58 %  
165 for 'normal' conditions to 73 % for periods impacted by transported pollution (Nolscher et al., 2012).

166 OH reactivity measurements have also been used to determine ozone production rates in southwestern Spain during  
167 the DOMINO campaign (Sinha et al., 2012) and in London during the ClearLo campaign (Whalley et al., 2015), and  
168 have been used in laboratory studies to assess our understanding of combustion systems (Nakashima et al., 2010)  
169 and atmospheric isoprene oxidation mechanisms (Nakashima et al., 2012; Nolscher et al., 2014).

170 Measurements of OH reactivity thus have a number of applications, and can be used to improve our understanding  
171 of atmospheric composition and chemistry. In this work we present the design and characterisation of an instrument  
172 using laser flash photolysis coupled with laser-induced fluorescence (LFP-LIF) to measure OH reactivity in the field  
173 and in chamber experiments.

174

175



## 176 2. Experimental

177 A schematic of the OH reactivity instrument is given in Figure 1. The instrument comprises a reaction cell (described  
178 in Section 2.1) and a detection cell (described in Section 2.2), with the two cells typically situated approximately 5 m  
179 above ground level on the roof of a shipping container housing the FAGE mobile laboratory during ambient  
180 measurements.

### 181 2.1 Reaction Cell

182 The reaction cell consists of a cylindrical stainless steel tube of 50 mm internal diameter and 85 cm in length.  
183 Ambient air is drawn through a stainless steel sampling line (50 mm internal diameter and 20 cm in length), enters  
184 the reaction cell at 90° to the air flow in the tube, and is drawn along the tube by an extractor fan mounted on the  
185 exit arm situated immediately prior to the OH detection cell (Section 2.2). The fan speed determines the flow rate of  
186 gas in the reaction cell, and is set to ensure a laminar flow of air through the cell with a Reynold's number less than  
187 2300. The flow rate of gas, determined by measurement of the flow velocity using a hot-wire anemometer (TSI Air  
188 velocity transducer 8455-150) or set by calibrated mass flow controllers during laboratory experiments (Sections 4, 6  
189 & 7) and measurements of instrument zeroes (Section 3.1), was in the range 12 to 14 slm, giving a residence time of  
190 7 to 8 s in the reaction cell and a Reynold's number of ~ 360 (Equation 4).

$$191 \quad R_e = \rho v D / \mu \quad (\text{Equation 4})$$

192 where  $R_e$  is the Reynold's number,  $\rho$  is the density of the gas,  $v$  is the mean gas velocity,  $D$  is the diameter of the flow  
193 tube and  $\mu$  is the dynamic viscosity of the gas.

194 Production of OH radicals within the reaction cell is achieved by the 266 nm laser photolysis of  $O_3$  in the presence of  
195 water vapour (reactions R4-R5). A flashlamp pumped Nd:YAG laser (Big Sky Laser CFR 200, Quantel USA) is used to  
196 generate laser light at 1064 nm, which is frequency doubled to 532 nm (lithium triborate,  $LiB_3O_5$ , doubling crystal)  
197 and then frequency doubled to generate the fourth harmonic 266 nm radiation (caesium lithium triborate,  $CsLiB_6O_{10}$ ,  
198 doubling crystal) with pulse energies of ~ 50 mJ, pulse length 8 ns, and beam diameter of 6.35 mm. The laser is  
199 operated with a Q-switch to modulate the intracavity losses and maximise the pulse energy.

200 The 266 nm laser head is situated adjacent to the reaction cell in order to minimise the footprint of the instrument  
201 when used in the field. The laser head is powered, controlled and water cooled by an Integrated Cooler and  
202 Electronics unit (Big Sky Laser ICE450, Quantel USA) which is housed within the FAGE shipping container and  
203 powered via an uninterruptible power supply (APC 1000VA, American Power Conversion by Schneider Electric).

204 Laser light exiting the laser head is directed into the reaction cell using two dielectrically coated 266 nm turning  
205 mirrors of 1" diameter (Thorlabs, NB1-K04). Immediately prior to the reaction cell, the 266 nm beam is expanded to  
206 a diameter of ~10 mm by a telescope incorporating a plano-concave lens (Thorlabs LC4252, focal length = -30 mm)  
207 and a plano-convex lens (Thorlabs LA4148, focal length = 50 mm) housed in a lens tube (SM1M20, Thorlabs) to  
208 increase the photolysis volume within the reaction cell. The photolysis laser enters the reaction cell through a fused  
209 silica window, initiating OH radical production.

210 Typically, there is sufficient production of OH in the instrument from reactions R4-R5 at ambient concentrations of  
211  $O_3$  and water vapour in order to measure a temporal decay of OH. At low ambient concentrations of  $O_3$  (< 10 ppb) or  
212 during laboratory tests (Sections 4, 6 and 7) and measurement of instrument zeroes (Section 3), the OH radical  
213 concentration in the reaction cell is increased by passing a small flow (0.5 slm) of humidified ultra-high purity air  
214 (BTCA 178, BOC Special Gases) across a low pressure Hg vapour lamp and mixed with the main sampled air flow (12 –  
215 14 slm) in the inlet to the reaction cell. The mixing ratio of ozone in the reaction cell is increased by ~ 50 ppb by this  
216 method (measured by an ozone analyser (Thermo Environmental Instruments Inc., 49C  $O_3$  Analyser) situated at the  
217 end of the reaction cell during laboratory tests). Given knowledge of the rate coefficient for reaction of OH with  $O_3$   
218 ( $k_{OH+O_3} = 7.3 \times 10^{-14} \text{ cm}^3 \text{ s}^{-1}$ , Atkinson et al., 2004), the chemical loss of OH resulting from the addition of 50 ppb  $O_3$  is <



219 0.1 s<sup>-1</sup> at 298 K, and any loss of VOCs in the reaction cell through reaction with ozone is extremely small given the  
220 reaction times involved.

## 221 2.2 OH Detection Cell

222 OH radicals in the reaction cell are monitored by laser-induced fluorescence (LIF) using the fluorescence assay by gas  
223 expansion (FAGE) technique. The LIF-FAGE detection cell has been described previously in detail (Ingham et al.,  
224 2009), thus only a brief description will be given here.

225 Initial experiments were conducted with the detection cell situated midway along the reaction cell (Figure 1a), and  
226 sampling at 90° to the direction of air flow along the reaction cell, in a similar design to that described by Sadanaga  
227 et al. (2004a) and Lou et al. (2010). However, the observed OH decays in such a configuration displayed  
228 biexponential behaviour, comprising a fast initial decay followed by a slower decay representative of the expected  
229 OH reactivity, as observed in previously described instruments (Sadanaga et al., 2004a; Lou et al., 2010). In this  
230 configuration, the photolysis laser is aligned such that the beam passes across the inlet to the detection cell without  
231 hitting the inlet (as shown in Figure 1a). The air sampled in to the detection cell thus likely contains air that has  
232 experienced the photolysis laser (containing elevated OH concentrations) and air that has not experienced the  
233 photolysis laser (containing only ambient OH concentrations), with this mixing of air potentially leading to an  
234 apparent increase in the initial OH decay rate owing to dilution of the air containing elevated OH concentrations with  
235 air containing lower OH concentrations. Once mixing of the air having experienced the photolysis laser with that  
236 outside the beam diameter has occurred sufficiently to give uniform OH concentrations in the reaction cell the  
237 observed OH decay will result from the chemical losses in the instrument, leading to biexponential decays. Such  
238 biexponential behaviour has been attributed to similar effects of non-homogeneous spatial distributions of OH near  
239 the inlet to the detection cell (Lou et al., 2010) and to local heating and turbulence of the gas flow caused by the  
240 photolysis laser (Sadanaga et al., 2004a). A comparison between sampling at 90° to the direction of air flow along  
241 the reaction cell and sampling along the axis of the direction of air flow from the centre of the reaction cell has been  
242 reported previously (Amedro et al., 2012), with different fitting procedures required to extract the OH reactivity for  
243 the different instrument configurations attributed to differences in physical effects such as diffusion which were  
244 more significant when sampling at 90° (Amedro et al., 2012).

245 Subsequent experiments in this work (including all those described below) were performed with the detection cell  
246 situated at the end of the reaction cell along the same axis as the direction of air flow to sample air directly from the  
247 centre of the reaction cell, as shown in Figure 1b. This configuration reduces the chance of sampling air into the  
248 detection cell that has not experienced the photolysis laser beam, and thus reduces the impact of physical effects  
249 such as diffusion. The observed OH signals in this instrument configuration are described by a single exponential  
250 decay, although biexponential decays can still be obtained if the photolysis laser is not correctly aligned along the  
251 axis of the reaction cell.

252 Air is sampled from the centre of the reaction cell through a pinhole of 0.8 mm diameter and 0.5 mm thickness into  
253 the aluminium detection cell, which consists of three orthogonal axes and is black anodised to minimise light  
254 scattering within the cell. The pressure in the cell is measured by a capacitance manometer (Sensotec Z/606-01ZA)  
255 and is maintained at ~ 1.5 Torr by a roots blower backed by a rotary pump (Leybold Vacuum SV200/WAU1001),  
256 resulting in an air flow of approximately 4 slm and a supersonic expansion of the air as it is drawn through the  
257 pinhole.

258 The probe laser consists of a Nd:YAG pumped Ti:Sapphire laser (Photonics Industries) which generates broadband  
259 radiation in the range 690-1000 nm. A diffraction grating is used to select radiation with  $\lambda = 924$  nm, which is  
260 frequency tripled through generation of the second harmonic at 462 nm followed by sum-frequency mixing of the  
261 462 nm radiation with that at 924 nm to produce the 308 nm light with a pulse repetition frequency (PRF) of 5 kHz,  
262 pulse length (full width half maximum (FWHM)) of 35 ns, laser line width (FWHM) of 0.065 cm<sup>-1</sup> and beam diameter  
263 of ~ 3mm (Bloss et al., 2003).





