

This is a repository copy of Self-consistent single mode investigations of the quasi-geostrophic convection-driven dynamo model.

White Rose Research Online URL for this paper: http://eprints.whiterose.ac.uk/133297/

Version: Accepted Version

Article:

Plumley, M, Calkins, MA, Julien, K et al. (1 more author) (2018) Self-consistent single mode investigations of the quasi-geostrophic convection-driven dynamo model. Journal of Plasma Physics, 84 (4). ARTN 735840406. ISSN 0022-3778

https://doi.org/10.1017/S0022377818000831

© Cambridge University Press 2018. This article has been published in a revised form in Journal of Plasma Physics https://doi.org/10.1017/S0022377818000831. This version is free to view and download for private research and study only. Not for re-distribution, re-sale or use in derivative works.

Reuse

Items deposited in White Rose Research Online are protected by copyright, with all rights reserved unless indicated otherwise. They may be downloaded and/or printed for private study, or other acts as permitted by national copyright laws. The publisher or other rights holders may allow further reproduction and re-use of the full text version. This is indicated by the licence information on the White Rose Research Online record for the item.

Takedown

If you consider content in White Rose Research Online to be in breach of UK law, please notify us by emailing eprints@whiterose.ac.uk including the URL of the record and the reason for the withdrawal request.



Self-consistent single mode investigations of the quasigeostrophic convection-driven dynamo model

Journal:	Journal of Plasma Physics				
Manuscript ID	PLA-2017-0218.R2				
Special Collection:	50 Years of Mean Field Electrodynamics				
Date Submitted by the Author:	29-Jun-2018				
Complete List of Authors:	Plumley, Meredith; University of Colorado Boulder, Applied Mathematics Calkins, Michael; University of Colorado Boulder, Physics Julien, Keith; University of Colorado Boulder, Applied Mathematics Tobias, Steve; University of Leeds, Department of Applied Mathematics				
Keywords:	Plasma Flows, Plasma Dynamics				

Perez

SCHOLARONE[™] Manuscripts

Under consideration for publication in J. Plasma Phys.

Self-consistent single mode investigations of the quasi-geostrophic convection-driven dynamo model

Meredith Plumley¹[†], Michael A. Calkins ², Keith Julien ¹, and Steven M. Tobias ³

¹Department of Applied Mathematics, University of Colorado, Boulder, CO 80309, USA ²Department of Physics, University of Colorado, Boulder, CO 80309, USA ³Department of Applied Mathematics, University of Leeds, Leeds LS2 9JT, UK

(Received xx; revised xx; accepted xx)

The quasi-geostrophic dynamo model (QGDM) is a multiscale, fully-nonlinear Cartesian dynamo model that is valid in the asymptotic limit of low Rossby number. In the additional limit of small magnetic Prandtl number investigated here, the QGDM is a self-consistent, asymptotically exact form of an α^2 large-scale dynamo. This article explores methods for simulating the multiscale QGDM and investigates how convection is altered by the magnetic field in the planetary regime of small Rossby number and small magnetic Prandtl number. At present, this combination is beyond the reach of direct numerical simulations. We use a simplified class of solutions whose horizontal structure is restricted to a periodic hexagonal lattice characterized by a single horizontal wavenumber (single mode). In contrast with previous kinematic investigations of the QGDM, the Lorentz force is included to study saturated, self-consistent dynamos. Two methodologies are used to assess handling the multiple time scales of the QGDM: a stiff, common-in-time approach where all time scales are converted to a single time variable and a heterogeneous multiscale modeling approach employing fast time averaging on the Reynolds, magnetic and buoyancy eddy fluxes that feed back onto the slow scales. The strategies produce consistent results and each illustrates self-similar dynamics as the time averaging window is increased. The properties of the convection are significantly altered by the dynamo-generated magnetic field. All solutions show a decrease in the overall heat transfer efficiency as compared to non-magnetic convection, suggesting that a change in length scale or flow planform plays a critical role in the enhanced heat transfer efficiency observed in previous dynamo studies. All dynamo solutions show a trend of increasing ohmic dissipation relative to viscous dissipation as the buoyancy forcing is increased.

1. Introduction

Planets and stars generate and sustain their magnetic fields via the so-called hydromagnetic dynamo process (Elsasser 1956), whereby the kinetic energy of fluid motion is converted into electromagnetic energy. Most natural dynamos are thought to be powered by buoyancy-driven convection resulting from thermal and/or compositional heterogeneities. The dynamics of natural dynamos are controlled and characterized by several non-dimensional parameters. Assuming the simplest case of an incompressible (i.e. Boussinesq) fluid layer of depth H driven solely by the thermal heterogeneities, the

strength of the buoyancy force is quantified by the Rayleigh number Ra,

$$Ra = \frac{g\gamma \Delta \Theta H^3}{\nu \kappa},\tag{1.1}$$

where γ is the thermal expansion coefficient, g is the magnitude of the gravitational acceleration (assumed to be constant), and $\Delta\Theta$ represents the temperature difference across the fluid layer. The viscosity and thermal diffusivity of the fluid are denoted by ν and κ , respectively, and their ratio defines the thermal Prandtl number

$$Pr = \frac{\nu}{\kappa}.$$
(1.2)

All convectively-driven planetary and stellar dynamos are estimated, or assumed, to be in the high Rayleigh number convective turbulence regime. For instance, values of $Ra > 10^{20}$ characterize both the Earth's liquid outer core and the solar convection zone (Schubert & Soderlund 2011; Ossendrijver 2003).

The system rotation vector $\boldsymbol{\Omega}$ and the helical fluid motions it induces are thought to play an important role in the generation of large-scale magnetic fields (Parker 1955). The dynamical influence of rotation on fluid motion is controlled by the Rossby and Ekman numbers, defined by

$$Ro = \frac{U}{2\Omega H}, \qquad E = \frac{\nu}{2\Omega H^2}, \tag{1.3}$$

respectively measuring the strength of inertial forces relative to the Coriolis force, and viscous forces relative to the Coriolis force. Here U is a characteristic flow speed and $\Omega = |\Omega|$. Systems in which $(E, Ro) \ll 1$ are said to be rapidly rotating. For both the Earth's liquid iron outer core and the Sun's outer convective layer, the Ekman number is small with $E < 10^{-14}$. Morever, the Rossby number based on observationally inferred speed and length scales for the core is $Ro \approx 10^{-6}$ (Finlay & Amit 2011). These values indicate that rotation strongly constrains fluid motions in the core. Estimating Ro in the solar convection zone continues to be an important goal of stellar physics research, with recent helioseismic studies inferring values of $Ro \approx 0.1$ near the top of the convection zone (Hanosage *et al.* 2012; Greer *et al.* 2015).

The fluid properties of natural dynamos are known to have a significant effect on the convective motions (Jones 2011). The thermal Prandtl number Pr is less than unity in both planets and stars. Liquid metals such as in the Earth's core are typically characterized by $Pr \approx 10^{-2}$ (Pozzo *et al.* 2013) and stellar plasmas have extremely small values of $Pr \approx 10^{-7}$ (Ossendrijver 2003). Laboratory experiments show that low Pr convection is strongly inertial, thus allowing for the presence of turbulent motions at significantly lower Rayleigh numbers than required to achieve turbulence in $Pr \approx 1$ fluids (King & Aurnou 2013; Aurnou *et al.* 2015). The magnetic Prandtl number, defined as

$$Pm = \nu/\eta, \tag{1.4}$$

where η is the magnetic diffusivity, ranges from $Pm \approx 10^{-3}$ in the outer regions of the Sun, down to $Pm \approx 10^{-6}$ in the interiors of stars and planets.

Modeling natural dynamos is difficult due to the broad range of scales that characterize these systems and the concomitant numerical stiffness of the governing equations. Moreover, these challenges are compounded by the fact that rotation often leads to flow anisotropy (Chandrasekhar 1961). Despite these hindrances, direct numerical simulation (DNS) investigations continue to be the most widely employed method for dynamo studies (e.g. Tobias *et al.* 2011; Jones 2011). However, these studies are limited to relatively modest parameter values owing to computational restrictions; the recent study Cambridge University Press

Single mode investigations of the QG dynamo model

of Schaeffer *et al.* (2017) has accessed turbulent dynamos in a spherical shell down to $E = 5 \times 10^{-8}$. Reduction strategies applied to the governing partial differential equations attempt to overcome modeling difficulties through the development of simplified equation sets that capture the fundamental physics while filtering dynamically unimportant phenomena. Multi-scale asymptotic methods are a particularly robust and powerful technique that yield equation sets with substantially reduced numerical stiffness; this approach has been utilized with significant success in the case of rapidly rotating hydrodynamics, where a leading order balance exists between the Coriolis and pressure gradient forces, otherwise known as geostrophy (Charney 1948, 1971). Perturbations about this balance yield so-called quasi-geostrophic (QG) dynamics, and simulations of QG dynamics continue to provide substantial insight into both the large-scale hydrostatic motions of the Earth's atmosphere and ocean (Pedlosky 1992), as well as non-hydrostatic convectively-driven flows that are relevant for planetary atmospheres and interiors, and the convecting regions of rapidly rotating stars (Julien *et al.* 2006; Aurnou *et al.* 2015).

Reduction strategies for dynamos were first carried out by Childress & Soward (1972), where a weakly nonlinear analysis first showed that Boussinesq QG convection in the plane-layer Rayleigh-Bénard geometry readily gives rise to dynamo action near the onset of fluid motion. This model was extended to anelastic atmospheres by Mizerski & Tobias (2013). The Childress-Soward model was recently generalized to the case of strongly forced convection and multiple length scales perpendicular to the rotation axis by Calkins et al. (2015); for brevity, we refer to this new model as the QG dynamo model (QGDM). To summarize the model, four distinct forms of the QGDM can be derived based on the magnitudes of the small-scale Reynolds number $Re_{\ell} = U\ell/\nu$ and the small-scale magnetic Reynolds number $Rm_{\ell} = U\ell/\eta$, where $\ell \ll H$ is the horizontal length scale of convection. The size of Re_{ℓ} dictates the importance of inertia relative to viscosity on the small convective scales, and Rm_{ℓ} determines the relative importance of advection of the small-scale magnetic field to ohmic diffusion (see also Calkins 2018; Calkins et al. 2017). Note that all cases are strongly nonlinear, as characterized by the large scale Reynolds number $Re = (H/\ell)Re_{\ell} \gg 1$. Though the precise form of the QGDM varies depending upon which case is considered, all of the forms are characterized as large-scale dynamo models since a magnetic field is generated on a scale that is asymptotically larger than the typical horizontal scale ℓ of convection. To date, only kinematic investigations of the QGDM have been carried out and have focused solely on the $Re_{\ell} = O(1), Rm_{\ell} \ll 1$ case (Calkins *et al.* 2016a, b); this particular form of the QGDM is relevant for understanding planetary interiors, such as the Earth's outer core, and has the mathematical advantage that the onset of dynamo action can be determined through an eigenvalue problem due to the simplified form of the small-scale induction equation.

The kinematic investigations of the QGDM have utilized both simplified, single mode solutions (Calkins *et al.* 2016*a*) and fully nonlinear, multimode solutions (Calkins *et al.* 2016*b*). Both of these investigations found that rapidly rotating dynamo action persists well beyond the onset of convection, though the efficiency of the dynamo (as measured, for instance, by the critical magnetic Reynolds number required for dynamo action) does decrease somewhat as the buoyancy forcing is increased and the flow becomes more turbulent. Nevertheless, the kinematic QG flows remain strongly helical and therefore lead to large-scale dynamo action for all parameters investigated. Both investigations found that the structure of the large-scale kinematic magnetic field is insensitive to changes in the flow regime and buoyancy forcing; it therefore remains to be determined if similar behavior is observed for self-consistent dynamos in which the Lorentz force is included.

In the present work, we extend the aforementioned kinematic studies (Calkins *et al.* Cambridge University Press

2016*a*,*b*) by investigating single mode solutions to the hydrodynamic QG model with the inclusion of the nonlinear Lorentz force. Single mode solutions exploit the form of the reduced equations to enforce a specific horizontal periodic structure (the planform) characterized by a single horizontal spatial wavenumber. As a result, the QGDM can be reduced from three spatial dimensions to only one vertical dimension. This single mode restriction produces unstable solutions, much akin to exact coherent structures obtained in shear flows (e.g. Waleffe 2001). Single mode solutions have proven useful in multiple contexts, including rotating, hydrodynamic convection (Bassom & Zhang 1994; Julien & Knobloch 1999; Sprague *et al.* 2006; Julien *et al.* 2016) and magnetoconvection (Julien *et al.* 1999; Matthews 1999). Further, given that single mode solutions are unstable solutions to the full problem, they can be utilized to reduce computation time for the full model by serving as initial conditions that are free of imbalances (e.g. Sprague *et al.* 2006). In this study, we focus on those planforms that exist on a hexagonal lattice, i.e., horizontal spatial structures constructed from the superposition of three roll planforms differing in orientation by 120° .

The QGDM is characterized by three distinct time scales: a fast convective timescale t; a slow intermediate mean magnetic field timescale $\tau = \epsilon^{3/2} t$, and a slow mean temperature timescale $T = \epsilon^2 t$, where the small parameter is defined as

$$\epsilon \equiv E^{1/3}.\tag{1.5}$$

This small parameter ϵ is the small-scale Rossby number. Though the separation of time scales provides physical insight into the problem of rotating convection-driven dynamos, it increases the complexity of solving the model numerically. One of the purposes of the present work is to evaluate various strategies for time-stepping the multi-scale problem with the simplified form of the equations under the single mode ansatz; the results will therefore provide a guide for multimode simulations of the QGDM. Section 2 contains a brief overview of the QGDM, defines the single mode solutions used, and covers the numerical methods that we use for the multiple time scales. Numerical results are presented in section 3, with conclusions given in section 4.

