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1 **Inorganically modified clay minerals: preparation, characterization, and arsenic adsorption**
2 **in contaminated water and soil**

3

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16

17 **Abstract**

18 The use of modified clay minerals for adsorbing arsenic (As) in contaminated soils is an
19 underexplored area of research. The adsorption behavior of As onto inorganically modified
20 smectite and kaolinite both in aqueous and soil media was studied. X-ray diffraction, infra-red
21 spectroscopy, scanning and transmission electron microscopy studies confirmed successful
22 modification of smectite through Fe-exchange and Ti-pillaring, and kaolinite through phosphate
23 binding. The modified smectites were more efficient than phosphate-bound kaolinite in adsorbing

24 As both in water and soil systems. Kinetic study revealed that the clay products reached adsorption
25 equilibrium within 3 h, and the data well fitted to the power function and simple Elovich equation
26 ($R^2 > 0.90$). The Freundlich isotherm model best described the As adsorption data ($R^2 > 0.86$) of
27 the modified clay products in both the systems. The Ti-pillared smectite exhibited the highest As
28 adsorption capacity ($156.54 \mu\text{g g}^{-1}$) in the aqueous medium, while the Fe-exchanged smectite was
29 the best material in the soil system ($115.63 \mu\text{g g}^{-1}$). The partition coefficient (K_d) and adsorption
30 efficiency (%) data also maintained the similar trend. Precipitation of As and binuclear complex
31 formation also took place in the soil system which made the metalloid non-labile as the time
32 passed. The inorganically modified clay products reported here hold a great potential to adsorb As
33 in contaminated groundwater, drinking water as well as soil.

34 **Key words:** Arsenic, inorganic modification, smectite, kaolinite, Freundlich isotherm, adsorption
35

36 **1. Introduction**

37 Arsenic (As) has become a major pollutant in soil and drinking water in many parts of the world.
38 Despite being relatively scarce in the natural environment (0.0005%; 20th abundant element in the
39 continental crust), arsenic is widely distributed over the globe (Gebreyowhannes, 2009). All over
40 the world where arsenic contamination in groundwater and its potentially severe human health
41 effects have been reported, the impact has been the highest in Bangladesh and the Bengal delta
42 basin of West Bengal, India (Chowdhury et al., 2000; Mukhopadhyay et al., 2002). More than 90%
43 of the total groundwater in West Bengal is affected by arsenic contamination (Sanyal and Nasar
44 2002). The contamination of water occurs due to the dissolution of minerals like arsenopyrites
45 from parent materials, geochemical reactions, biological activities and/or from anthropogenic
46 sources such as the leaching of manmade arsenic compounds from smelting of metal ores, and

47 wood preservatives (Shevade and Ford, 2004). There are two chemical hypotheses, namely
48 arsenopyrite oxidation hypothesis (Mandal et al., 1996) and ferric oxyhydroxide reduction
49 hypothesis (Bhattacharyya et al., 1997), which explains the widespread arsenic occurrence in the
50 groundwater of Bengal delta basin and Bangladesh. The latter hypothesis proved more consistent
51 according to some literature (Sanyal 1999; Aziz et al., 2016). According to this hypothesis, anoxic
52 condition of the aquifers caused the mobilization of arsenic from arsenic bearing minerals into the
53 groundwater.

54 Buildup of arsenic in soil due to the use of contaminated groundwater for irrigation has led to a
55 global environmental concern, especially for rice production and food security in South Asia (India
56 and Bangladesh) (Das et al., 2008; Khan et al., 2009). In addition, drinking of arsenic contaminated
57 groundwater is a direct health threat to the people in this region and other parts of the world (Liu
58 et al., 2002; Ng et al., 2003; Das et al., 2011; Sarkar et al., 2016). Consumption of arsenic
59 contaminated drinking water may cause kidney, urinary tract, liver, skin, and rectum cancers in
60 humans (Pontius, 1994). Non-carcinogenic diseases related to arsenic exposure are hypertension,
61 diabetes mellitus, cerebrovascular and cardiovascular systems, and dysfunction of respiratory
62 system (Thomas et al., 2001).

63 Arsenic is dominantly present in soils in inorganic forms which are a function of redox potential
64 and pH of the medium. Arsenate (As (V)) is the major arsenic species in surface water over the
65 pH range of 5 to 12 (Zeng, 2004). Arsenite (As (III)) is mostly found under reducing conditions,
66 and thus is the most dominant species in groundwater over the pH range of 2 to 9 (Zeng, 2004).

67 Anion exchange, adsorption, reverse osmosis, coagulation, co-precipitation and solvent extraction
68 are commonly known methods for removing arsenic from aqueous systems (Mohan and Pittman,
69 2007; Jadhav et al., 2015; Shakoor et al., 2016; Vithange et al., 2017). Various adsorbents

70 including activated carbon, biochar, agricultural and industrial byproducts, zeolite and clay
71 minerals, metal oxides and hydroxides, nanomaterials and resins were reported to remove arsenic
72 from contaminated water (Mohan and Pittman, 2007; Sharma et al., 2014). However, available
73 technologies for removing arsenic from soils are very limited. Phytoremediation of As from soil
74 could be carried out through plants such as *Pteris vittata* (Mandal et al., 2012; 2017; Fayiga and
75 Saha, 2016; Niazi et al., 2016). Ghosh et al. (2012) used farm yard manure and compost to make
76 chelation of arsenic with humate compounds and immobilize it in soils. While the influence of
77 organic matter on the mobility and bioavailability of arsenic species are ambiguous (Suda and
78 Makino, 2016; Wang et al., 2016), some inorganic amendments such as iron compounds,
79 phosphates, alkaline compounds, gypsum, biosolid, red mud, fly ash and clay minerals were found
80 effective in immobilizing As in soils (Lombi et al., 2004; Kumpiene et al., 2008; Miretzky and
81 Cirelli, 2010; Lee et al., 2011; Lim et al., 2016). However, many of these materials are expensive
82 and still do not have standard protocol of practical soil application.

83 Due to their low cost, worldwide distribution and superior physico-chemical properties (e.g., high
84 specific surface area, ion-exchange capacity, mechanical stability and lamellar structure), clay
85 minerals (with or without modification) have also found widespread research attention in the
86 remediation of metals and metalloids in water and soil (Sarkar et al., 2012; 2013; Sun et al., 2013;
87 Perelomov et al., 2016; Kumararaja et al., 2017). In addition, such materials are ideal for one-time
88 use requiring no regeneration.

89 Clay minerals modified with different inorganic ions like Fe, Al, Ti, organic acids, mineral acids,
90 polymers, surfactants and nanoparticles were found promising for As remediation in contaminated
91 water (Li et al., 2007; Boddu et al., 2008; Doušová et al., 2009; Akpomie and Dawodu 2016;
92 Sarkar et al., 2016). Ti-pillared smectite was considered in the present investigation due to its high

93 adsorption affinity towards arsenic in aqueous system as it forms a polyhydroxy stable
94 complexation with arsenic. Similarly the Fe-exchanged smectite forms stable iron-arsenate which
95 enhanced arsenic adsorption efficiency in aqueous system. However, use of these modified clay
96 minerals for the adsorption of As in contaminated soils is still an underexplored area of research.
97 There is an urgent need to develop cost-effective methods for the removal/immobilization of As
98 in contaminated soils. Therefore, the present study aimed to prepare and characterize three
99 different inorganically modified clay minerals (Fe-exchanged and Ti-pillared smectite, and
100 phosphate-bound kaolinite), and compare their As adsorption behavior in contaminated water and
101 soil.

102

103 **2. Materials and Methods**

104 2.1 Chemicals and clay minerals

105 Potassium dihydrogen phosphate (KH_2PO_4), iron sulphate heptahydrate ($\text{FeSO}_4 \cdot 7\text{H}_2\text{O}$), calcium
106 chloride dihydrate ($\text{CaCl}_2 \cdot 2\text{H}_2\text{O}$) and other chemicals were of analytical grade and purchased from
107 Merck Millipore, Mumbai, India. Sodium arsenate heptahydrate ($\text{Na}_2\text{HAsO}_4 \cdot 7\text{H}_2\text{O}$), and titanium
108 chloride (TiCl_4) were purchased from Sigma-Aldrich Chemicals Pvt. Ltd., New Delhi, India. The
109 As stock solution (1000 mgL^{-1}) was prepared using $\text{Na}_2\text{HAsO}_4 \cdot 7\text{H}_2\text{O}$ in 0.01 M CaCl_2 in double
110 distilled water. The working solutions of As were freshly prepared by diluting the stock solution
111 in 0.01 M CaCl_2 solution. 0.1 N HNO_3 and 0.1 N NaOH were used to adjust the pH of As solutions
112 as necessary.

113 The smectite and kaolinite samples in the form of bentonite and kaolin were purchased respectively
114 from S D Fine-Chem Limited, Mumbai, India and Molychem, Mumbai, India. The clay samples
115 contained 88% smectite (kaolinite and quartz as impurities) and 86% kaolinite (Fe-oxides and

116 quartz as impurities), respectively, as the main mineral composition. The average particle size of
117 unmodified, Fe-exchanged and Ti-pillared smectite was 158.10, 169.30 and 174.60 nm,
118 respectively, while this values was 189.40 and 171.70 nm, respectively, for unmodified and
119 phosphate-bound kaolinite. In order to reduce the cost of production, the raw materials were used
120 as received without any purification.