264 A reference fluorescence cell containing a heated nichrome wire filament and humidified air at ~2 Torr to produce a  
265 constant stable source of OH radicals from dissociation of water vapour is used to facilitate tuning the probe laser to  
266 the precise wavelength required for the desired OH transition. Approximately 1 mW of the 308 nm laser light is used  
267 for this purpose, with ~9 mW used to make measurements of OH reactivity and a further ~13 mW remaining to  
268 make measurements of ambient OH, HO<sub>2</sub> and RO<sub>2</sub> concentrations in a separate instrument (see, for example,  
269 Whalley et al., 2015).

270 The probe laser, reference cell and pumps are all situated inside the shipping container. The ~9 mW of the 308 nm  
271 laser light used to measure OH reactivity is passed to the detection cell on the roof of the shipping container via an  
272 anti-reflective coated optical fibre with an angled and polished end (Oz Optics, QMMJ-55-UVVIS-200/240-3-30-AR2-  
273 SP, length = 30 m) through a baffled side-arm at 90° to the air flow. The probe laser light exits the detection cell  
274 through a baffled side-arm and is directed onto a photodiode (New Focus Large Area Photoreceiver 2032) to  
275 measure the laser power to enable normalisation of fluorescence signals for fluctuations in laser power. For a recent  
276 intercomparison at the SAPHIR chamber, the OH reactivity instrument, comprising the reactor flowtube and OH  
277 fluorescence and associated equipment was placed in the shipping container itself.

278 Fluorescence from electronically excited OH radicals resulting from excitation of the Q<sub>1</sub>(1) A<sup>2</sup>Σ<sup>+</sup> (v'=0) – X<sup>2</sup>Π<sub>3/2</sub> (v''=0)  
279 transition at 308 nm is collimated by a symmetrical biconvex collimating lens (Melles-Griot, focal length = 50 mm at λ  
280 = 546.1 nm, diameter = 50 mm) and focused onto the photocathode of a channeltron photomultiplier tube (PMT)  
281 (Perkin Elmer C 943P) by two plano-convex focusing lenses (UQG Optics Ltd., focal length = 75 mm at λ = 250 nm,  
282 diameter = 50 mm). A narrow band UV interference filter (Barr Associates Inc., FWHM bandwidth of 8 ± 1.6 nm  
283 centred at 309 ± 1 nm with a peak transmission of > 50 % at 308 nm and a blocking factor of 10<sup>6</sup> at other  
284 wavelengths) is situated between the excitation region in the detection cell and the PMT to minimise detection of  
285 scattered solar photons. The solid angle from which fluorescence is collected is effectively doubled through the use  
286 of a spherical concave mirror coated for high UV reflectance which is mounted in the detection cell opposite the  
287 side-arm bearing the PMT. Discrete photon signals on the PMT are processed using a multi-channel scaler photon  
288 counting card (Becker and Hickl, PMS 400, minimum bin width of 250 ns) in the computer used to control the  
289 instrument.

### 290 **2.3 Instrument Control**

291 A digital delay pulse generator (Stanford Research Systems DG535) produces a 5 kHz TTL (transistor-transistor logic)  
292 pulse to trigger the Ti:Sapphire laser and a second delay generator (Stanford Research Systems DG535) which  
293 subsequently triggers the gating of the PMT detector for the reactivity instrument and a third digital delay pulse  
294 generator (Berkeley Nucleonics Corporation 555) to trigger the 266 nm photolysis laser and the photon counting  
295 card at the specified pulse repetition frequency in synchronisation with the 308 nm probe laser. A personal  
296 computer is used to automate data collection, with analogue signals from measurements of the pressure in the  
297 detection cell and the power of the 308 nm probe laser at the photodiode attached to the detection cell digitised by  
298 an A/D card (Measurement Computing, PCI-DAS 1200). Electrical power to all parts of the instrument is supplied via  
299 an uninterruptible power supply (APC 1000VA).

### 300 **2.4 Data Acquisition**

301 Data acquisition is initiated by triggering of the photon counting card, with a background signal measured for 100 ms  
302 before triggering of the 266 nm photolysis laser and production of OH in the reaction cell. To avoid saturation of the  
303 PMT resulting from detection of the 308 nm laser pulse itself, the PMT is gated off at the onset of the 308 nm laser  
304 pulse (35 ns FWHM) until ~100 ns after the laser pulse, thereby preventing detection of any reflected or scattered  
305 laser light. The fluorescence signal is typically collected for 1 μs following each 308 nm probe laser pulse. Repeated  
306 measurements of the OH fluorescence signal are taken for 900 ms following each 266 nm photolysis laser pulse,  
307 during which time the OH concentration and hence the fluorescence signal will decay to the background level. Under  
308 normal conditions this occurs within ~300 ms of the photolysis laser pulse, although this is of course dependent  
309 upon the magnitude of the OH reactivity, and may be longer. The pulse repetition frequency of the 308 nm probe



310 laser (5 kHz) results in measurement of the OH fluorescence signal every 200  $\mu$ s throughout the measurement  
 311 period. The data collection cycle, as illustrated in Figure 2 is typically repeated every 1 s. Experiments, both in the  
 312 laboratory and in the field, in which the PRF of the photolysis laser was varied between 0.1 and 1 Hz showed no  
 313 effect on the observed OH reactivity.

314 When measurements of OH reactivity are made alongside those of ambient OH concentrations, the acquisition of OH  
 315 reactivity data is linked to measurements of ambient OH concentrations owing to the dual use of the 308 nm  
 316 excitation laser. Under such circumstances, measurements are taken on an approximate 7 min cycle, with a 5 min  
 317 'online' period during which the 308 nm laser is at the precise wavelength to excite the OH transition, followed by a  
 318 1 min 'offline' period during which the wavelength of the laser is moved to a nearby wavelength at which the OH  
 319 transition is not excited in order to enable measurement of a background signal for determination of ambient OH  
 320 concentrations (see, for example, Whalley et al. (2010)). Approximately 1 min is then required to scan the laser  
 321 wavelength over the OH transition to find the maximum OH fluorescence signal in the reference cell (Section 2.2).  
 322 OH reactivity measurements are thus taken during the 5 min online period, and data from successive measurement  
 323 cycles during each online period are co-added to improve the signal-to-noise ratio. Figure 3 shows typical OH decays  
 324 derived from the co-addition of data recorded throughout 5 min online periods during the Clean Air for London  
 325 (ClearLo) campaign in summer 2012.

326 Measurements of OH reactivity may also be made independently of any other use of the 308 nm probe laser, in  
 327 which case the timescale over which successive measurement cycles are co-added may be selected as desired, with  
 328 the laser periodically scanned over the OH transition to ensure that the maximum OH signal is obtained.

### 329 2.5 Calibration of the FAGE Detection Cell

330 Calibration of the detection cell, although not strictly necessary for measurements of OH reactivity, is required to  
 331 ensure that pseudo-first-order conditions are met in the reaction cell (i.e. combined concentrations of OH sinks are  
 332 in excess over the OH concentration) and provides a means to determine any potential interferences from  
 333 production of OH via ambient HO<sub>2</sub> + NO in the reaction cell (from ambient HO<sub>2</sub> which may survive the sampling inlet,  
 334 and any HO<sub>2</sub> generated following oxidation of OH sinks in the instrument) and to monitor potential changes in  
 335 instrument sensitivity with time.

336 The calibration procedure has been described in detail by Commane et al. (2009). Production of OH (and HO<sub>2</sub>) is  
 337 achieved through reaction R1 (and HO<sub>2</sub> through reaction R2) by passing a turbulent flow of humidified ultra-high  
 338 purity air (BTCA 178, BOC Special Gases) across a low pressure mercury vapour lamp to photolyse water vapour at  $\lambda$   
 339 = 184.9 nm.



341 The concentration of OH is given by Equation 5:

$$342 \quad [\text{OH}] = [\text{H}_2\text{O}] \sigma_{\text{H}_2\text{O}} \varphi_{\text{OH}} F \delta t \quad (\text{Equation 5})$$

343 where  $\sigma_{\text{H}_2\text{O}}$  is the absorption cross-section of H<sub>2</sub>O at 184.9 nm  $(7.1 \pm 0.2) \times 10^{-20} \text{ cm}^2$  (Cantrell et al., 1997; Creasey et  
 344 al., 2000),  $\varphi_{\text{OH}}$  is the quantum yield for OH production ( $\varphi_{\text{OH}} = 1$ ),  $F$  is the photon flux of the mercury lamp at 184.9 nm  
 345 and  $\delta t$  is the residence time in the photolysis region. The product  $F \delta t$  is determined by N<sub>2</sub>O actinometry (Commane  
 346 et al., 2009), with  $F$  varied by changing the current supplied to the lamp, and  $\delta t$  controlled by the flow rate of the gas  
 347 used in the calibration. The concentration of water vapour in the flow is determined by diverting a small known flow  
 348 of the air to a dew point hygrometer (CR4, Buck Research Instruments).

349 The calibration for OH was conducted over a range of mercury lamp fluxes, giving a calibration factor ( $C_{\text{OH}}$ ) of  $(2.13 \pm$   
 350  $0.27) \times 10^{-8} \text{ counts s}^{-1} \text{ molecule}^{-1} \text{ cm}^3 \text{ mW}^{-1}$  and indicating an initial OH concentration of  $\sim 10^9 \text{ cm}^{-3}$  produced in the  
 351 reaction cell by the 266 nm photolysis laser during measurements of OH reactivity in ambient air. The instrumental





352 limit of detection for OH radicals was determined to be  $\sim 10^7 \text{ cm}^{-3}$ , enabling observation of sufficient changes in OH  
 353 radical concentrations in the reaction cell to allow measurements of ambient OH reactivity.

