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## Lidar soundings of the mesospheric nickel layer using Ni(<sup>3</sup>F) and Ni(<sup>3</sup>D) transitions 2

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## Key Points:

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7	- First observation of the mesospheric Ni layer from meta-stable $\mathrm{Ni}(^{3}\mathrm{D})$ state re-
8	vealed peak densities of $\sim 280 - 450 \mathrm{cm}^{-3}$
9	• Compared to Fe and their respective abundance in CI-Chondrites, Ni is depleted
10	by a factor of about two
11	• Observations hint at faster-than-expected conversion of Ni into ions and neutral
12	reservoir molecules

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#### 13 Abstract

During six nights between January and March 2018 we observed the mesospheric Ni layer 14 by lidar from Kühlungsborn, Germany (54°N, 12°E). For most of the soundings we uti-15 lized for the first time a transition from the low-lying excited  $Ni(^{3}D)$  state at 341 nm. 16 For additional soundings we used the ground-state  $Ni({}^{3}F)$  transition at 337 nm, giving 17 similar results but a worse signal-to-noise ratio. We observed nightly mean Ni peak den-18 sities between  $\sim 280$  and  $450 \text{ cm}^{-3}$  and column abundances between  $3.1 \cdot 10^8 \text{ cm}^{-2}$  and 19  $4.9 \cdot 10^8 \,\mathrm{cm}^{-2}$ . Comparing with iron densities we get a Fe/Ni ratio of 38, which is a fac-20 tor of 2 larger than the ratio in CI-Chondrites and factor of 32 larger than the Fe/Ni ra-21 tio observed by the only previous measurement of mesospheric Ni [Collins et al., 2015]. 22 The underabundance of Ni compared to CI-Chondrites suggests that Ni is more efficiently 23 sequestered as Ni<sup>+</sup> or neutral reservoir species than Fe. 24

#### <sup>25</sup> 1 Introduction

The mesospheric metal layers produced by meteoric ablation between 80 and 110 km 26 were first detected  $\sim 90$  years ago by Na airglow observations [Slipher, 1929]. Subsequently, 27 resonance lidars have been used for active, altitude resolved soundings [Bowman et al., 28 1969], and most recently satellite-borne spectrometers, detecting scattered sunlight or 29 nightglow, have provided a near global coverage [e.g. Gumbel et al., 2007]. So far the al-30 kali metals sodium, lithium and potassium [e.g. Qian and Gardner, 1995; Jegou et al., 31 1980; Eska et al., 1998], alkaline earth metals calcium [e.g. Granier et al., 1985; Gerding 32 et al., 2000] and magnesium [e.g. Correira et al., 2008; Langowski et al., 2015] as well 33 as the transition metal iron [e.g. Bills and Gardner, 1990; Alpers et al., 1993; Chu et al., 34 2011] have been investigated. The metal ions Ca<sup>+</sup> and Mg<sup>+</sup> have also been observed by 35 ground-based lidar [e.g. Gerding et al., 2000] and from space [Langowski et al., 2015], re-36 spectively. Just as in the case of FeO chemiluminescence produced by the reaction Fe 37 and O<sub>3</sub> [Saran et al., 2011; Unterguggenberger et al., 2017], chemiluminescence from NiO 38 has also been detected [Evans et al., 2011]. Very recently, a first detection of atomic Ni 39 by lidar was published by *Collins et al.* [2015] (in the following abbreviated as CLM2015). 40 The major source of these metals is the ablation of interplanetary dust particles origi-41 nating from comets and asteroids [Carrillo-Sánchez et al., 2016]. However, it has been 42 clear even from early soundings that the relative abundances of the metals in the me-43 sospheric layers can be quite different from their relative abundances in CI-Chondrites 44

[Plane, 1991; Gerding et al., 1999; Raizada et al., 2004; Höffner and Friedman, 2005; Yi
et al., 2009]. Clearly, factors including differential ablation, chemistry and dynamics play
important roles in determining the relative metal concentrations and their temporal and
spatial variations [e.g. Plane et al., 2015].

For a metallic species in the mesosphere to be detectable by lidar from the ground, 49 the product of resonance backscatter cross section and metal density must be large enough 50 to produce a sufficient signal-to-noise ratio (SNR). Furthermore, the wavelength needs 51 to be in the transmission range of the atmosphere (i.e. at wavelengths longer than the 52  $\sim 300 \,\mathrm{nm}$  cut-off caused by the stratospheric O<sub>3</sub> layer). For nickel, the expected signal 53 should be weak, because the relative Ni abundance should be low based on their chon-54 dritic abundances, around 1/18 that of Fe [Asplund et al., 2009]. In fact, the first and 55 so far only observations of the Ni layer are reported by CLM2015 from Chatanika, Alaska 56  $(65^{\circ}N, 147^{\circ}W)$ . The peak Ni concentration was found to be  $\sim 16,000 \text{ cm}^{-3}$ . This is sur-57 prisingly large, within a factor of 2 of the Fe concentration measured at the same loca-58 tion and time of year. 59

In early 2018 we performed Ni observations during six nights at Kühlungsborn, Germany (54°N, 12°E), using Ni resonance transitions at 337 nm for ground-state Ni(<sup>3</sup>F), and at 341 nm to probe the low-lying Ni(<sup>3</sup>D) state (CLM2015 used only the 337 nm transition). In Section 2 we describe the lidar setup and the selection of resonance lines. The observations are presented in Section 3, followed by a discussion of the results (Section 4).