121

122 2.2 Soil sample

123 Arsenic contaminated soil sample was collected (0-15 cm depth, order: Inceptisol) from Mitrapur,
124 West Bengal, India (22.9981° N and 88.6121° E). The collected soil sample was air dried under
125 shade and ground with mortar and pestle. The ground sample was passed through 2-mm sieve for
126 further analysis purposes. The As content and other physico-chemical characteristics of the soil
127 are given in Table 1.

128

129 2.3 Preparation of inorganically modified clay minerals

130 *2.3.1 Preparation of iron-exchanged smectite*

131 The smectite sample was treated with 0.1 M FeSO₄ solution in double distilled water (Te et al.,
132 2015). In brief, 20 g of dried smectite was mixed with 200 mL of 0.1 M FeSO₄ in a 250 mL conical
133 flask and stirred for 24 h in a mechanical shaker with a speed of 200 rpm at room temperature. The
134 solid smectite was separated using centrifugation technique (4000 rpm for 5 min) and dried at
135 105°C for 24 h. The dried smectite sample was further heated to 350°C in a muffle furnace for 3 h
136 and cooled to room temperature. Furthermore, the sample was washed with double distilled water
137 until no reddish colour appeared upon addition of 1:10 phenanthroline (negative test for Fe) and
138 dried at 60°C for overnight in a hot air oven.

139

140 2.3.2 Preparation of Ti- pillared smectite

141 At first, Na-smectite was prepared by adding 0.1 M NaOH drop by drop to a smectite suspension
142 (10% w/v) under continuous stirring in a magnetic stirrer for 16 h. The saturated particles were
143 separated through centrifugation (5000 rpm for 10 min) and washed with double distilled water
144 until it became chloride free (no white precipitation upon the addition of a drop of 0.1 N AgNO₃
145 solution). Then Na-smectite was dried at 80°C for 4 h for further use. Ti-pillaring of Na-smectite
146 followed the procedure of hydrolysis of TiCl₄ in reaction with HCl as described by Na et al. (2010).
147 In short, 100 mL of TiCl₄ was dissolved in 200 mL of 3 M HCl solution under vigorous stirring
148 for 2 h. The resultant solution was mixed with a 0.5 M HCl solution (1:2.5 ratio), and the pillaring
149 agent was obtained. Na-smectite (10 g) was slowly added to 490 mL of double distilled water
150 under stirring at 2000 rpm for 1 h. Then the mixture was aged for 12 h at 25°C. The pillaring agent
151 was then added into this mixture over a period of 1 h under vigorous stirring at about 3000 rpm.
152 Afterwards, the mixture was transferred to a bottle and treated in a water bath at 80°C for 6 h. The
153 product obtained was then filtered and washed with distilled water, dried, ground to powder.

154

155 2.3.3 Preparation of phosphate-bound kaolinite

156 The kaolinite sample was modified with KH₂PO₄ solution following the procedure of Amer et al.
157 (2010). Kaolinite (100 g) was equilibrated with 1 L of 200 mgL⁻¹ KH₂PO₄ in a rotary orbital shaker
158 at 150 rpm for 24 h. The sample was then washed several times with double distilled water to
159 remove excess phosphate from the mineral surface. Phosphate in the washout solution was found
160 negative, and the clay sample was dried in the oven at 105°C for 24 h.

161

162 2.4 Characterization

163 The cation exchange capacity (CEC) of the clay minerals (modified and unmodified) was
164 determined by $\text{CaCl}_2\text{-MgCl}_2$ method (Alexiades and Jackson, 1965) and specific surface area was
165 calculated using EGME (ethylene glycol monoethyl ether) adsorption method (Heilman et al.,
166 1965).

167 X-ray diffraction analyses of the powdered samples were performed using a Philips PW1710 X-
168 ray diffractometer with a monochromatic $\text{Cu-K}\alpha$ ($\lambda = 1.5418 \text{ \AA}$) source operating at 40 kV and 20
169 mA. The diffraction pattern recorded from 4° to 30° with a scan rate of $1.5^\circ 2\theta \text{ min}^{-1}$.

170 Fourier transform infra-red spectroscopy (FTIR) of the samples was carried out in the form of
171 KBr pellets using Bruker ALPHA, FTIR/ATR system (typically 64 scans, resolution 4 cm^{-1}), and
172 samples were scanned in $4,000\text{-}600 \text{ cm}^{-1}$ region. Scanning electron microscopy (SEM) images of
173 the clay products were taken using a VEGA3 LM scanning electron microscope (Tescan Orsay
174 Holding Instrument, Czech Republic) equipped with backscattered electron (BSE) and secondary
175 electron (SE) detectors.

176 Transmission electron microscopy (TEM) images of all the modified clay products were taken
177 using a JEOL TEM model JEM-1101 (JEOL Ltd., Japan).

178

179 2.5 Adsorption studies

180 2.5.1 Adsorption kinetics

181 The kinetics of As adsorption was performed at room temperature ($25\pm 1^\circ\text{C}$) and pH 6.2
182 (maintained with 0.1 N NaOH and 0.1 N HNO_3) with an As (arsenate) concentration of $50 \mu\text{g mL}^{-1}$.
183 Each type of unmodified and modified clay mineral (0.5 g) (solid: solution=1:20) was
184 equilibrated in triplicate with the As solution in plastic centrifuge tubes (50 mL) on an end over

185 end shaker (225 rpm). Two blanks (without adsorbent and water without As) were taken as
186 controls. The samples were equilibrated for different time intervals (20, 40, 60, 80, 100, 120, 140,
187 160 and 180 min). After equilibration, the clay samples were centrifuged at 10,000 rpm for 10 min
188 using a REMI-24 research centrifuge (REMI, Mumbai, India). The supernatant was immediately
189 filtered through 0.45- μm membrane filter. Final volume of the filtrate was made up using 2%
190 HNO_3 to avoid any damage to the interface on the instrument. The concentration of As in the
191 filtrate was analyzed against a standard As solution provided by the National Institute of Standards
192 and Technology (NIST) (purchased from Merck Millipore, Mumbai, India) using an inductively
193 coupled plasma mass spectrophotometer (ICP-MS) (Perkin Elmer NexION 300X, USA). The
194 amount of As adsorbed at time t (q_t) ($\mu\text{g g}^{-1}$) was calculated from the mass balance relationship
195 using Eq. 1:

$$196 \quad q_t = \frac{(C_0 - C_t) v}{m} \quad (\text{Eq. 1})$$

197 where, C_0 and C_t are the concentrations of As ($\mu\text{g mL}^{-1}$) at time zero and t, respectively; v is the
198 volume of As solution (mL); and m is the mass (g) of adsorbent. At equilibration, q_t was called as
199 the equilibrium adsorption and C_t was known as the equilibrium concentration.

200

201 *2.5.2 Adsorption isotherm in aqueous system*

202 Various clay adsorbents (1 g) were weighed and taken in plastic centrifuge tubes (50 mL). Arsenic
203 solutions (20 mL) of various concentrations (2.5 to 50 $\mu\text{g mL}^{-1}$ in 0.01 M CaCl_2) were then added
204 into the tubes. The suspensions were equilibrated for 24 h on an end over end shaker (225 rpm) to
205 ensure the equilibrium. All the experiments were conducted in triplicate at room temperature (25
206 $\pm 1^\circ\text{C}$) and pH 6.2 (maintained with 0.1 N HNO_3 and 0.1 N NaOH). After equilibration, the
207 supernatants were separated and analyzed for As concentrations by ICP-MS as described earlier.

208 Blank tests under the same conditions revealed no As adsorption on the tube wall during the
209 reaction period. The amount of As adsorbed was calculated using Eq. 1, and the As adsorption
210 efficiency (%) was calculated using Eq. 2:

$$211 \text{ Adsorption efficiency (\%)} = \frac{(C_0 - C_e)}{C_0} \times 100 \quad (\text{Eq. 2})$$

212 where, C_0 is the As concentration ($\mu\text{g mL}^{-1}$) at time zero; C_e is the equilibrium As concentration
213 ($\mu\text{g mL}^{-1}$) at time t.

214

215 *2.5.3 Adsorption isotherm in soil system*

216 The As contaminated soil (1.5 g) was amended with each type of clay products at 0.25% (w/w)
217 application rate in plastic centrifuge tubes. Then, 30 mL aqueous solutions of As (2.5 to 50 $\mu\text{g mL}^{-1}$
218 in 0.01 M CaCl_2) were added to the soil. Further experimental procedures and conditions were
219 maintained similar to those described earlier. All experiments were conducted in triplicate. Two
220 blank tests were conducted with soil (absence of clay products) and without soil (absence of both
221 soil and clay product). No As was adsorbed on the wall of the tubes. The amount of As adsorbed
222 by the clay products was calculated by subtracting the amount adsorbed by the soil (absence of
223 clays) from the amount adsorbed by the soil amended with the clay products using the mass balance
224 relationship (Eq. 1). Similarly, the As adsorption efficiency (%) of different clay products in the
225 soil system was calculated using Eq. 2.