354

### 355 3. Determination of OH Reactivity

356 The observed pseudo-first-order rate coefficient for OH loss ( $k_{\text{loss}}$ ) is determined by least-squares fitting Equation 6 to  
 357 the time-resolved OH decay:

$$358 \quad S_{\text{OH},t} = S_{\text{OH},0} \exp(-k_{\text{loss}}t) + b \quad (\text{Equation 6})$$

359 where  $S_{\text{OH},t}$  is the fluorescence signal at time  $t$  after firing of the 266 nm photolysis laser,  $S_{\text{OH},0}$  is the fluorescence  
 360 signal at time zero (i.e. immediately following firing of the 266 nm laser and production of OH in the reaction cell),  
 361  $k_{\text{loss}}$  is the observed rate coefficient for loss of the fluorescence signal,  $t$  is the time since firing of the 266 nm  
 362 photolysis laser and  $b$  is the background fluorescence signal measured by the PMT averaged for the 1 s prior to firing  
 363 of the photolysis laser (typically zero). Values for  $S_{\text{OH},0}$  and  $k_{\text{loss}}$  are permitted to vary in the fitting process. Figure 3  
 364 shows typical fits of Equation 6 to measurements of OH reactivity made in ambient air.

365 The value for  $k_{\text{loss}}$  determined from the fit contains a contribution from  $k'_{\text{OH,obs}}$ , the rate coefficient for OH loss owing  
 366 to chemical losses of OH in the reaction cell (the OH reactivity), and  $k_{\text{phys}}$ , the instrument 'zero' corresponding to the  
 367 rate coefficient for physical losses of OH owing to diffusion out of the sampling volume and heterogeneous losses on  
 368 the walls on the reaction cell. The chemical loss of OH in the reaction cell is thus given by Equation 7, and in order to  
 369 determine the OH reactivity from measurements of  $k_{\text{loss}}$  it is therefore essential to characterise  $k_{\text{phys}}$  (Section 3.1).

$$370 \quad k'_{\text{OH,obs}} = k_{\text{loss}} - k_{\text{phys}} \quad (\text{Equation 7})$$

371 At low ambient concentrations of ozone (< 10 ppb) and in laboratory experiments (Sections 4, 6 and 7) and  
 372 measurements of  $k_{\text{phys}}$ , it was necessary to add a small flow of humidified air containing a constant mixing ratio of  
 373 ozone ( $\sim 50$  ppb) to the main air flow sampled in order to produce sufficient OH radicals in the reaction cell. This  
 374 'non-ambient' ozone added to the reaction cell results in a small loss of OH owing to the reaction of  $\text{O}_3$  with OH, but  
 375 is expected to be  $< 0.1 \text{ s}^{-1}$  at 298 K (Section 2.1). However, addition of the small ozone-containing air flow (0.5 slm)  
 376 to the sampled flow of ambient air (12 slm) does require a correction for the dilution of the ambient air flow, such  
 377 that the OH reactivity ( $k'_{\text{OH}}$ ) is given by Equation 8:

$$378 \quad k'_{\text{OH}} = k'_{\text{OH,obs}} (1+f) \quad (\text{Equation 8})$$

379 where  $f$  is the dilution factor of the ambient air flow, given by the ratio of the small ozone-containing flow rate to the  
 380 total flow rate of the air in the reaction cell ( $\sim 0.04$  for the conditions used in this work). Potential errors arising from  
 381 errors in measurements of  $k_{\text{phys}}$  and  $f$  have been included in overall reported errors for  $k'_{\text{OH}}$ .

#### 382 3.1 Determination of $k_{\text{phys}}$

383 Determination of  $k_{\text{phys}}$  is critical to the evaluation of the true OH reactivity from observations of the total OH loss rate  
 384 in the instrument (Equations 7 & 8), and requires the measurement of the OH loss rate in the absence of any  
 385 chemical removal processes such that  $k_{\text{loss}}$  is equal to  $k_{\text{phys}}$ . To minimise the chemical losses of OH in the reaction cell  
 386 (and thus to minimise  $k'_{\text{OH,obs}}$ ) the loss of OH in the instrument is measured in ultra-high purity air (BTCA 178, BOC  
 387 Special Gases) passed through scrubbers (Gatekeeper Gas Purifiers) to remove  $\text{H}_2$ , CO and  $\text{CO}_2$  to sub-ppb levels.  
 388 Despite the use of scrubbed ultra-high purity air, low levels of residual VOCs can remain in the air, leading to  
 389 chemical losses. Such residual VOCs in the scrubbed ultra-high purity air have been quantified by gas-  
 390 chromatography and their contributions ( $< 1 \text{ s}^{-1}$ ) to the observed OH loss subtracted.

391 Furthermore, humidification and addition of a small amount of  $\text{O}_3$  to the ultra-high purity air are necessary for the  
 392 production of OH in the instrument during experiments to determine  $k_{\text{phys}}$ . Approximately 50 ppb of  $\text{O}_3$  is added to



393 ensure production of sufficient OH, leading to a chemical loss of  $< 0.1 \text{ s}^{-1}$  at 298 K through the reaction of OH with O<sub>3</sub>.  
 394 Moreover, despite the use of purified water for humidification, obtained using a water purification system (PURELAB  
 395 flex PRIPLB0163, Elga LabWater, Veolia Water Solutions & Technologies), impurities in the water can lead to  
 396 significant chemical losses for OH and the components in the purification system must remain uncontaminated in  
 397 order to ensure accurate determinations of  $k_{\text{phys}}$ .

398 Determination of  $k_{\text{phys}}$  in the laboratory and in the field for the ClearLo campaign in London in 2012 gave an average  
 399 value  $(1.1 \pm 1.0) \text{ s}^{-1}$  and  $(1.25 \pm 0.42) \text{ s}^{-1}$  for the campaign in York in 2014 (Section 5).

400

#### 401 **4. Instrumental validation via measurements of $k_{\text{OH}+\text{CO}}$ and $k_{\text{OH}+\text{CH}_4}$**

402 As a real time technique, the accuracy of the time axis during which the OH decay is obtained is determined by the  
 403 accuracy of the delay generators used to trigger the lasers and other delays (as described in Section 2.4 and shown in  
 404 Figure 2), which should be absolute within 1 ps. Hence the method should be absolute in terms of the time  
 405 separation between points in the decay. However, owing to various reasons, for example the appropriateness of the  
 406 function used to fit the decay, or any recycling of OH from oxidation products (for example the reaction of HO<sub>2</sub> with  
 407 NO), it is prudent to characterise the instrument through the use of known concentrations of reactants for which the  
 408 rate coefficient with OH is also well known. In order to validate measurements of ambient OH reactivity, the well-  
 409 known rate coefficients for reactions of OH with CO and CH<sub>4</sub> were both measured under pseudo-first-order  
 410 conditions using the instrumental setup described above. Ultra-high purity air (BTCA 178, BOC Special Gases) was  
 411 mixed with an excess of either CO (5 % in N<sub>2</sub>, BOC Special Gases) or CH<sub>4</sub> (BOC, CP grade, 99.5%), producing a main  
 412 flow of 11.5 slm with known concentrations of CO or CH<sub>4</sub>, prior to mixing with a small flow of humidified air (0.5 slm)  
 413 containing  $\sim 50$  ppb O<sub>3</sub> generated by passing the air flow across a mercury vapour lamp.

414 Figure 4 shows the OH reactivity, determined by fitting Equation 6 to the OH decay and subtracting  $k_{\text{phys}}$  (Section 3),  
 415 for a series of CO and CH<sub>4</sub> concentrations. The bimolecular rate coefficients for OH + CO ( $k_{\text{OH}+\text{CO}}$ ) and OH + CH<sub>4</sub>  
 416 ( $k_{\text{OH}+\text{CH}_4}$ ), determined at 298 K from the relationships  $k'_{\text{OH}} = k_{\text{OH}+\text{CO}}[\text{CO}]$  and  $k'_{\text{OH}} = k_{\text{OH}+\text{CH}_4}[\text{CH}_4]$ , were found to be  $(2.4 \pm$   
 417  $0.2) \times 10^{-13} \text{ cm}^3 \text{ s}^{-1}$  and  $(6.4 \pm 0.6) \times 10^{-15} \text{ cm}^3 \text{ s}^{-1}$ , respectively (errors are  $1\sigma$ ). The values for  $k_{\text{OH}+\text{CO}}$  and  $k_{\text{OH}+\text{CH}_4}$   
 418 determined here are in agreement with the literature values of  $(2.3 \pm 0.6) \times 10^{-13} \text{ cm}^3 \text{ s}^{-1}$  and  $(6.4 \pm 1.3) \times 10^{-15} \text{ cm}^3 \text{ s}^{-1}$   
 419 at 298 K (Atkinson et al., 2004), respectively, providing confidence in measurements of ambient OH reactivity.

420

#### 421 **5. Field measurements**

422 The laser flash photolysis OH reactivity instrument was deployed at the North Kensington measurement site (51° 31'  
 423 N, 0° 12' W) during the Clean Air for London (ClearLo) summer campaign in July and August 2012 (Bohnenstengel et  
 424 al., 2014), with near-continuous measurements made from the 21<sup>st</sup> July to 18<sup>th</sup> August 2012, alongside FAGE  
 425 measurements of OH, HO<sub>2</sub> and RO<sub>2</sub> radical concentrations. Measurements of O<sub>3</sub>, CO, NO, NO<sub>2</sub>, HONO, VOCs and  
 426 aerosol mass and composition were also made at the site during the campaign.

427 Figure 5 shows the full time series of measured OH reactivity for the campaign. The observed reactivity was highest  
 428 for air masses that had previously passed over central London (24<sup>th</sup>-27<sup>th</sup> July (Julian days 206-209) and 8<sup>th</sup>-10<sup>th</sup>  
 429 August (Julian days 221-223)), with a maximum reactivity of  $116 \text{ s}^{-1}$  recorded during rush hour on 24<sup>th</sup> July 2012.  
 430 Measurements taken on the 25<sup>th</sup> July 2012 (Julian day 207) are shown in Figure 6 to highlight the capability of the  
 431 instrument, and the average diurnal profile for the campaign is shown in Figure 7. A peak reactivity, on average, of  
 432  $\sim 27 \text{ s}^{-1}$  was observed during morning rush hour, with a minimum of  $\sim 15 \text{ s}^{-1}$  during the afternoon and a second peak  
 433 during evening rush hour. Detailed analysis of these data, including model calculations using the Master Chemical  
 434 Mechanism constrained to observed concentrations of long-lived species, is described by Whalley et al. (2015), and  
 435 the data are shown here only to highlight the instrument capability.



