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Seco, I., Schellart, A.N., Gomez-Velentin, M. et al. (1 more author) (2018) Prediction of organic combined sewer sediment release and transport. *Journal of Hydraulic Engineering*, 144 (3). ISSN 0733-9429

[https://doi.org/10.1061/\(ASCE\)HY.1943-7900.0001422](https://doi.org/10.1061/(ASCE)HY.1943-7900.0001422)

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Prediction of organic combined sewer sediment release and transport

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1 ABSTRACT

2 Accurate predictions of sediment loads released by sewer overflow discharges are important for being
3 able to provide protection to vulnerable receiving waters. These predictions are sensitive to the
4 estimated sediment characteristics and on the site-conditions of in-pipe deposit formation. Their
5 application without a detailed analysis and understanding of the “initial conditions” under which in-
6 sewer deposits were formed normally results in very poor estimations. In this study, in-sewer sediment
7 samples deposited during dry-periods in a combined sewer system were collected, and their properties
8 assessed. Parameters in the sediment transport relationship first proposed by Skipworth for in-pipe
9 deposits were estimated based on simulating the in-pipe deposit formation conditions in laboratory
10 erosion tests. The measured parameters were then used to simulate sediment transport through a small
11 combined sewer network for a number of rain events for which rainfall, hydraulic and water quality
12 data were available. Results showed that the model of Skipworth can provide good predictions of the
13 sediment loads released from such in-sewer deposits. The experimentally-derived calibration
14 parameters used with Skipworth’s model allowed for a realistic simulation of the in-sewer sediment
15 behaviour and so can be used to accurately estimate the sediment load released from combined sewer
16 systems during rainfall events.

17 Keywords

18 In-sewer erosion; quality modelling; organic sediments; in-sewer sediment transport

19 **Nomenclature**

20 The following symbols are used in this paper:

A_s [m ²]	Sediment surface exposed to erosion	M [g s ⁻¹ m ⁻²], b [-]	Calibrated transport parameters
C_{SS} [g l ⁻¹]	Suspended solids concentration	V_s [m ³]	Sediment volume
d [mm]	Cumulative depth of erosion	ρ_s [kg m ⁻³]	Sediment bulk density
d_e [mm]	Sediment eroded depth per time step	ρ_m [kg m ⁻³]	Sediment-water mixture density
d' [mm]	Thickness of the upper sediment layer of the deposit	τ_b [N m ⁻²]	Applied bed shear stress
d'' [mm]	Thickness of the surficial layer eroded during consolidation period	τ_c [N m ⁻²]	Critical shear stress
d_{50} [mm]	Characteristic particle size	τ_{cs} [N m ⁻²]	Critical surficial shear stress
E [kg m ⁻² s ⁻¹]	Erosion rate	τ_{cu} [N m ⁻²]	Critical shear stress of the underlying layer

21

22 **INTRODUCTION**

23 Existing software packages for the hydraulic modelling of sewer network systems generally show
 24 good predictive performance. However, the simulation of water quality processes in sewer system
 25 network models has been less reliable (e.g. Ashley *et al.*, 1999; Kanso *et al.*, 2005) and sewer flow
 26 water quality data are generally less available (e.g. Willems 2010).

27 Water quality modelling in combined sewer systems predicts sediment and pollutant loads for time
 28 varying flows. Research has shown that a significant contribution of suspended sediment originates
 29 from the release and re-suspension of sediment from in-sewer deposits during the initial period of
 30 storms (Ahyerre and Chebbo 2002; Ashley *et al.* 2004; Gromaire-Mertz *et al.* 2001; Saul and
 31 Thornton 1989; Tait *et al.* 2003a). The rapid suspension of previously deposited in-pipe sediment has
 32 been observed in releases from combined sewer overflows during intense rainfall events. This
 33 phenomenon has been termed a *first foul flush* (Gupta and Saul 1996) The *first-flush* phenomenon
 34 (Obermann *et al.* 2009) is often observed in regions with a semi-arid climate, such as in Mediterranean
 35 catchments which are characterized by dry-weather periods followed by intense storm events. The
 36 high variability of the flow regime of the rivers in these regions are also strongly dependent on the
 37 seasonal rainfall, this can result in a quite limited dilution capacity of the natural receiving waters (Prat

38 and Munné 2000) thus, in areas of water scarcity, *first flush* can cause a very significant impact. In the
39 Mediterranean region where the case study catchment is based, it is therefore important to achieve
40 reliable predictions of sediment and pollutants loads that can reach the receiving waters through
41 combined sewer overflows (CSOs) during intense rainfall events. An improved prediction of sediment
42 loads could allow for action to better manage pollutants that are released and are known to generate
43 high oxygen demand in receiving waters. Most sediment transport research has been focussed on
44 sediment movement in rivers. The findings resulted in predictive relationships, empirically calibrated,
45 and developed from observations of the movement of mainly granular sediments. The application of
46 existing granular based fluvial transport models, such as Ackers (1984, 1991) and May (1993),
47 modified to simulate erosion and transport of granular and organic sediments through piped sewer
48 systems does not perform well (Ashley *et al.* 2004; McIlhatton *et al.* 2005; Schellart *et al.* 2008b; De
49 Sutter *et al.* 2003). Considering the additional processes that can occur in sewer sediment deposits, the
50 use of sediment transport relationships originally developed for fluvial environments and granular
51 sediment can be reasonably questioned.

52 Biochemical transformation processes, interactions between particles, and microbiological activity can
53 have a significant influence on the resistance to erosion of in-pipe deposits (Banasiak and Tait 2008;
54 McIlhatton *et al.* 2005; Sakrabani *et al.* 2005; Seco *et al.* 2014b; Vollertsen and Hvitved-Jacobsen
55 2000). The available sediment transport relationships for cohesive deposits oversimplify the process
56 occurring in sewers (Freni *et al.* 2008; Mannina *et al.* 2012; Schellart *et al.* 2010).