226

227 *2.5.4 Partition coefficient (K_d)*

228 The K_d value (partition coefficient) of As adsorption on different modified clay minerals in the
229 aqueous and soil system was calculated using Eq. 3:

$$230 \quad K_d = q_e / C_e \quad (\text{Eq. 3})$$

231 where, K_d (mL g^{-1}) is the partition coefficient of As adsorption; q_e is the amount of As adsorbed
232 per unit mass of adsorbent at equilibrium ($\mu\text{g g}^{-1}$); C_e is the equilibrium As concentration ($\mu\text{g mL}^{-1}$).
233

234

235 2.6 Statistical analyses

236 The kinetic and isotherm data of As adsorption were fitted to model equations (power function and
237 simple Elovich model for adsorption kinetic study, and Freundlich model for adsorption isotherm
238 study) by nonlinear regression using least squares method. The correlation coefficients (R^2)
239 obtained from the regressions analyses were used to evaluate the applicability of the model
240 equations.

241

242 **3. Results and Discussion**

243 3.1 Characterization of clay products

244 *3.1.1 General properties*

245 Selected physico-chemical properties of the unmodified and modified clay minerals are given in
246 Table 2. The raw smectite was alkaline in nature (pH 8.20), while the kaolinite was neutral in
247 reaction (pH 6.75). After reactions of smectite with FeSO_4 and TiCl_4 , the products' pH values
248 decreased drastically to 3.93 and 5.95, respectively. This was because of saturation of the clay
249 mineral with cations (Fe^{2+} and Ti^{4+}) and also hydrolysis in the case of Ti-pillaring in the presence
250 of HCl. The pH of the kaolinite changed only slightly (pH 6.50) following the modification with
251 phosphate. The specific surface area (SSA) of the unmodified smectite was $202.69 \text{ m}^2 \text{ g}^{-1}$, whereas
252 it had increased up to 2-folds in Fe-exchanged smectite and Ti-pillared products (485.62 and
253 $437.06 \text{ m}^2 \text{ g}^{-1}$, respectively). The SSA of kaolinite also increased by five times from 18.40 to 89.08

254 $\text{m}^2 \text{g}^{-1}$ after the modification with phosphate (Adebowale et al., 2006). The SSA increase in the
255 clay products was likely due to the removal of impurities from the clay minerals (Rusmin et al.,
256 2016). Additionally, the replacement of exchangeable cations, particularly with Fe^{2+} and Ti^{4+} ,
257 might have exposed the clay edges and increased the SSA (Akpomie and Dawodu, 2016). The
258 cation exchange capacity (CEC) of the unmodified smectite and kaolinite was 118.50 and 22.25
259 $\text{cmol (p}^+) \text{kg}^{-1}$, respectively, whereas the value was 115.75, 105.75 and 40.50 $\text{cmol (p}^+) \text{kg}^{-1}$ for
260 Fe-exchanged smectite, Ti-pillared smectite and phosphate-bound kaolinite, respectively. The
261 CEC of smectite decreased after the exchange and pillaring reactions because the guest cations
262 and polyhydroxy titanium might have blocked the adsorption sites (Karamanis et al., 1997).
263 However, CEC of kaolinite was increased due to ligand adsorption of phosphate on to the clay
264 mineral surface (Ioannou and Dimirkou, 1997).

265

266 *3.1.2. X-ray diffraction (XRD)*

267 The XRD patterns of the unmodified and modified clay products are given in Fig. 1. The $d(001)$
268 value (basal spacing) was calculated from the corresponding first order reflection of the 001 plane.
269 A successful modification of the interlayer environment of the smectite was confirmed by the
270 deviation in the 2θ angle of the corresponding 001 reflection. The $d(001)$ value of the unmodified
271 smectite was 14.24 \AA (Fig. 1; Table 2) which was comparable to the $d(001)$ value of a typical
272 montmorillonite. Both the Fe-exchanged and Ti-pillaring modifications resulted in an increase of
273 the $d(001)$ value (interlayer expansion) (Fig. 1). The increase in $d(001)$ value of Fe-exchanged and
274 Ti-pillared smectite was 2.11 and 2.73 \AA , respectively (Table 2), which was in alignment with
275 previous reports (Karamanis et al., 1997; El Miz et al., 2014). Conversely, the $d(001)$ values of the

276 kaolinite and phosphate-bound kaolinite were 7.25 and 7.13 Å, respectively (Fig. 1), indicating
277 that phosphate might have attached only on broken edges of the kaolinite (Bidwell et al., 1970).

278

279 *3.1.3 Functional group analysis*

280 The FTIR bands at 3642 and 3452 cm⁻¹ (Fig. 2) in the spectra of unmodified smectite were
281 attributed to –OH bending mode of the adsorbed water (Farmer, 1974). Two other bands at 2435
282 and 1654 cm⁻¹ were ascribed to the basic layer silicate structure of smectite (Madejova and
283 Komadel, 2001). The bands at 1037 and 959 cm⁻¹ were due to the stretching vibration of Si-O-Si
284 groups and bending vibration of Al-O(OH)-Al groups, respectively (Madejova and Komadel,
285 2001). In the FTIR spectrum of Ti-pillared smectite (Fig. 2), the band at 3642 cm⁻¹ shifted to 3696
286 cm⁻¹, and the band due to vibration of water molecules in interlayer shifted to 3397 cm⁻¹ (Zhang et
287 al., 2008). Similarly, in the spectrum of Fe-exchanged smectite (Fig. 2), the band at 3642 shifted
288 to 3690 cm⁻¹, and the water molecules band moved to 3394 cm⁻¹. The shifting of bands from 3696
289 cm⁻¹ to the region of 3390 cm⁻¹ in both of the modified clay products suggested the H-bond
290 formation (Zhang et al., 2008). The presence of bands around 2419 cm⁻¹, 2412 cm⁻¹ and 1698 cm⁻¹
291 in both spectra of the modified clay products proved that their layer silicate structures were
292 preserved. The band around 1037 cm⁻¹ indicated Si-O-Si bonding in both the modified smectites.
293 The band at 962 cm⁻¹ in the spectrum of Fe-exchanged smectite confirmed the presence of Al-O-
294 Al bonding which was minimized in the case of Ti-pillared smectite. The presence of new bands
295 at 621cm⁻¹ in the spectrum of Ti-pillared smectite and at 634 cm⁻¹ in Fe-exchanged smectite might
296 be due to the formation of Si-O-Ti and Si-O-Fe bonds, respectively (Zhang et al., 2008). In both
297 unmodified and modified kaolinites, the appearance of the bands around 3690 and 3680 cm⁻¹,
298 which were attributed to the stretching vibration of inner surface hydroxyl groups (Fig. 2)

299 (Unuabonah et al., 2008; Amer et al., 2010). The band at 3551 cm^{-1} was attributed to the OH
300 stretching vibration of inner hydroxyl group (Unuabonah et al., 2007). The band at 3551 cm^{-1} in
301 the spectrum of kaolinite was splitted to 3503 and 3464 cm^{-1} in the phosphate-bound product (Fig.
302 2). The bands at 2414 and 1717 cm^{-1} in the unmodified kaolinite attributed to the typical layer
303 silicate structures and were found unchanged after the phosphate modification. The bands at 1035 ,
304 916 and 697 cm^{-1} in kaolinite were ascribed to Si-O-Si bond, Al-O(OH)-Al bond and Al-O bond,
305 respectively (Unuaboanh et al., 2008). Only the band for Si-O-Si bond shifted from 1035 to 1027
306 cm^{-1} after phosphate modification (Amer et al., 2010).

307

308 *3.1.4 SEM characterization*

309 The surface of the unmodified smectite was smooth and crystalline in structure (Fig. 3). However,
310 after modification, the Ti-pillared and Fe-exchanged smectite became swollen, fluffy and slightly
311 clustered (Akpomie and Dawodu, 2016). Swollenness and fluffiness occurred likely due to
312 increase in $d(001)$ values of the clay mineral following modification. These changes were very
313 clear in Ti-pillared and Fe-exchanged smectite (Fig. 3). The SEM image of the unmodified
314 kaolinite was very tiny and thin plate like structural surface (Fig. 3). The kaolinite also became
315 little bit aggregated in its surface morphology after the phosphate modification (Fig. 3)
316 (Unuabonah et al., 2008).

317

318 *3.1.5. TEM characterization*

319 The approximate shape and size of the clay products were determined through TEM images (Fig.
320 4). The shape of smectite particle appeared as a clear crystalline structure with an approximate
321 length 41.82 nm (in one dimension) (Fig. 4). But after modification, the shapes of Ti-pillared and

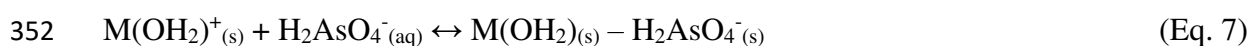
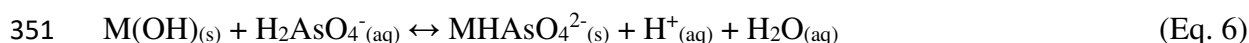
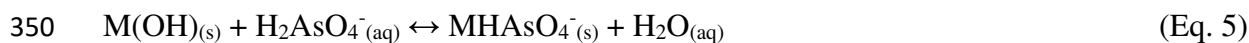
322 Fe-exchanged smectite (Fig. 4) particles became cloudy, fluffy, swollen and slightly clustered due
323 to increase in basal spacing (Seki and Yurdakoc, 2007), as was also observed in SEM images. The
324 approximate one dimensional length of Ti-pillared and Fe-exchanged smectite particles was 47.50
325 and 44.87 nm, respectively. The unmodified kaolinite particles appeared as a clear hexagonal, thin
326 plate like structure (Fig. 4) with approximate length of 275.78 nm (Ma and Eggleton, 1999). The
327 shape of phosphate-bound kaolinite (Fig. 4) did not change significantly, but remained clustered.
328 The approximate length of phosphate-bound kaolinite particle was 150.59 nm. This huge variation
329 in the length of kaolinite minerals before and after the phosphate modification might be due to
330 their irregular shaped particle distribution (Hassan and Abdu, 2014). Therefore, the TEM images
331 substantiated the SEM surface morphology results discussed earlier.