436 Field measurements of OH reactivity have also been made at a site at the University of York (53° 56' N, 1° 02' W)  
437 from the 19<sup>th</sup> May to 16<sup>th</sup> June 2014, approximately 3 km south-east of the centre of York and 2 km west of a major  
438 road, with a small wooded area immediately to the east, and thus subject to anthropogenic emissions and local  
439 biogenic emissions. Figure 8 shows the average diurnal during this period shown. The observed OH reactivity was  
440 typically lower than that observed during the ClearLo campaign, with a maximum in the diurnal average of  $\sim 6 \text{ s}^{-1}$ .  
441 Measurements of O<sub>3</sub>, CO, NO, NO<sub>2</sub>, VOCs and were also made at the site during this period, alongside measurements  
442 by a new instrument coupling an OH reactor to measurements of VOCs by gas chromatography with time of flight  
443 mass spectrometry (GC-TOFMS) to aid identification of any 'missing' OH reactivity. Detailed analysis of the results  
444 will be given in future publications.

445

## 446 6. Chamber measurements

447 The field instrument described above has also been modified in order to interface to the Highly Instrumented  
448 Reactor for Atmospheric Chemistry (HIRAC) to enable measurements of OH reactivity during VOC oxidation under  
449 controlled conditions. For complex reaction mechanisms, the oxidation pathway followed will have a characteristic  
450 time-evolution of the reactivity as secondary products are generated, and measurement of OH reactivity and  
451 comparison with a model prediction provides greater constraint for experimental determination of the mechanism.

452 HIRAC is a 2.25 m<sup>3</sup> stainless steel chamber equipped with UV photolysis lamps to initiate photochemistry and a  
453 comprehensive suite of analytical instrumentation, including gas chromatography (GC), Fourier transform infrared  
454 (FT-IR) spectroscopy, cavity ringdown spectroscopy (CRDS) and LIF-FAGE for radical measurements. Photolysis lamps  
455 within the chamber enable initiation of photochemistry, and experiments can be conducted at temperatures  
456 between 203 and 343 K and pressures up to 760 Torr (Glowacki et al., 2007; Malkin et al., 2010; Winiberg et al.,  
457 2015).

458 Gas is sampled from HIRAC through ½ " PTFE tubing at a flow rate of 1 slm and diluted with 5 slm of ultra-high purity  
459 air immediately on exiting the chamber, then diluted further with 9 slm of humidified ultra-high purity air and 1 slm  
460 of ultra-high purity air passed over a low pressure Hg lamp in order to generate O<sub>3</sub>, giving a total flow of 16 slm and  
461 hence a dilution factor of 1:16. The diluted gas flow, containing  $\sim 45 \text{ ppb O}_3$ , is then directed into the reaction cell of  
462 the OH reactivity instrument, with instrument operation and analysis as described in Sections 2 and 3 (including the  
463 correction of observed reactivity for dilution of sampled gas from the chamber using Equation 8, which is significant  
464 for these experiments in order to avoid measurement of high reactivities (Section 7) and to reduce the volume of gas  
465 removed from the chamber for the reactivity measurements).

466 Experiments were conducted to verify the sampling procedure by filling HIRAC with air containing known  
467 concentrations of a reactive gas, with a well-characterised rate coefficient for reaction with OH, followed by  
468 measurement of the OH reactivity in the chamber. Determination of the pseudo-first-order rate coefficients  
469 describing the OH loss for each of the given reactive gas concentrations in the chamber enabled determination of  
470 the bimolecular rate coefficient for reaction of the reactive gas with OH for comparison with literature values, as for  
471 experiments described in Section 4. Figure 4c shows the bimolecular plots for experiments in which the chamber  
472 was filled with *n*-butanol (*n*-C<sub>4</sub>H<sub>9</sub>OH) in air at total pressures of 760 Torr and temperatures of 298 K. A bimolecular  
473 rate coefficient of  $(8.5 \pm 0.1) \times 10^{-12} \text{ cm}^3 \text{ s}^{-1}$  was obtained, in comparison to the literature value of  $(8.5 \pm \frac{3.5}{2.5}) \times 10^{-12}$   
474  $\text{cm}^3 \text{ s}^{-1}$  (Atkinson et al., 2004), thus indicating the validity of the sampling procedure. No dependence of the  
475 observed reactivity on the total flow rate through the instrument was observed between 10 and 22 slm.

476 The coupling of OH reactivity measurements to chamber studies will enable detailed assessment of our  
477 understanding of the chemistry of secondary products in complex oxidation mechanisms by providing increased  
478 constraint on oxidation budgets during chamber experiments, and will be explored further in future work.

479



## 480 7. Effects of averaging time and future improvements to sampling

481 Measurements have also been made in HIRAC to investigate the effect of the averaging time on the measured OH  
482 reactivity. Figure 9 shows the observed OH reactivities, for a given set of conditions, as a function of the averaging  
483 time, showing successful measurements with an averaging time of 10 s and indicating the potential for further  
484 improvements for future integration of ambient OH reactivity observations with flux measurements.

485 Experiments described in this work using known concentrations of reactive gases have been able to reproduce  
486 recommended literature values for known rate coefficients, indicating the validity of the technique described here  
487 over the dynamic ranges investigated. However, recent work in Leeds has shown that, at higher reactivities,  
488 observed kinetics can be influenced by sampling issues related to the transport time of sampled gas from the pinhole  
489 nozzle to the point at which fluorescence is excited and detected in the FAGE cell, leading to underestimations of  
490 very high reactivities (Stone et al., 2016). Successful measurements in HIRAC with known concentrations of CH<sub>4</sub> at  
491 reactivities > 150 s<sup>-1</sup> (Section 6), although scattered owing to the small number of time points over which fast decays  
492 can be measured, indicate that such effects should be minimal for the instrument described in this work, even for  
493 the highest reactivities observed during the ClearLo campaign (> 100 s<sup>-1</sup>). However, experiments incorporating OH  
494 reactivity measurements in chamber studies, such as those described in Section 6, must also ensure that the gas  
495 sampled from the chamber has been sufficiently diluted so as to avoid the measurement of high reactivities directly.

496 Future work will incorporate a new inlet designed to minimise the distance between the pinhole nozzle and the  
497 point of excitation fluorescence and detection, ideally such that detection occurs within the supersonic jet formed  
498 on expansion of the gas as it flows through the pinhole. The new inlet will not only increase the dynamic range over  
499 which reactivity measurements can be made, but sampling within the supersonic jet will also lead to increased  
500 signal-to-noise and enable further reductions in the averaging time required to achieve adequate signal-to-noise for  
501 measurements with high time resolution.

502

## 503 8. Conclusions and Outlook

504 In this work we present the design and characterisation of an instrument to make field and chamber measurements  
505 of OH reactivity by laser flash photolysis (LFP) coupled with laser-induced fluorescence (LIF) using the FAGE  
506 technique. The LFP-LIF reactivity instrument, its operation and data analysis have been described in detail. Ambient  
507 reactivity measurements obtained during field campaigns in London, UK, and York, UK, have been presented, and  
508 will be discussed further in future work. The instrument has also been coupled to an atmospheric chamber, and  
509 preliminary results have been shown to demonstrate the potential for reactivity measurements during future  
510 chamber experiments.

511 Reactivity measurements have been made using an averaging time of 10 s, indicating potential for integration of  
512 ambient OH reactivity observations with flux measurements. Future development of the instrument will increase  
513 the dynamic range over which measurements can be made and will enable reduced averaging times owing to  
514 improvements in the signal-to-noise ratio.