2. Model description

We provide only a brief summary of the main features of the QGDM since the detailed derivation is given in Calkins *et al.* (2015). The basis of the multiscale model derivation is the Boussinesq form of the equations for convection-driven magnetohydrodynamics. A multiple scales expansion is employed in the axial and temporal coordinates, i.e.

$$\partial_z \to \partial_z + \epsilon \partial_Z, \qquad \partial_t \to \partial_t + \epsilon^{3/2} \partial_\tau + \epsilon^2 \partial_T.$$
 (2.1)

The large-scale vertical coordinate is defined by $Z = \epsilon z$, where z is the small vertical coordinate over which the Proudman-Taylor theorem (Proudman 1916; Taylor 1923; Greenspan 1968) is satisfied. Each dependent variable is separated into mean and fluctuating components; for instance, a scalar f is decomposed according to

$$f = \overline{f} + f', \tag{2.2}$$

where the overbar is defined as an average over the fast spatial (\mathbf{x}) and temporal (t) scales according to

$$\overline{f}(Z,\tau,T) = \lim_{\Gamma,\mathcal{V}\to\infty} \frac{1}{\Gamma\mathcal{V}} \int_{\Gamma,\mathcal{V}} f(\mathbf{x},Z,t,\tau,T) \, d\mathbf{x} \, dt, \quad \overline{f'} \equiv 0.$$
(2.3)
Cambridge University Press

Here Γ is a time interval between the fast convective and mean magnetic time scales and \mathcal{V} is the volume in small spatial coordinates. Upon separating the governing equations into mean and fluctuating components, all dependent variables are represented as a perturbation expansion in powers of $\epsilon^{1/2}$. The aspect ratio of the convection is defined by $H/\ell = \epsilon^{-1}$ where, again, ℓ is the horizontal scale of the convection.

The horizontal velocity $\mathbf{u}_{\perp} = (u, v)$ in the QGDM is geostrophically balanced, such that $\hat{\mathbf{z}} \times \mathbf{u}_{\perp} = -\nabla_{\perp} \psi$, where ψ is the geostrophic stream function. The horizontal velocity components can then be written as $\mathbf{u}_{\perp} = (-\partial_y \psi, \partial_x \psi) = \nabla^{\perp} \psi$. Equations are developed for the time evolution of the vertical vorticity $\nabla_{\perp}^2 \psi$, vertical velocity w, temperature $\theta = \overline{\theta} + \epsilon \theta'$ and magnetic field $\mathbf{B} = \overline{\mathbf{B}} + \epsilon^{1/2} \mathbf{b}'$. The mean and fluctuating magnetic field vectors are defined by $\overline{\mathbf{B}} = (\overline{B}_x, \overline{B}_y)$ and $\mathbf{b}' = (b'_x, b'_y, b'_z)$. For brevity, hereafter, we include the prime superscript only on variables that possess non-trivial mean values (e.g. θ and \mathbf{B}). The final set of reduced equations, non-dimensionalized using the horizontal viscous diffusion timescale ℓ^2/ν , is given by

$$D_t^{\perp} \nabla_{\perp}^2 \psi - \partial_Z w = \overline{\mathbf{B}} \cdot \nabla_{\perp} j_z' + \nabla_{\perp}^4 \psi, \qquad (2.4)$$

$$D_t^{\perp} w + \partial_Z \psi = \frac{Ra}{Pr} \theta' + \overline{\mathbf{B}} \cdot \nabla_{\perp} b'_z + \nabla_{\perp}^2 w, \qquad (2.5)$$

$$D_t^{\perp}\theta' + w\partial_Z \overline{\theta} = \frac{1}{Pr} \nabla_{\perp}^2 \theta', \qquad (2.6)$$

$$\partial_T \overline{\overline{\theta}}^{\tau} + \partial_Z \overline{(w\theta')}^{\tau} = \frac{1}{Pr} \partial_{ZZ} \overline{\overline{\theta}}^{\tau}, \qquad (2.7)$$

$$\partial_{\tau}\overline{\vartheta} + \partial_{Z}(\overline{w\theta'} - \overline{\overline{w\theta'}}^{\tau}) = 0, \qquad (2.8)$$

$$\partial_{\tau} \overline{\mathbf{B}} = \widehat{\mathbf{z}} \times \partial_{Z} \overline{\boldsymbol{\mathcal{E}}} + \frac{1}{\widetilde{Pm}} \partial_{ZZ} \overline{\mathbf{B}}, \qquad (2.9)$$

$$0 = \overline{\mathbf{B}} \cdot \nabla_{\perp} \left(\nabla_{\perp}^2 \psi \right) + \frac{1}{\widetilde{Pm}} \nabla_{\perp}^2 j'_z, \qquad (2.10)$$

$$0 = \overline{\mathbf{B}} \cdot \nabla_{\perp} w + \frac{1}{\widetilde{Pm}} \nabla_{\perp}^2 b'_z, \qquad (2.11)$$

The various operators are defined by $D_t^{\perp} = (\partial_t + \boldsymbol{u} \cdot \nabla_{\perp})$ and $\nabla_{\perp} = (\partial_x, \partial_y, 0)$, and the additional overbar with superscript τ appearing in the mean heat equation denotes further averaging over the τ timescale. The mean temperature is now decomposed according to $\overline{\theta} = \overline{\overline{\theta}}^{\tau} + \epsilon^{1/2} \overline{\vartheta}$. We note that in the QGDM advection and diffusion act only over the small horizontal scales due to the spatial anisotropy (Calkins *et al.* 2015). The vertical component of the fluctuating current density is $j'_z = \partial_x b'_y - \partial_y b'_x$ and the mean electromotive force (emf) is defined by

$$\overline{\mathcal{E}} = \overline{\mathbf{u} \times \mathbf{b}'}.\tag{2.12}$$

The asymptotically reduced Rayleigh number is defined by

$$\widetilde{Ra} = E^{4/3} Ra. \tag{2.13}$$

In the present work we restrict the analysis to the $Re_{\ell} = O(1)$, $Rm_{\ell} \ll 1$ form of the QGDM. Since $Rm_{\ell} = PmRe_{\ell}$, this implies that the model can also be characterized by a small magnetic Prandtl number. In this limit the fluctuating induction equation loses material advection and the asymptotically reduced magnetic Prandtl number is defined as

$$Pm = \epsilon^{1/2} \widetilde{Pm}, \tag{2.14}$$
 Cambridge University Press

such that $\widetilde{Pm} = O(1)$.

For the small Pm limit considered here, the emf can be written as $\overline{\mathcal{E}}_i = \alpha_{ij}\overline{B}_j$, where the pseudo-tensor α_{ij} is

$$\alpha_{ij} = \widetilde{Pm} \left(\frac{\overline{w \nabla_{\perp}^{-2} \partial_{xx} \psi - \partial_x \psi \nabla_{\perp}^{-2} \partial_x w}}{\overline{w \nabla_{\perp}^{-2} \partial_{xy} \psi - \partial_y \psi \nabla_{\perp}^{-2} \partial_x w}} - \frac{\overline{w \nabla_{\perp}^{-2} \partial_{xy} \psi - \partial_x \psi \nabla_{\perp}^{-2} \partial_y w}}{\overline{w \nabla_{\perp}^{-2} \partial_{yy} \psi - \partial_y \psi \nabla_{\perp}^{-2} \partial_y w}} \right).$$
(2.15)

The above relation follows upon eliminating the fluctuating electromagnetic fields from the mean induction equations with the use of equations (2.10), (2.11) and (2.12). Since both components of the mean magnetic field are induced by an α -effect, the two small Pm forms of the QGDM (see Calkins 2018) can be considered asymptotically-exact α^2 dynamos (e.g. Moffatt 1978). For sufficiently long simulation times, it is expected, and observed by Calkins *et al.* (2016*b*), that α_{ij} will become symmetric; it was further observed that $\alpha_{12} \approx \alpha_{21} \approx 0$ as $t \to \infty$.

We employ constant temperature boundary conditions such that the mean temperature at the bottom and top boundaries is given by

$$\overline{\theta} = 1 \text{ at } Z = 0, \ \overline{\theta} = 0 \text{ at } Z = 1.$$
 (2.16)

Impenetrable boundary conditions are imposed on the velocity field,

$$w = 0$$
 at $Z = 0, 1.$ (2.17)

The electromagnetic boundary conditions can be either electrically insulating or perfectly conducting. For the present work we impose the former such that,

$$\overline{\mathbf{B}} = 0 \text{ at } Z = 0, 1. \tag{2.18}$$

Finally, we note that no explicit boundary conditions need to be imposed on either the fluctuating temperature or the fluctuating magnetic field since their respective equations require that they satisfy the boundary conditions

$$\theta' = 0, \, \boldsymbol{b}' = 0 \text{ at } Z = 0, 1.$$
 (2.19)

2.1. Single mode theory

The single mode method represents the dependent variables as products of a horizontal planform and vertical structure functions. Single mode expansions for the hydrodynamic variables are given by the general form

$$(\psi, w, \theta') = \sum_{j=1}^{n} \left[\Psi_j(t, Z, \tau, T), W_j(t, Z, \tau, T), \Theta_j(t, Z, \tau, T) \right] h_j(x, y) + c.c..$$
(2.20)

On a periodic hexagonal lattice the sum consists of three roll solutions (n = 3) whose orientations differ by 120°, with the direction of each roll characterized by unit vectors $\hat{\mathbf{m}}_1 = (1,0), \hat{\mathbf{m}}_2 = (1/2, \sqrt{3}/2)$ and $\hat{\mathbf{m}}_3 = \hat{\mathbf{m}}_1 - \hat{\mathbf{m}}_2 = (1/2, -\sqrt{3}/2)$. Here $h_j(x, y) = \exp[ik(\hat{\mathbf{m}}_j \cdot \mathbf{x})]$, satisfying the horizontal planform equation $\nabla_{\perp}^2 h_j = -k^2 h_j$ with wavenumber magnitude k. While multiple planform options exist and have been investigated previously in asymptotic dynamo investigations (e.g., square planforms in Soward 1974), only the hexagonal planform is included here to limit the number of choices. In particular, the hexagonal lattice was chosen for its property of admitting the effects of advective nonlinearity. Different planforms lead to small differences in the functional form of the single mode model, but they result in qualitatively similar Cambridge University Press

Single mode investigations of the QG dynamo model

dynamics. We consider two distinct single mode approaches: the single amplitude (SA) model and the mixed amplitude (MA) model.

2.1.1. Single Amplitude (SA) Model

The single amplitude (SA) model uses an equal amplitude approximation on each component of the superposition (2.20). This gives, for example, $|\Psi_1| = |\Psi_2| = |\Psi_3|$ or $|\Psi_1| = 0, |\Psi_2| = |\Psi_3|$ such that (2.20) may be collapsed to the real-valued solution

$$(\psi, w, \theta') = [\Psi(t, Z, \tau, T), W(t, Z, \tau, T), \Theta(t, Z, \tau, T)] h(x, y).$$
(2.21)

Here h(x, y) satisfies the planform equation, with $h = \sum_{j=1}^{3} h_j$, so that the hexagonal lattice takes the form

$$h(x,y) = \sqrt{2/3} \left[\cos(kx) + \cos\left(\frac{k}{2}x + \frac{\sqrt{3}}{2}ky\right) + \cos\left(\frac{k}{2}x - \frac{\sqrt{3}}{2}ky\right) \right], \quad (2.22)$$

where the normalization factor of $\sqrt{2/3}$ ensures that $\overline{h^2} = 1$.

Since mean quantities are independent of x and y, the mean temperature and magnetic field are given by

$$\overline{\theta} = \overline{\theta}(Z, \tau, T), \quad \overline{\mathbf{B}} = \left[\overline{B}_x(Z, \tau, T), \overline{B}_y(Z, \tau, T), 0\right].$$
(2.23)

Using the low-Pm induction equations (2.10) and (2.11) and defining $\mathbf{b}_{\perp}' = \nabla^{\perp} \Phi$, leads to

$$\nabla_{\perp}^2 \Phi = -\widetilde{Pm} \Psi(\overline{B}_x h_x + \overline{B}_y h_y), \qquad (2.24)$$

$$\nabla_{\perp}^2 b'_z = -\widetilde{Pm} W(\overline{B}_x h_x + \overline{B}_y h_y), \qquad (2.25)$$

such that $j'_z = \nabla^2_\perp \Phi$.