332

333 *3.2. Arsenic adsorption kinetics*

334 The initial adsorption of As on Ti-pillared smectite was slightly lesser than Fe-exchanged and
335 unmodified smectite (Fig. 5). The amount of As adsorbed were 448.0, 536.0 and 479.2 $\mu\text{g g}^{-1}$ by
336 Ti-pillared, Fe-exchanged and unmodified smectite at initial 20 min of adsorption (Fig. 5), which
337 suggested that Fe-exchanged smectite required less contact time to remove As from aqueous
338 system than the other two clay products. The exchanged iron (Fe) in the interlayer of smectite
339 probably formed iron hydroxides and oxides through oxidation and hydration processes during
340 reaction with water in the interlayer and the oxygen atoms at the end of each sheet of the clay
341 mineral (Stucki et al., 2002). The faster adsorption mechanism was in agreement with previous
342 reports using crystalline hydrous ferric oxide (Manna et al., 2003). Moreover, Fe had greater
343 affinity to adsorb As species than Ti (Lenoble et al., 2002). The Ti-pillared smectite started rapid
344 adsorption after a slow start. But from 40 minutes onwards, the Fe-exchanged smectite exhibited

345 more or less uniform adsorption rate and the trend continued throughout the rest of the reaction
346 period. In aqueous medium, Ti-pillared smectite adsorbed As by chemisorption mechanism
347 (Lenoble et al., 2002; Na et al., 2010). The chemisorption reaction mechanism of arsenate
348 adsorption onto Ti-pillared smectite in aqueous system is as follow:



353 where, $\text{M(OH)}_{(s)}$ represents a hydroxy group of the Ti-pillared smectite; $\text{H}^+_{(aq)}$ is the solution pH;
354 $\text{MHAsO}_4_{(s)}$ and $\text{MAso}_4_{(s)}$ the inner sphere complex, and $\text{M(OH}_2^+) - \text{H}_2\text{AsO}_4_{(s)}$ the outer sphere
355 complex of arsenate.

356 After 3 h of reaction, Fe-exchanged smectite indicated a higher As adsorption ($582.8 \mu\text{g g}^{-1}$) than
357 unmodified ($530.8 \mu\text{g g}^{-1}$) and Ti-pillared smectite ($490.0 \mu\text{g g}^{-1}$). The arsenate species (H_2AsO_4^-)
358 in aqueous phase is mostly dominant in between pH 2.2 to 6.98. Surface of the Fe-exchanged
359 smectite became more positively charged due to acidic nature of Fe-exchanged smectite and the
360 medium, which formed stable iron arsenate (Te et al., 2015). This might be a probable mechanism
361 of As adsorption by Fe-exchanged smectite. On the other hand, phosphate-bound kaolinite also
362 adsorbed a lesser amount of As ($454.0 \mu\text{g g}^{-1}$) than unmodified kaolinite ($467.6 \mu\text{g g}^{-1}$) initially at
363 20 min of reaction (Fig. 5), which indicated kaolinite's greater initial As removal capacity than the
364 modified product. It was likely due to the presence of some positive surface charges on broken
365 edges of the kaolinite as a function of mild acidic reaction (pH 6.2). It was reported that As
366 adsorption was regulated by its type of solution species, especially As (V) and surface charge (Xu
367 et al., 1988; Manning and Goldberg, 1996). But from 60 min onwards, unmodified kaolinite started

368 adsorbing considerably less amount than the phosphate-bound kaolinite. At equilibration,
369 phosphate-bound kaolinite adsorbed ($502.2 \mu\text{g g}^{-1}$) more than the unmodified kaolinite ($495.0 \mu\text{g g}^{-1}$).
370 The results demonstrated that all these clay minerals and their products increased As removal
371 at equilibrium (3 h). A similar inference was drawn by several researchers (Zeng, 2004; Mohapatra
372 et al., 2007; Na et al., 2010) who used different types of smectite and kaolinite. However, among
373 all the clay minerals, Fe-exchanged smectite exhibited the maximum As adsorption ($582.8 \mu\text{g g}^{-1}$)
374 and highest As removal capacity.

375 Two adsorption kinetic models (power function and simple Elovich) were used to fit the kinetic
376 data by nonlinear regression (Sparks, 1989). These models are expressed respectively by Eq. 8 and
377 9:

378 Power function equation: $q = a t^b$ (Eq. 8)

379 where, q = adsorbed As ($\mu\text{g g}^{-1}$) at time t (min); a and b are the constants; b is positive and less than
380 1.

381 Elovich equation: $q = a + b \ln t$ (Eq. 9)

382 where, q = adsorbed As ($\mu\text{g g}^{-1}$) at time t (min); a is intercept and b is the slope. The simple Elovich
383 parameters were estimated without using the origin ($q = 0, t = 0$).

384 On the basis of R^2 values (>0.90), the kinetics of As adsorption onto different clay minerals could
385 be described by both the model equations (Table 3). A similar kind of inference about well fitted
386 kinetics model of As adsorption was drawn by Zeng (2004) using a Fe (III)-Si binary oxide
387 adsorbent. The higher value of 'a' in the power function equation suggested the higher adsorbed
388 amount of adsorbate with time. Similarly, in simple Elovich equation, high value of slope 'b' and
389 intercept 'a' implied higher adsorption rate with time (Sparks, 1989).

390

391 3.3. Adsorption isotherm study in aqueous system

392 The unmodified smectite was a poor adsorbent of As in aqueous system because of its low
393 adsorption efficiency (48.56%) (Table 4). The Fe-exchanged and Ti-pillared smectites gave 73.64
394 and 78.82% sorption efficiency, respectively. The adsorption efficiencies were 66.07 and 79.05%
395 in unmodified kaolinite and phosphate-bound kaolinite, respectively. High adsorption efficiency
396 of modified smectite can be explained by high specific surface area and more availability of
397 exchangeable sites. Earlier, Grygar et al. (2007) reported that As adsorption was enhanced due to
398 the presence of high amount of Fe in Fe-exchanged smectite. Ti-pillared smectite gave the highest
399 adsorption efficiency among all the smectites. Owing to the expansion of the interlayer after Ti-
400 pillaring. The polyhydroxy titanium cations led to the generation of many active OH⁻ ions. Later
401 on, As species formed complexation with OH⁻ ions of smectite in the aqueous system (Na et al.,
402 2010). The adsorption efficiency of unmodified kaolinite was higher than unmodified smectite.
403 The edges of kaolinite might have provided some positive surface charges due to the presence of
404 mild acidic aqueous medium (pH 6.2) and As could form an aqua complexation (Mohapatra et al.,
405 2007). The phosphate-bound kaolinite accounted for higher adsorption efficiency likely due to its
406 higher surface area and the ligand exchange between phosphate and solution As species. Like the
407 adsorption efficiency, average partition coefficient [K_d (mLg⁻¹)] followed a similar trend. The Ti-
408 pillared smectite exhibited the maximum K_d value (175.77) followed by Fe-exchanged smectite
409 (121.89), phosphate-bound kaolinite (109.46), unmodified kaolinite (85.20) and unmodified
410 smectite (67.30) (Table 4). Ramesh et al. (2007) reported that the K_d value increased for As
411 adsorption because of the increased surface area and availability of more exchangeable sites in
412 polymeric Fe-modified montmorillonite. Therefore, variations in K_d values of different modified
413 clay products might be supported by the reasons explained for the adsorption efficiencies. Simply,

414 higher the sorption efficiency, higher is the partition coefficient. However, in general, with the
415 increasing equilibrium concentration of As, the adsorption amount decreased onto the clay
416 products in the aqueous system.

417 Freundlich adsorption isotherm has more flexibility. Any changes in adsorption behavior can be
418 best described by Freundlich isotherm. Data for As adsorption onto different clay minerals in
419 aqueous system (Fig. 6) was fitted to the Freundlich isotherm model only (Eq. 10) due to surface
420 heterogeneity of the modified clays and high significant coefficient of determination (R^2).

$$421 \quad q = K C^{1/n} \quad (\text{Eq. 10})$$

422 where, q is the amount of adsorbed As on different clay minerals at equilibrium ($\mu\text{g g}^{-1}$), and C is
423 the As equilibrium concentration in solution ($\mu\text{g mL}^{-1}$), K and (1/n) are the Freundlich constants.
424 K represents the adsorption of As at equilibrium concentrations, and (1/n) indicates the degree to
425 which adsorption is a function of As concentration (Table 5).

426 The 1/n values signify the degree of the intensity of adsorption. The 1/n values for As adsorption
427 on the modified clay minerals were less than unity indicating L-type isotherms (Giles et al., 1960).
428 L-type isotherms are characterized by the decrease in the adsorption at higher aqueous
429 concentrations of the solute. This suggests a greater competition for adsorption sites which become
430 limited as solute concentration in solution increases.