515

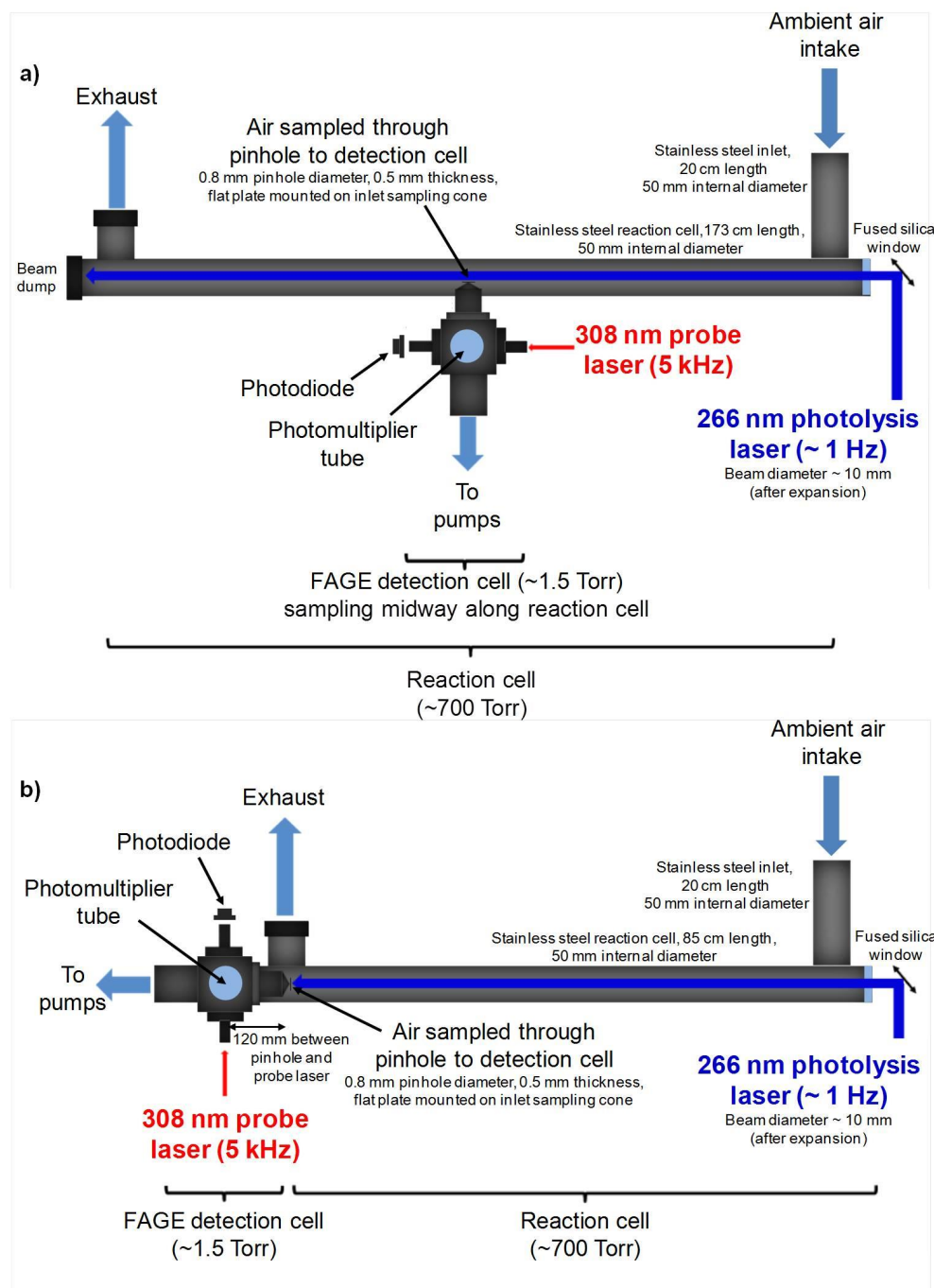
## 516 Acknowledgements

517 This work was supported by the National Environment Research Council (NERC) under grants NE/H003193/1,  
518 NE/J008990/1 and NE/L010798/1. DS is grateful to NERC for the award of an Independent Research Fellowship  
519 (NE/L010798/1). We are also grateful to the National Centre for Atmospheric Science (NCAS), which is funded by  
520 NERC, for ongoing support.

521



522 Figures

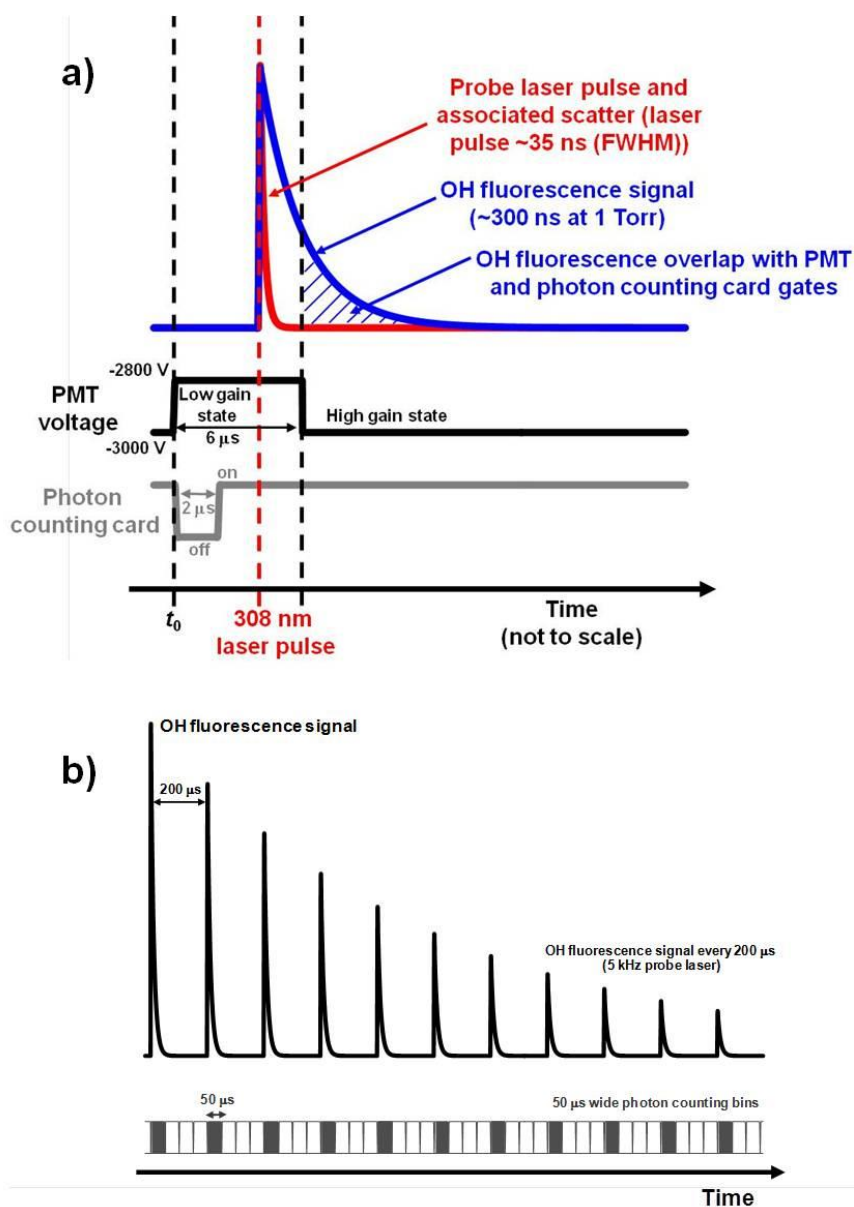


523

524 Figure 1: Schematic of the laser flash photolysis laser-induced fluorescence OH reactivity instrument for a)  
525 experiments in which the FAGE detection cell was situated midway along the reaction cell, resulting in sampling at  
526 90° to the air flow and leading to biexponential OH decays; b) experiments and field measurements in which the  
527 FAGE detection cell was situated at the end of the reaction cell and sampled from the centre of the photolysed  
528 volume, leading to OH decays described by a single exponential. All laboratory and field measurements shown in  
529 this work were obtained with the instrument configuration as shown in panel (b).



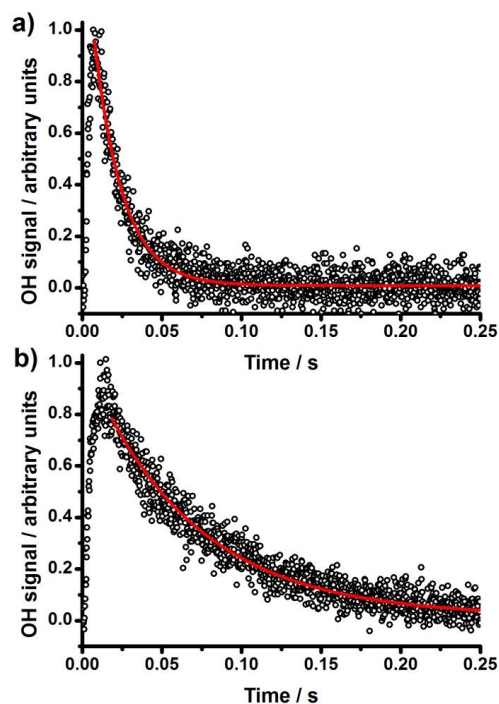
530



531

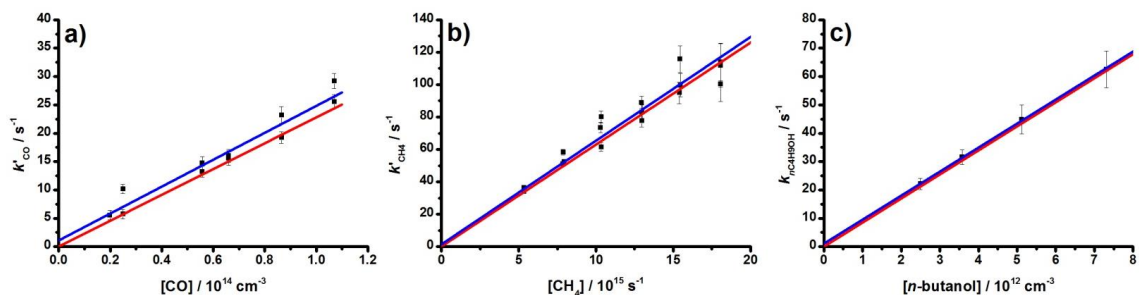
532 Figure 2: a) Schematic to illustrate the Stanford Research Systems delay generator controlled gate timing of the PMT  
 533 detector and photon counting card in the OH reactivity instrument. The blue hatched region indicates the overlap  
 534 between the OH fluorescence signal and the photon counting gates; b) Schematic to illustrate the photon counting  
 535 bin structure used to collect OH fluorescence photons after each 308 nm probe laser pulse (5kHz pulse repetition  
 536 frequency). Four 50  $\mu$ s wide photon counting bins cover the time period between each 308 nm laser pulse, but only  
 537 the bins immediately after the laser pulse collect any fluorescence photons (shaded bins), and only the photon  
 538 counts from these bins are used to construct the OH decay.





539

540 Figure 3: Typical OH time profiles following photolysis of ambient air (mixed with a small flow of  $N_2/O_2/O_3/H_2O$ )  
 541 observed during the Clean Air for London (ClearLo) campaign (black points) with fits of Equation 6 (red lines) to the  
 542 LIF data to determine  $k'_{OH}$  for data recorded a) during a polluted period on 25<sup>th</sup> July 2012 ( $k'_{OH} = (46.6 \pm 0.2) s^{-1}$ ) and  
 543 b) during a cleaner period on 7<sup>th</sup> August 2012 ( $k'_{OH} = (13.9 \pm 0.1) s^{-1}$ ). Time zero is defined as the time at which  
 544 photolysis occurs. Decays represent data co-added throughout 5 min periods.