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#### 2 Lidar setup and selection of resonance lines

In order to produce laser emission at the Ni resonance wavelengths at  $\lambda_{air} = 336.9563$  nm 66 and  $\lambda_{air} = 341.4764 \,\mathrm{nm}$  (see Fig. S1 in the Supporting Information (SI)), we used an excimer-67 pumped (XeCl) dye laser with a repetition rate of 30 pps, as described in Gerding et al. 68 [2000], but here we combined it with a different receiver. As in CLM2015, we used p-69 Terphenyl (PTP) dye dissolved in p-Dioxane for operation at 337 nm or 341 nm [Brack-70 mann, 1994]. The oscillator of the dye laser was equipped with an intracavity etalon to 71 limit the spectral bandwidth to 0.4 pm. Using two amplifier stages the dye laser pulse 72 energy was up to 4 mJ at 337 nm, and 14 mJ at 341 nm. For some soundings a waveme-73 ter (High Finesse WS6-200) was used for wavelength calibration and adjustment. The 74 backscattered light was collected by a 78 cm telescope and guided to the detection bench 75

-3-

<sup>76</sup> by a quartz fiber. The detector was equipped with a 29 nm (full width at half maximum)
<sup>77</sup> interference filter (IF) with ~85% transmission at the two wavelengths, and a Hamamatsu
<sup>78</sup> R7600U-200 photomultiplier tube with ~40% quantum efficiency. The comparatively wi<sup>79</sup> deband IF therefore required lidar soundings during moonless nights, but the high trans<sup>80</sup> mission together with the PMT specification resulted in a large SNR.

The first successful observations of the Ni layer above Kühlungsborn were made 81 with the Ni(<sup>3</sup>D) transition at  $\lambda_{air} = 341.4764$  nm. This wavelength is closer to the emis-82 sion maximum of the dye than the Ni(<sup>3</sup>F) transition at  $\lambda_{air} = 336.9563$  nm, yielding a 83 better wavelength stability, less broadband emission, and larger laser power. Later on, 84 we tuned the laser to the  $Ni(^{3}F)$  resonance line at  $337 \,\mathrm{nm}$ . Comparisons of the soundings 85 at both transitions are presented in Section 3.3. The (effective) differential backscatter 86 cross sections are calculated as described by Fricke and von Zahn [1985] and Chu and 87 Papen [2005] using oscillator strengths of 0.12 for Ni( $^{3}D$ ) and 0.024 for Ni( $^{3}F$ ) [Kramida 88 *et al.*, 2018]. This yields  $\sigma_{res} = 1.08 \cdot 10^{-17} \,\mathrm{m^2/sr}$  ( $\sigma_{res} = 2.12 \cdot 10^{-18} \,\mathrm{m^2/sr}$ ) and  $\sigma_{Ray} = 3.57 \cdot 10^{-31} \,\mathrm{m^2/sr}$ 89  $(\sigma_{Ray}=3.75\cdot10^{-31}\,\mathrm{m^2/sr})$  for resonance and Rayleigh backscatter at 341 nm (337 nm), 90 respectively. Numbers are given for 0.4 pm laser full width at half maximum (FWHM, 91 assuming Lorentz shape) and 200 K atmospheric temperature at  $\sim 90$  km altitude in win-92 ter. Later on we used a corrected effective cross section for 337 nm, see Section 4. Note 93 that the  $Ni(^{3}D)$  resonance cross section is about five times larger than that for  $Ni(^{3}F)$ . 94 The Ni densities at altitude z are calculated as usual by the equation 95

$$\rho_{Ni}(z) = \rho_{air}(z_R) \cdot \sigma_{Ray} / \sigma_{res} \cdot N_{res}^{corr}(z) / N_{Ray}^{corr}(z_R)$$

with  $N_{res}^{corr}$  the range-corrected resonance count rate,  $N_{Ray}^{corr}(z_R)$  the range-corrected 97 Rayleigh count rate at the reference altitude,  $\rho_{air}(z_R)$  air density at the reference alti-98 tude taken from NRLMSISE-00 [Picone et al., 2002], and  $z_R$  reference altitude chosen 99 as 50 km, i.e. avoiding corrections for stratospheric aerosol backscatter and ozone ab-100 sorption. An additional factor is then applied which takes account of the thermal po-101 pulations of the  ${}^{3}D_{3}$  and  ${}^{3}F_{4}$  states used here. At typical upper mesospheric tempera-102 ture of 200 K these are 15.1 and 84.9%, respectively. The calculation of these fractions 103 is explained in the SI. The fraction of  ${}^{3}D_{3}$  is temperature dependent, and given by the 104 expression  $0.570 \cdot exp(-265.8/T)$ . Simultaneous temperature soundings with the IAP 105 RMR lidar [Gerding et al., 2016] revealed temperatures varying with time between 180 106 and 220 K in the peak region of the Ni layer at 85 km (not shown). Climatological data 107