431 To obtain a meaningful comparison of their K values, the 1/n values for all modified clay minerals
432 should nearly be equivalent. The K ($\mu\text{g g}^{-1}$) value or As adsorption capacity was the maximum in
433 Ti-pillared smectite (156.54), followed by Fe-exchanged smectite (127.63), phosphate-bound
434 kaolinite (124.43), unmodified kaolinite (93.82) and unmodified smectite (66.03). The higher K
435 value might be explained by the respective K_d values. The higher partition coefficient (K_d)
436 suggested higher As adsorption capacity (K) by the modified clay products. The other reason might

437 be the surface complexation which is one of the stable bond formations that occurred in the case
438 of Ti-pillared smectite. The R^2 obtained for Freundlich model fitting were more than 0.90 for all
439 modified clays except the unmodified smectite and kaolinite (Table 5). Arsenate and arsenite
440 adsorption on iron oxide-coated sand and ferrihydrite was also best described earlier by the
441 Freundlich isotherm model (Thirunavukkarasu et al., 2001). The Freundlich isotherm model fitted
442 the experimental data very well due to the heterogeneous distribution of active sites on the
443 modified clay mineral surfaces. The $1/n$ values of all the clay minerals were less than 1 (Table 5),
444 which suggested that the adsorption process was favorable (Treybal, 1980). Similar results of
445 Freundlich model fitting and $1/n$ values less than unity for the adsorption of As (V), As (III) and
446 organic As species on polymeric Al/Fe-modified montmorillonite was reported by Ramesh et al.
447 (2007). Among all the smectites, Ti-pillared smectite adsorbed the maximum amount of As,
448 followed by Fe-exchanged and unmodified smectite (Fig. 6). The phosphate-bound kaolinite
449 adsorbed more As than unmodified kaolinite (Fig. 6). The reason might be that As has more spatial
450 compatibility with the adsorption sites of modified clay minerals because of their high surface
451 area. Therefore, the highest As adsorption capacity (K) onto the modified clay minerals in aqueous
452 system maintained the order: Ti-pillared smectite > Fe-exchanged smectite > phosphate-bound
453 kaolinite > unmodified kaolinite > unmodified smectite.

454

455 *3.4 Adsorption isotherm study in soil system*

456 Very few studies so far focused on the adsorption of As by modified clay minerals in contaminated
457 soils (Sarkar et al., 2012). The adsorption behavior of different modified clay minerals in the soil
458 system was slightly different from that of the aqueous system (Table 4). The adsorption
459 efficiencies were in the order: Fe-exchanged smectite > Ti-pillared smectite > phosphate-bound

460 kaolinite > unmodified kaolinite > unmodified smectite (Table 4). The maximum adsorption
461 efficiency of the Fe-exchanged smectite could be explained by higher surface area and higher
462 positive charge of Fe-exchanged smectites due to reduction in pH (Table 2). Similarly, the
463 increased surface area and ligand exchange reaction imparted greater adsorption efficiency by
464 phosphate-bound kaolinite than the unmodified kaolinite. Phosphate and arsenate ions could
465 compete non-specifically for ligand exchange and complexation sites on kaolinite when present
466 simultaneously in the soil system (Goldberg, 2002). But following modification with phosphate,
467 the kaolinite adsorbed more As than the unmodified counterpart. Earlier, Violante and Pinga
468 (2002) also reported a higher affinity of kaolinite to adsorb arsenate than phosphate. The
469 adsorption mechanism of Ti-pillared smectite would be the same as discussed in the aqueous
470 system. Partition coefficient [K_d (mLg^{-1})] values also followed the same trend of adsorption
471 efficiencies (Table 4). The Fe-exchanged smectite had the maximum K_d (100.13) value (Table 4).
472 However, adsorption efficiencies (except in unmodified smectite) and K_d values for all clay
473 products became lower in the soil system than the aqueous system because the former system was
474 more heterogeneous than the latter. In soil system several factors (e.g., clay type and content,
475 amorphous Fe content, total Fe content, soil pH and organic carbon) (Table 1) (Rivaz-Perez et al.,
476 2015) along with competitive ions might have influenced arsenic adsorption behavior on modified
477 clay minerals.

478 Arsenic adsorption data in the soil system were also fitted to the Freundlich equation (Eq. 10)
479 having $R^2 > 0.85$. Like the aqueous system, here also the $1/n$ values were less than unity (Table 5).
480 The Freundlich coefficient K (μgg^{-1}) was the highest in Fe-exchanged smectite (115.63) followed
481 by other clay products and maintained the same trends of adsorption efficiency and partition
482 coefficient (K_d) (Table 5). The pH of contaminated soil was 6.49, which was almost equal to the

483 pH value (6.2) maintained in the solid-solution interaction experiment. Mohapatra et al. (2007)
484 reported that montmorillonite and kaolinite adsorbed the maximum amount of arsenate at pH 6.0
485 and 5.0, respectively. As a result, the modified kaolinite accounted for a lesser K value for As
486 adsorption than the modified smectite.

487 Fig. 7 depicted some irregular or discontinuous adsorption trend followed in the soil system. With
488 the increase in solute concentration, there were breaks and reverse turns in the adsorption curves
489 for both smectite and kaolinites. The Fe-exchanged smectite and unmodified kaolinite exhibited a
490 more continuous increasing trend of adsorption than the other clay products (Fig. 7). The
491 deviations were greater at relatively higher equilibrium concentration of As. Ti-pillared smectite,
492 unmodified smectite and phosphate-bound kaolinite started showing the reverse breaks after 14.3,
493 22.74 and 12.11 $\mu\text{g mL}^{-1}$ equilibrium concentrations, respectively. This discontinuous trend of
494 solute concentration could be explained by the formation of binuclear As complexes and As
495 precipitation in the soil system (Datta 2002). In a similar fashion, phosphorus solution
496 concentration sharply decreased after attaining a maximum value due to the precipitation of
497 phosphorus in soil (Datta 2002). The non-labile P formation took place due to binuclear complex
498 formation and precipitation (Datta 2002). The precipitated phosphorus became crystallized and
499 non-labile with time (Datta 2002). Arsenic is also analogous to phosphorus (Marschner, 1995),
500 and probably followed the same mechanism in the soil system. Therefore, binuclear complexes
501 and precipitation might have taken place on the surfaces of these modified clay products in soil as
502 a function of soil constituents and competitive anions. Thus, precipitation or binuclear complex
503 formation of As by these modified clay products was quite effective to immobilize As from labile
504 form to non-labile form in the soil.

505

506 **4. Conclusions**

507 Modifications of smectite and kaolinite through Fe-exchange, Ti-pillaring and phosphate binding
508 were confirmed by XRD, FTIR, SEM and TEM characterization techniques. The smectite products
509 (Fe-exchanged and Ti-pillared) were more efficient in As adsorption than phosphate-bound
510 kaolinite in both soil and aqueous systems. While the Ti-pillared smectite was the most suitable
511 product for As adsorption in the aqueous system due to surface complex formation, the Fe-
512 exchanged smectite showed the best effectiveness in the soil system because of its high surface
513 area ($485.62 \text{ m}^2 \text{ g}^{-1}$) and higher affinity to As. The phosphate-bound kaolinite adsorbed As to a
514 greater extent than unmodified kaolinite through ligand exchange mechanism in both the systems.
515 One of the most important properties exhibited by some of the clay minerals was As precipitation
516 and/or binuclear complex formation in the soil system. This made As non-labile in soil with time
517 passed, which was most desirable and effective in remediating the metalloid. Thus, the application
518 of these modified clay products hold a great potential to immobilize As not only in contaminated
519 groundwater and drinking water, but also in contaminated soil.

520

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528

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Tables

Table 1: Physico-chemical properties of arsenic contaminated soil (Mitrapur, West Bengal, India)

Properties	Values
Clay (%)	26.5
Sand (%)	23.2
Silt (%)	50.3
Textural Class	Silty clay loam
pH (1:2.5)	6.49
EC (dS m ⁻¹)	0.26
Organic carbon (g kg ⁻¹)	4.50
Amorphous Fe (%)	0.29
Total Fe (%)	1.31
Total As (mg kg ⁻¹)	14.1
Olsen extractable As (mg kg ⁻¹)	3.6
CEC [cmol (p ⁺) kg ⁻¹]	24.7

Table 2: Characteristics of unmodified and modified clay minerals

Clay minerals	pH (1:2.5)	Specific surface area (m ² g ⁻¹)	Cation exchange capacity [cmol (p ⁺) kg ⁻¹]	d(001) (Å)
Unmodified smectite	8.20	202.69	118.50	14.24
Fe-exchanged smectite	3.93	485.62	115.75	16.35
Ti-pillared smectite	5.95	437.06	105.75	16.97
Unmodified kaolinite	6.75	18.40	22.25	7.25
Phosphate-bound kaolinite	6.50	89.08	40.50	7.13

Table 3: Estimated kinetic model parameters for arsenic adsorption on unmodified and modified clay minerals in aqueous medium

Models	Parameter	Ti-pillared smectite	Fe-exchanged smectite	Unmodified smectite	Phosphate-bound kaolinite	Unmodified kaolinite
Power function	a	398.66	475.78	408.56	397.2	435.18
	b	0.039	0.038	0.051	0.047	0.026
	R ²	0.92	0.95	0.97	0.98	0.96
Simple Elovich	a	392.88	469.12	398.2	389.43	432.76
	b	18.31	21.04	25.56	22.35	12.48
	R ²	0.92	0.94	0.97	0.98	0.96