57 The deposit erosion methodology developed by Skipworth *et al.* (1999) links the sediment erosion rate
58 to critical shear stress levels related to different layers within the sediment deposit in pipes. The
59 methodology is derived from laboratory observations obtained from the erosion and transport of
60 cohesive-like synthetic sediment previously deposited in a pipe and subjected to steady flow
61 conditions.

62 Results obtained by Skipworth *et al.* (1999) and later verified by Rushford *et al.* (2003), confirm that
63 their methodology improves prediction of the transport rate of cohesive sediment. The potential for
64 improvements in the prediction of sediment erosion rates when using Skipworth's model, can only be

65 attained if realistic values for the calibration parameters of the deposit erosion model can be obtained.
66 In this study field data is used to test this type of deposit erosion to assess its utility for modelling
67 sediment releases from sewer system overflows during intense rainfall events.

68 The determination of shear stress at the threshold of motion (τ_c) exerted on the sediment bed surface is
69 crucial in the evaluation of the release of sediments from layered deposits, however, this threshold is
70 difficult to determine in-situ. McIhatton *et al.* (2005) and Oms *et al.* (2008) reported observed values
71 of τ_c in the range between 0.15 and 0.85 N/m² for in-sewer sediment deposits in combined sewer
72 systems in Dundee (Scotland) and in Paris (France) respectively.

73 Highly-organic sediment deposits can be observed in combined sewer systems serving highly
74 urbanized areas found in the Mediterranean region where high levels of catchment imperviousness are
75 common. Additionally, large fluctuations in combined sewer flows are associated with semiarid
76 climates and this pattern of variation can have an effect on the sediment accumulation-flushing cycles
77 found in sewer networks. The main aim of this paper was to examine the suspended sediment load
78 evolution that can be discharged into natural watercourses from CSOs activated during intense rain
79 events. The accurate estimation of the sediment discharge pattern will help in quantifying the impact
80 of CSOs on receiving waters. With this aim, the study had the following objectives: to evaluate the
81 process of mobilization from in-sewer sediment deposits, and to validate Skipworth's deposit
82 relationship in a particular catchment under realistic rainfall conditions.

83 To achieve these objectives the empirical deposit and transport parameters were estimated based on
84 laboratory observations. The performed tests allowed the analysis of the erosion behaviour of highly-
85 organic sediment sampled from a real sewer network. Previous investigations on the erodibility of
86 highly organic sediment (Seco *et al.* 2014a) provided key knowledge on the properties of sediment
87 recovered from the same combined sewer system. The experimental and analytical procedures were
88 modified based on the results obtained in the earlier study. Controlled environmental temperature
89 conditions were now established. An intermediate Dry-Weather Period (DWP) between the formerly
90 established 16 and 64 hours was also implemented to obtain a deeper comprehension on the process
91 that influence erosion rate evolution. The results obtained from the laboratory experiments reported in

92 this work allow for the assessment of the calibration parameters involved in the deposit-erosion model
93 proposed by Skipworth *et al.* (1999).The use of real sewer sediments for the determination of the
94 transport parameters allowed for the verification of the application of the Skipworth in-pipe deposit
95 model at a network scale.

96

97 **METHODS**

98 **Study site location and description**

99 The field study site is situated in the south-east of Spain, in the city of Granollers (35 km north of
100 Barcelona, Spain). The local rainfall pattern is irregularly distributed throughout the year and
101 characterised by dry-weather periods often longer than a week followed by single storm events. A
102 small urban catchment in Granollers was selected for the study, covering an area of approximately 10
103 hectares (Fig. 1). The land use is mainly residential and commercial, with a high population density of
104 150 inh/ha. The area has a significant presence of commercial food activity. The catchment surface
105 displays a high degree of imperviousness that reaches almost 100% in some zones, with an average
106 imperviousness of 84% over the whole catchment. Given the highly impervious conditions of the
107 catchment, and the limited existence of soil areas, inorganic sediments are a minor contribution during
108 storm runoff (Gómez-Valentín *et al.* 2015).

109 The urban area has a gravity driven combined sewer system composed of circular concrete pipes with
110 diameters ranging from 300 to 1000 mm. General characteristics of the catchment and the combined
111 sewer network are given in Table 1.

112 **Hydrological, hydraulic and water quality monitoring**

113 Flow rates, water quality data and rainfall data were collected during storm events. The purpose of the
114 monitoring programme was to obtain field data to validate the reported modelling work. The layout
115 and the operation of the case study sewer network is similar to that of many other combined systems
116 throughout Europe and the eastern coast of the USA. The results of the study are therefore expected to

117 be widely applicable. The monitoring programme was carried out over an 18-month period. The
118 events of interest were selected based on two threshold conditions: a rainfall depth which will produce
119 enough runoff to increase water depths and velocities in the sewer network and also have sufficient
120 flow to produce a measurable resuspension of sediments previously deposited inside the network, and
121 an antecedent DWP sufficient to produce enough sediment accumulation for the detection of
122 increasing pollutant loads at the outlet of the analysed catchment. Precipitation depth of 5mm and
123 antecedent DWP of the order of several days were established as thresholds. Events that experienced
124 major disruptions during flow recording or water quality sampling were discarded. After pre-
125 processing, four rainfall events satisfying these conditions remained; see events 1 to 4 in Table 2. For
126 these events, physical samples for water quality analysis were collected at the outlet of the catchment
127 simultaneously with rainfall data and flow data. Two additional events where no satisfactory water
128 quality data were recorded (events 5 and 6 in Table 2) were used to calibrate the network
129 hydrodynamic model.