The derivation of the single amplitude model with (2.21) and (2.23) is included in appendix A.1 with details regarding the treatment of the mean equations (2.7)–(2.9) in appendix A; the complete set of equations expressed over the convective time scale for the vertical structure functions is given by

$$\partial_t \Psi + \frac{1}{k^2} \partial_Z W = -\frac{\overline{Pm}}{2} \left[(\overline{B}_x^{\mathcal{V}})^2 + (\overline{B}_y^{\mathcal{V}})^2 \right] \Psi - k^2 \Psi, \qquad (2.26)$$

$$\partial_t W + \partial_Z \Psi = \frac{Ra}{Pr} \Theta - \frac{Pm}{2} \left[(\overline{B}_x^{\mathcal{V}})^2 + (\overline{B}_y^{\mathcal{V}})^2 \right] W - k^2 W, \tag{2.27}$$

$$\partial_t \Theta + W \partial_Z \overline{\theta}^{\mathcal{V}} = -\frac{k^2}{Pr} \Theta, \qquad (2.28)$$

$$\epsilon^{-2}\partial_t \overline{\theta}^{\mathcal{V}} + \partial_Z (W\Theta) = \frac{1}{Pr} \partial_{ZZ} \overline{\theta}^{\mathcal{V}}, \qquad (2.29)$$

$$\epsilon^{-3/2}\partial_t \overline{B}_x^{\mathcal{V}} = -\widetilde{Pm}\,\partial_Z \Big(\Psi W \,\overline{B}_y^{\mathcal{V}}\Big) + \frac{1}{\widetilde{Pm}}\partial_{ZZ} \overline{B}_x^{\mathcal{V}},\tag{2.30}$$

$$\epsilon^{-3/2} \partial_t \overline{B}_y^{\mathcal{V}} = \widetilde{Pm} \, \partial_Z \left(\Psi W \, \overline{B}_x^{\mathcal{V}} \right) + \frac{1}{\widetilde{Pm}} \partial_{ZZ} \overline{B}_y^{\mathcal{V}}. \tag{2.31}$$

Given this single mode formulation, the emf in the mean induction equations can be written as $\overline{\mathcal{E}}_{i}^{\mathcal{V}} = \alpha_{ij}\overline{B}_{j}^{\mathcal{V}}$, where

$$\alpha_{ij} = \widetilde{Pm} \begin{bmatrix} \Psi W & 0\\ 0 & \Psi W \end{bmatrix}$$
(2.32)

M. Plumley, M. A. Calkins, K. Julien and S. M. Tobias

is a diagonal and symmetric matrix.

As discussed earlier, the single mode ansatz for the hydrodynamic fields results in no advective nonlinear terms in (2.26) - (2.28). The only nonlinear hydrodynamic response is due to the buoyancy flux in (2.29). In the absence of feedback from the Lorentz force (when $\overline{\mathbf{B}} = 0$), these equations produce exact solutions to the hydrodynamic problem (Julien & Knobloch 1999). From (2.24) and (2.25), we note that while the power of any generated magnetic fluctuations remains in spectral wavenumber k, the horizontal periodic structure differs from the assumed pattern h(x, y) of the hydrodynamic quantities. Thus, only the component of the Lorentz force that projects onto the planform h(x, y) is captured in the present single mode formulation.

2.1.2. Mixed amplitude (MA) model

The mixed amplitude (MA) method allows for different amplitudes of the various components of the planform function. Soward (1974), for instance, used this approach for both square and hexagonal planforms. Without loss of generality, we find that the MA method can be reduced to the real-valued ansatz

$$(\psi, w, \theta') = \sum_{j=1}^{n} \left[\Psi_j(t, Z, \tau, T), W_j(t, Z, \tau, T), \Theta_j(t, Z, \tau, T) \right] h_j(x, y).$$
(2.33)

For the particular case of a hexagonal lattice we have n = 3 with the planform functions $h_j(x, y)$ given by

$$h_1 = \sqrt{\frac{2}{3}} \cos(kx), \quad h_2 = \sqrt{\frac{2}{3}} \cos\left(\frac{1}{2}kx + \frac{\sqrt{3}}{2}ky\right), \quad h_3 = \sqrt{\frac{2}{3}} \cos\left(\frac{1}{2}kx - \frac{\sqrt{3}}{2}ky\right).$$
(2.34)

Here the functions h_j are chosen as components of the original planform function h(x, y) in (2.22), and we note that $\overline{h_j^2} = 1/3$ for j = 1, 2 and 3. From this point the derivation of the single mode equations proceeds in a similar way to the SA model, except that each equation is multiplied by h_1 , h_2 and h_3 separately and then horizontally averaged. The orthogonality of each planform function results in $\overline{h_1h_2} = \overline{h_1h_3} = \overline{h_2h_3} = 0$, such that separate equations for each amplitude can be derived. Following this procedure results in a total of 9 equations governing the fluctuating quantities for the MA model:

$$\partial_t \Psi_1 + \frac{1}{k^2} \partial_Z W_1 = -\widetilde{Pm} (\overline{B}_x^{\mathcal{V}})^2 \Psi_1 - k^2 \Psi_1, \qquad (2.35)$$

$$\partial_t \Psi_2 + \frac{1}{k^2} \partial_Z W_2 = -\widetilde{Pm} \left(\frac{1}{4} (\overline{B}_x^{\mathcal{V}})^2 + \frac{\sqrt{3}}{2} \overline{B}_x^{\mathcal{V}} \overline{B}_y^{\mathcal{V}} + \frac{3}{4} (\overline{B}_y^{\mathcal{V}})^2 \right) \Psi_2 - k^2 \Psi_2, \qquad (2.36)$$

$$\partial_t \Psi_3 + \frac{1}{k^2} \partial_Z W_3 = -\widetilde{Pm} \left(\frac{1}{4} (\overline{B}_x^{\mathcal{V}})^2 - \frac{\sqrt{3}}{2} \overline{B}_x^{\mathcal{V}} \overline{B}_y^{\mathcal{V}} + \frac{3}{4} (\overline{B}_y^{\mathcal{V}})^2 \right) \Psi_3 - k^2 \Psi_3, \qquad (2.37)$$

$$\partial_t W_1 + \frac{1}{2\sqrt{2}} k^2 (\Psi_2 W_3 - \Psi_3 W_2) + \partial_Z \Psi_1 = \frac{\widetilde{Ra}}{Pr} \Theta_1 - \widetilde{Pm} \, (\overline{B}_x^{\mathcal{V}})^2 W_1 - k^2 W_1, \qquad (2.38)$$

$$\partial_t W_2 + \frac{1}{2\sqrt{2}} k^2 (\Psi_3 W_1 - \Psi_1 W_3) + \partial_Z \Psi_2 = \frac{\widetilde{Ra}}{\frac{\widetilde{Ra}}{Pr}} \Theta_2 - \widetilde{Pm} \left(\frac{1}{4} (\overline{B}_x^{\nu})^2 + \frac{\sqrt{3}}{2} \overline{B}_x^{\nu} \overline{B}_y^{\nu} + \frac{3}{4} (\overline{B}_y^{\nu})^2 \right) W_2 - k^2 W_2,$$
(2.39)

Single mode investigations of the QG dynamo model

$$\partial_t W_3 + \frac{1}{2\sqrt{2}}k^2(\Psi_1 W_2 - \Psi_2 W_1) + \partial_Z \Psi_3 =$$
(2.40)

$$\frac{\widetilde{Ra}}{Pr}\Theta_3 - \widetilde{Pm}\left(\frac{1}{4}(\overline{B}_x^{\nu})^2 - \frac{\sqrt{3}}{2}\,\overline{B}_x^{\nu}\overline{B}_y^{\nu} + \frac{3}{4}(\overline{B}_y^{\nu})^2\right)W_3 - k^2W_3,\tag{2.41}$$

$$\partial_t \Theta_1 + \frac{1}{2\sqrt{2}}k^2(\Psi_2 \Theta_3 - \Psi_3 \Theta_2) + W_1 \partial_Z \overline{\theta}^{\mathcal{V}} = -\frac{k^2}{Pr}\Theta_1, \qquad (2.42)$$

9

$$\partial_t \Theta_2 + \frac{1}{2\sqrt{2}} k^2 (\Psi_3 \Theta_1 - \Psi_1 \Theta_3) + W_2 \partial_Z \overline{\theta}^{\mathcal{V}} = -\frac{k^2}{Pr} \Theta_2 , \qquad (2.43)$$

$$\partial_t \Theta_3 + \frac{1}{2\sqrt{2}} k^2 (\Psi_1 \Theta_2 - \Psi_2 \Theta_1) + W_3 \partial_Z \overline{\theta}^{\mathcal{V}} = -\frac{k^2}{Pr} \Theta_3.$$
(2.44)

The mean temperature equation becomes

$$\epsilon^{-2}\partial_t\overline{\theta}^{\mathcal{V}} + \frac{1}{3}\partial_Z\left(W_1\Theta_1 + W_2\Theta_2 + W_3\Theta_3\right) = \frac{1}{Pr}\partial_{ZZ}\overline{\theta}^{\mathcal{V}}.$$
(2.45)

Similar expressions to (2.24) and (2.25) can be found for the MA case. The two components of the mean induction equation are

$$\epsilon^{-3/2} \partial_t \overline{B}_x^{\mathcal{V}} = -\partial_Z \overline{\mathcal{E}}_y^{\mathcal{V}} + \frac{1}{\widetilde{Pm}} \partial_{ZZ} \overline{B}_x^{\mathcal{V}}, \qquad (2.46)$$

$$\epsilon^{-3/2} \partial_t \overline{B}_y^{\mathcal{V}} = \partial_Z \overline{\mathcal{E}}_x^{\mathcal{V}} + \frac{1}{\widehat{Pm}} \partial_{ZZ} \overline{B}_y^{\mathcal{V}}, \qquad (2.47)$$

where the emf terms are given by

$$\overline{\mathcal{E}}_{y}^{\mathcal{V}} = \widetilde{Pm} \left[\frac{1}{2} \overline{B}_{y}^{\mathcal{V}} \left(W_{2} \Psi_{2} + W_{3} \Psi_{3} \right) + \frac{\sqrt{3}}{6} \overline{B}_{x}^{\mathcal{V}} \left(W_{2} \Psi_{2} - W_{3} \Psi_{3} \right) \right], \qquad (2.48)$$

$$\overline{\mathcal{E}}_{x}^{\mathcal{V}} = \widetilde{Pm} \left[\frac{1}{6} \overline{B}_{x}^{\mathcal{V}} \left(4W_{1}\Psi_{1} + W_{2}\Psi_{2} + W_{3}\Psi_{3} \right) + \frac{\sqrt{3}}{6} \overline{B}_{y}^{\mathcal{V}} \left(W_{2}\Psi_{2} - W_{3}\Psi_{3} \right) \right].$$
(2.49)

Setting $\overline{\boldsymbol{\mathcal{E}}}_{i}^{\mathcal{V}} = \alpha_{ij}\overline{B}_{j}^{\mathcal{V}}$ produces

$$\alpha_{ij} = \widetilde{Pm} \begin{bmatrix} \frac{1}{2} (W_2 \Psi_2 + W_3 \Psi_3) & \frac{\sqrt{3}}{6} (W_2 \Psi_2 - W_3 \Psi_3) \\ \frac{\sqrt{3}}{6} (W_2 \Psi_2 - W_3 \Psi_3) & \frac{1}{6} (4W_1 \Psi_1 + W_2 \Psi_2 + W_3 \Psi_3) \end{bmatrix}.$$
 (2.50)

Like the SA model, the MA model yields a symmetric α_{ij} . Moreover, statistically $\overline{W_1\Psi_1} \approx \overline{W_2\Psi_2} \approx \overline{W_3\Psi_3}$, such that α_{ij} becomes identical in structure to the SA model. While the model is not necessarily time-averaged to ensure that the off-diagonal terms are zero, the averaged profiles $(W_2\Psi_2, W_3\Psi_3)$ are similar such that those terms are small relative to the diagonal entries.

As shown by equations (2.35)-(2.41), for the MA case the Lorentz force now completely projects onto the horizontal lattice. However, the ansatz of non-equal amplitude results in non-zero horizontal advection; these nonlinear interactions generate higher order spatial harmonics that are not considered in the present work. We can therefore consider the current MA model as a truncated Galerkin expansion, as was done by Soward (1974) for the weakly nonlinear case. The MA model is consistent with the SA model in the previous section through the cancellation of extra nonlinear terms when $A_1 = A_2 = A_3 = A_{SA}$, where A_i denotes the three components of the vertical function in the MA case and A_{SA} is the amplitude of the SA model.

2.2. Key Definitions

Several quantities are used to analyze the results. A common and useful measure which characterizes the efficiency of heat transfer through the fluid layer is the Nusselt number Nu, which is defined as the ratio of convective heat flux to total heat flux and can be written as

$$Nu = Pr W\Theta - \partial_Z \overline{\theta}, \qquad (2.51)$$

where averaging in the vertical direction gives

$$Nu = 1 + Pr \langle W\Theta \rangle. \tag{2.52}$$

Here the vertical average is denoted by angled brackets, i.e.

$$\langle f \rangle = \int_0^1 f dZ. \tag{2.53}$$

The Nusselt number, as defined in (2.51) and (2.52), can be formally obtained from the mean temperature equation under the assumption of temporal stationarity and realized by averaging over all time scales (Julien *et al.* 2012*b*). In the following we focus solely on (2.52) and find it informative to report both instantaneous and time-averaged values of the Nusselt number, respectively denoted as Nu(t) and \overline{Nu} . Given the independence of Nu on (fast) spatial coordinates, fast time-averaging can be trivially expressed using the overbar notation denoting fast time and space averaging. A similar approach is used below for time averaging the quadratic expressions of single mode amplitudes.