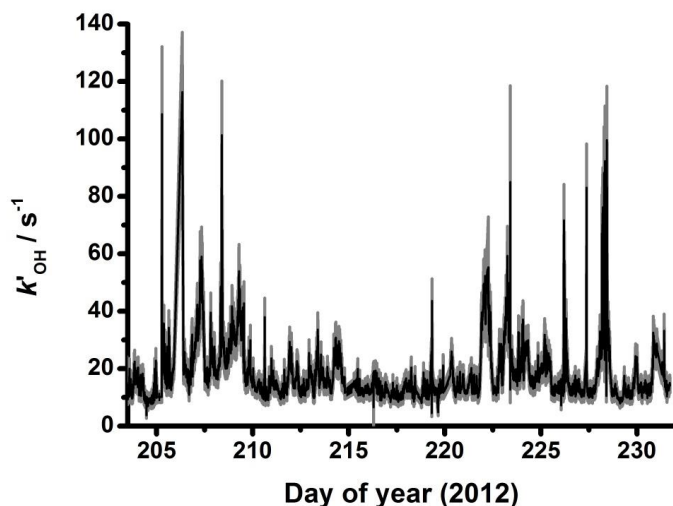


545

546 Figure 4: Bimolecular plots of pseudo-first-order rate coefficients describing OH loss ( $k'_{OH}$ ) against known  
 547 concentrations of reactive gases during laboratory tests (black points) with best fit lines (blue) and literature values  
 548 (red) for a) CO at 298 K; b)  $CH_4$  at 298 K; c)  $n$ -butanol ( $n-C_4H_9OH$ ) in chamber experiments at 298 K. Literature values  
 549 are taken from Atkinson et al. (2004). Note that values for  $k_{phys}$  have been subtracted from the observed rate  
 550 coefficients describing OH loss (Equation 7), and corrections for dilution have been applied (Equation 8). Errors are  
 551  $1\sigma$ .



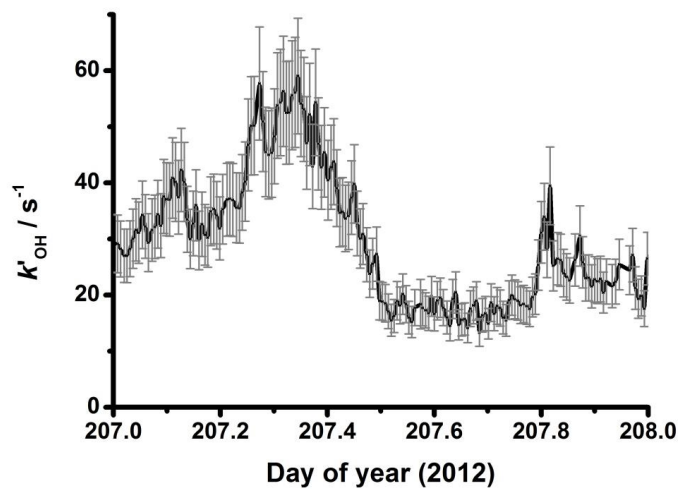
552



553

554 Figure 5: Time series of OH reactivity observed during the Clean Air for London (ClearfLo) campaign (21<sup>st</sup> July to 18<sup>th</sup>  
555 August 2012). Uncertainties (represented by the shaded area) represent 1 $\sigma$  combined uncertainties from the fits to  
556 observed OH decays (Equation 6), determinations of  $k_{\text{phys}}$  (Equation 7) and uncertainties in the dilution factor,  $f$   
557 (Equation 8).

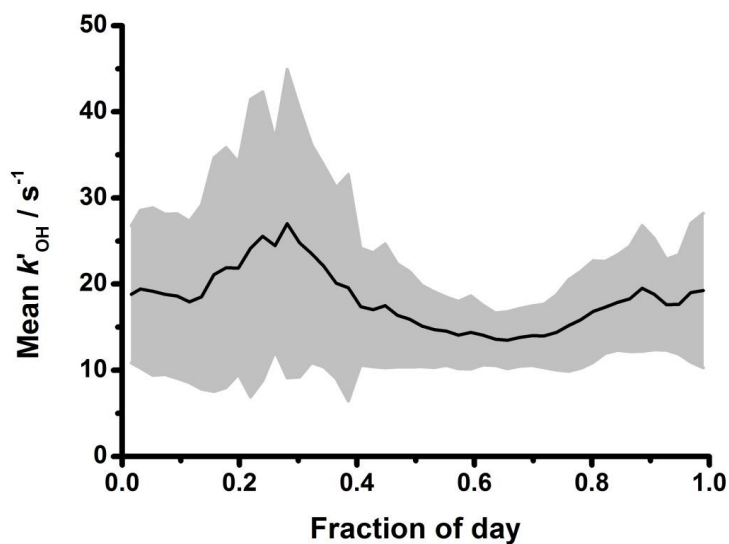
558



559

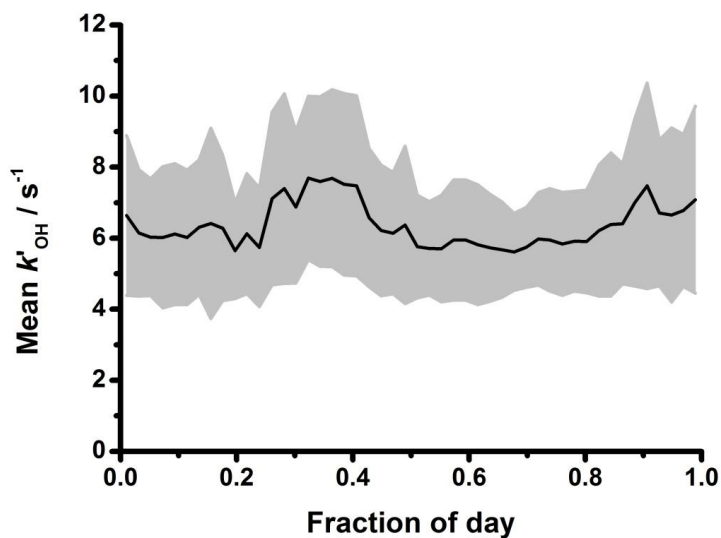
560 Figure 6: Time series of OH reactivity observed during the Clean Air for London (ClearfLo) campaign on the 25<sup>th</sup> July  
561 2012 (Julian day 207). Error bars represent 1 $\sigma$  combined uncertainties from the fits to observed OH decays  
562 (Equation 6), determinations of  $k_{\text{phys}}$  (Equation 7) and uncertainties in the dilution factor,  $f$  (Equation 8).

563



564

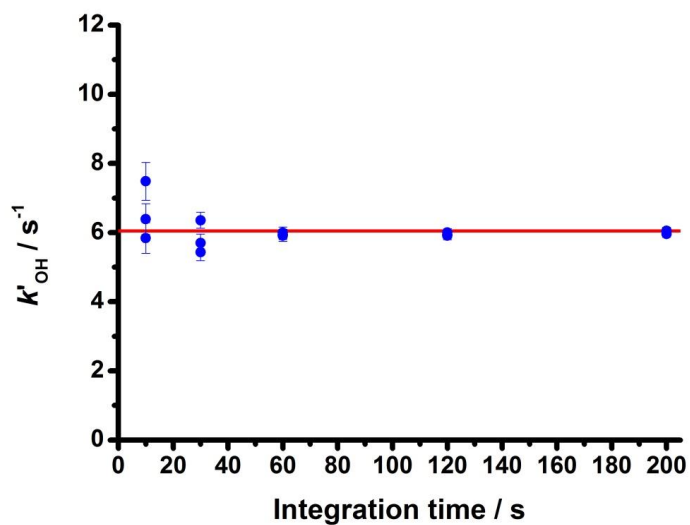
565 Figure 7: Diurnal average for observed OH reactivity (15 min averages) during the Clean Air for London (ClearLo)  
566 campaign (21<sup>st</sup> July to 18<sup>th</sup> August). Shaded areas represent the measurement variability from day-to-day during the  
567 campaign.



568

569 Figure 8: Diurnal average for observed OH reactivity (15 min averages) during in York (19<sup>th</sup> May to 16<sup>th</sup> June 2014).  
570 Shaded areas represent the measurement variability from day-to-day.

571



572

573 Figure 9: Observed OH reactivities, for a given gas composition (for which the expected OH reactivity is shown in  
574 red), as a function of averaging time, obtained using 15 mW of 308 nm probe laser power.