130 Flow was continuously monitored using an automatic portable flowmeter (HACH-Lange, Sigma 950
131 model). The instrument was provided with a bubbler water level sensor and a doppler velocity sensor,
132 and the flowrate was then calculated. The water samples were collected during rainfall with an
133 automatic sampler (HACH-Lange Sigma SD900 model). The sampler was equipped with a peristaltic
134 high speed pump taking in 1000 ml in 2 minutes through a tube with a strainer at the end, followed by
135 a cleaning cycle that takes another 2 minutes. An increase in flowrate compared to the dry weather
136 flow pattern triggered the collection of water quality samples. Due to the high imperviousness of the
137 catchment, it was expected that the runoff rapidly releases and washes off sediments from the surface
138 and erodes them from inside the network. The highest sampling frequency was therefore set at 5
139 minutes for the first 15 minutes of a rainfall event and then less frequently for a total of 2 hours.
140 Following the trigger at $t=0$, samples were taken at: 0, 5, 10, 15, 30, 45, 60, 90, 120 minutes. The
141 established sampling frequency was intended to focus on the beginning of a storm event in order to
142 analyse the occurrence of a first flush pollutant phenomenon.

143

144 **Deposited sediment characteristics and behaviour**

145 *Sediment deposit sampling and analysis*

146 A batch of 3 kg of in-sewer sediment was manually collected, directly from the invert of a 600 mm
147 pipe with 0.002 m/m slope upstream of a diameter reduction (from 600m to 400 mm). According to
148 the local operators, sediments deposit formations were typically observed in this section after
149 prolonged dry-periods. The collection was conducted during dry-weather flows when the water depth
150 was less than 5 cm. The deposited sediments were collected manually, immediately refrigerated at 4
151 °C and then transported within 48 hours to Sheffield in UK, where the analysis and erosion tests were
152 performed. Upon arrival in Sheffield the sediment temperature was found to be 4.7 °C, after which the
153 sediments were immediately stored in a refrigerator at 4 °C. Despite the destruction in the layer
154 structure of the deposit during collection, no alterations were believed to have taken place in the
155 physical characteristics of the sediments, while biological activity and microbiological decomposition
156 of the sediment samples were inhibited by the low temperatures during the storage and transport
157 procedures. Thus, for physical characterization the collected sediment were considered representative
158 of the deposit formed in the invert of the original sewer pipe during dry-weather periods.

159 Analysis and sediment preservation follows the *Standard Methods for the Examination of Water and*
160 *Wastewater* (APHA *et al.* 2005). A summary of the sediment characteristics is shown in Table 3.

161 The sediments were characterized for organic content, which is defined as the proportion between the
162 volatile solids (VS) and the total dry mass of sediments (TS) (section 2540E, *Standard Method*). An
163 average of 95 % ± 2 of VS/TS rate was obtained. The density of the deposit was assessed using the
164 displacement principle method. The presence of fat, oil and greases was established through visual
165 observation of the sediment. The characteristic particle diameter d_{50} was obtained following the British
166 Standards (BS 1796-1:1989. Test sieving) for the gross part (>1 mm), while the fine part (< 1 mm) was
167 performed by laser diffraction method (ISO 13320:2009 Particle size analysis. Laser diffraction
168 methods) using a Mastersizer 2000, Malvern instrument Ltd. Figure 2 shows the particle size
169 distribution curve of the collected sediment samples.

170 *Laboratory erosion test procedure*

171 The laboratory tests were carried out with a sample of sewer sediment deposit, placed in a device
172 called an erosionmeter (developed by Liem *et al.* 1997). The erosionmeter consists of a vertical
173 perspex tube provided with a centrally located propeller, and vertical vanes to reduce lateral
174 circulation, and a container for the sediment deposit. By applying an angular velocity to the water
175 column a reasonably uniform shear stress is exerted over the sediment surface. Six vertically spaced
176 outlets are used to sample the sediment eroded from the bed that remained suspended in the water
177 column. The samples were analysed later for TSS following the *Standard Methods for the Examination*
178 *of Water and Wastewater* (2005). A detailed description of the equipment and calibration process is
179 given in Seco *et al.* (2014a).

180 The preparation of the samples follows a defined procedure with the intention of establishing
181 repeatable conditions and to simulate the dry weather flow conditions found in the case study sewer.
182 The whole batch of collected disturbed sediment deposit was thoroughly mixed and separated into
183 individual samples. The container with the individual sediment sample was then carefully filled with
184 water and left for 72 hours at 4 °C, in a phase of quiescent physical consolidation where the biological
185 reactions were retarded by the low temperature. After the pre-consolidation phase the sample was
186 placed in the bottom of the erosionmeter and allowed to assimilate to 20°C. Aerobic conditions were
187 set by supplying air to the supernatant water. An oxygenated environment in a gravity sewer network
188 is likely to be produced under conditions of varying flows (Hvitved-Jacobsen *et al.* 2013). A low bed
189 shear stress (0.15 N/m²), similar to that found during dry weather flows in the system, was applied
190 over the bed by slowly rotating the propeller. By applying a low bed shear stress it was intended to
191 simulate the dynamic consolidation conditions at which sediment deposits were subjected in sewers
192 during periods of sediment deposition between rain events (DWP). Additionally, the low velocity of
193 the propeller ensures a continuous mixing and creates a uniform environment regarding water
194 temperature and dissolved oxygen (DO) levels. The results from this study focus therefore on the
195 erosion and transport of sediments subjected to aerobic conditions at 20°C during the depositional
196 DWP prior to a storm, and the tests were carried out in a temperature controlled room. Four different

197 DWP durations between 16 and 64 hours were considered to simulate the consolidation process
198 thought to be present in the actual sewer system. The DWP durations were in the order of magnitude
199 of several days for two reasons: firstly, although there are longer DWPs in the catchment, the average
200 DWP throughout the 18 month field monitoring period was 3 days, secondly, as described in Seco *et*
201 *al.* (2014b), the sediments were quite biologically active and it was assumed that during DWP the
202 upper sediment layers are continuously being biodegraded as well as replenished with fresh sediments
203 originating from the dry weather flow. The critical threshold of motion at the solid-fluid interface of
204 the resulting deposit was then assessed by step-wise increase of the propeller speed. The erosion phase
205 of the tests was then performed by increasing the applied shear stress in a stepwise fashion. Samples
206 were collected from the water column at steady erosion state conditions (Parchure and Mehta 1985) at
207 each step of applied bed shear stress, which lasts 45 minutes (Schellart *et al.* 2005; Tait *et al.* 2003b).
208 The eroded material and resultant erosion rate was calculated from the measured suspended sediment
209 (SS) concentration of the collected samples. These data are reported below and were used in the
210 calibration of the erosion model described below.