for higher altitudes has been published by Gerding et al. [2008], showing that similar tem-108 peratures can be expected for the whole range of the Ni layer and the whole Jan-March 109 period. For these temperatures the fraction of  $Ni(^{3}D_{3})$  varies from 13.1 to 17.0%. Not 110 having direct temperature measurements available for the whole altitude range, we as-111 sume a constant fraction of 15.1%. Additionally, for density calculations the relaxation 112 of the  ${}^{3}F$  transition via a 380.7 nm emission is acknowledged with 7% probability (see 113 Figure S1 in the SI). All other relevant relaxations are within the transmission range of 114 the IF, in particular at  $339.3 \,\mathrm{nm}$  and  $347.3 \,\mathrm{nm}$  for the excitation of Ni(<sup>3</sup>F), and at  $339.1 \,\mathrm{nm}$ 115 for the excitation of  $Ni(^{3}D)$  [cf. Kramida et al., 2018]. 116

- Overall we expect a similar resonance signal for the Ni(<sup>3</sup>D) transition compared to the Ni(<sup>3</sup>F) transition used by CLM2015 if the laser power is the same at both wavelengths. Photon counts were collected in 200 m bins and further integrated to 1 km in order to improve the SNR. Statistical uncertainties of the Ni density profiles are calculated based on Poisson statistics for these 1 km bins.
- Similar to CLM2015 we did not have an absolute wavelength reading available for 122 most of the soundings. Only for the soundings in March 2018 were we able to measure 123 the wavelength of the pulsed light in the lab (see below). Wavelength adjustment using 124 only the laser keypad interface has an unknown bias of 30-50 pm. Therefore, for the first 125 successful soundings in January 2018, we rapidly scanned the dye laser over a wide range 126 of  $\sim 100 \,\mathrm{pm}$  and continuously checked the backscatter signal as well as recording the data 127 for later analysis. After a coarse adjustment the final wavelength was found in a detailed 128 scan across  $\sim 1 \,\mathrm{pm}$  based on the normalized resonance backscatter after integration of 129 4000 laser pulses ( $\sim 2.5$  min) and 0.2 pm wavelength steps per profile. For the next soun-130 ding nights typically only the fine tuning needed to be repeated, because we kept the la-131 ser electronics running continuously and the laser temperature stabilized. During the soun-132 dings, the wavelength was checked about once per hour to avoid wavelength drifts due 133 to thermal adjustment of the laser resonator. Similarly, the bandwidth of the pulsed la-134 ser was checked about once per hour by inspection of the transmission of an external mo-135 nitoring etalon. 136
- In March 2018 we measured the true laser wavelength with the WS6-200 wavelength meter. This wavemeter has an absolute accuracy of 0.2 pm in our wavelength range, which is in the range of the bandwidth of the pulsed laser. The optimal wavelength was cross-

-5-

checked by the atmospheric return. We used the WS6-200 both for the observation of the Ni( $^{3}$ D) transition at 341 nm as well as for the Ni( $^{3}$ F) transition at 337 nm.

### <sup>142</sup> 3 Observations at 341 nm and 337 nm

Nickel soundings at Kühlungsborn were made successfully during six nights in Ja nuary to March 2018. Further off-resonance soundings were made beforehand for initial
 tests. In this section we present examples of the raw data obtained at 341 nm and 337 nm
 as well as the calculated density profiles for all data.

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### 3.1 Raw data with and without background 8/9 January 2018

First observations of the mesospheric Ni layer above Kühlungsborn were made du-148 ring the night 7/8 January 2018. The laser was operated at the resonance wavelength 149  $\lambda_{air} = 341.4764$  nm for 40 min (72,000 laser pulses). In the following night, we observed 150 the Ni layer for  $\sim 2.5$  h at the same transition. Figure 1 shows the integrated raw data 151 profile with and without the background count rate. Above 50 km the molecular back-152 scatter (Rayleigh signal) is visible, decreasing with altitude due to decreasing air den-153 sity. Above  $\sim 78 \,\mathrm{km}$  the signal clearly increases due to the additional resonance backs-154 catter. The nickel layer can be observed up to  $\sim 100$  km. At higher altitudes the constant 155 background count rate is due to detector noise and sky background ( $\sim 1400 \text{ counts/km}$ ). 156 After background subtraction the Ni layer is even more clearly identified. Above 78 km 157 the profile shows initially a superposition of the comparatively intense Rayleigh signal 158 and the resonance signal. At the altitude of the maximum of the Ni layer,  $\sim 200 \text{ photons/km}$ 159 are counted from the Rayleigh signal and 2400 photons/km are due to resonance back-160 scatter. We extrapolate the range-corrected Rayleigh signal above 76 km with a norma-161 lized nightly NRLMSISE-00 density profile [Picone et al., 2002], and subtract this data 162 to get a pure resonance count rate (not shown). 163