The initial As addition was 50 µg As mL⁻¹, solid: solution = 1: 20, and pH maintained at 6.2

Table 4: Average partition coefficient (K_d) and adsorption efficiency of arsenic for different unmodified and modified clay minerals in aqueous and soil systems

Clays	K _d (mL g ⁻¹) in aqueous system	K _d (mL g ⁻¹) in soil system	Adsorption efficiency in aqueous system (%)	Adsorption efficiency in soil system (%)
Unmodified smectite	67.30	47.15	48.56	53.80
Fe-exchanged smectite	121.89	100.13	73.64	72.08
Ti-pillared smectite	175.77	85.45	78.82	68.58
Unmodified kaolinite	85.20	54.48	66.07	58.36
Phosphate-bound kaolinite	109.46	70.70	79.05	66.90

Table 5: Freundlich isotherm model constants for arsenic adsorption on different unmodified and modified clay minerals in aqueous and soil systems

Adsorption system	Parameters	Ti-pillared smectite	Fe-exchanged smectite	Unmodified smectite	Phosphate-bound kaolinite	Unmodified kaolinite
Aqueous	K ($\mu\text{g/g}$ ($\text{mL}/\mu\text{g}$) ^{1/n})	156.54	127.63	66.03	124.43	93.82
	1/n	0.46	0.46	0.39	0.62	0.50
	R ²	0.91	0.93	0.48	0.92	0.65
Soil	K ($\mu\text{g/g}$ ($\text{mL}/\mu\text{g}$) ^{1/n})	104.19	115.63	65.54	95.05	68.65
	1/n	0.48	0.49	0.46	0.49	0.53
	R ²	0.86	0.85	0.63	0.85	0.66

Figures

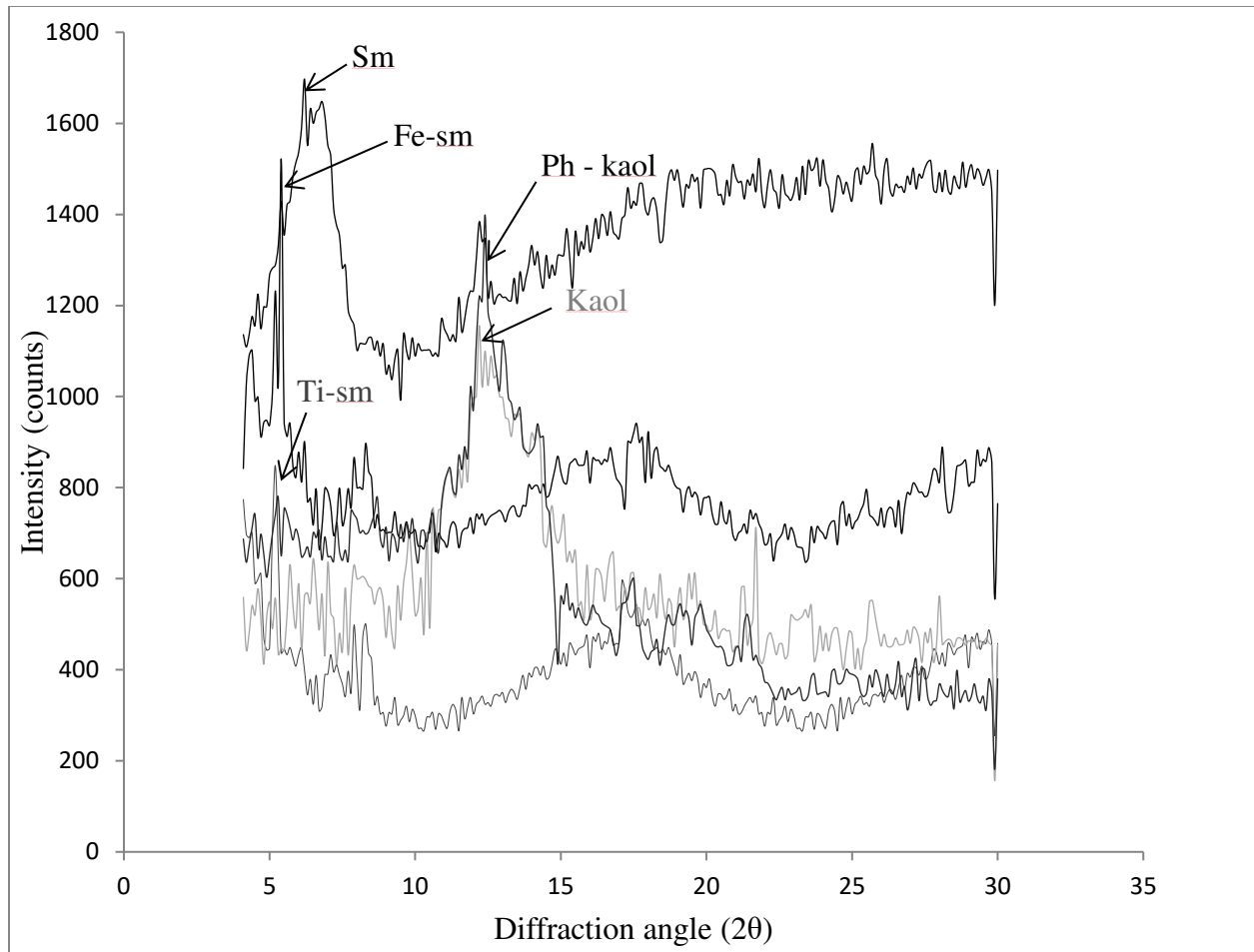


Fig. 1 X-ray diffraction patterns of unmodified smectite (Sm), Ti-pillared smectite (Ti-sm), Fe-exchanged smectite (Fe-sm), unmodified kaolinite (Kaol) and phosphate-bound kaolinite (Ph-kaol)

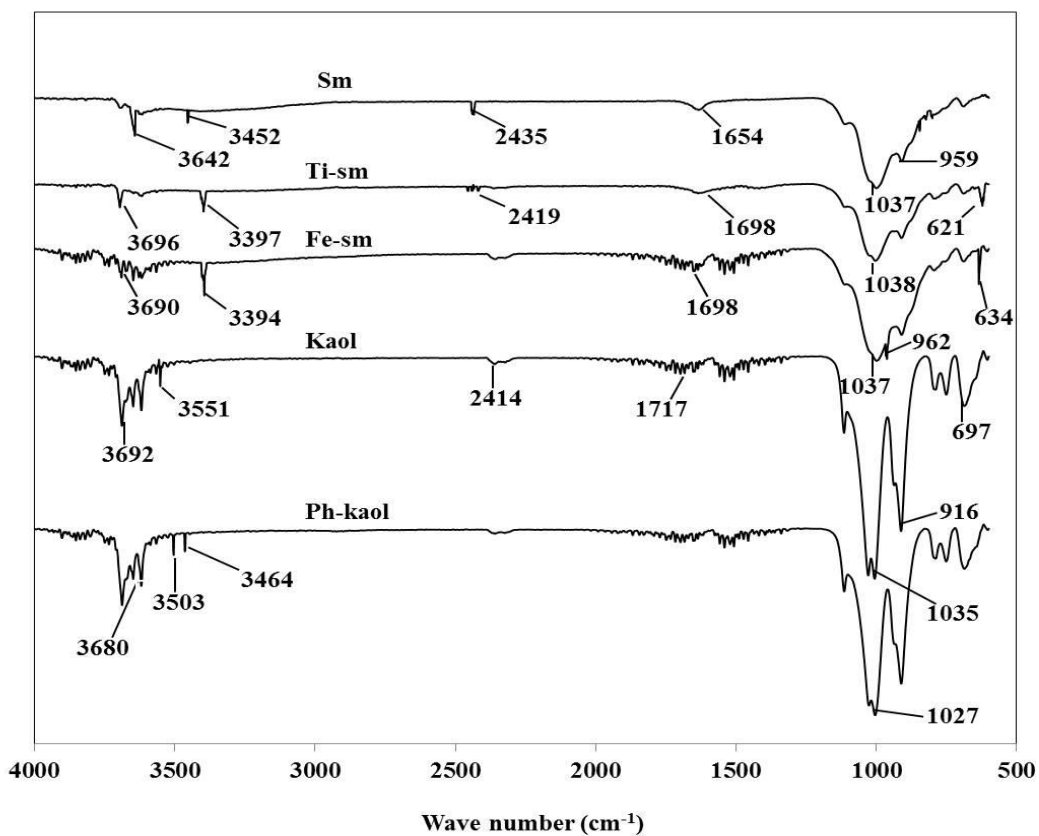


Fig. 2 FT-IR spectra of unmodified smectite (Sm), Ti-pillared smectite (Ti-sm), Fe-exchanged smectite (Fe-sm), unmodified kaolinite (Kaol) and phosphate-bound kaolinite (Ph-kaol)