211

212 **Modelling sediment transport in a field study catchment**

213 *Hydrodynamic modelling*

214 The *SWMM5* (Storm Water Management Model) software package was selected for the rainfall-runoff
215 and hydrodynamic modelling through the combined sewer system in the study case. The hydrological
216 model (Fig. 1) is defined based on a sub-catchment delineation established from topographic data of
217 the catchment drainage areas and of the combined sewer network complemented by in-situ
218 observations to complete information about impervious-pervious surfaces and their drainage
219 characteristics. The hydrodynamic network model is directly related to the sewer network system
220 information provided by the local sewerage company; it comprises 57 pipes and manholes, and 42
221 sub-catchments in a 10 hectare area. Flow measurements were performed at the outlet of the studied
222 catchment, using the equipment and procedures described above.

223 A calibration and validation process of the hydrodynamic model was performed by comparing
224 simulated with measured flow rates during several rainfall events. Model calibration was carried out
225 using rainfall events 5 and 6 (Table 2). Subsequently, the model was validated by applying
226 independent data sets corresponding to events 2 and 3. The relative errors of total runoff volume range
227 from 1 % to 10 % for the analysed events, which are indicated in Table 4. The relative error of peak
228 flow is between 2 % and 10 % and the difference in the elapsed time to reach the peak flow range
229 from 2 to 8 minutes. The goodness of fit obtained can be observed in Fig. 3 and Table 4.

230

231 *Sediment erosion model of Skipworth et al. (1999)*

232 The methodology proposed by Skipworth et al. (1999) is based on the concept of a bed structure with
233 different layers, in which each layer displays a different resistance to erosion.

234 The simulation method proposed by Skipworth *et al.* (1999) is based on an excess shear stress
235 relationship to predict the sediment erosion rate for estuarine deposits so-called Ariathurai-
236 Partheniades equation (1) (Ariathurai 1974, as referenced by McAnally and Mehta 2000)

$$E = M \cdot \left(\frac{\tau_b - \tau_c}{\tau_c} \right) \quad (1)$$

237 where E is the erosion rate in $\text{kg/m}^2/\text{s}$ for the applied bed shear stress τ_b [N/m^2] and τ_c [N/m^2] is the
238 critical shear stress, M is a transport parameter used as a calibration factor that has the same units as E
239 and is equal to the erosion rate when $\tau_b = 2 \cdot \tau_c$.

240 By examining the erosion rate over time, Skipworth concluded that in-pipe deposits showed a weaker
241 upper layer transitioning to a stronger underlying layer. It was later observed, also verified by
242 Schellart *et al.* (2005) and Seco *et al.* (2014a), that the organic content, oxygen availability and length
243 of the consolidation period have an influence on the subsequent erosion resistance of the deposited
244 layers. The sketch in Fig. 4 shows the variation of the erosional resistance with depth for cohesive-like
245 sediment deposits. At the upper layer, the erosional strength increases in depth from a surface
246 erosional strength (τ_{cs}) until a value of deposit strength (τ_{cu}). Once the thickness of the upper layer (d')
247 is exceeded and the lower layer is reached, the deposit has an almost uniform resistance to erosion.

248 Skipworth *et al.* (1999) proposed a power law shown in equation (2), that represents the depth
249 variation of the shear stress necessary to erode the upper weak layer.

$$\tau_c = \begin{cases} \left[\left(\frac{d}{d'} \right)^{1/b} \cdot (\tau_{cu} - \tau_{cs}) \right] + \tau_{cs} & \text{for } 0 \leq d \leq d' \\ \tau_c = \tau_{cu} & \text{for } d > d' \end{cases} \quad (2)$$

250 Where d is the cumulative depth of erosion, d' represents the thickness of the upper layer (Fig. 4), b is
251 a calibration parameter which describes the rate of change in bed strength with depth. The factor M is
252 also a model calibration parameter. Due to the high dependency on the sediment bed properties, the
253 values of M , b , d' , τ_{cs} and τ_{cu} must be empirically determined to obtain a realistic prediction of
254 sediment erosion and transport.

255 *Coupling of a sediment transport model and SWMM5*

256 In order to analyse the performance of this model for predicting sediment release in a combined sewer
257 network under time-varying hydraulic conditions, the erosion relationship of Skipworth was coded
258 using MATLAB and then coupled with a sediment transport network model also coded in MATLAB.
259 This code was based on the concept of a model previously used by Schellart *et al.* (2008a), which
260 simulates the transport of sediment eroded from in-pipe deposits, based on hydraulic parameters
261 simulated by an uncoupled hydrodynamic sewer network model, and assuming conservation of
262 sediment mass between sediment advection, released sediment and the sediment stored in the in-pipe
263 deposits. Predictions from the calibrated SWMM5 hydraulic model were used as inputs for the
264 sediment erosion and transport model coded in MATLAB. The linked modelling structure is shown in
265 Fig. 5.