In the absence of magnetic fields, the heat transfer \overline{Nu} is also related to the viscous dissipation rate ε_u^c of the system through the power integral (e.g. Chandrasekhar 1961)

$$\underbrace{k^2 \left\langle \overline{W^2} \right\rangle + k^4 \left\langle \overline{\Psi^2} \right\rangle}_{\varepsilon_u^c} = \frac{\widehat{Ra}}{Pr^2} \left(\overline{Nu}^c - 1 \right). \tag{2.54}$$

The above expression is obtained by multiplying the vorticity and vertical momentum equations by $-\psi$ and w, respectively, and averaging over all space and time to find $\varepsilon_u = \langle (\overline{\nabla_{\perp}^2 \Psi})^2 \rangle + \langle (\overline{\nabla_{\perp} W})^2 \rangle$. While this expression is derived for the SA model, the MA model follows a similar derivation with the three components of each variable, such that $\varepsilon_u = 1/3 \sum_{j=1}^3 \langle (\overline{\nabla_{\perp}^2 \Psi_j})^2 \rangle + \langle (\overline{\nabla_{\perp} W_j})^2 \rangle$.

The inclusion of the magnetic field and the associated Lorentz force leads to non-zero ohmic dissipation ε_B , defined as $\varepsilon_B = [\langle \overline{(j'_z)^2} \rangle + \langle \overline{(\nabla_\perp b'_z)^2} \rangle]/\overline{Pm}$, such that the power integral for dynamos becomes

$$\underbrace{k^2 \left\langle \overline{W^2} \right\rangle + k^4 \left\langle \overline{\Psi^2} \right\rangle}_{\varepsilon_u} + \underbrace{\frac{\widetilde{Pm}}{2} \left\langle \left(\overline{B}_x^2 + \overline{B}_y^2 \right) \overline{\left(k^2 \Psi^2 + W^2 \right)} \right\rangle}_{\varepsilon_B} = \frac{\widetilde{Ra}}{Pr^2} \left(\overline{Nu} - 1 \right).$$
(2.55)

Dynamo action therefore implies $\varepsilon_B > 0$. Comparison of the quantities in (2.54) and (2.55) can distinguish differences in the energy balances between non-magnetic convection and convection-driven dynamos. The thermal dissipation rate ε_{Θ} , obtained by multiplying the fluctuating heat equation by θ' , defines an additional power integral as

$$\varepsilon_{\Theta} = \langle \overline{(\partial_Z \overline{\theta})^2} \rangle + k^2 \langle \overline{\Theta^2} \rangle = \overline{Nu} , \qquad (2.56)$$

and one can relate these quantities, with only a dependence on the inputs \widetilde{Ra} and Pr, as

$$\varepsilon_u^c = \frac{\widetilde{Ra}}{Pr^2} (\varepsilon_\Theta^c - 1), \quad \varepsilon_u + \varepsilon_B = \frac{\widetilde{Ra}}{Pr^2} (\varepsilon_\Theta - 1).$$
 (2.57)

Single mode investigations of the QG dynamo model 11

The mean magnetic energy and the fluctuating kinetic energy are defined by, respectively,

$$E_M = \frac{1}{2} \left\langle \overline{\mathbf{B}}^{\mathcal{V}} \cdot \overline{\mathbf{B}}^{\mathcal{V}} \right\rangle, \qquad (2.58)$$

$$E_K = \frac{1}{2} \langle \overline{\mathbf{u} \cdot \mathbf{u}}^{\mathcal{V}} \rangle \,. \tag{2.59}$$

For E_M and E_K , we use the same convention as with Nu and denote the instantaneous result as $(E_M(t), E_K(t))$ and the time-averaged values as simply $(\overline{E}_M, \overline{E}_K)$. We also find it useful to characterize time-varying quantities with an appropriately defined root-meansquare (rms) value; for some scalar quantity f this is defined by

$$f_{rms} = \left[\overline{\left(f - \overline{f}^{\mathcal{A}} \right)^2} \right]^{1/2}, \qquad (2.60)$$

where $\overline{(\cdot)}^{\mathcal{A}}$ denotes an average over the horizontal area of the domain \mathcal{A} . To increase the convergence rate of statistics we exploit the expected vertical symmetries of the profiles about the midplane (Z = 0.5); a similar strategy was used in the kinematic investigation of Calkins *et al.* (2016*b*).

We find it useful to examine the kinetic helicity, which is known to be conducive to large-scale dynamo action (Moffatt 1978). The mean helicity $\overline{\mathcal{H}}$ and relative helicity $\overline{\mathcal{H}}_R$ are defined as

$$\overline{\mathcal{H}} = \overline{\mathbf{u} \cdot \boldsymbol{\zeta}}, \qquad \overline{\mathcal{H}}_R = \frac{\overline{\mathcal{H}}}{\overline{\mathbf{u}^2}^{1/2} \overline{\boldsymbol{\zeta}^2}^{1/2}}, \qquad (2.61)$$

where $\boldsymbol{\zeta} = \langle \partial_y w, -\partial_x w, (\partial_x v - \partial_y u) \rangle$ is the vorticity vector. For the QGDM, the helicity is

$$\overline{\mathcal{H}} = 2\overline{w\zeta}, \qquad (2.62)$$

where ζ is the vertical vorticity. For the SA model these expressions reduce to

$$\overline{\mathcal{H}} = -2k^2 \overline{\Psi} \overline{W}, \qquad \overline{\mathcal{H}}_R = \frac{-2k\Psi W}{k^2 \overline{\Psi}^2 + \overline{W}^2}, \qquad (2.63)$$

and, recall that since Ψ and W are functions only of Z, the overbar reduces to an average in time when used in conjunction with these quantities. For the SA model, $\overline{\mathcal{H}}$ is directly proportional to the emf and α_{ij} entries.

2.3. Numerical methods

All equations are temporally discretized with the second order, semi-implicit Runge– Kutta Crank–Nicolson two-stage scheme (e.g. Peyret 2013). The vertical spatial dimension in the QGDM is discretized using Chebyshev polynomials, and vertical derivatives are handled efficiently with the quasi-inverse technique of Julien & Watson (2009). The boundary conditions are imposed with the tau method (Peyret 2013). The resulting system of equations is solved with sparse LU solvers at each timestep. Details of the simulations are provided in Table 1. Resolutions were verified *a posteriori* with the requirement that at least 6 physical gridpoints are present within the thermal boundary layer. The code was rigorously benchmarked by setting the temporal and spatial dependence of assumed (known) solutions, generating forcing functions in each equation from those predefined solutions, and testing the timestepper against the known solutions.

M. Plumley, M. A. Calkins, K. Julien and S. M. Tobias

2.4. Timestepping considerations

Timestepping the mean temperature equation introduces an issue with the length of simulations. For instance, the time required for the mean temperature to have diffused through the vertical layer is $O(E^{-2/3}t)$. The simulation length therefore becomes computationally prohibitive as E is lowered even for this simplified 1D problem; for the Earth's outer core, where $E = O(10^{-15})$, this requires a simulation to be carried out to a fast convective time $t = O(10^{10})$ for one large-scale magnetic diffusion time. However, in the purely hydrodynamic regime the mean temperature in the single mode system achieves a steady state where $\partial_T \overline{\theta}^{T} = 0$ (Julien & Knobloch 1998). Adopting this as an *a priori* constraint resolves the issue associated with the simulation length by significantly reducing the time required to reach the steady state.

In the present work, the existence of magnetic fields and the associated mean magnetic evolution timescale $\tau = \epsilon^{3/2}t$ represent an added complexity in pursuing the aforementioned *a priori* constraint. However, this constraint may be achieved as follows (see Appendix A for further details). The composite mean temperature equation from which the multiscale decomposition (2.7) and (2.8) are deduced is given by the $O(\epsilon^2)$ ordered term

$$\partial_T \overline{\overline{\theta}}^{\tau} + \partial_\tau \overline{\theta} + \partial_Z (\overline{w\theta'}) = \frac{1}{Pr} \partial_{ZZ} \overline{\overline{\theta}}^{\tau} .$$
(2.64)

On defining the composite temperature $\overline{\theta}(Z, \tau, T) = \overline{\overline{\theta}}^{\tau}(Z, T) + \epsilon^{1/2}\overline{\vartheta}(Z, \tau, T)$, we obtain

$$\partial_T \overline{\theta} + \partial_Z (\overline{w\theta'}) = \frac{1}{Pr} \partial_{ZZ} \overline{\theta} + O(\varepsilon^{1/2}).$$
(2.65)

An instantaneous evolution equation for the mean temperature may be obtained by commuting the differential operators of (2.65) into the kernel of the averaging operator (2.3). The instantaneous temperature equation is obtained by constraining the kernel to be zero (details of this can be found in appendix A), leaving

$$\epsilon^{-2}\partial_t \overline{\theta}^{\mathcal{V}} + \partial_Z \left(\overline{w\theta'}^{\mathcal{V}} \right) = \frac{1}{Pr} \partial_{ZZ} \overline{\theta}^{\mathcal{V}}, \qquad (2.66)$$

where $\overline{f}^{\mathcal{V}}$ denotes averaging over small spatial scales. The single mode mean temperature equations (2.29) in the SA model and (2.45) in the MA model are deduced from this equation. Invoking the steady state approximation $\epsilon^{-2}\partial_t \overline{\theta}^{\mathcal{V}} = 0$ implies

$$\partial_Z \left(\overline{w\theta'}^{\mathcal{V}}\right) = \frac{1}{Pr} \partial_{ZZ} \overline{\theta}^{\mathcal{V}} \quad \text{and} \quad Nu(t,\tau,T) = Pr\left(\overline{w\theta'}^{\mathcal{V}}\right) - \partial_Z \overline{\theta}^{\mathcal{V}}.$$
 (2.67)

Time averaging these instantaneous relations gives the relations for intermediate and slow time scale variations over t, τ

$$\partial_Z \left(\overline{w\theta'} \right) = \frac{1}{Pr} \partial_{ZZ} \overline{\theta} \quad \text{and} \quad Nu(\tau, T) = Pr\left(\overline{w\theta'} \right) - \partial_Z \overline{\theta} ,$$
 (2.68)

$$\partial_Z \left(\overline{\overline{w\theta'}}^{\tau}\right) = \frac{1}{Pr} \partial_{ZZ} \overline{\overline{\theta}}^{\tau} \quad \text{and} \quad Nu(T) = Pr\left(\overline{\overline{w\theta'}}^{\tau}\right) - \partial_Z \overline{\overline{\theta}}^{\tau}.$$
 (2.69)

The *a priori* assumption $\partial_T \overline{\overline{\theta}}_0^{\tau} = 0$ is compatible with (2.69) and indicates that Nu(T) = Nu is a stationary value. Moreover, the above results also indicate that the variation of the heat transport about the stationary value is given by

$$Nu(\tau,T) - Nu = Pr\left(\overline{w\theta'} - \overline{\overline{w\theta'}}^{\tau}\right) + O(\varepsilon^{1/2}).$$
(2.70)

This variation in the convective heat flux can admit oscillations with a frequency on Cambridge University Press

Single mode investigations of the QG dynamo model

the mean inductive timescale. Thus, as E is lowered, oscillations that occur should show longer periods. Notably, (2.70) implies no *a priori* bounds for the amplitude of the variation.

Similar to the analysis and discussion of the mean temperature timescale, a large number of fast convective timesteps are required to reach a mean magnetic diffusion time $\tau = E^{1/2}t$ because of the scale separation between the fast convective time and the mean magnetic evolution time. For an Ekman number of $E = 10^{-6}$, a fast convective time of $t = O(10^3)$ is required to reach $\tau = O(1)$. With a timestep size of $\Delta t = 10^{-5}$, which is characteristic of many of our cases detailed in Table 1, this yields a total of 10^8 timesteps. However, to reduce computation time, the simulations are considered converged if statistically stationary results are obtained for the Nusselt number and magnetic energy upon time averaging.

2.5. Multiscale timestepping strategies

Obtaining numerical solutions to the QGDM requires specialized numerical techniques given the presence of multiple time scales in the governing equations. In particular, temporal discretization of the QGDM implies that averaging must also be performed in a discretized fashion, and the primary issue then becomes determining the appropriate averaging timescale, or window. Several methods are in current use for obtaining numerical solutions to temporally multiscale equations, though the choice of the method employed depends strongly on the detailed form of the equations and significant testing is required to determine which method best suits the particular problem (e.g. E *et al.* 2007; Haut & Wingate 2014). Below we discuss the two different techniques that are used in the present work.

2.5.1. The stiff method

One method for treating the different time scales numerically is to directly simulate the time derivatives in (2.4) - (2.11) with their form written in terms of ϵ and the convective time derivative. This requires commuting time averaging with differential operators as outlined in Appendix A. The single mode approximation then results in the SA and MA models given in sections (2.1.1) and (2.1.2). Recall that we apply a steady state approximation for the mean heat equation so no time derivative remains in the mean temperature equation. Since the derivation of the QGDM requires $\epsilon \ll 1$, the small ϵ parameter in the governing equations leads to a numerically stiff problem and we refer to the use of this common-in-time technique simply as the stiff method.