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#### 3.2 Ni density profiles observed at 341 nm wavelength

The integrated, range-corrected and Rayleigh-subtracted backscatter profiles are used to calculate a mean Ni density profile. The statistical uncertainty is taken as the square root of the original count rate (with 1 km resolution) assuming Poisson statistics. In Figure 2 the Ni density profile on the evening 8 Jan 2018 is presented by the blue line,

Fig.2

Fig.1



Figure 1. Integrated raw data profile for the evening 8 January 2018 before and after back ground subtraction. The altitude resolution is set to 1 km.

with the dotted line showing the statistical uncertainty. The Ni layer extends from 78 km 171 to more than 100 km altitude. The maximum is observed at 84 km with a peak density 172 of  $\sim 280 \,\mathrm{cm}^{-3}$ , and a vertical column abundance of  $3.3 \cdot 10^8 \,\mathrm{cm}^{-2}$ . Fig. 2 also shows the 173 other four Ni density profiles measured using the  $Ni(^{3}D)$  resonance at 341 nm. These pro-174 files were obtained from observational periods ranging from 0.65 to 3.5 h. The peak den-175 sities vary between  $\sim 280 \,\mathrm{cm}^{-3}$  (8 Jan and 20 Mar 2018) and  $\sim 450 \,\mathrm{cm}^{-3}$  (7 Jan 2018). 176 The lower edges of the Ni layer vary by as much as 5 km between the nights. The layer 177 shape often differs from an ideal Gaussian and the peak height varies by  $\sim 3 \,\mathrm{km}$ . 178

The long sounding and high signal level on 8/9 January 2018 reveals the tempo-183 ral evolution of the Ni layer for the first time. Figure 3 shows that during the 3.5 h of 184 lidar observation the layer was highly variable. The peak density varies by up to 50%185 and the peak altitude partly changes by 2 km within only  $\sim 15 \text{ min}$ . The lower edge of 186 the layer ascends slowly by 2 km from 78 to 80 km. On the topside of the layer the va-187 riability is larger and occurs on shorter scales. Overall, even though this is only a sin-188 gle observation, the variability of the Ni layer seems to be larger compared to the Fe layer, 189 which is surprising given that both are transition metals and thus might to a first ap-190 proximation be expected to behave similarly. 191

Fig.3



Figure 2. Nickel density profiles for all measurements using the Ni( ${}^{3}$ D) transition at 341 nm. Uncertainties (dotted lines or error bars at the layer maximum) calculated for the original resolution of 1 km. Profiles are smoothed by 7-point Hann windows for visualization only. Numbers in brackets denote vertical column abundances in units of  $10^{8}$  cm<sup>-2</sup>.



Figure 3. Temporal evolution of the nickel layer during the 3.5 h long sounding on 8 Jan 2018.
Data are plotted with a 5-profile sliding average (20,000 pulses) and vertical smoothing using a
7-bin Hann window.

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### 3.3 Observations at 337 and 341 nm

After successful nickel soundings using the  $Ni(^{3}D)$  transition at 341 nm we tuned the laser to the  $Ni(^{3}F)$  transition at 337 nm in order to make soundings at the same wavelength as CLM2015. Because of the lower laser pulse energy the Rayleigh signal is weaker here. As mentioned above, the backscatter cross section of the  $Ni(^{3}F)$  transition is smaller, but the fraction of atoms in the ground state is larger. Overall, a similar signal to that at 341 nm would be expected. Making use of the WS6-200 wavemeter we were able to adjust the laser to the Ni(<sup>3</sup>F) transition and receive a resonance signal from the
mesopause region. The first soundings were performed for 2.5 h in the night 18/19 March
2018. Nickel densities were found to be lower than for Jan/Feb observations (not shown).

Two nights later on 20/21 March 2018 we made comparative soundings at both wa-205 velengths sequentially. First, we set the laser to the  $Ni({}^{3}F)$  resonance at 337 nm and col-206 lected data between 19:04 and 21:22 UTC (236,000 pulses). After that we tuned the la-207 ser back to 341 nm and recorded data for another hour (21:41–22:43 UTC). Figure 4 (top) 208 shows the integrated raw data profiles for both periods after background subtraction and 209 range correction. For 337 nm the uncorrected profile with background is also displayed. 210 Similar to the raw data shown by CLM2015, the 337 nm resonance backscatter is quite 211 weak and hardly visible on top of the background. (Note the stronger smoothing of the 212 CML2015 data.) In contrast to this, the 341 nm resonance signal was similar to the soun-213 ding on 8 Jan 2018, taking the shorter integration time into account (see Fig. 1). The 214 Rayleigh signals at 337 and 341 nm reflect the differences in laser power (1:3) and inte-215 gration time (2:1), but still the resonance signal at 337 nm is smaller then expected from 216 341 nm. 217