266 *Performance evaluation criteria*

267 The goodness of fit between observed and simulated suspended sediment (SS) concentration values
268 was evaluated by using the following criteria: the sum of squared errors *SSE* (Eq. (3)); the percent
269 peak error *PE* (Eq. (4)); and the Nash-Sutcliffe efficiency (NSE) (Eq. 5) where $C_{SS,m,i}$, $C_{SS,s,i}$ are the SS
270 concentration measure and simulated at time i respectively, and $C_{SS,peak}$ is the concentration peak,
271 defined as the maximum SS concentration value of the event.

272 NSE values range between 1 for a perfect fit and $-\infty$.

$$SSE = \sum_{i=1}^n (C_{SS,m,i} - C_{SS,s,i})^2 \quad (3)$$

$$PE = \frac{(C_{SS,m,peak} - C_{SS,s,peak})}{C_{SS,m,peak}} \cdot 100 \quad (4)$$

$$NSE = 1 - \frac{\sum_{i=1}^n (C_{SS,m,i} - C_{SS,s,i})^2}{\sum_{i=1}^n (C_{SS,m,i} - \overline{C_{SS,m,i}})^2} = 1 - \frac{SSE}{\sum_{i=1}^n (C_{SS,m,i} - \overline{C_{SS,m,i}})^2} \quad (5)$$

273

274 **RESULTS AND DISCUSSION**

275 **Assessment and optimization of transport parameters based on laboratory results**

276 The values of the calibration parameters of the equation proposed by Skipworth (Eq. 1 and 2) can be
277 derived from analysis of the data obtained from laboratory erosion tests.

278 The determination of the erosional strength with depth is derived from each time step application of
279 increased shear stress linked with the stable SS concentration measured ($C_{SS,m}$) at the end of each time
280 step. The relationship between applied shear stress and erosion rate is shown in Fig. 6, for tests carried
281 out under aerobic conditions and for different durations of antecedent dry-weather period. The errors
282 in the determination of the applied shear stress ($\pm 0.07 \text{ N/m}^2$) derived from the erosionmeter
283 calibration process were also represented (refer to Seco *et al.* (2014a) for more detail). Through a
284 regression analysis a series of best fit trend functions were obtained (Fig. 6).

285 *Assessment of parameters τ_{cs} , τ_{cu} , d'' and d'*

286 At the end of each time step during the erosion test, the mass of sediment obtained from the SS sample
287 concentration can be translated to a sediment erosion depth (d_e), and so it is possible to link the deposit
288 properties to the applied shear stress (τ_b). The bulk density of the bed formed by collected sewer
289 organic-cohesive sediment is of 1310 kg/m^3 ($\pm 146 \text{ kg/m}^3$). Sediment bed density was assumed to
290 remain constant during the test since the duration of the erosion test is relatively short compared to any

291 consolidation processes that can produce significant changes in density of the deposit structure due to
292 excess pore water effects.

293 The applied shear stress against the depth of erosion is shown in Fig. 7.

294 During the antecedent DWP simulated in the tests, the erosion meter was set to exert $\tau_{DWP}=0.15 \text{ N/m}^2$
295 on the sediment bed. This τ_{DWP} value was estimated by examination of the bed shear stress value at the
296 outlet pipe predicted during DWF in the case study network.

297 It was noticed that during all DWP tested, a near constant and thin surficial layer was eroded at the end
298 of the consolidation period. The depth of this eroded layer can be assessed from the sample of the
299 sediment concentration at the end of DWP (Eq. (6)), which allows establishment of the value of a
300 parameter d'' as the observed value 1.25 mm (standard deviation SD = 0.13 mm). There were no
301 significant changes observed in the depth of the eroded layer with different DWP durations. Hence, it
302 is assumed that the value of the critical shear stress at the surface layer τ_{cs} can be considered equal to
303 the applied shear stress during the antecedent DWP (0.15 N/m^2). This means that the τ_{cs} and d'' can be
304 considered independent of the length of the DWP when consolidation of the sediment deposit takes
305 place.

$$d_e = \left(C_{SS} \cdot \frac{V_s}{A_s} \right) \cdot \frac{1}{\rho_s} \quad (6)$$

306 Following the profile of sediment resistance against erosion shown in Fig. 4, the value of τ_{cu} would be
307 obtained when the resistance strength becomes uniform with depth. The experimental tests, however,
308 did not achieve a completely uniform resistance against erosion. Therefore, the thickness of the upper
309 layer of sediments (d') is estimated by assuming that a gradient of 0.03 ($\Delta\tau_b/\Delta d$) practically marks the
310 transition between the upper layer (d') and the lower more uniform layer. Fig. 8 (a) shows the values
311 of d' and τ_{cu} estimated from the erosion tests performed after different consolidation periods, a dot
312 marks the estimated transition point below which the τ_{cu} is assumed to be sensibly constant. In Fig. 8a,
313 the errors in the assessment of the sediment depth of erosion ($\pm 6 \text{ mm}$) and the accuracy of the applied
314 shear stress ($\pm 0.07 \text{ N/m}^2$ after, Seco *et al.* 2014a) are indicated by shaded error bands. From this plot it
315 can be observed that after 24 hours of consolidation, the increase in the resistance against erosion of
316 the sediment bed is not significant.

317 *Determination of the values adopted by the model parameters b and M*

318 In order to apply Eq. 1 and Eq. 2, the values of the parameters M and b need to be determined. An
319 optimization for calibration parameters b and M is therefore performed by comparing the calculated
320 erosion rate E_c against measured erosion rate E_m , given the applied shear stress τ_b . This optimization
321 was carried out by varying both parameters at the same time, in order to obtain a minimum value for
322 the root mean square error $RMSE$ (Eq. (7)).

$$RMSE = \sqrt{(E_c - E_m)^2} \quad (7)$$

323 The ranges in which the values of the parameters b and M were varied during the optimization were
324 initially assumed to be those determined by Skipworth and Rushforth and presented in Table 5.
325 However, this did not lead to a minimum, hence the range of variation for the b -parameter was
326 increased to 0.025 and 1 (with increments of 0.025), and for the M -parameter varying from 0.05 and 2
327 (with increments of 0.05).