2.5.2. The heterogeneous multiscale method

A common method for the numerical treatment of multiscale equations is the socalled heterogeneous multiscale method (HMM) (E *et al.* 2007). This method, which incorporates time scale separation, has previously been applied to asymptotic equations for wind-driven oceanic flows (e.g., Malecha *et al.* 2014). In the present context, the HMM consists of timestepping the mean equations on specified large intervals of the fast timescale. We refer to this as the multiscale method. Specifically, the mean magnetic field is updated every S_B convective timesteps. Thus, the time discretization for the mean magnetic field is given by $\Delta \tau = S_B \Delta t$. The nonlinear terms in the mean magnetic field equations are averaged over all the fast convective timesteps in each $\Delta \tau$ time interval. Appropriate values for S_B can be related back to the Ekman number, $E = \epsilon^3$, by recalling the relation between the slow and fast time scales such that

$$S_B = O\left(\epsilon^{-3/2}\right) = O\left(E^{-1/2}\right). \tag{2.71}$$

M. Plumley, M. A. Calkins, K. Julien and S. M. Tobias

2.6. Stiff vs Multiscale

We find that the stiff and multiscale methods lead to identical results when relation (2.71) is satisfied, though minor differences can appear when converting S_B to an integer value. That both methods yield the same solutions is encouraging and allows for either method to be used, yet there are a few important points to note. The multiscale method requires fewer linear solves than the stiff method, since the mean magnetic field equations are only solved after a specified number of fast timesteps. In general, however, we find that the multiscale method requires (for numerical stability) slightly smaller convective timestep sizes than the stiff method. For instance, the SA model with $\tilde{Ra} = 40$, Pr = 1, and $S_B = 10$ requires $\Delta t \leq 8 \times 10^{-3}$ for the multiscale method, whereas the stiff method is numerically stable for $\Delta t \leq 1 \times 10^{-2}$. At the slightly larger Rayleigh number with $\tilde{Ra} = 60$, the multiscale method requires $\Delta t \leq 3.5 \times 10^{-3}$ for numerical stability while the stiff method can use $\Delta t \leq 6 \times 10^{-3}$. Given that the stiff method requires 2 extra matrix solves roughly every step (for S_B large), the resulting computational cost is comparable but slightly favors the stiff method for the $\tilde{Ra} = 60$ case.

The difference in permissible timestep sizes for the two methods indicates that the stiff method becomes more efficient for use at higher \widetilde{Ra} . This benefit could be related to the averages used for the nonlinear terms in the mean magnetic equations in the multiscale method, whereas in the stiff technique the nonlinear terms are updated at each timestep. We note that there are a multitude of different HMM techniques, including some that decouple Δt and $\Delta \tau$ and focus solely on the slow dynamics. In such cases, the fast dynamics are run until a stationary state producing stable averages is achieved (see Malecha *et al.* 2014). This approach enables larger numerical timesteps for the slow dynamics, such that the multiscale method becomes more efficient. Since we are interested in both continuous fast and slow dynamics and want to understand how they are coupled, we choose to keep our timestepping linked through these parameters (E or S_B), rather than a condition based on a minimal time to acquire a stationary state. In the remainder of this work, we report results from the multiscale method only, since both methods yield the same results.

3. Results

We have carried out a broad suite of simulations (runs) of the SA model, the MA model and the single-amplitude, non-magnetic, hydrodynamic model (setting $\mathbf{B} = 0$ in the SA model). Various details of the simulations are provided in table 1. For all cases presented we fix the wavenumber k to be that corresponding to the critical wavenumber k_c for the onset of convection. In the present work we restrict the simulations to Pr = 1and therefore use the wavenumber $k = k_c = 1.3048$ associated with steady convection (Chandrasekhar 1961). For $Pr \lesssim 0.68$ the onset of convection is oscillatory. In preliminary simulations with Pr = 0.1 we found that the timestepping and averaging methods employed in the present work were unsatisfactory in the sense that no convergence of statistics was observed as the time-averaging window of the mean magnetic field (S_B) was varied. We attribute this behavior to the periodic dynamics that occur for low Prandtl number fluids; capturing accurate statistics requires a more sophisticated averaging procedure since the position (in time) of the averaging window has a significant influence on the resulting average. In particular, significant errors in averaged quantities of periodic signals are introduced when averaging over non-integer multiples of the fundamental period of oscillation. Identifying the period of oscillation a priori is not possible for nonlinear dynamo solutions. For these reasons we leave the analysis of low **Cambridge University Press**

Model	\widetilde{Ra}	Pr	\widetilde{Pm}	\widetilde{Pm}_c	S_B	N_Z	Δt	\overline{Nu}	\overline{E}_K	\overline{E}_M	N_{BL}
SA	40	1	0.7	0.24	10^{2}	80	5×10^{-4}	11.3	65.6	3.14	8
SA	60	1	0.7	0.14	10^{2}	80	1×10^{-4}	17.7	109	5.96	7
\mathbf{SA}	80	1	0.7	0.10	10^{1}	100	1×10^{-4}	24.1	163	8.85	8
\mathbf{SA}	80	1	0.7	0.10	10^{2}	100	1×10^{-4}	24.8	162	8.29	8
\mathbf{SA}	80	1	0.7	0.10	10^{3}	100	1×10^{-4}	24.9	163	8.25	8
\mathbf{SA}	100	1	0.7	0.074	10^{2}	100	1×10^{-5}	31.8	222	10.3	7
\mathbf{SA}	200	1	0.7	0.074	10^{2}	150	1×10^{-5}	70.9	591	21.8	8
MΔ	40	1	0.7		10^{1}	80	5×10^{-5}	0.20	72 /	1 30	8
MΔ	40	1	0.7	_	10^{10}	80	5×10^{-5}	10.3	03.3	0.650	8
MA	40	1	0.7	_	10^{3}	80	5×10^{-5} 5×10^{-5}	10.0	84 3	0.000	8
MA	60	1	0.7	_	10^{2}	80	5×10^{-5}	13.3	148	1.84	7
MA	80	1	0.7	-	10^{2}	100	5×10^{-5}	14.3	194	2.45	8
MA	100	1	0.7		10^{2}	100	1×10^{-5}	15.5	241	3.33	7
MA	200	1	0.7		10^{2}	150	1×10^{-5}	21.4	411	10.5	8
\mathbf{SA}	20	1	1	0.68	10^{2}	100	1×10^{-4}	4.73	19.0	0.375	17
\mathbf{SA}	20	1	2	0.68	10^{2}	100	1×10^{-4}	2.97	5.97	1.03	22
\mathbf{SA}	20	1	5	0.68	10^{2}	100	1×10^{-4}	1.68	1.15	0.900	32
\mathbf{SA}	20	1	20	0.68	10^{2}	100	1×10^{-4}	1.06	0.0681	0.305	no BL
\mathbf{SA}	100	1	0.1	0.074	10^{2}	140	1×10^{-4}	79.5	2143	2.17	6
\mathbf{SA}	100	1	0.3	0.074	10^{2}	110	5×10^{-5}	46.0	561	11.9	6
\mathbf{SA}	100	1	0.5	0.074	10^{2}	100	5×10^{-5}	36.4	315	11.2	7
\mathbf{SA}	100	1	1	0.074	10^{2}	100	1×10^{-5}	29.5	166	9.85	7

Single mode investigations of the QG dynamo model

Table 1: Details for the simulations presented. N_Z is the number of Chebyshev points in the vertical direction and N_{BL} is the approximate number of Chebyshev points within the time averaged thermal Θ boundary layer. The time discretization for the mean magnetic field is given by $\Delta \tau = S_B \Delta t$. The timesteps result in converged solutions, although they are not optimized or adapted during the runs, meaning larger timesteps are likely possible. The critical magnetic Prandtl number \widetilde{Pm}_c indicates the \widetilde{Pm} above which a kinematic dynamo can exist (calculated for the SA kinematic model in Calkins *et al.* 2016*a*). Additional runs with varying \widetilde{Pm} were calculated, but the energetics are unchanged from those reported in the paper.

Prandtl number dynamos for future work. Simulations with higher Pr = 10 displayed qualitatively similar results to the Pr = 1 case, albeit with higher \widetilde{Ra} transitions to unsteady dynamics due to the increased importance of viscosity for higher Pr. Additional Pr = 1 runs with fixed $\widetilde{Ra} = 20,100$ and varied \widetilde{Pm} show similar energetics.

3.1. Determination of the emf averaging window, S_B

We find a posteriori that the influence of changing the averaging window $S_B \Delta t$ in the mean induction equation on the resulting dynamo solutions depends on several factors: the model used (SA or MA); the type of convection (steady or oscillatory); and the value of \widetilde{Ra} (indicating flow supercriticality). To simplify the presentation, only the fast convective time t is used in plots presenting time-dependent data.

The simplest solutions that we find are those that correspond to steady convection. Cambridge University Press



Figure 1: Plots illustrating the influence of the mean magnetic field averaging length S_B on both (a) the Nusselt number Nu and (b) $\langle \overline{B}_x^2(t) \rangle^{1/2}$. The interval length S_B denotes the number of fast convective timesteps over which the mean magnetic field equations are held constant and the nonlinear terms in the mean induction equations are averaged. All results were obtained with $\widetilde{Ra} = 20$, Pr = 1 and $\widetilde{Pm} = 1$. The time in (b) has been scaled by $S_B/10^4$ to highlight the temporal self-similarity of all the runs.



Figure 2: Comparison of (a) the Nusselt number Nu(t) and (b) the magnetic energy $E_M(t)$ versus the fast convective time t, for three different values of S_B using the single amplitude (SA) model. The simulation parameters are $\widetilde{Ra} = 80$, Pr = 1 and $\widetilde{Pm} = 0.7$. The oscillations that are observed in Nu decrease in amplitude and increase in period as S_B is increased.

Steady convection is found at sufficiently low combinations of \widetilde{Ra} and \widetilde{Pm} values and yields steady values of the magnetic energy E_M for all interval lengths S_B . In general, the values of $(\widetilde{Ra}, \widetilde{Pm})$ for which these steady solutions are achieved varies. For example, $\widetilde{Ra} = 20$ produces steady solutions for all $\widetilde{Pm} \leq 20$, whereas $\widetilde{Ra} = 100$ is only steady for runs with $\widetilde{Pm} \leq 0.4$. Although E_M is constant for steady convection solutions, changing S_B affects the periodicity of the individual components $\langle \overline{B}_x^2(t) \rangle^{1/2}, \langle \overline{B}_y^2(t) \rangle^{1/2}$. While we label these as steady results due to the fact that $\overline{B}_x^2 + \overline{B}_y^2$ reaches a constant value, the Cambridge University Press





Figure 3: Results illustrating the influence of S_B for the mixed amplitude (MA) model. The quantities shown are (a) Nusselt number Nu(t); (b) kinetic energy $E_K(t)$; (c) magnetic energy $E_M(t)$; and (d) magnetic energy $E_M(t)$ with rescaled time. The parameters are $\tilde{R}a = 40$, Pr = 1 and $\tilde{P}m = 0.7$ for $S_B = 10^1$, 10^2 and 10^3 . While all of the Nu and E_K results appear similar, there is a visible difference in the magnetic energy results between the $S_B = 10^1$ results and the larger S_B cases. Plot (d) highlights similarities in E_M for the two largest values of S_B , where the $S_B = 10^2$ run time has been scaled by 10 and begins after the initial transient.

dynamos are oscillatory for all cases. An example of these steady solutions is illustrated in figure 1, where the Nusselt number Nu settles to a constant value for all S_B . However, as E is decreased (and S_B is increased), the result requires a longer simulation time to reach equilibrium. The runs presented in figure 1b have been rescaled to demonstrate self-similarity of the solutions. Each run at a given S_B has been multiplied by the factor needed to match the timescale of the $S_B = 10^4$ run; for example, the $S_B = 10^2$ run has been scaled by 10^2 . For these steady cases, the choice of S_B is irrelevant since timeaveraged quantities such as the emf and Nusselt number do not change with time.

The results show more complex dependence on S_B when the underlying convection is unsteady. We find that the value of S_B becomes critical to capturing self-similar behavior, which we consider to be indicative of capturing physically correct behavior. Therefore, the results are deemed 'converged' if self-similar behavior is identified, where self-similar behavior implies reaching the same statistically steady results for \overline{Nu} and \overline{E}_M . Unsteady convective dynamics appear as \widetilde{Ra} is increased, where the precise value of \widetilde{Ra} at which unsteady motions occur depends on both Pr and \widetilde{Pm} . Above the transition to unsteady Cambridge University Press



M. Plumley, M. A. Calkins, K. Julien and S. M. Tobias

Figure 4: The instantaneous magnetic energy $E_M(t)$ and kinetic energy $E_K(t)$ for the SA model (a,b) and MA model (c,d) for various values of Ra. Oscillations appear as Ra is increased in the SA model, losing periodicity for Ra = 200, and all the MA model results show large fluctuations. The parameters are Pr = 1, Pm = 0.7 and $S_B = 10^2$. Different vertical and horizontal axes ranges are used to more clearly display the MA results.

motions, the hydrodynamic component of these flows becomes oscillatory with a single frequency for sufficiently small Ra. Additional frequencies in both the convection and mean magnetic field are then generated as Ra is increased further until chaotic convection is observed for sufficiently high reduced Rayleigh number. An example of a case with periodic oscillations is shown in Figure 2 for several values of S_B . We note that (2.70) implies variability in the heat transfer on the τ timescale; as such an ϵ -dependent increase is observed in the period of oscillations in the Nusselt number on an $O(\epsilon^{-3/2})$ scale with a constant of proportionality ~ 1/20 (for the $S_B = 10^3$ run). Additionally, the magnetic energy shows self-similar behavior in time for the longer $S_B = 10^2, 10^3$ runs, with consistent averages despite the fact that the amplitudes of the oscillations in E_M decrease with increasing S_B , settling to a constant value for the $S_B = 10^3$ run. For Pr = 10 we observe the same behavior in the dynamics on S_B . The Pr = 10 case also displays both the transition to unsteady convection for higher Ra, and the requirement of longer S_B intervals for self-similar behavior of the runs.