However, before calculating the Ni density for the sounding at 337 nm, some techni-223 cal limitations need to be acknowledged. In contrast to the sounding at 341 nm we as-224 sume here a 33% fraction of Amplified Spontaneous Emission (ASE) and a de-tuning of 225 the laser wavelength of 0.1 pm (cf. Section 4). The resulting density profiles are shown 226 in Figure 4 (bottom). For most of the altitudes they agree within their uncertainties. The 227 general structure of the profiles is similar with a maximum around 85 km and a "shoul-228 der" between 90 and  $95 \,\mathrm{km}$ . Nevertheless, the first profile shows about 30% lower den-229 sities than the second profile. These differences are within the temporal variability of the 230 Ni layer displayed in Fig. 3. 231

#### 232

### 4 Discussion and Conclusion

<sup>233</sup> We have detected atomic Ni as a layer in the middle atmosphere with peak den-<sup>234</sup> sities between  $\sim 280$  and  $450 \text{ cm}^{-3}$  and column abundances between  $3.1 \cdot 10^8 \text{ cm}^{-2}$  and <sup>235</sup>  $4.9 \cdot 10^8 \text{ cm}^{-2}$ . There are numerous publications on observations of the mesospheric Na, <sup>236</sup> K, Mg, Ca and Fe layers, differences of the relative metal abundance to their Chrond-<sup>237</sup> ritic ratios, and the related chemistry (see, e.g., the review by *Plane et al.* [2015]). In



Figure 4. top: Raw data profile of Ni soundings on 20 Mar 2018 with background (dashed) and after background subtraction and additional range correction (solid). Dotted red/blue line: Normalized NRLMSISE-00 density profile used for Rayleigh subtraction. Blue: 337 nm, red: 341 nm, black: background level of 337 nm data. bottom: Ni density profiles calculated from the raw data. Uncertainties and smoothing as in Fig. 2.

contrast, before the present study there is only one previous observation of the nickel layer published by CLM2015, showing peak densities of  $\sim 16,000 \,\mathrm{cm}^{-3}$  and column densities of  $2.7 \cdot 10^{10} \,\mathrm{cm}^{-2}$ . These numbers are a factor of 50–70 larger than the densities derived here. In the following we discuss our observations and potential reasons for the disagreement with CLM2015.

For the first attempts to detect the Ni layer using the Ni(<sup>3</sup>F) transition at 337 nm, the dye laser was operated without the intracavity etalon that we used later to limit the laser linewidth. This configuration was in fact similar to the setup used by CLM2015. Unfortunately we failed to detect a resonance signal while scanning the laser wavelength.

-10-

We noted a large variability of the Rayleigh signal level during the first trials, presuma-247 bly due to changing atmospheric humidity that affects the aerosol properties, i.e. visi-248 bility in the UV. The signal level decreased in some nights by a factor of 100, while the 249 simultaneous soundings of the RMR lidar at 532 nm [Gerding et al., 2016] essentially sho-250 wed no change. During the nights of normal signal level we detected the Rayleigh sig-251 nal well into the altitude of the metal layer (cf. Fig. 4), i.e. much higher than observed 252 by CLM2015. Therefore, the lidar soundings in the present study should have a much 253 higher SNR. 254

It turned out that the laser performed much better at 341 nm, i.e. at the resonance transition of the low-lying Ni(<sup>3</sup>D) metastable state. We were able to assemble the intracavity etalon, narrowing the laser linewidth. Based on the laser adjustment made, it was possible later to tune the laser to 337 nm even with the intracavity etalon. Finally, the wavemeter assured a much better wavelength tuning of laser, and also weak resonance backscatter was sufficient for Ni density measurements.

We now consider several reasons why our soundings might have yielded Ni densi-261 ties that are so much smaller than CLM2015. First, the laser might not have been tu-262 ned to the center of the resonance line. However, we carefully checked the wavelength 263 by manual changes in steps of  $0.2 \,\mathrm{pm}$ . We clearly identified the largest backscatter counts 264 from the metal layer and then set the wavelength appropriately. These checks were re-265 peated regularly, so we estimate the potential systematic error to be only 0.2 pm, resulting 266 in a potential underestimation of the true density by at most 25%. Second, the laser out-267 put may have had a large fraction of broadband emission that contributed to the Ray-268 leigh signal but not to the resonance signal. However, we regularly checked the band-269 width of the laser and the fraction of broadband emission throughout the sounding by 270 means of an external etalon. In later soundings, the wavemeter provided some additi-271 onal evidence that most of the laser emission was narrowband. Third, the overlap be-272 tween laser beam and telescope field of view might have decreased with altitude. Howe-273 ver, the overlap was regularly checked at 30 km altitude where the SNR is large. For hig-274 her altitudes we checked the overlap based on the nightly integrated data by compari-275 son with the RMR lidar signal obtained simultaneously. No significant difference was found. 276 We do not have any indications for further systematic errors due to laser performance 277 and therefore assess an additional underestimation of the Ni density being smaller than 278

279 280 the potential error due to incorrect tuning of the laser wavelength. Finally, we consider that the Ni densities published here are reliable within the stated uncertainties.