328 The optimization results produced a narrow range of values for b (Fig. 9 a) where the mean value
329 obtained is $b = 0.125$ (SD = 0.071). Regarding the value of the parameter M , the variation is wider
330 (Fig. 9 b). However, a relationship between the value adopted by the M -parameter and the applied
331 shear stress for each test could be observed, and this trend changes with the length of the DWP
332 analysed. Thus, it can be suggested that a weak relation exists between the duration of the
333 consolidation period and the parameter M (coefficient of proportionality between 0.51 and 0.74). The
334 optimised values for b and new ranges found for M and the other parameters involved in the
335 calculation or erosion rate are included in Table 5.

336 Fig. 8 indicates that after 24 hours of consolidation the resistance against erosion throughout the depth
337 of the deposit stabilized. Based on that finding, the values of the sediment transport parameters b and
338 M that were used for the network sediment transport model were those average values obtained in the
339 tests with DWP longer than 24 hours. A linear relationship (Eq. (8)) was implemented for the
340 evaluation of the M -parameter for each applied shear stress (τ_b) during the simulations, valid for values
341 of τ_b higher than 0.40 N/m². For lower values of τ_b the value of M was constant and equal to 0.05.

$$M = 0.725 \cdot \tau_b - 0.0487 \quad ; \quad \tau_b > 0.40 \text{ N/m}^2 \quad (8)$$

342 **Modelling Sediment Transport in the Case Study Catchment**

343 Hydrodynamic predictions were obtained from the calibrated SWMM5 model for the four rainfall
344 events 1 to 4 from Table 2. These predictions were input into the sediment transport model using
345 Skipworth's erosion relationship calibrated with the case study sediment. Initial conditions for the
346 available in-pipe sediment deposits were set to a 5 cm deep sediment deposit, as this allowed for
347 analysis of sediment transport not to be limited by the availability of sediment in the simulations (i.e.
348 after all the simulations there was still sediment left in each pipe). This ensured that the initial model
349 boundary conditions did not impact on the model predictions. A selection of computation time-steps
350 were examined and were seen to influence the simulated erosion rate. A time step higher than 1 minute
351 started to reduce the peak values of sediment concentration; hence a time-step of 20 seconds was used.

352 In this study, based on previous research (Ahyerre and Chebbo 2002; Gromaire-Mertz *et al.* 2001; Tait
353 *et al.* 2003a) it was hypothesised that the sediment transport inside pipes due to incoming rainfall
354 runoff does not include significant sediment wash-off from catchment surfaces, and that the main
355 source of suspended sediment is re-erosion of previously deposited in-pipe sediments.

356 *Sensitivity analysis*

357 A sensitivity analysis of some parameters of the erosion model was carried out by applying controlled
358 variations of their values in a valid range. In particular, the effect and influence of the bed porosity and
359 the bulk density were estimated. Porosity and bulk density were both included in the model in order to
360 calculate the volume of eroded sediments, which enables an update of the remaining sediment deposit
361 depth available for erosion. Porosity of the sediments was initially assumed as 0.20 based on initial
362 measurements ($p = 0.215 \pm 0.05$ performed by desiccation of fresh samples at 105°C during 24 hours).
363 During the sensitivity analysis, the porosity values were changed over the range 0.10 to 0.30, as sewer
364 deposits with fats and greases have been observed to have porosity ranging from 0.10 to 0.24 (Keener
365 *et al.* 2008). No significant influence on the eroded sediment depth evolution was observed under
366 porosity variation. Results obtained by using the event ID 2 are shown as an example in Fig. 10 (a).

367 Less than 8 % of variation in sediment concentration peak and around 10 % in sediment mass
368 mobilized was simulated, compared to simulation results obtained with $p = 0.20$.

369 The effects of changes in the sediment bulk density in the assessed range of variation for the local
370 sediments (1066 – 1458 kg/m³; average 1310 kg/m³) were also verified (Fig. 10 (b)). For event ID 2
371 shown as an example, variation from values calculated with the average sediment bulk density were
372 found between 1.5 to 6.4% regarding maximum sediment concentration, and between 9.4 and 16%
373 regarding total mass of sediment mobilized.

374 The greatest influence on the sediment transport loads is exerted by the hydraulic conditions. The
375 remobilization of sediments is directly related to the hydraulics that determined the boundary shear
376 stress values.

377 *Model results and performance*

378 The performance of the coupled SWMM5 and the calibrated Skipworth model (Fig. 5) was tested by
379 comparing measured versus modelled sediment peak concentrations and calculating NSE (Eq. 5).
380 Performance of the sediment transport model was analysed in the periods for which SS concentration
381 was measured and the obtained values are shown in Table 6.

382 Unfortunately, the total mass of sediment could not be considered for testing model performance
383 because of the adopted sampling strategy, addressed mainly to collect the first flush by including a
384 sampling collection for a total of 120 minutes which in most cases covered the first part of the rainfall
385 event duration.

386 Fig. 11 shows the sediment transport loads evolution assessed by the proposed model which is based
387 on the relationship of Skipworth with calibrated parameters. The SS concentration values obtained
388 were represented as an average value over the pumping interval (pumping-cleaning cycle in sample
389 collection).

390 During the rain event 1 (Fig. 11 a), the first phase of runoff arriving to the outlet of the catchment
391 generates an increase in water depth that was lower than the threshold water depth established for the
392 start of the operation of the automatic sampling collection. Thus, the first SS peak that can be observed

393 in the modelling results (Fig. 11a) were not covered by the measured SS data. Collected SS
394 concentration data corresponds instead with a second simulated peak when greater flow rates triggered
395 the collection of samples. It can be observed that there is a slight delay (6 minutes) between the
396 sediment concentration peak time measured and simulated during the event. It can be hypothesised
397 that this could be due to the 4 minutes delay between observed and measured peak flow. The 4
398 minutes delay observed at Fig. 11 (b) between simulated and measured C_{SS} for the event 2 might also
399 be linked with delays in the hydrodynamic results (8 minutes delay between observed and measured
400 peak flow from Table 4).