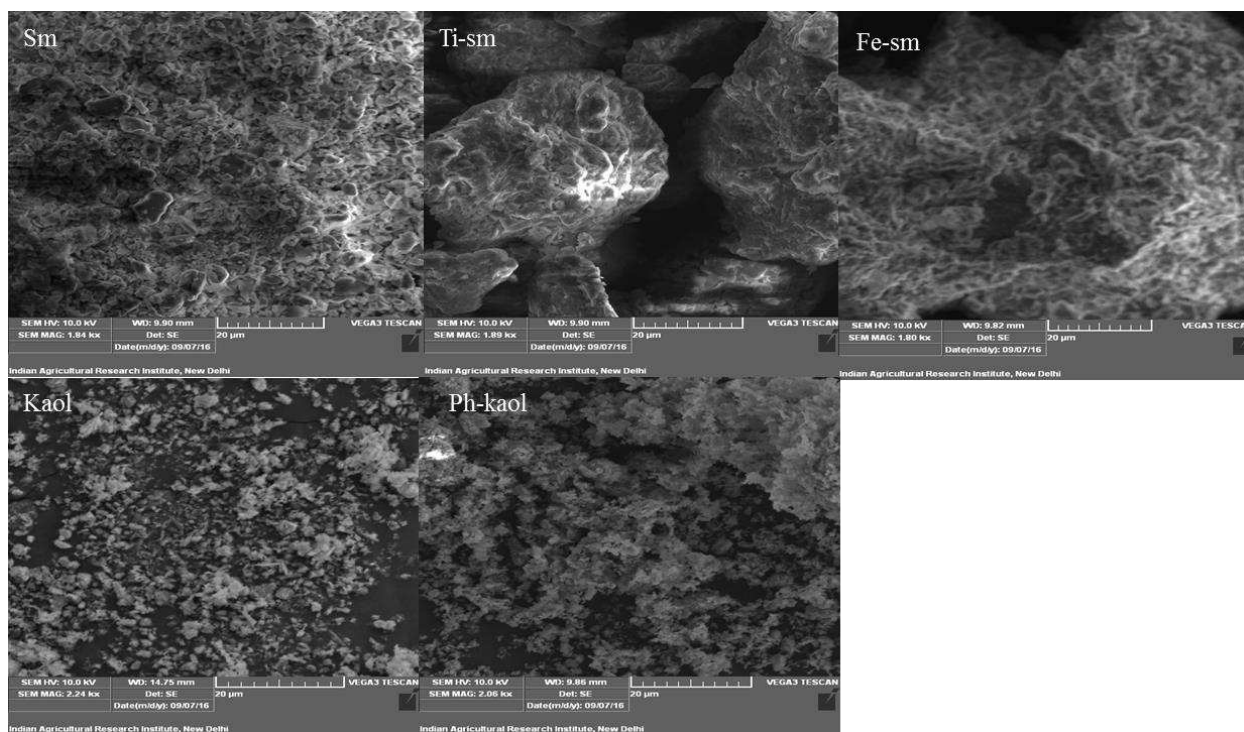


Fig. 3 SEM images of unmodified smectite (Sm), Ti-pillared smectite (Ti-sm), Fe-exchanged smectite (Fe-sm), unmodified kaolinite (Kaol) and phosphate-bound kaolinite (Ph-kaol)

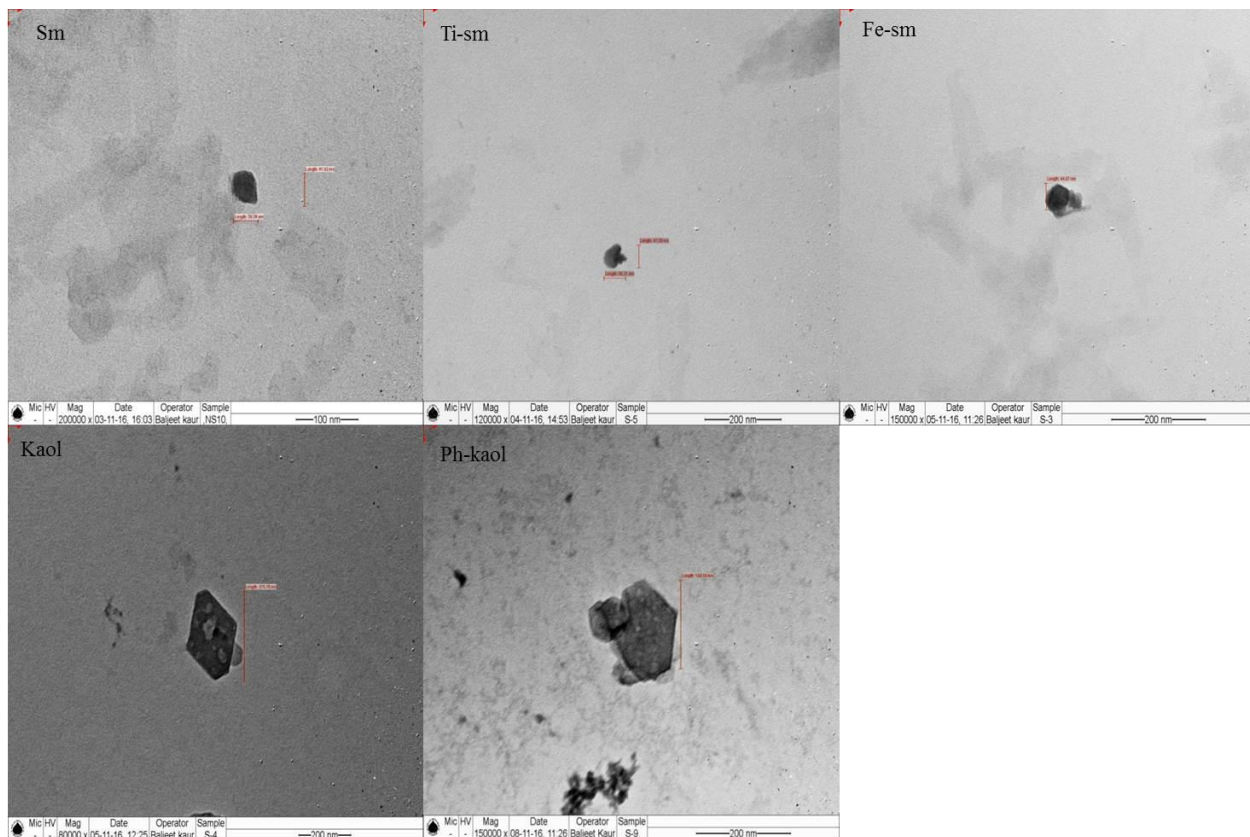


Fig. 4 TEM images of unmodified smectite (Sm), Ti-pillared smectite (Ti-sm), Fe-exchanged smectite (Fe-sm), unmodified kaolinite (Kaol) and phosphate-bound kaolinite (Ph-kaol)

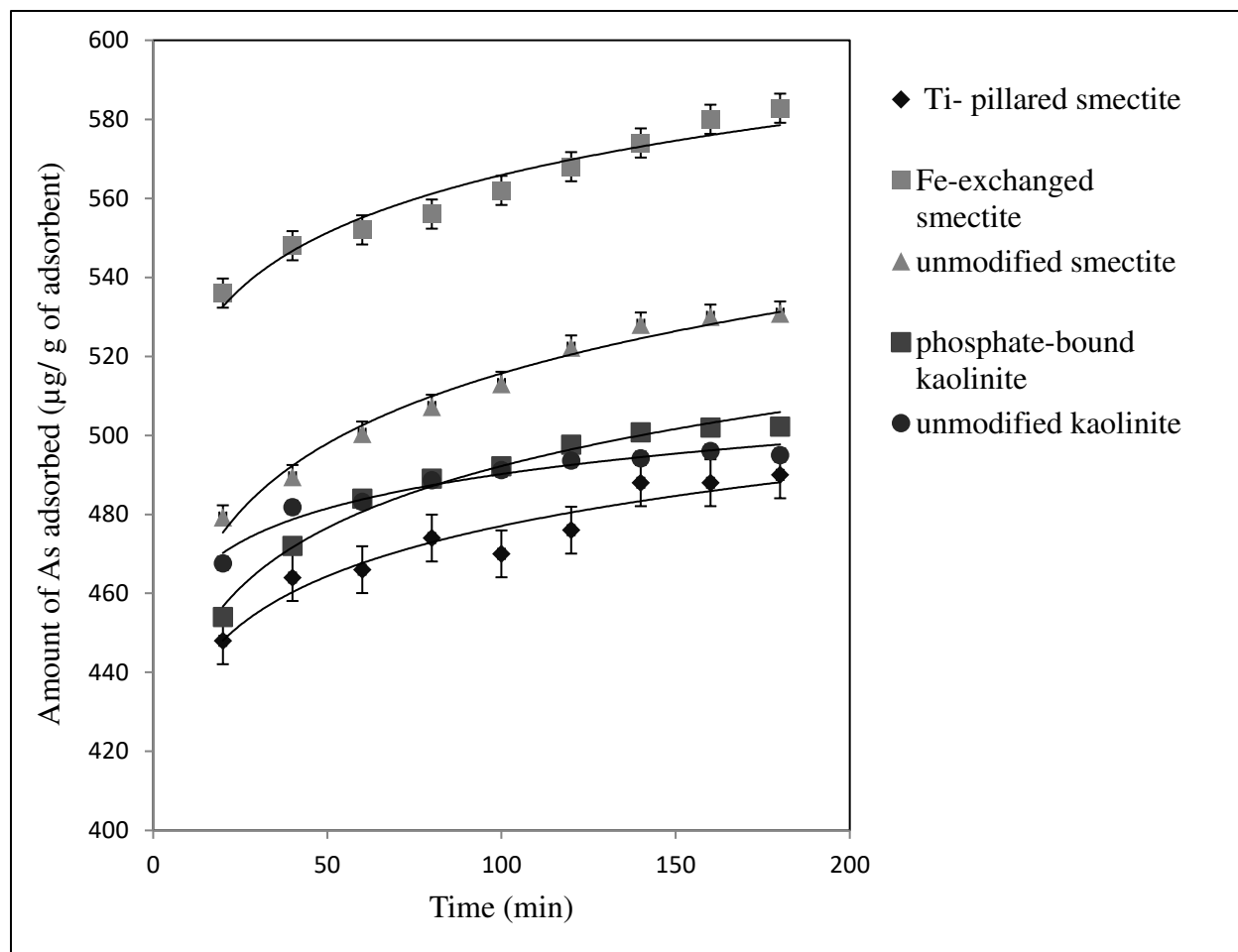


Fig. 5 Kinetics of arsenic adsorption onto unmodified, Ti-pillared, Fe-exchanged smectites, and unmodified and phosphate-bound kaolinities in aqueous system (the fitting curves represent the power function equation).