As mentioned above, the laser performance at 337 is worse compared to 341 nm 281 which is closer to the fluorescence maximum of the PTP dye at 343 nm [Brackmann, 1994]. 282 Therefore, operation of the laser at 337 nm is more sensitive to misalignment etc. Indeed, 283 during the soundings at 337 nm we noticed a poorer contrast of the ring system produ-284 ced by the external monitoring etalon. Also the wavemeter indicated a larger fraction 285 of broadband emission, although this is difficult to quantify because of the design of the 286 instrument. Regarding a potential offset of the laser wavelength with respect to the re-287 sonance line it should be noted that the spectral resolution of the wavemeter WS6-200 288 is  $0.2 \,\mathrm{pm}$  (at 337 nm), and a potential offset of the laser of  $0.1 \,\mathrm{pm}$  is still smaller than 289 the expected accuracy of manual scanning of the laser based on atmospheric return. Ta-290 king both effects into account (33% broadband emission and 0.1 pm offset), we estimate 291 an effective differential backscatter cross section at  $337 \,\mathrm{nm}$  of  $1.42 \cdot 10^{-18} \,\mathrm{m^2/sr}$ . This is 292 already taken into account in the density calculations (see Fig. 4, bottom). 293

One unexpected result of the present study is that a much better lidar resonance 294 signal is achieved by monitoring the metastable  ${}^{3}D$  state at 341 nm. This is the first ex-295 ample of a metal resonance lidar where this is the case. It is therefore worth considering 296 whether the  ${}^{3}D$  and  ${}^{3}F$  states are likely to be in thermal equilibrium, which we have as-297 sumed when calculating their relative populations (see the SI). Like the other meteoric 298 metals [Plane et al., 2015], Ni should exist in a fast chemical steady state with NiO, con-299 trolled by the reactions of Ni with  $O_3$  and NiO with O. The latter reaction could pro-300 duce Ni in a non-Boltzmann population initially. However, taking a recently measured 301 rate coefficient  $k(\text{Ni} + \text{O}_3 \rightarrow \text{NiO} + \text{O}_2) = (6.5 \pm 0.7) \cdot 10^{-10} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1} [\text{T. P. Man-$ 302 gan, University of Leeds, pers. comm.] and a typical  $O_3$  concentration at 85 km of 5. 303  $10^8 \,\mathrm{cm}^{-3}$ , the e-folding time for Ni conversion to NiO will be  $\sim 3.1 \,\mathrm{s}$ . During this time 304 the Ni atom will experience on the order of  $10^5$  collisions with air molecules and, given 305 that the separation of the  ${}^{3}D$  and  ${}^{3}F$  states is only 204.8 cm<sup>-1</sup>, it is very likely that these 306 states will be fully equilibrated. 307

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Finally, we consider the Ni vertical column abundance of  $4.0 \cdot 10^8 \text{ cm}^{-2}$  in the context of the well-studied mesospheric Fe layer. A typical Fe column abundance in January-March at mid-latitudes is  $1.5 \cdot 10^{10} \text{ cm}^{-2}$  [Kane and Gardner, 1993], which implies a Fe/Ni

-12-

ratio around 38. This ratio is about a factor of 2 larger than the Fe/Ni ratio of 18 in CI-Chondrites [Asplund et al., 2009]. Since Ni mainly resides in meteorites as Ni-Fe-S grains which melt at a lower temperature than the Fe-containing silicate phase [Levasseur-Regourd et al., 2018], it is unlikely that Ni will ablate less efficiently than Fe. Therefore the 2fold depletion of Ni indicates that Ni is more efficiently sequestered as Ni<sup>+</sup> or neutral reservoir species, compared with Fe. Studies of the relevant Ni kinetics are currently underway at Leeds, in an attempt to understand this.

The Fe/Ni ratio of 1.2 published by CLM2015, i.e. a factor of 22 smaller than the 318 CI ratio, is even more difficult to account for as those authors recognized. There is no 319 obvious explanation for the factor of 50-70 discrepancy (in absolute density, or factor 320 of  $\sim 40$  in Fe/Ni) between our observations and those of CLM2015. We cannot exclude 321 latitudinal differences, but these would have to be surprisingly large: typically, metal abun-322 dances increase by no more than a factor of 2 between mid- and high-latitudes [Feng et al., 323 2013; Langowski et al., 2015]. Future soundings, using the comparatively easy-to-reach 324 transition at 341 nm suggested here, may help to resolve this discrepancy. Preferably, 325 these should be conducted simultaneously with co-located Fe soundings. 326

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### 334 References

- Alpers, M., T. Blix, S. Kirkwood, D. Krankowsky, F.-J. Lübken, S. Lutz, and U. von
- Zahn (1993), First simultaneous measurements of neutral and ionized iron densi-
- ties in the upper mesosphere, J. Geophys. Res., 98, 275–283.
- Asplund, M., N. Grevesse, A. J. Sauval, and P. Scott (2009), The chemical compo-
- sition of the Sun, Annual Review of Astronomy and Astrophysics, 47(1), 481–522,
- doi:10.1146/annurev.astro.46.060407.145222.