401 Both the NSE values and visual analysis of the pollutographs (Fig. 10) indicated a good fit between
402 simulated and observed data for events 1 and 2, a reasonable fit for event 3 and a poor fit for event 4.
403 Lower total precipitation and lower rainfall intensity for the event 4 might influence the predicted
404 results since the lower shear stresses generated in the SWMM model are very close to the anticipated
405 surface threshold shear stress of the water sediment interface.

406 Fig. 12 shows that for the events 1 and 2 the applied bed shear stress (τ_b) observed at the outlet of the
407 analysed sewer system reaches values higher than the critical value of the deeper layer (τ_{cu}).
408 Meanwhile much lower values of applied shear stress are observed for the events 3 and 4. In these
409 events the shear stress does not even reach the level at which the superficial layer (d') is fully eroded.
410 This indicated that for rainfall events in which the shear stress is low and for thin surface layers in
411 which the shear stress threshold changes quickly, such calibrated models struggle to accurately
412 simulate erosion rates.

413

414 **CONCLUSIONS**

415 **Transport parameters assessment**

416 Based on the laboratory findings for the highly organic sewer sediments collected in this study, it can
417 be confirmed that the critical shear stress values can be linked to the sediment bed depth, and hence

418 the values of the parameters d' , τ_{cs} , τ_{cu} , b and M , depend on the characteristics of the sediment and on
419 the structure of the in-pipe deposit.

420 From the analysis of the results obtained regarding the performance of the parameters it can be
421 suggested that the variation of the parameter M might be dependent on other sediment characteristics,
422 such as the median particle size (d_{50}) of the eroded sediments. The range of values adopted by b and M
423 might be also dependent on the density of the sediment eroded.

424 The sediment erosion and transport model performed well for three out of four rainfall events for
425 which flow and suspended sediment data were collected in the case study catchment. It predicted the
426 peak SS concentrations in these events with a Nash-Sutcliffe efficiency ranging from 0.73 to 0.85.
427 However, it needs to be stressed that the collection of the sewer sediment samples for the laboratory
428 analysis is practically difficult and assumptions had to be made in the design of the consolidation
429 periods to simulate deposition conditions in the sewer environment in the laboratory. The design of the
430 laboratory consolidation conditions may have an influence on the estimation of the values of the
431 calibration parameters used in the sediment erosion and transport model. Furthermore, temporal and
432 spatial variability of the sediment characteristic in the system might introduce a level of uncertainty
433 that was not examined, as the laboratory tests were all completed using samples collected at a single
434 location on a single day.

435 Because of site-specific sewer sediment characteristics, the parameters involved in the sediment
436 erosion model must be determined using local sediments. Performing erosion tests in the laboratory
437 gives the possibility of assessing the necessary parameters to deliver a more reliable prediction of in-
438 sewer transport and erosion.

439 Results from the assessment of the critical shear stress through the erosion tests confirmed the
440 structure of the sediment deposit model proposed by Skipworth regarding the existence of a weak
441 upper layer and increasing resistant erosional strength with depth through the bed. A power law trend
442 was found to describe the variation of the erosional resistance against the depth of the deposit.
443 Furthermore, the values obtained in the present work for the critical shear stress τ_c , varying from 0.15
444 up to 1.4 N/m² (depending on the consolidation period for a deposit of 30mm depth), are in the range

445 found from previous in-situ and laboratory work with real sewer sediments carried out by McIlhatton *et*
446 *al.* (2005) and Oms *et al.* (2008) who reported values in the range between 0.15 and 0.85 N/m².

447 The results from erosion tests also suggested that the behaviour of newly-deposited surficial sediments
448 subject to dynamic consolidation for up to around 24 hours show an increasing resistance against
449 erosion, and when the period of consolidation exceeds the 24 hours; any further increase in resistance
450 becomes insignificant (Fig. 8).

451 Further research is needed to identify a more direct relationship between the parameter *b* and *M* with
452 the sediment characteristics.

453 **Sediment transport modelling application**

454 For the case study described in this paper it was verified that the initial conditions regarding sediment
455 deposit properties and hydraulic parameters are indeed relevant in the prediction of SS loads released
456 and mobilized from in-sewer pipes during rainfall events. The large variation in the nature and
457 behaviour of the deposited sediments, the highly variable hydraulic conditions, and the complexities of
458 the processes occurring in-sewer makes a calibration process and validation against locally measured
459 data essential.

460 The predictive capacity of the sediment transport model proposed by Skipworth *et al.* (1999) was
461 verified with NSE between 0.85 and 0.73 for three out of four events. The indicated performance on
462 the results is directly related to an adequate assessment of the values of the transport parameters
463 considering the local sediment characteristics, and to an adequate calibration of the hydraulic model
464 using locally measured rainfall and flow data.

465 Following the analysis of the simulation results it can be observed that the rapid change in SS
466 concentrations is due to the quick response of the system influenced by a high level of imperviousness
467 in the catchment as well as the pattern of rainfall. It was concluded that reducing the sampling
468 frequency at the beginning of the event is desirable so as to be able to capture with more detail the
469 highly variable start of the pollutograph. Sampling interval adjustments will depend on the catchment
470 characteristics and concentration time on the case study. As an alternative, the on-line probes that can

471 make indirect measurements of the SS concentration could be used to obtain data with a higher
472 temporal resolution. The locally calibrated data can then be directly compared with the temporal
473 pattern of the SS concentration prediction.