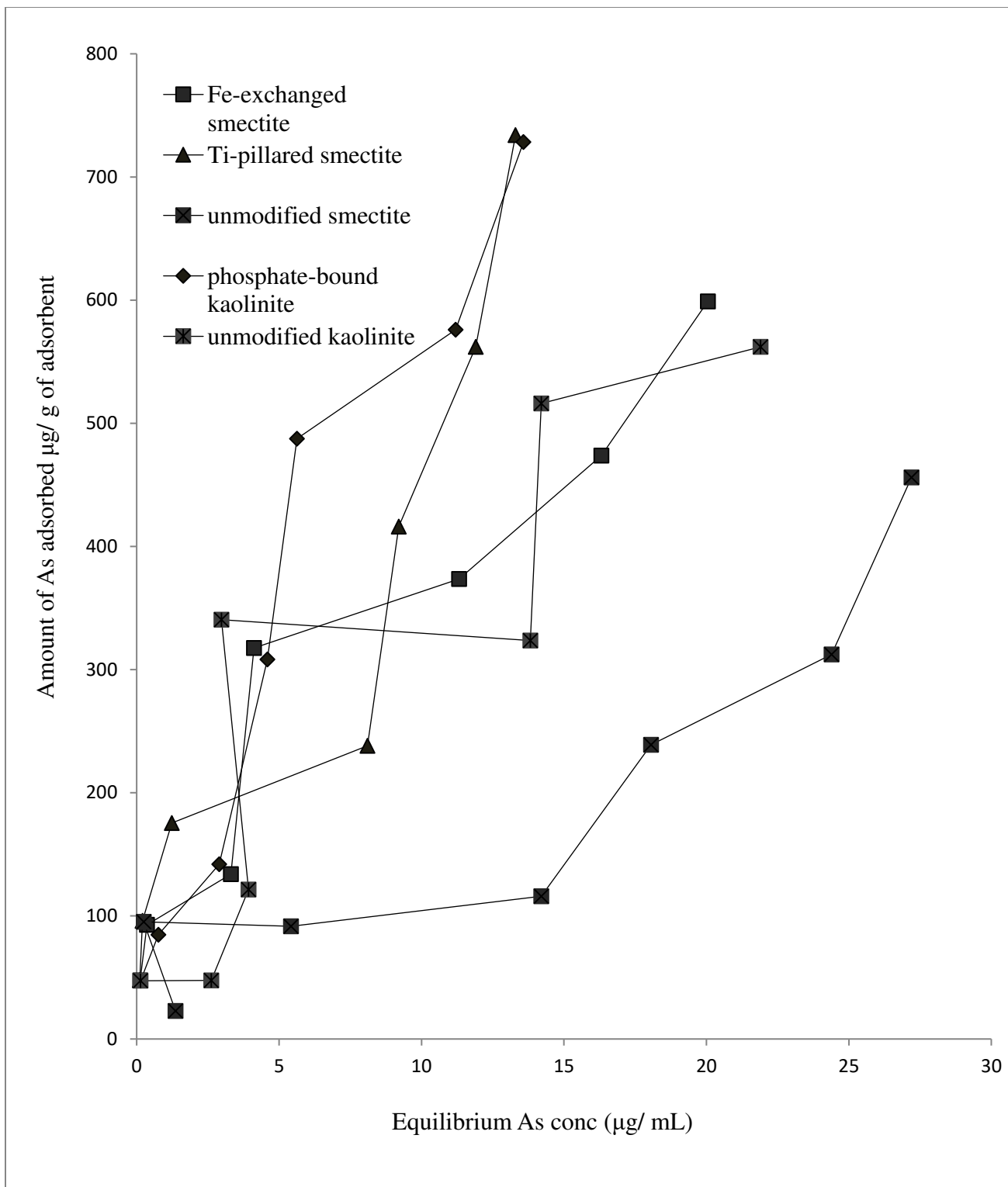


Fig. 6 Arsenic adsorption isotherm onto unmodified, Ti-pillared and Fe-exchanged smectites, and unmodified and phosphate-bound kaolinites in aqueous system (the curves were best fitted in Freundlich adsorption isotherm).

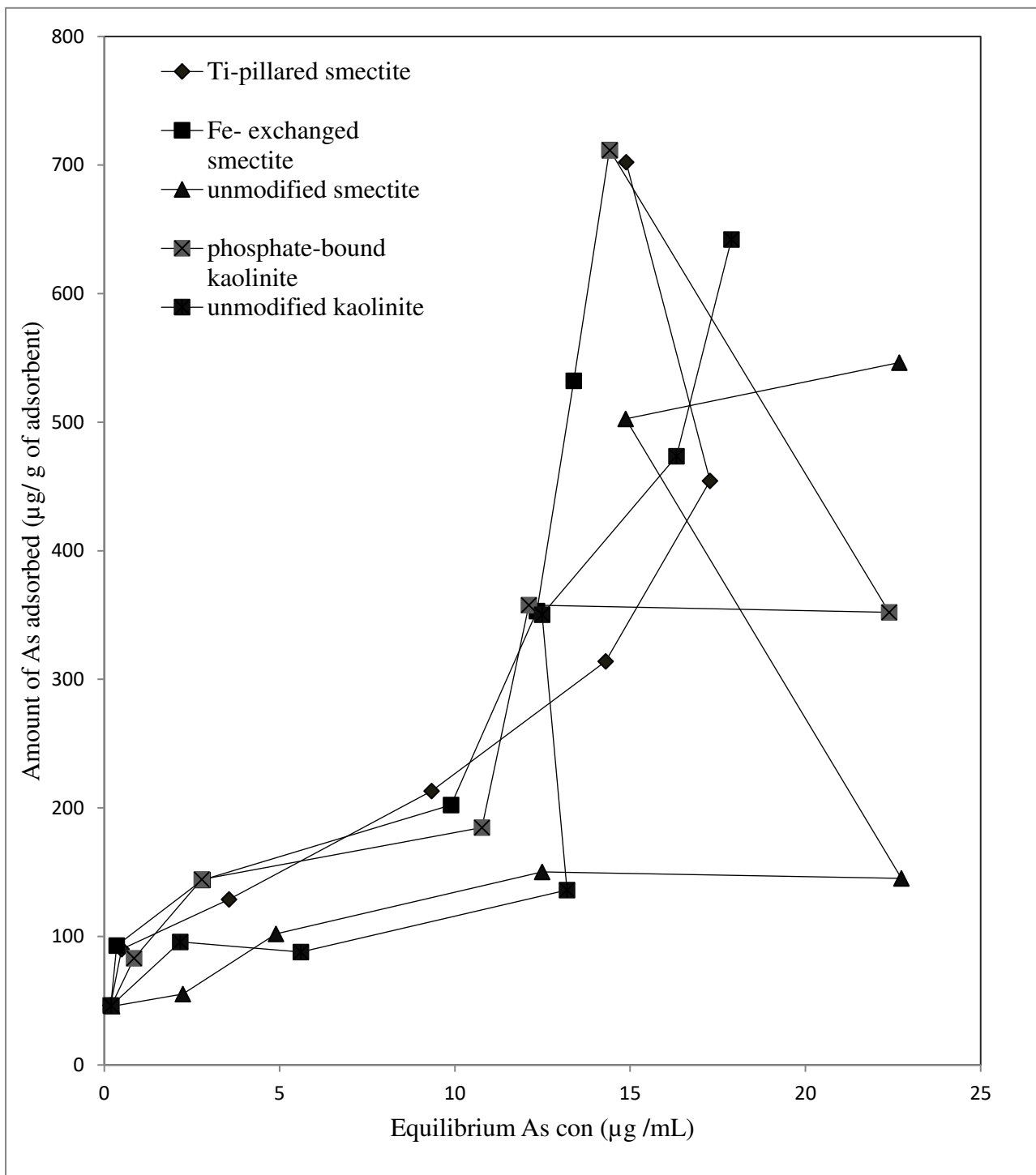


Fig. 7 Arsenic adsorption isotherm onto unmodified, Ti-pillared and Fe-exchanged smectites, and unmodified and phosphate-bound kaolinites in soil system (the curves were best fitted in Freundlich adsorption isotherm).