In comparison to the SA model, the results from the MA are more physically rich in both their time dependence and spatial (Z) structure (figure 3). Whereas the Nu(t) and $E_K(t)$ results shown in figure 3a and 3b, respectively, exhibit similar behavior for the **Cambridge University Press**



Figure 5: Vertical profiles for various rms quantities of the SA model with $\widetilde{Pm} = 0.7$, Pr = 1, $S_B = 10^2$ and various \widetilde{Ra} ; (a) vertical velocity; (b) fluctuating temperature; (c) vertical fluctuating current density; and (d) vertical fluctuating magnetic field. To better display the thermal boundary layers only the lower half of the Θ profile is shown in (b). Profiles in (c,d) are averaged in time for all \widetilde{Ra} and the W, Θ profiles for $\widetilde{Ra} = 80, 100, 200$ are averaged in time.

three values of S_B , figure 3c shows that significant differences in the magnetic energy E_M appear for different values of S_B . These results illustrate that it is insufficient to rely on the behavior of a single quantity when determining if the results converge, or exhibit selfsimilar behavior, with increasing S_B . Figure 3d shows the magnetic energy for only the two largest values of S_B over scaled time; here we observe similar temporal dependence and amplitudes of E_M for the two cases, suggesting that $S_B = 10^2$ is sufficient to capture the correct dynamics. We attribute these differences to a lack of sufficient temporal scale separation in the sense that the equivalent value of E is too large, as was also observed in the $\widetilde{Ra} = 80$ SA case shown in figure 2. The spatial structure of the solutions confirms this as the time averaged vertical profiles of the two longer cases agree, while the $S_B = 10$ result shows slight differences in the interior structure.

In all of the results that follow we set $S_B = 10^2$ (see table 1). This value of S_B was chosen because lower S_B runs take less time to reach a statistically stationary state, and Cambridge University Press



Figure 6: Vertical profiles for various rms quantities of the MA model with Pm = 0.7, Pr = 1, $S_B = 10^2$ and various \widetilde{Ra} ; (a) vertical velocity; (b) fluctuating temperature; (c) vertical fluctuating current density; and (d) vertical fluctuating magnetic field. To better display the thermal boundary layers only the lower half of the Θ profile is shown in (b). All of the profiles have been time averaged.

all of the tested cases show that the $S_B = 10^2$ results are similar to the higher S_B results as the time-averaged quantities agree with higher S_B metrics.

3.2. Increasing the convective supercriticality

We examine the effect of increasing the convective supercriticality over the range of $20 \leq \widetilde{Ra} \leq 200$ (see table 1) for a fixed value of \widetilde{Pm} . The magnetic energy E_M and kinetic energy E_K , shown in figure 4, illustrate the oscillations that are found in dynamo solutions as \widetilde{Ra} is increased for the SA model (upper row) and the MA model (lower row). Given that E_M and Nu are tied through the $\overline{B}_x^2 + \overline{B}_y^2$ term that contributes to the Lorentz force in the Ψ and W equations, oscillations in the convective dynamics naturally lead to oscillations in the magnetic field. Evident from the $\widetilde{Ra} = 80,100$ results in 4(a),(b) is the inverse relationship between E_M and E_K .

Profiles showing the vertical structure of the fluctuating quantities in the SA model are given in figure 5 for a range of Rayleigh numbers. As expected, both W and b'_z show Cambridge University Press

20



Figure 7: Comparison of (a) the vertical velocity profile W and (b) the horizontal structure of the full w = W(Z) h(x, y) field for the MA model at the midplane. Plot (a) shows the vertical profiles for the two classes of dynamo solutions; the separate amplitudes $(W_{1,2,3})$ of the MA model are shown in gray, and the symmetrized rms profiles for the MA model and the SA are shown by the solid red and solid black curves, respectively. The horizontal structure of w is shown in plot (b), where the hexagonal structure shows slight variations in the shape arising from the different amplitudes in each direction. The parameters are $\widetilde{Ra} = 100$, Pr = 1, $\widetilde{Pm} = 0.7$ and $S_B = 10^2$.



Figure 8: Profiles of the rms x-component of the mean magnetic field vector for (a) the SA model and (b) the MA model. The parameters are $\widetilde{Pm} = 0.7$, Pr = 1 and $S_B = 10^2$.

a growing amplitude with \widetilde{Ra} , and the fields Θ and j'_z show the development of boundary layers that decrease in width with increasing \widetilde{Ra} . In comparison to non-magnetic convection, the magnitude of all hydrodynamic quantities (W, Ψ, Θ) is diminished due to the conversion of kinetic energy into magnetic energy.

Typical profiles of MA dynamo solutions are shown in figure 6a,b. The solutions do not settle to an equal amplitude, e.g. $W_1 = W_2 = W_3$, even when using initial conditions with each component equal to the SA solution. Thus, we observe that the subspace associated with the SA model is unstable. However, due to rotational invariance in the horizontal plane, it is expected that statistically these amplitudes should agree when adequate Cambridge University Press



Figure 9: Relative helicity $\overline{\mathcal{H}}_R$ profiles for (a) non-magnetic ($\overline{\mathbf{B}} = 0$) convection and (b),(c) dynamo solutions. The parameters are $\widetilde{Pm} = 0.7$, Pr = 1, $S_B = 10^2$ and various \widetilde{Ra} . The line markings for each \widetilde{Ra} as listed in (a) also apply for (b),(c). All profiles are anti-symmetric with respect to Z so that only the lower half of each profile is displayed.

averaging is performed such that $W_1 \approx W_2 \approx W_3$. Figure 7a shows comparisons of the W profiles for the SA and MA dynamo models, including the individual components and the rms profile. The horizontal structure is displayed for the midplane of $W_1(Z)h_1(x,y) + W_2(Z)h_2(x,y) + W_3(Z)h_3(x,y)$ in figure 7b where the hexagonal structure is evident, although deviating slightly from a perfect hexagonal pattern.

The time averaged profiles of the x-component of the mean magnetic field are shown in Figure 8. The mean magnetic field also forms boundary layers that progressively thin with increasing \widetilde{Ra} . The mean magnetic field oscillates in time for all of the cases investigated here; however, as \widetilde{Ra} is increased, a broad range of frequencies become excited which gives way to a complex temporal dependence even for these relatively simple single mode convection solutions.

Figure 9 shows the influence of Ra on the helicity for both (a) non-magnetic convection solutions and (b,c) magnetic solutions. As the thermal forcing increases, the anti-symmetric profiles form boundary layers which move closer to the boundaries and become narrower than the non-magnetic counterparts. These results are a clear indication that the convection is significantly modified by the presence of the magnetic field.

3.3. Dissipation and heat transfer

Comparison of these runs with non-magnetic convection reveals lower Nusselt numbers and larger magnitude interior mean temperature gradients for a fixed value of \widetilde{Ra} (see figure 10a,b). The non-magnetic case approaches an isothermal interior according to the scaling law $|\partial_Z \overline{\theta}| \sim \widetilde{Ra}^{-1}$ (Sprague *et al.* 2006), while the SA dynamos show a diminished exponent and the MA dynamos saturate at large Rayleigh numbers. Similarly, the kinetic energy is diminished from non-magnetic rotating convection single mode solutions, as shown in figure 10c. Given that these results can be tied to the equality of power integrals in (2.57), the exchange and balance of energy can be understood through a comparison of the three figures. Since $Nu = \varepsilon_{\Theta}$ and $2k^2 E_K = \varepsilon_u$, a scaling relation of $\varepsilon_{\Theta}^c \widetilde{Ra}_{E_K}^a = \alpha \alpha$ is expected for the non-magnetic case where $\varepsilon_B = 0$. Scaling $2k^2 E_K = \varepsilon_u \sim \widetilde{Ra}_{E_K}^a$ and $Nu = \varepsilon_{\Theta} \sim \widetilde{Ra}_{Nu}^{\beta}$; figures 10a,c confirm the balance with $(\widetilde{Ra}_{Nu})^{1.5} \widetilde{Ra}^1 = (\widetilde{Ra}_{E_K})^{2.5}$. For the magnetic case, there must be a balance between the three terms in (2.57) since $\varepsilon_B \neq 0$. The SA model shows ε_u and ε_{Θ} are diminished from the non-magnetic case with scalings of $\widetilde{Ra}_{E_K}^{1.5}$ and $\widetilde{Ra}_{Nu}^{1.2}$ so that the non-magnetic balance is violated. This Cambridge University Press



Figure 10: Heat transfer quantities and volume-integrated energies for the two classes of dynamo solutions and non-magnetic ($\overline{\mathbf{B}} = 0$) SA results: (a) Nusselt number Nu; (b) mean temperature gradient at the vertical midplane $\partial_Z \overline{\theta}(Z = 0.5)$; (c) kinetic energy E_K ; and (d) fraction of ohmic dissipation $f_{ohm} = \varepsilon_B/(\varepsilon_B + \varepsilon_u)$. Results for the SA model are shown by circles, the MA model results are shown by triangles and non-magnetic convection results are denoted with a square. Here Pr = 1, $\widehat{Pm} = 0.7$ and $S_B = 10^2$. The $\widetilde{Ra} = 20$ results closely align with the non-magnetic result, as it is barely above the critical threshold for dynamo action and the magnetic energy remains small. All other \widetilde{Ra} results included are well above the threshold for dynamo action, with critical $\widetilde{Pm_c}$ values reported in column 5 of table 1.

implies from (2.57) that ε_B must scale around $\widetilde{Ra}_B^{2.2}$. The calculated scaling of the ohmic dissipation (not shown) is approximately $\varepsilon_B \sim \widetilde{Ra}_B^{2.5}$, indicating that ohmic dissipation is dominant over viscous dissipation in the fluid layer for high \widetilde{Ra} . Figure 10d shows the fraction of ohmic dissipation, where it is observed that ε_B dominates over ε_u at large \widetilde{Ra} for both the SA and MA models.

The mean temperature profiles for a range of Ra are shown in figure 11 for the magnetic SA case (plot a) and the magnetic MA case (plot b). The SA cases resemble that of non-magnetic convection in which the mean temperature profiles become increasingly isothermal in the fluid interior and show the development of thin thermal boundary layers as the Rayleigh number is increased. In contrast, the MA cases show a more Cambridge University Press



Figure 11: Mean temperature profiles θ for the (a) SA and (b) MA dynamo solutions and a range of \widetilde{Ra} .

complicated dependence on Ra, where the interior mean temperature gradient saturates at large Ra, though the thermal boundary layers continue to thin with increasing Ra. Both dynamo solutions highlight the decreased efficiency of thermal mixing in the interior relative to the non-magnetic case. While it is known that the mean temperature gradient for non-magnetic single mode solutions does not saturate with increasing Ra, multimode simulations have demonstrated saturation for $Pr \leq 1$ (Julien *et al.* 2012*a*).

The dissipation relations for non-magnetic convection and the dynamo can be compared by setting \widetilde{Ra}/Pr^2 fixed and equal in both expressions (2.54) and (2.55). These expressions can then be rearranged and equated to indicate that

$$\frac{\varepsilon_u}{Nu-1} + \frac{\varepsilon_B}{Nu-1} = \frac{\varepsilon_u^c}{Nu^c - 1}.$$
(3.1)

Given that $\varepsilon_B \ge 0$, this can be written as an inequality

$$\frac{\varepsilon_u}{Nu-1} \leqslant \frac{\varepsilon_u^c}{Nu^c - 1}.$$
(3.2)

To achieve this balance, there are three possibilities: (i) $\varepsilon_u \leq \varepsilon_u^c$ and $Nu \geq Nu^c$, (ii) $\varepsilon_u < \varepsilon_u^c$ and $Nu < Nu^c$ or (iii) $\varepsilon_u > \varepsilon_u^c$ and $Nu > Nu^c$. Here we use the bold < symbols to denote that the quantity must increase or decrease an amount large enough to compensate for an opposing increase or decrease; for instance, in the third case the relative increase in ε_u (which increases the fraction) must be compensated by the relative increase in Nu to make the inequality in (3.2) hold.

In table 2, the decrease in both ε_u and Nu, relative to the non-magnetic case, can be seen over the entire range of \widetilde{Ra} , indicating that these runs are in case (ii). Upon substituting these values in to (3.2), we find the relationship is satisfied. Equality in (3.2) is only achieved when $\varepsilon_B \approx 0$ as for the $\widetilde{Ra} = 20$, $\widetilde{Pm} = 0.7$ case. These results indicate that in these single mode models dynamo action acts to throttle the heat flux through the layer.