575

576

577

578

579

580

581

582 **References**

- 583 Amedro, D., Miyazaki, K., Parker, A., Schoemaceker, C., and Fittschen, C.: Atmospheric and kinetic studies of OH and  
584 HO<sub>2</sub> by the FAGE technique, *J. Environ. Sci.*, 24, 1, 78-86, 2012
- 585 Atkinson, R., Baulch, D.L., Cox, R.A., Crowley, J.N., Hampson, R.F., Hynes, R.G., Jenkin, M.E., Rossi, M.J., and Troe, J.:  
586 IUPAC Subcommittee for Gas Kinetic Data Evaluation, <http://iupac.pole-ether.fr/>, *Atmos. Chem. Phys.*, 4, 1461-1738,  
587 2004
- 588 Bloss, W.J., Gravestock, T.J., Heard, D.E., Ingham, T., Johnson, G.P., and Lee, J.D.: Application of a compact all solid-  
589 state laser system to the in situ detection of atmospheric OH, HO<sub>2</sub>, NO and IO by laser-induced fluorescence, *J.*  
590 *Environ. Monit.*, 5, 21–28, 2003
- 591 Bohnenstengel, S.I., Belcher, S.E., Aiken, A.C., Allan, J.D., Allen, G., Bacak, A., Bannan, T.J., Barlow, J.F., Beddows,  
592 D.C.S., Bloss, W.J., Booth, A.M., Chemel, C., Coceal, O., Di Marco, C.F., Mavendra, D.K., Faloon, K.H., Fleming, Z.,  
593 Furger, M., Gietl, J.K., Graves, R.R., Green, D.C., Grimmond, C.S.B., Halios, C., Hamilton, J.F., Harrison, R.M., Heal,  
594 M.R., Heard, D.E., Helfter, C., Herndon, S.C., Holmes, R.E., Hopkins, J.R., Jones, A.M., Kelly, F.J., Kotthaus, S.,  
595 Langford, B., Lee, J.D., Leigh, R., Lewis, A.C., Lidster, R.T., Lopez-Hilfiker, F., McQuaid, J.B., Mohr, C., Monks, P.S.,  
596 Nemitz, E., Ng, N.L., Percival, C.J., Prévôt, A.S.H., Ricketts, H.M.A., Sokhi, R., Stone, D., Thornton, J.A., Tremper, A.H.,  
597 Valach, A.C., Visser, S., Whalley, L.K., Williams, L.R., Xu, L., Young, D.E., and Zotter, P.: Meteorology, air quality, and  
598 health in London: The ClearLo project, *Bull. Am. Meteor. Soc.*, 96, 5, 779-804, [http://dx.doi.org/10.1175/BAMS-D-](http://dx.doi.org/10.1175/BAMS-D-12-00245.1)  
599 [12-00245.1](http://dx.doi.org/10.1175/BAMS-D-12-00245.1), 2015
- 600 Cantrell, C.A., Shetter, R.E., Calvert, J.G., Eisele, F.L. and Tanner, D.J.: Some considerations of the origin of nighttime  
601 peroxy radicals observed in MLOPEX, *J. Geophys. Res.-Atmos.*, 102, 15899–15913, 1997
- 602 Chatani, S., Shimo, N., Matsunaga, S., Kajii, Y., Kato, S., Nakashima, Y., Miyazaki, K., Ishii, K., and Ueno, H.: Sensitivity  
603 analyses of OH missing sinks over Tokyo metropolitan area in the summer of 2007, *Atmos. Chem. Phys.*, 9, 8975–  
604 8986, 2009
- 605 Commane, R., Floquet, C.F.A., Ingham, T., Stone, D., Evans, M.J., and Heard, D.E: Observations of OH and HO<sub>2</sub> radicals  
606 over West Africa, *Atmos. Chem. Phys.*, 10, 8783-8801, 2010
- 607 Creasey, D.J., Heard, D.E., and Lee, J.D.: Absorption Cross-section measurements of O<sub>2</sub> and H<sub>2</sub>O at 185 nm.  
608 Implications for the calibration of field instruments to measure OH, HO<sub>2</sub> and RO<sub>2</sub> radicals, *Geophys. Res. Lett.*, 27,  
609 1651-1654, 2000
- 610 Di Carlo, P., Brune, W.H., Martinez, M., Harder, H., Leshner, R., Ren, X., Thornberry, T., Carroll, M.A., Young, V.,  
611 Shepson, P. B., Reimer, D., Apel, E., and Campbell, C.: Missing OH Reactivity in a Forest: Evidence for Unknown  
612 Reactive Biogenic VOCs, *Science*, 304, 722–725, 2004
- 613 Dolgorouky, C., Gros, V., Sarda-Estève, R., Sinha, V., Williams, J., Marchand, N., Sauvage, S., Poulain, L., Sciare, J., and  
614 Bonsang, B.: Total OH reactivity measurements in Paris during the 2010 MEGAPOLI winter campaign, *Atmos. Chem.*  
615 *Phys.*, 12, 9593– 9612, doi:10.5194/acp-12-9593-2012, 2012
- 616 Edwards, P.M., Evans, M.J., Furneaux, K.L., Hopkins, J., Ingham, T., Jones, C., Lee, J.D., Lewis, A.C., Moller, S.J., Stone,  
617 D., Whalley, L.K., and Heard, D.E.: OH reactivity in a South East Asian tropical rainforest during the Oxidant and  
618 Particle Photochemical Processes (OP3) project, *Atmos. Chem. Phys.*, 13, 9497–9514, doi:10.5194/acp-13-9497-  
619 2013, 2013
- 620 Fuchs, H., Hofzumahaus, A., Rohrer, F., Bohn, B., Brauers, T., Dorn, H.-P., Haseler, R., Holland, F., Kaminski, M., Li, X.,  
621 Lu, K., Nehr, S., Tillmann, R., Wegener, R., and Wahner, A.: Experimental evidence for efficient hydroxyl radical  
622 regeneration in isoprene oxidation, *Nature Geoscience*, 6, 1023-1026, 2013



- 623 Glowacki, D.R., Goddard, A., Hemavibool, K., Malkin, T.L., Commane, R., Anderson, F., Bloss, W.J., Heard, D.E.,  
624 Ingham, T., Pilling, M.J., and Seakins, P.W.: Design and initial results from a Highly Instrumented Reactor for  
625 Atmospheric Chemistry, *Atmos. Chem. Phys.*, 7, 20, 5371-5390, 2007
- 626 Hansen, R.F., Griffith, S.M., Dusanter, S., Rickly, P.S., Stevens, P.S., Bertman, S.B., Carroll, M.A., Erickson, M.H., Flynn,  
627 J.H., Grossberg, N., Jobson, B.T., Lefer, B.L., and Wallace, H.W.: Measurements of total hydroxyl radical reactivity  
628 during CABINEX 2009 – Part 1: Field measurements, *Atmos. Chem. Phys.*, 14, 2923-2937, 2014
- 629 Ingham, T., Goddard, A., Whalley, L.K., Furneaux, K.L., Edwards, P.M., Seal, C.P., Self, D.E., Johnson, G.P., Read, K.A.,  
630 Lee, J.D., and Heard, D.E.: A flow-tube based laser-induced fluorescence instrument to measure OH reactivity in the  
631 troposphere, *Atmos. Meas. Tech.*, 2, 465-477, 2009
- 632 Kim, S., Guenther, A., Karl, T., and Greenberg, J.: Contributions of primary and secondary biogenic VOC to total OH  
633 reactivity during the CABINEX (Community Atmosphere-Biosphere INTERactions Experiments)-09 field campaign,  
634 *Atmos. Chem. Phys.*, 11, 8613-8623, doi:10.5194/acp-11-8613-2011, 2011
- 635 Kovacs, T.A. and Brune, W.H.: Total OH Loss Rate Measurement, *J. Atmos. Chem.*, 39, 105-122, 2001
- 636 Kovacs, T.A., Brune, W.H., Harder, H., Martinez, M., Simpas, J.B., Frost, G.J., Williams, E., Jobson, T., Stroud, C.,  
637 Young, V., Fried, A., and Wert, B.: Direct measurements of urban OH reactivity during Nashville SOS in summer 1999,  
638 *J. Environ. Monitor.*, 5, 68-74, 2003
- 639 Lee, J.D., Young, J.C., Read, K.A., Hamilton, J.F., Hopkins, J.R., Lewis, A.C., Bandy, B.J., Davey, J., Edwards, P., Ingham,  
640 T., Self, D.E., Smith, S.C., Pilling, M.J., and Heard, D.E.: Measurement and calculation of OH reactivity at a United  
641 Kingdom coastal site, *J. Atmos. Chem.*, 64, 53-76, 2009
- 642 Lou, S., Holland, F., Rohrer, F., Ku, K., Bohn, B., Brauers, T., Chang, C.C., Fuchs, H., Haseler, R., Kita, K., Kondo, Y., Li,  
643 X., Shao, M., Zeng, L., Wahner, A., Zhang, Y., Wang, W., and Hofzumahaus, A.: Atmospheric OH reactivities in the  
644 Pearl River Delta – China in summer 2006: measurements and model results, *Atmos. Chem. Phys.*, 10, 11243-11260,  
645 2010
- 646 Lu, K.D., Hofzumahaus, A., Holland, F., Bohn, B., Brauers, T., Fuchs, H., Hu, M., Haseler, R., Kita, K., Kondo, Y., Li, X.,  
647 Lou, S.R., Oebel, A., Shao, M., Zeng, L.M., Wahner, A., Zhu, T., Zhang, Y.H., and Rohrer, F.: Missing OH source in a  
648 suburban environment near Beijing: observed and modelled OH and HO<sub>2</sub> concentrations in summer 2006, *Atmos.*  
649 *Chem. Phys.*, 13, 1057-1080, 2013
- 650 Malkin, T.L., Goddard, A., Heard, D.E., and Seakins, P.W.: Measurements of OH and HO<sub>2</sub> yields from the gas phase  
651 ozonolysis of isoprene, *Atmos. Chem. Phys.*, 10, 3, 1441-1459, 2010
- 652 Mao, J., Ren, X., Brune, W.H., Olson, J.R., Crawford, J.H., Fried, A., Huey, L.G., Cohen, R.C., Heikes, B., Singh, H.B.,  
653 Blake, D.R., Sachse, G.W., Diskin, G.S., Hall, S.R., and Shetter, R.E.: Airborne measurement of OH reactivity during  
654 INTEX-B, *Atmos. Chem. Phys.*, 9, 163-173, 2009
- 655 Mao, J., Ren, X., Chen, S., Brune, W.H., Chen, Z., Martinez, M., Harder, H., Lefer, B., Rappenglück, B., Flynn, J., and  
656 Leuchner, M.: Atmospheric oxidation capacity in the summer of Houston 2006: Comparison with summer  
657 measurements in other metropolitan studies, *Atmos. Environ.*, 44, 4107-4115, 2010
- 658 Martinez, M., Harder, H., Kovacs, T.A., Simpas, J.B., Bassis, J., Leshner, R., Brune, W.H., Frost, G.J., Williams, E.J.,  
659 Stroud, C.A., Jobson, B.T., Roberts, J.M., Hall, S.R., Shetter, R.E., Wert, B., Fried, A., Alicke, B., Stutz, J., Young, V.L.,  
660 White, A.B., and Zamora, R.J.: OH and HO<sub>2</sub> concentrations, sources, and loss rates during the Southern Oxidants  
661 Study in Nashville, Tennessee, summer 1999, *J. Geophys. Res. Atmos.*, 108, D19, 4617, 1-17,  
662 doi:10.1029/2003JD003551, 2003