- <sup>341</sup> Bills, R. E., and C. S. Gardner (1990), Lidar observations of mesospheric Fe and
- sporadic Fe layers at Urbana, Illinois, *Geophys. Res. Lett.*, 17, 143–146, doi:
   10.1029/GL017i002p00143.
- Bowman, M. R., A. J. Gibson, and M. C. W. Sandford (1969), Atmospheric sodium measured by a tuned laser radar, *Nature*, 221, 456–457, doi:10.1038/221456a0.
- Brackmann, U. (1994), Lambdachrome laser dyes, Lambda Physik GmbH, Goettingen, Germany.
- Carrillo-Sánchez, J. D., D. Nesvorný, P. Pokorný, D. Janches, and J. M. C. Plane
  (2016), Sources of cosmic dust in the Earth's atmosphere, *Geophys. Res. Lett.*,
- $_{350}$  43(23), 11,979-11,986, doi:10.1002/2016GL071697.
- <sup>351</sup> Chu, X., and G. C. Papen (2005), *Resonance fluorescence lidar for measurements of* <sup>352</sup> *the middle and upper atmosphere*, CRC Press, Boca Raton, USA.
- <sup>353</sup> Chu, X., Z. Yu, C. S. Gardner, C. Chen, and W. Fong (2011), Lidar observations
- of neutral Fe layers and fast gravity waves in the thermosphere (110-155 km) at McMurdo (77.8°S, 166.7°E), Antarctica, *Geophys. Res. Lett.*, 38, L23807, doi:
- 10.1029/2011GL050016.
- <sup>357</sup> Collins, R. L., J. Li, and C. M. Martus (2015), First lidar observation of the mesospheric nickel layer, *Geophys. Res. Lett.*, 42, 665–671, doi:10.1002/2014GL062716.
- <sup>359</sup> Correira, J., A. C. Aikin, J. M. Grebowsky, W. D. Pesnell, and J. Burrows (2008),
- Seasonal variations of magnesium atoms in the mesosphere-thermosphere, *Geophys. Res. Lett.*, 35, L06103, doi:10.1029/2007GL033047.
- Eska, V., J. Höffner, and U. von Zahn (1998), Upper atmosphere potassium layer
  and its seasonal variations at 54° N, J. Geophys. Res., 103 (A12), 29,207–29,214,
  doi:10.1029/98JA02481.
- Evans, W. F. J., R. L. Gattinger, A. L. Broadfoot, and E. J. Llewellyn (2011), The
  observation of chemiluminescent NiO\* emissions in the laboratory and in the
  night airglow, Atmos. Chem. Phys., 11(18), 9595–9603, doi:10.5194/acp-11-95952011.
- <sup>369</sup> Feng, W., D. R. Marsh, M. P. Chipperfield, D. Janches, J. Höffner, F. Yi, and
- J. M. C. Plane (2013), A global atmospheric model of meteoric iron, J. Geophys.
   *Res.*, 118(16), 9456–9474, doi:10.1002/jgrd.50708.
- Fricke, K., and U. von Zahn (1985), Mesopause temperatures derived from probing
- the hyperfine structure of the D2 resonance line of sodium by lidar, J. Atmos.

374	Solar-Terr. Phys., $47(5)$ , $499 - 512$ , doi:10.1016/0021-9169(85)90116-3.
375	Gerding, M., M. Alpers, J. Höffner, and U. von Zahn (1999), Simultaneous K
376	and Ca lidar observations during a meteor shower on March 6-7, 1997, at
377	Kühlungsborn, Germany, J. Geophys. Res., 104, 24,689–24,698.
378	Gerding, M., M. Alpers, U. von Zahn, R. J. Rollason, and J. M. C. Plane (2000),
379	Atmospheric Ca and Ca $^+$ layers: Midlatitude observations and modeling, J. Geop-
380	hys. Res., $105(A12)$ , 27,131–27,146, doi:10.1029/2000JA900088.
381	Gerding, M., J. Höffner, J. Lautenbach, M. Rauthe, and FJ. Lübken (2008), Sea-
382	sonal variation of nocturnal temperatures between 1 and 105 km altitude at $54^{\circ}\mathrm{N}$
383	observed by lidar, Atmos. Chem. Phys., 8, 7465–7482.
384	Gerding, M., M. Kopp, J. Höffner, K. Baumgarten, and FJ. Lübken (2016), Me-
385	so spheric temperature soundings with the new, daylight-capable IAP RMR lidar,
386	Atmos. Meas. Tech., $9(8)$ , 3707–3715, doi:10.5194/amt-9-3707-2016.
387	Granier, C., J. P. Jegou, and G. Megie (1985), Resonant lidar detection of Ca
388	and $Ca^+$ in the upper atmosphere, <i>Geophys. Res. Lett.</i> , 12, 655–658, doi:
389	10.1029/GL012i010p00655.
390	Gumbel, J., Z. Y. Fan, T. Waldemarsson, J. Stegman, G. Witt, E. J. Llewel-
391	lyn, CY. She, and J. M. C. Plane (2007), Retrieval of global mesospheric so-
392	dium densities from the Odin satellite, Geophys. Res. Lett., $34(4)$ , L04813, doi:
393	10.1029/2006GL028687.
394	Höffner, J., and J. S. Friedman (2005), The mesospheric metal layer topside: Exam-
395	ples of simultaneous metal observations, J. Atmos. Solar-Terr. Phys., 67(13), 1226
396	-1237, doi:10.1016/j.jastp.2005.06.010.
397	Jegou, JP., ML. Chanin, G. Mégie, and J. E. Blamont (1980), Lidar mea-
398	surements of atmospheric lithium, Geophys. Res. Lett., 7(11), 995–998, doi:
399	10.1029/GL007i011p00995.
400	Kane, T. J., and C. S. Gardner (1993), Structure and seasonal variability of the
401	nighttime mesospheric Fe layer at midlatitudes, J. Geophys. Res., 98, 16, doi:
402	10.1029/93JD01225.
403	Kramida, A., Yu. Ralchenko, J. Reader, and and NIST ASD Team
404	(2018), NIST Atomic Spectra Database (ver. 5.5.6), [Online]. Available:
405	https://physics.nist.gov/asd [2018, September 20]. National Institute of
406	Standards and Technology, Gaithersburg, MD.