Experimental studies of non-rotating convection in the presence of an imposed vertical magnetic field report a decrease in the heat transfer from the non-magnetic case (Cioni *et al.* 2000; Aurnou & Olson 2001). However, DNS results from self-consistent rotating dynamos have found that the heat transfer increases in the presence of magnetic forces Cambridge University Press

24

Single mode investigations of the QG dynamo model

\widetilde{Ra}	$\begin{vmatrix} \overline{\mathbf{B}} = 0\\ \overline{Nu^c} \end{vmatrix}$	ε^c_u	$\frac{\mathrm{SA}}{Nu}$	ε_u	ε_B	f_{ohm}	$\left \frac{\mathrm{MA}}{Nu} \right $	ε_u	ε_B	f_{ohm}
$20 \\ 40$	$5.3584 \\ 19.177$	87.1645 727.08	$5.3584 \\ 11.3346$	$87.1645 \\ 223.27$	$O(10^{-7})$ 190.10	$\begin{array}{c} 0 \\ 0.46 \end{array}$	$5.36 \\ 10.3$	$87.1645 \\ 317.68$	$\begin{array}{c} 0 \\ 54.32 \end{array}$	$\begin{array}{c} 0 \\ 0.15 \end{array}$
60	37.703	2202.2	17.7433	370.45	634.09	0.63	13.3	503.94	234.06	0.32
100	83.126	4003.3 8212.6	$^{24.8}_{31.8}$	548.91 757.1	1302.4 2412	$0.71 \\ 0.76$	14.5 15.5	820.60	405.42 629.39	$0.38 \\ 0.43$
200	223.64	44528	70.7	1988	11951	0.86	22.8	1438.6	2931.6	0.67

Table 2: Various quantities from the power integral relations given in equations (2.54) and (2.57) for a range of \widetilde{Ra} with $\widetilde{Pm} = 0.7$, Pr = 1, k = 1.3048 and $S_B = 10^2$. Here ε_u^c is the total viscous dissipation for non-magnetic convection, and ε_u and ε_B are the total viscous and ohmic dissipations for dynamo solutions. The fraction of ohmic dissipation is defined by $f_{ohm} = \varepsilon_B/(\varepsilon_B + \varepsilon_u)$. Note that for $\widetilde{Ra} = 20$, the critical magnetic Prandtl number is $\widetilde{Pm}_c \approx 0.65$.



Figure 12: The Nusselt number (a) and magnetic energy (b) results for an MA run with $\widetilde{Ra} = 40$, Pr = 1, $\widetilde{Pm} = 0.7$ and $S_B = 10^3$ showing the time variability and intermittent bursts of increased E_M in the runs. For this run, the mean magnetic diffusion time is $t = O(10^3)$.

relative to non-magnetic simulations (Stellmach & Hansen 2004; Christensen & Aubert 2006). The DNS dynamo model results can therefore be classified as case (iii), given the higher reported kinetic energies as well. One possible explanation for the discrepancy between the change in Nu of the present asymptotic model and the DNS dynamo results is that the fixed horizontal structure considered here eliminates this phenomenon. Our results indicate that the restructuring of the convective planform in the presence of a magnetic field is crucial for enhancing the heat transfer (Chandrasekhar 1961; Fearn 1979).

3.4. Intermittency

Figure 12 shows the heat transfer and magnetic energy results from a MA model run with $\widetilde{Ra} = 40$, Pr = 1, $\widetilde{Pm} = 0.7$ and $S_B = 10^3$ that displays intermittent changes with Cambridge University Press

M. Plumley, M. A. Calkins, K. Julien and S. M. Tobias

time. While the results show large variations for the majority of the time, intermittent bursts of nearly periodic patterns appear at some intervals. Whereas the Nusselt number shown in figure 12a is characterized by smaller fluctuations during these periods, we observe large variations (bursts) in the magnetic energy (figure 12b). This behavior is generally consistent with our findings that the heat transfer is reduced in the presence of strong magnetic field for these single mode solutions. The bursts in activity appear over roughly $1/5S_B = 1/5E^{-1/2}$ diffusion times. Throughout these bursts, the averaged Nusselt number and kinetic energy values decrease and the averaged magnetic energy increases in comparison to averages outside of these intermittent periods. The vertical structure functions show more variation between them during these bursts as well. Similar intermittent behavior was observed in a recent DNS investigation (Bushby *et al.* 2018), though the DNS results were found for parameter values disparate from those considered here.

4. Conclusion

This study investigates the dynamics of the QGDM, which is a fully nonlinear dynamo model that is valid in the limit of low Rossby and low magnetic Prandtl numbers. This regime is relevant to the dynamics of Earth's liquid outer core and planetary interiors and cannot currently be accessed through direct numerical simulation or laboratory experiments. Using the single mode technique of assuming periodic horizontal planforms characterized by a single horizontal wavenumber k, the problem is reduced from three spatial dimensions to a single dimension in the vertical. While single mode solutions are known to be unstable in multimode simulations, they provide a useful comparison for more realistic dynamics (Sprague *et al.* 2006). The dimensional simplification allows for a detailed examination of time-stepping methods to handle the three intrinsic time scales within the QGDM and permits an investigation of self-consistent dynamo action resulting from a specialized class of solutions.

The mean temperature equation is handled instantaneously on the fast convective timescale to avoid the computationally demanding length of the vertical diffusion times that would be needed. This is achieved by factoring the time averaging over t from the mean temperature equation (see Appendix A). The mean induction equation is then tested with two multiscale methods: the common-in-time stiff approach where the induction timescale is converted to the convective timescale, i.e. $\partial_{\tau} = \epsilon^{-3/2} \partial_t$; and the heterogenous multiscale approach, where the mean induction equations are updated at intervals of fast convective time specified by $S_B \Delta t$ and use the fast-time averaging of Reynolds, convective and magnetic stresses from that interval. We find that both of these methods provide consistent results. Moreover, varying the time interval S_B over which the induction equation is averaged shows that self-similar behavior is reached for longer averaging windows, indicating that sufficient scale separation of the equations is necessary for this model. The dynamics, especially for the MA model, show strong time dependence and smooth results rely on extensive time averaging of the solutions, despite reaching statistically steady heat transfer and magnetic energy results.

Although the single mode solutions represent a highly simplified class of solutions, they nevertheless yield complex spatiotemporal dynamics. For sufficiently large Rayleigh numbers, both classes of solutions (SA and MA) yield chaotic dynamos with a broad range of time scales. Moreover, in contrast to previous kinematic investigations of the QGDM, the present work demonstrates that the dynamo-generated magnetic field has a strong influence on the structure and dynamics of the convection. The helicity of the Cambridge University Press

flow is strongly modified and gives rise to magnetic fields that differ substantially in both spatial structure and temporal behavior from the kinematic solutions.

All the dynamo solutions found in the present study show a decrease in the heat transfer relative to non-magnetic solutions at the same Rayleigh number. In contrast, DNS investigations (e.g. Stellmach & Hansen 2004; Christensen & Aubert 2006) generally find the opposite effect whereby the heat transfer is increased in the presence of magnetic field. There is no formal proof that either case should be preferred. It is likely that our decreased heat transfer results are a consequence of the single mode, fixed horizontal planform solutions assumed in the present work. Our results therefore suggest that the restructuring of the convective planform (the single wavenumber, hexagonal horizontal structure assumed in the ansatz) plays a critical role in allowing enhanced heat transfer. This restructuring of the dominant horizontal wavenumber, or a combination of these two effects.

A detailed investigation of the dissipation mechanisms in the single mode solutions shows that at sufficiently large Rayleigh numbers ohmic dissipation dominates over viscous dissipation. The solutions show a monotonic increase of the fraction of ohmic dissipation f_{ohm} with increasing \tilde{Ra} . In general, both plane layer (Tilgner 2014) and spherical (Christensen & Aubert 2006) DNS studies find a non-monotonic dependence of f_{ohm} on the thermal forcing. At present, it is unclear why such behavior occurs, though we speculate that it may be related to the loss of rotational dominance in some cases due to the finite values of the Ekman numbers employed in DNS. Our results show that viscous dissipation plays a minor role in the heat transfer for increased forcing. Thus, the model is not categorically viscously dominated, despite the presence of a single, viscously-selected horizontal scale that induces the dynamo.

By construction, the QGDM is designed to operate in the regime of low Ro and low Pm, a regime currently challenging for DNS and laboratory experiments to probe. Accordingly, this model provides results that cannot truly be compared with current direct numerical simulations, which remain limited to higher Pm values. Connections between the results of the asymptotic model and DNS must come from simulations of an alternative derivation of the QGDM, which features a Pm = O(1) and is not an α^2 dynamo (Calkins *et al.* 2015). However, some qualitative comparisons are made, for instance, in regards to the intermittent behavior of the magnetic energy. This extended regime that the QGDM can access is made possible by the filtering of fast inertial waves and fast Alfvén waves that are of secondary importance to the comparatively slow magneto-convective dynamics resulting from dynamo action. Ultimately, this study provides a foundation for future explorations of the model, where all dynamically active scales are permitted, through diagnostics for time integration strategies, requirements of self-similar behavior and indications that the magnetic field strongly influences the convective flow.

Acknowledgements This work was supported by the NASA Earth and Space Science Fellowship program (MP) and the National Science Foundation under award numbers DMS-1317666 and EAR-1620649 (MAC, KJ).

Appendix A. Derivation of the instantaneous mean temperature equation

The mean temperature equations (2.7) and (2.8) are deduced from the equation

$$\partial_T \overline{\overline{\theta}}^{\tau} + \partial_\tau \overline{\theta} + \partial_Z (\overline{w\theta'}) = \frac{1}{Pr} \partial_{ZZ} \overline{\overline{\theta}}^{\tau} .$$
 (A1)

occurring at $O(\epsilon^2)$ in the asymptotic development (Calkins *et al.* 2015). For clarity, we repeat here that the overbar $\overline{}$ from (2.3) denotes averaging over the small spatial variables and fast convective time t, and $\overline{}^{\tau}$ denotes averaging over mean magnetic time τ . Setting $\overline{\theta}(Z,\tau,T) = \overline{\overline{\theta}}^{\tau}(Z,T) + \epsilon^{1/2}\overline{\theta}(Z,\tau,T)$, (A 1) can be expressed as a composite equation

$$\epsilon^{-1/2}\partial_{\tau}\overline{\theta} + \partial_{Z}\left(\overline{w\theta'}\right) = \frac{1}{Pr}\partial_{ZZ}\overline{\theta} + O(\epsilon^{1/2}).$$
(A 2)

Decomposing the averaging into its spatial and temporal components gives

$$\epsilon^{-1/2}\partial_{\tau}\frac{1}{\Gamma}\int_{0}^{\Gamma}\overline{\theta}^{\mathcal{V}}dt + \partial_{Z}\frac{1}{\Gamma}\int_{0}^{\Gamma}\overline{(w\theta')}^{\mathcal{V}}dt = \frac{1}{Pr}\partial_{ZZ}\frac{1}{\Gamma}\int_{0}^{\Gamma}\overline{\theta}^{\mathcal{V}}dt + O(\epsilon^{1/2}).$$
(A3)

Here $\overline{\cdot}^{\mathcal{V}}$ denotes averaging over the small spatial scales and Γ is an interval in time, where $O(t) \ll \Gamma \ll O(\tau)$. Given that Γ is not a function of τ , T or Z, the integral and derivatives can be reordered in each term.

$$\frac{1}{\Gamma} \int_0^\Gamma \left[\epsilon^{-1/2} \partial_\tau \overline{\theta}^{\mathcal{V}} + \partial_Z \overline{(w\theta')}^{\mathcal{V}} = \frac{1}{Pr} \partial_Z \overline{\theta}^{\mathcal{V}} \right] dt + O(\epsilon^{1/2}) \,. \tag{A4}$$

A similar analysis follows for the mean induction equations. Choosing the strong constraint that the kernel must be zero, we obtain the instantaneous mean temperature equation

$$\epsilon^{-2}\partial_t \overline{\theta}^{\mathcal{V}} + \partial_Z \overline{(w\theta')}^{\mathcal{V}} = \frac{1}{Pr} \partial_{ZZ} \overline{\theta}^{\mathcal{V}}.$$
 (A 5)

This equation expresses the slowly varying nature of the mean temperature and, following Sprague *et al.* (2006), the stationary state approximation is achieved on setting $\epsilon^{-2}\partial_t \overline{\theta}^{\nu} = \partial_T \overline{\theta}^{\nu} = 0.$

A.1. Single Mode Model Derivations

Utilizing the ansatz for the single amplitude (SA) model, the Lorentz term in the vorticity equation and the Lorentz force in the vertical momentum equation become

$$\overline{\mathbf{B}} \cdot \nabla_{\perp} j'_{z} = -\widetilde{Pm} \left(\partial_{x}^{2} h \, \overline{B}_{x}^{2} + 2 \, \partial_{x} \partial_{y} h \, \overline{B}_{x} \overline{B}_{y} + \partial_{y}^{2} h \, \overline{B}_{y}^{2} \right) \Psi, \tag{A 6}$$

$$\overline{\mathbf{B}} \cdot \nabla_{\perp} b'_{z} = \frac{Pm}{k^{2}} \left(\partial_{x}^{2} h \, \overline{B}_{x}^{2} + 2 \, \partial_{x} \partial_{y} h \, \overline{B}_{x} \overline{B}_{y} + \partial_{y}^{2} h \, \overline{B}_{y}^{2} \right) W. \tag{A7}$$

Multiplication by h and horizontal averaging (denoted by the overbar) then gives

$$\overline{h\,\overline{\mathbf{B}}\cdot\nabla_{\perp}j_{z}'} = -\widetilde{Pm}\left(\overline{h\partial_{x}^{2}h}\,\overline{B}_{x}^{2} + 2\,\overline{h\partial_{x}\partial_{y}h}\,\overline{B}_{x}\overline{B}_{y} + \overline{h\partial_{y}^{2}h}\,\overline{B}_{y}^{2}\right)\Psi,\tag{A8}$$

$$\overline{h\,\overline{\mathbf{B}}\cdot\nabla_{\perp}b_{z}'} = \frac{Pm}{k^{2}} \left(\overline{h\partial_{x}^{2}h}\,\overline{B}_{x}^{2} + 2\,\overline{h\partial_{x}\partial_{y}h}\,\overline{B}_{x}\overline{B}_{y} + \overline{h\partial_{y}^{2}h}\,\overline{B}_{y}^{2}\right)W.$$
(A9)