- 663 Michoud, V., Hansen, R.F., Locoge, N., Stevens, P.S., and Dusanter: Detailed characterizations of the new Mines  
664 Douai comparative reactivity method instrument via laboratory experiments and modelling: Atmos. Meas. Tech., 8,  
665 3537-3553, doi:10.5194/amt-8-3537-2015, 2015
- 666 Mogensen, D., Smolander, S., Sogachev, A., Zhou, L., Sinha, V., Guenther, A., Williams, J., Nieminen, T., Kajos, M. K.,  
667 Rinne, J., Kulmala, M., and Boy, M.: Modelling atmospheric OH-reactivity in a boreal forest ecosystem, Atmos. Chem.  
668 Phys., 11, 9709– 9719, doi:10.5194/acp-11-9709-2011, 2011
- 669 Nakashima, Y., Kamei, N., Kobayashi, S., and Kajii, Y.: Total OH reactivity and VOC analyses for gasoline vehicular  
670 exhaust with a chassis dynamometer, Atmos. Environ., 44, 4, 468-475, 2010
- 671 Nakashima, Y., Tsurumaru, H., Imamura, T., Bejan, I., Wenger, J.C., and Kajii, Y.: Total OH reactivity measurements in  
672 laboratory studies of the photooxidation of isoprene, Atmos. Environ., 62, 243-247, 2012
- 673 Nakashima, Y., Kato, S., Greenberg, J., Harley, P., Karl, T., Turnipseed, A., Apel, E., Guenther, A., Smith, J., and Kajii, Y.:  
674 Total OH reactivity measurements in ambient air in a southern Rocky mountain ponderosa pine forest during  
675 BEACHON-SRM08 summer campaign, Atmos. Environ., 85, 1-8, 2014
- 676 Nölscher, A.C., Williams, J., Sinha, V., Custer, T., Song, W., Johnson, A.M., Axinte, R., Bozem, H., Fischer, H., Pouvesle,  
677 N., Phillips, G., Crowley, J.N., Rantala, P., Rinne, J., Kulmala, M., Gonzales, D., Valverde-Canossa, J., Vogel, A.,  
678 Hoffmann, T., Ouwersloot, H.G., Vilà-Guerau de Arellano, J., and Lelieveld, J.: Summertime total OH reactivity  
679 measurements from boreal forest during HUMPPA-COPEC 2010, Atmos. Chem. Phys., 12, 8257–8270,  
680 doi:10.5194/acp-12-8257-2012, 2012
- 681 Nolscher, A.C., Butler, T., Auld, J., Veres, P., Munoz, A., Taraborrelli, D., Vereecken, L., Lelieveld, J., and Williams, J.:  
682 Using total OH reactivity to assess isoprene photooxidation via measurement and model, Atmos. Environ., 89, 453-  
683 463, 2014
- 684 Parker, A.E., Amedro, D., Schoemaeker, C., and Fittschen, C.: OH radical reactivity measurements by FAGE, Environ.  
685 Eng. and Management J., 10, 1, 107-114, 2011
- 686 Ren, X., Harder, H., Martinez, M., Leshner, R.L., Oligier, A., Shirley, T., Adams, J., Simpasa, J.B., and Brune, W.H.: HO<sub>x</sub>  
687 concentrations and OH reactivity observations in New York City during PMTACS-NY2001, Atmos. Environ., 37, 3627–  
688 3637, 2003
- 689 Ren, X., Brune, W.H., Mao, J.Q., Mitchell, M.J., Leshner, R.L., Simpasa, J.B., Metcalf, A.R., Schwab, J.J., Cai, C.X., Li, Y.Q.,  
690 Demerjian, K.L., Felton, H.D., Boynton, G., Adams, A., Perry, J., He, Y., Zhou, X. L., and Hou, J.: Behavior of OH and  
691 HO<sub>2</sub> in the winter atmosphere in New York city, Atmos. Environ., 40, 252–263, 2006a
- 692 Ren, X., Brune, W.H., Oligier, A., Metcalf, A.R., Simpasa, J.B., Shirley, T., Schwab, J.J., Bai, C., Roychowdhury, U., Li, Y.,  
693 Cai, C., Demerjian, K.L., He, Y., Zhou, X., Gao, H., and Hou, J.: OH, HO<sub>2</sub>, and OH reactivity during the PMTACS-NY  
694 Whiteface Mountain 2002 campaign: Observations and model comparison, J. Geophys. Res., 111, D10S03,  
695 doi:10.1029/2005JD006126, 2006b
- 696 Sadanaga, Y., Yoshino, A., Watanabe, K., Yoshioka, A., Wakazono, Y., Kanaya, Y., and Kajii, Y.: Development of a  
697 Measurement System of OH reactivity in the atmosphere using a laser-induced pump and probe technique, Rev. Sci.  
698 Instrum., 75, 2648– 2655, 2004a
- 699 Sadanaga, Y., Yoshino, A., Kato, S., Yoshioka, A., Watanabe, K., Miyakawa, Y., Hayashi, I., Ichikawa, M., Matsumoto,  
700 J., Nishiyama, A., Akiyama, N., Kanaya, Y., and Kajii, Y.: The importance of NO<sub>2</sub> and volatile organic compounds in the  
701 urban air from the viewpoint of the OH reactivity, Geophys. Res. Lett., 31, L08102, doi:10.1029/2004gl019661,  
702 2004b



- 703 Shirley, T.R., Brune, W.H., Ren, X., Mao, J., Leshner, R., Cardenas, B., Volkamer, R., Molina, L.T., Molina, M.J., Lamb, B.,  
704 Velasco, E., Jobson, T., and Alexander, M.: Atmospheric oxidation in the Mexico City Metropolitan Area (MCMA)  
705 during April 2003, *Atmos. Chem. Phys.*, 6, 2753–2765, 2006
- 706 Sinha, V., Williams, J., Crowley, J.N., and Lelieveld, J.: The Comparative Reactivity Method – a new tool to measure  
707 total OH Reactivity in ambient air, *Atmos. Chem. Phys.*, 8, 2213–2227, 2008
- 708 Sinha, V., Williams, J., Diesch, J.M., Drewnick, F., Martinez, M., Harder, H., Regelin, E., Kubistin, D., Bozem, H.,  
709 HosaynaliBeygi, Z., Fischer, H., Andrés-Hernández, M.D., Kartal, D., Adame, J.A., and Lelieveld, J.: Constraints on  
710 instantaneous ozone production rates and regimes during DOMINO derived using in-situ OH reactivity  
711 measurements, *Atmos. Chem. Phys.*, 12, 7269–7283, doi:10.5194/acp-12-7269-2012, 2012
- 712 Stone, D., Blitz, M., Ingham, T., Onel, L., Medeiros, D.J. and Seakins, P.W.: An instrument to measure fast gas phase  
713 radical kinetics at high temperatures and pressures, submitted 2016
- 714 Whalley, L.K., Furneaux, K.L., Goddard, A., Lee, J.D., Mahajan, A., Oetjen, H., Read, K.A., Kaaden, N., Carpenter, L. J.,  
715 Lewis, A. C., Plane, J. M. C., Saltzman, E. S., Wiedensohler, A., and Heard, D. E.: The chemistry of OH and HO<sub>2</sub> radicals  
716 in the boundary layer over the tropical Atlantic Ocean, *Atmos. Chem. Phys.*, 10, 1555–1576, 2010
- 717 Whalley, L.K., Edwards, P.M., Furneaux, K.L., Goddard, A., Ingham, T., Evans, M.J., Stone, D., Hopkins, J.R., Jones, C.E.,  
718 Karunaharan, A., Lee, J.D., Lewis, A.C., Monks, P.S., Moller, S.J., and Heard, D.E.: Quantifying the magnitude of a  
719 missing hydroxyl radical source in a tropical rainforest, *Atmos. Chem. Phys.*, 11, 7223–7233, 2011
- 720 Whalley, L.K., Stone, D., Bandy, B., Dunmore, R., Hamilton, J.F., Hopkins, J., Lee, J.D., Lewis, A.C., and Heard, D.E.:  
721 Atmospheric OH reactivity in central London: observations, model predictions and estimates of in-situ ozone  
722 production, *Atmos. Chem. Phys. Discuss.*, 15, 31247–31286, 2015
- 723 Winiberg, F.A.F., Smith, S.C., Bejan, I., Brumby, C.A., Ingham, T., Malkin, T.L., Orr, S.C., Heard, D.E., and Seakins, P.W.:  
724 Pressure-dependent calibration of the OH and HO<sub>2</sub> channels of a FAGE HO<sub>x</sub> instrument using the Highly Instrumented  
725 Reactor for Atmospheric Chemistry (HIRAC), *Atmos. Meas. Tech.*, 8, 2, 523–540, 2015
- 726 Yoshino, A., Sadanaga, Y., Watanabe, K., Kato, S., Miyakawa, Y., Matsumoto, J., and Kajii, Y.: Measurement of total  
727 OH reactivity by laser-induced pump and probe technique – comprehensive observations in the urban atmosphere  
728 of Tokyo, *Atmos. Environ.*, 40, 7869–7881, 2006
- 729 Yoshino, A., Nakashima, Y., Miyazaki, K., Kato, S., Suthawaree, J., Shimo, N., Matsunaga, S., Chatani, S., Apel, E.,  
730 Greenberg, J., Guenther, A., Ueno, H., Sasaki, H., Hoshi, J., Yokota, H., Ishii, K., and Kajii, Y.: Air quality diagnosis from  
731 comprehensive observations of total OH reactivity and reactive trace species in urban central Tokyo, *Atmos.*  
732 *Environ.*, 49, 51–59, 2012
- 733