-15-

407	Langowski, M. P., C. von Savigny, J. P. Burrows, W. Feng, J. M. C. Plane, D. R.
408	Marsh, D. Janches, M. Sinnhuber, A. C. Aikin, and P. Liebing (2015), Global
409	investigation of the Mg atom and ion layers using SCIAMACHY/Envisat obser-
410	vations between 70 and 150 km altitude and WACCM-Mg model results, $Atmos.$
411	Chem. Phys., $15(1)$ , 273–295, doi:10.5194/acp-15-273-2015.
412	Levasseur-Regourd, AC., J. Agarwal, H. Cottin, C. Engrand, G. Flynn, M. Fulle,
413	T. Gombosi, Y. Langevin, J. Lasue, T. Mannel, S. Merouane, O. Poch, N. Tho-
414	mas, and A. Westphal (2018), Cometary dust, Space Science Reviews, 214(3), 64,
415	doi:10.1007/s11214-018-0496-3.
416	Picone, J. M., A. E. Hedin, D. P. Drob, and A. C. Aikin (2002), NRLMSISE-00 em-
417	pirical model of the atmosphere: Statistical comparison and scientific issues, $J$ .
418	Geophys. Res., $107(A12)$ , 1468, doi:10.1029/2002JA009430.
419	Plane, J. M. C. (1991), The chemistry of meteoric metals in the Earth's upper
420	atmosphere, International Reviews in Physical Chemistry, $10(1)$ , 55–106, doi:
421	10.1080/01442359109353254.
422	Plane, J. M. C., W. Feng, and E. C. M. Dawkins (2015), The mesosphere and
423	metals: Chemistry and changes, Chemical Reviews, $115(10)$ , $4497-4541$ , doi:
424	10.1021/cr500501m.
425	Qian, J., and C. S. Gardner (1995), Simultaneous lidar measurements of mesosp-
426	heric Ca, Na, and temperature profiles at Urbana, Illinois, J. Geophys. Res., 100,
427	7453–7461, doi:10.1029/94JD02748.
428	Raizada, S., C. A. Tepley, D. Janches, J. S. Friedman, Q. Zhou, and J. D. Mathews
429	(2004), Lidar observations of Ca and K metallic layers from Arecibo and compari-
430	son with micrometeor sporadic activity, J. Atmos. Solar-Terr. Phys., 66, 595–606,
431	doi:10.1016/j.jastp.2004.01.030.
432	Saran, D. V., T. G. Slanger, W. Feng, and J. M. C. Plane (2011), FeO emission in
433	the mesosphere: Detectability, diurnal behavior, and modeling, J. Geophys. Res.,
434	116, D12303, doi:10.1029/2011JD015662.
435	Slipher, V. M. (1929), Emissions in the spectrum of the light of the night sky, <i>Publ.</i>
436	Astron. Soc. Pac., 41, 262265.
437	Unterguggenberger, S., S. Noll, W. Feng, J. M. C. Plane, W. Kausch, S. Kimeswen-
438	ger, A. Jones, and S. Moehler (2017), Measuring FeO variation using astrono-
439	mical spectroscopic observations, Atmos. Chem. Phys., 17(6), 4177–4187, doi:

-16-

# 440 10.5194/acp-17-4177-2017.

- 441 Yi, F., C. Yu, S. Zhang, X. Yue, Y. He, C. Huang, Y. Zhang, and K. Huang (2009),
- $_{442}$  Seasonal variations of the nocturnal mesospheric Na and Fe layers at 30°N, J.
- 443 Geophys. Res., 114 (D1), D01301, doi:10.1029/2008JD010344.