474 Improved *first flush* prediction is required to better manage the pollution events on receiving natural
475 watercourse pollution through CSOs. The sediment modelling provided a better fit for the three largest
476 rainfall events, indicating that more research may be needed in defining how exactly the weak layer at
477 the very top of the in-sewer deposits erodes.

478

479 **ACKNOWLEDGEMENT**

480 The authors would like to acknowledge the managers from *Consorti per a la Defensa de la conca del*
481 *Besòs* and the technical staff from *Drenatges Urbans del Besòs* for funding and selfless collaboration
482 in the fieldwork. Also to the technical staff of the *University of Sheffield* for providing the facilities
483 necessary to carry out the laboratory work from where the results reported in this study were obtained.
484 The corresponding author wishes to thanks *AGAUR* (pre-doctoral research grant No. IUE/2644/2010)
485 for the financial support.

486

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- 583

584 **FIGURE CAPTION LIST**

585 Fig 1. Location of the study urban catchment (on the left) adapted from official cartographic data
586 (Institut Cartogràfic i Geològic de Catalunya 2017) and layout of the combined sewer network (on the
587 right) and catchment subdivision for the hydrodynamic and quality modelling.

588 Fig 2. Particle size distribution in raw sewage deposited sediments at Granollers, Spain. PSD
589 performed with standard sieve ($>1\text{mm}$ sub-sample) and laser diffraction analysis ($<1\text{mm}$ sub-
590 fraction).

591 Fig 3. Comparison between measured and calibrated hydrograph for rain events taken as examples.

592 Fig 4. Variation of the erosional resistance of the sediment deposit in a depth profile (after Skipworth,
593 1999).

594 Fig 5. Scheme of the simplified network sediment transport module coded in MATLAB.

595 Fig 6. Erosion rate against applied shear stress. Measured data, error in measurement and regression
596 function found.

597 Fig 7. Sediment bed depth strength against applied shear stress. Measured data from erosion tests and
598 trend.

599 Fig 8. Bed strength profile in depth of the sediment layer.

600 Fig 9. Variation on the parameters b and M values against applied shear stress for all the dry-period
601 tested.

602 Fig 10. Influence of the variation of characteristic sediment parameters on the evolution of sediment
603 concentration over time for event ID 2.

604 Fig 11. Sediment transport loads evolution. Measured and simulation values based on the relationship
605 of Skipworth (1999) with adapted transport parameters assessed for high organic sediments.

606 Fig 12. Applied and critical bed shear stress evolution and sediment bed depth evolution during
607 erosion process for the different rain events analysed.

608 **Table 1.** General characteristics for the catchment and combined sewer network of the study site.

catchment		combined sewer network	
Area	10.1 ha	average wastewater flow at outlet	24 m ³ /h
surface slopes	between 0.5 and 2.15 %	total length of pipes	2.2 km
% impermeability	between 77 and 93%	pipe diameters	300 to 1000 mm

609

610 **Table 2.** Rainfall events registered in the study site and used for the sediment transport modelling validation

registered data	ID	Date	total rainfall depth [mm]	maximum intensity [mm/h]	duration [minute]	antecedent dry-weather period length [days]
rainfall, flow and quality	1	17/09/2010	19.0	36.2	130	28
	2	31/05/2011	26.2	33.5	315	16
	3	24/10/2011	6.4	37.0	80	39
	4	13/07/2011	11.1	18.2	235	6
rainfall and flow	5	09/10/2010	33.5	36.6	605	21
	6	12/03/2011	71.6	18.2	1130	22

611

612 **Table 3.** Characteristics of sediments used by Skipworth *et al.* (1999), Rushforth (2001) and (Seco *et al.* 2014a)
 613 experimentation and in this work.

sediment type		characteristic particle size d_{50} [mm]	sediment density [kg/m ³]	organic content [%]
Sewer sediment from urban catchment in Granollers. Spain	(Seco <i>et al.</i> 2014a)	0.31(± 0.16)	1310 (± 146)	74 (VS/TSS)
	(batch used in this work)		1313 (± 95)	95 (VS/TSS) ±2
Crushed olivestone	(Skipworth <i>et al.</i> 1999) and (Rushforth 2001)	0.047	1445	100

614

615 **Table 4.** Relative errors used as goodness of fit measured flow rate with simulated flows during rain events.

Errors	Calibration events		Validation events	
	Rain event ID 5	Rain event ID 6	Rain event ID 2	Rain event ID 3
	09/10/2010	12/03/2011	31/05/2011	24/10/2011
Relative error of total runoff volume [%]	10 %	1 %	6 %	5 %
Relative error of peak flow [%]	10 %	2 %	7 %	8 %
time to 1rst peak error [min]	2	2	8	4

616

617 **Table 5.** Comparison of the values of transport parameters obtained from previous experimental studies
 618 (Rushforth 2001; Skipworth et al. 1999) and the values obtained in this study.

Parameter		values obtained in this study	Skipworth et al. (1999)		Rushforth (2001) (validation of Skipworth model)
			1:500 slope	1:1000 slope	
Material used		Sewer sediments	Crushed Olivestone		Crushed Olivestone
M	[g/s/m ²]	0.5 - 1.5	2.0	0.35-0.65	0.73
b	[-]	0.125	0.45		0.93
d'	[mm]	32 - 64	7	3.8	7.2
τ_{cs}	[N/m ²]	0.15	0.20	0.10	0.07
τ_{cu}	[N/m ²]	1.07 – 1.38	0.50	0.20	0.37

619

620 **Table 6.** Performance evaluation between observed and simulated suspended sediment transport evolution.

Rain event	ID 1 17/09/2010	ID 2 31/05/2011	ID 3 24/10/2011	ID 4 13/07/2011
Relative error of peak in sediment concentration	14.4%	1.1%	38.3%	89.1%
<i>NSE</i> - Nash-Sutcliffe efficiency	0.80	0.85	0.73	-0.18

621