The vorticity and momentum equations can then be written as

$$-k^{2}\partial_{t}\Psi - \partial_{Z}W = \widetilde{Pm}\left[\overline{(\partial_{x}h)^{2}}\overline{B}_{x}^{2} + 2\overline{(\partial_{x}h\partial_{y}h)}\overline{B}_{x}\overline{B}_{y} + \overline{(\partial_{y}h)^{2}}\overline{B}_{y}^{2}\right]\Psi + k^{4}\Psi, \quad (A\ 10)$$
$$\partial_{t}W + \partial_{Z}\Psi = \frac{\widetilde{Ra}}{Pr}\Theta - \frac{\widetilde{Pm}}{k^{2}}\left[\overline{(\partial_{x}h)^{2}}\overline{B}_{x}^{2} + 2\overline{(\partial_{x}h\partial_{y}h)}\overline{B}_{x}\overline{B}_{y} + \overline{(\partial_{y}h)^{2}}\overline{B}_{y}^{2}\right]W - k^{2}W, \quad (A\ 11)$$

where use has been made of the following identities

$$\overline{h\partial_x^2 h} = -(\partial_x h)^2, \qquad (A\,12)$$

$$\overline{h\partial_x\partial_yh} = -\overline{(\partial_xh\partial_yh)},\tag{A 13}$$

$$\overline{h\partial_y^2 h} = -\overline{(\partial_y h)^2}.$$
(A 14)

Upon noting that $\overline{\partial_x h \partial_y h} = 0$ for the single mode solutions, the complete set of equations is given by

$$\partial_t \Psi + \frac{1}{k^2} \partial_Z W = -\widetilde{Pm} \left[c_1 \,\overline{B}_x^2 + c_2 \,\overline{B}_y^2 \right] \Psi - k^2 \Psi, \tag{A 15}$$

$$\partial_t W + \partial_Z \Psi = \frac{\widehat{Ra}}{Pr} \Theta - \widetilde{Pm} \left[c_1 \,\overline{B}_x^2 + c_2 \,\overline{B}_y^2 \right] W - k^2 W, \tag{A16}$$

$$\partial_t \Theta + W \partial_Z \overline{\theta} = -\frac{k^2}{Pr} \Theta, \qquad (A\,17)$$

$$\partial_T \overline{\theta} + \partial_Z \left(\overline{W} \overline{\Theta} \right) = \frac{1}{Pr} \partial_{ZZ} \overline{\theta},$$
 (A 18)

$$\partial_{\tau}\overline{B}_{x} = -2\widetilde{Pm}c_{2}\,\partial_{Z}\left(\overline{\Psi}\overline{W}\,\overline{B}_{y}\right) + \frac{1}{\widetilde{Pm}}\partial_{ZZ}\overline{B}_{x},\tag{A 19}$$

$$\partial_{\tau}\overline{B}_{y} = 2\widetilde{Pm}c_{1}\partial_{Z}\left(\overline{\Psi W}\overline{B}_{x}\right) + \frac{1}{\widetilde{Pm}}\partial_{ZZ}\overline{B}_{y}.$$
 (A 20)

Here (A 18) is obtained from (A 5) and all overbars now denote area averaging over the horizontal spatial variables x, y. The averaged values of the derivatives depend on the planform chosen and are denoted as $(\overline{\partial_x h})^2 = c_1 k^2$ and $(\overline{\partial_y h})^2 = c_2 k^2$. For the hexagonal planform, $c_1 = c_2 = 1/2$. A similar analysis results in the MA model. The steady state approximation is achieved on setting $\partial_T \overline{\theta} = 0$.

REFERENCES

- AURNOU, J. & OLSON, P. 2001 Experiments on Rayleigh–Bénard convection, magnetoconvection and rotating magnetoconvection in liquid gallium. J. Fluid Mech. 430, 283–307.
- AURNOU, J. M., CALKINS, M. A., CHENG, J. S., JULIEN, K., KING, E. M., NIEVES, D., SODERLUND, K. M. & STELLMACH, S. 2015 Rotating convective turbulence in earth and planetary cores. *Phys. Earth Planet. Int.* 246, 52–71.
- BASSOM, A. P. & ZHANG, K. 1994 Strongly nonlinear convection cells in a rapidly rotating fluid layer. *Geophys. Astrophys. Fluid Dyn.* 76, 223–238.
- BUSHBY, P. J., KÄPYLÄ, P. J., MASADA, Y., BRANDENBURG, A., FAVIER, B., GUERVILLY, C. & KÄPYLÄ, M. J. 2018 Large-scale dynamos in rapidly rotating plane layer convection. *Astron. Astron. Astrophy.*.

CALKINS, M. A. 2018 Quasi-geostrophic dynamo theory. Phys. Earth Planet. Int. 276, 182–189.

- CALKINS, M. A., JULIEN, K. & TOBIAS, S. M. 2017 Inertia-less convectively-driven dynamo models in the limit of low Rossby number. *Phys. Earth Planet. Int.* 266, 54–59.
- CALKINS, M. A., JULIEN, K., TOBIAS, S. M. & AURNOU, J. M. 2015 A multiscale dynamo model driven by quasi-geostrophic convection. J. Fluid Mech. 780, 143–166.

- CALKINS, M. A., JULIEN, K., TOBIAS, S. M., AURNOU, J. M. & MARTI, P. 2016a Convectiondriven kinematic dynamos at low Rossby and magnetic Prandtl numbers: single mode solutions. *Phys. Rev. E.* 93, 023115.
- CALKINS, M. A., LONG, L., NIEVES, D., JULIEN, K. & TOBIAS, S. M. 2016b Convection-driven kinematic dynamos at low Rossby and magnetic Prandtl numbers. *Phys. Rev. Fluids.* 1, 083701.
- CHANDRASEKHAR, S. 1961 Hydrodynamic and Hydromagnetic Stability. Oxford: Oxford University Press.
- CHARNEY, J. G. 1948 On the scale of atmospheric motions. *Geofys. Publ.* 17, 3–17.
- CHARNEY, J. G. 1971 Geostrophic turbulence. J. Atmos. Sci. 28 (6), 1087–1095.
- CHILDRESS, S. & SOWARD, A. M. 1972 Convection-driven hydromagnetic dynamo. Phys. Rev. Lett. 29 (13), 837–839.
- CHRISTENSEN, U. R. & AUBERT, J. 2006 Scaling properties of convection-driven dynamos in rotating spherical shells and application to planetary magnetic fields. *Geophys. J. Int.* 166 (1), 97–114.
- CIONI, S., CHAUMAT, S. & SOMMERIA, J. 2000 Effect of a vertical magnetic field on turbulent Rayleigh-Bénard convection. *Phys. Rev. E* **62** (4), R4520.
- E, W., ENGQUIST, B., LI, X., REN, W. & VANDEN-EIJNDEN, E. 2007 Heterogeneous multiscale methods: a review. *Commun. Comput. Phys* **2** (3), 367–450.
- ELSASSER, W. M. 1956 Hydromagnetic dynamo theory. Rev. Mod. Phys. 28 (2), 135.
- FEARN, D. R. 1979 Thermally driven hydromagnetic convection in a rapidly rotating sphere. Proc. R. Soc. A **369**, 227–242.
- FINLAY, C. C. & AMIT, H. 2011 On flow magnitude and field-flow alignment at Earth's core surface. *Geophys. J. Int.* 186, 175–192.
- GREENSPAN, H. P. 1968 The Theory of Rotating Fluids. Cambridge, United Kingdom: Cambridge University Press.
- GREER, B. J., HINDMAN, B. W., FEATHERSTONE, N. A. & TOOMRE, J. 2015 Helioseismic imaging of fast convective flows throughout the near-surface shear layer. Astrophys. J. Lett. 803 (2), L17 (5 pp).
- HANOSAGE, S. M., DUVALL, T. L. & SREENIVASAN, K. R. 2012 Anomalously weak solar convection. Proc. Nat. Acad. Sci. 109 (30), 11928–11932.
- HAUT, T. & WINGATE, B. 2014 An asymptotic parallel-in-time method for highly oscillatory pdes. *SIAM J. Sci. Comput.* **36** (2), A693–A713.
- JONES, C. A. 2011 Planetary magnetic fields and fluid dynamos. Annu. Rev. Fluid Mech. 43, 583–614.
- JULIEN, K., AURNOU, J., CALKINS, M., KNOBLOCH, E., MARTI, P., STELLMACH, S. & VASIL, G. 2016 A nonlinear model for rotationally constrained convection with Ekman pumping. J. Fluid Mech. 798, 50–87.
- JULIEN, K. & KNOBLOCH, E. 1998 Strongly nonlinear convection cells in a rapidly rotating fluid layer: the tilted *f*-plane. J. Fluid Mech. 360, 141–178.
- JULIEN, K. & KNOBLOCH, E. 1999 Fully nonlinear three-dimensional convection in a rapidly rotating layer. *Phys. Fluids* **11** (6), 1469–1483.
- JULIEN, K., KNOBLOCH, E., MILLIFF, R. & WERNE, J. 2006 Generalized quasi-geostrophy for spatially anisotropic rotationally constrained flows. J. Fluid Mech. 555, 233–274.
- JULIEN, K., KNOBLOCH, E., RUBIO, A. M. & VASIL, G. M. 2012a Heat transport in low-Rossby-number Rayleigh-Bénard convection. Phys. Rev. Lett. 109, 254503.
- JULIEN, K., KNOBLOCH, E. & TOBIAS, S. M. 1999 Strongly nonlinear magnetoconvection in three dimensions. *Physica D* 128, 105–129.
- JULIEN, K., RUBIO, A. M., GROOMS, I. & KNOBLOCH, E. 2012b Statistical and physical balances in low Rossby number Rayleigh-Bénard convection. Geophys. & Astrophys. Fluid Dynam. 106 (4-5), 392–428.
- JULIEN, K. & WATSON, M. 2009 Efficient multi-dimensional solution of PDEs using Chebyshev spectral methods. J. Comput. Phys. 228, 1480–1503.
- KING, E. M. & AURNOU, J. M. 2013 Turbulent convection in liquid metal with and without rotation. Proc. Nat. Acad. Sci. 110 (17), 6688–6693.
- MALECHA, Z., CHINI, G. & JULIEN, K. 2014 A multiscale algorithm for simulating spatiallyextended langmuir circulation dynamics. J. Comput. Phys. 271, 131 – 150.

- MATTHEWS, P. C. 1999 Asymptotic solutions for nonlinear magnetoconvection. J. Fluid Mech. **387**, 397–409.
- MIZERSKI, K. A. & TOBIAS, S. M. 2013 Large-scale convective dynamos in a stratified rotating plane layer. *Geophys. Astrophys. Fluid Dyn.* 107, 218–243.
- MOFFATT, H. K. 1978 Magnetic Field Generation in Electrically Conducting Fluids. Cambridge: Cambridge University Press.
- OSSENDRIJVER, M. 2003 The solar dynamo. The Astron. Astrophys. Rev. 11, 287–367.
- PARKER, E. N. 1955 Hydromagnetic dynamo models. Astrophys. J. 122, 293–314.
- PEDLOSKY, J. 1992 Geophysical Fluid Dynamics. New York, NY: Springer.
- PEYRET, R. 2013 Spectral methods for incompressible viscous flow. Springer Science & Business Media.
- POZZO, M., DAVIES, C. J., GUBBINS, D. & ALFÉ, D. 2013 Transport properties for liquid silicon-oxygen-iron mixtures at Earth's core conditions. *Phys. Rev. B* 87, 014110.
- PROUDMAN, J. 1916 On the motion of solids in a liquid posessing vorticity. Proc. R. Soc. Lond. 92, 408–424.
- SCHAEFFER, N., JAULT, D., NATAF, H.-C. & FOURNIER, A. 2017 Turbulent geodynamo simulations: a leap towards Earth's core. *Geophys. J. Int.* 211, 1–29.
- SCHUBERT, G. & SODERLUND, K. M. 2011 Planetary magnetic fields: Observations and models. Phys. Earth Planet. Int. 187 (3), 92–108.
- SOWARD, A. M. 1974 A convection-driven dynamo: I. the weak field case. *Phil. Trans. R. Soc. Lond. A* 275, 611–646.
- SPRAGUE, M., JULIEN, K., KNOBLOCH, E. & WERNE, J. 2006 Numerical simulation of an asymptotically reduced system for rotationally constrained convection. J. Fluid Mech. 551, 141–174.
- STELLMACH, S. & HANSEN, U. 2004 Cartesian convection driven dynamos at low Ekman number. *Phys. Rev. E* **70** (5), 056312.
- TAYLOR, G. I. 1923 Experiments on the motion of solid bodies in rotating fluids. Proc. R. Soc. Lond. 104, 213–218.
- TILGNER, A. 2014 Magnetic energy dissipation and mean magnetic field generation in planar convection-driven dynamos. *Phys. Rev. E* **90**, 013004.
- TOBIAS, S. M., CATTANEO, F. & BOLDYREV, S. 2011 *Ten Chapters in Turbulence*, chap. MHD Dynamos and Turbulence. Cambridge University Press.
- WALEFFE, F. 2001 Exact coherent structures in channel flow. J. Fluid Mech. 435, 93-102.

