

The Atlanto-Pacific multidecade oscillation and its imprint on the global temperature record

Melinda Nagy¹ · Kristóf Petrovay^{1,2} · Robert Erdélyi^{2,3}

Received: 14 October 2015 / Accepted: 17 April 2016 / Published online: 3 June 2016
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Abstract Results from a multiregression analysis of the global and sea surface temperature anomalies for the period 1950–2011 are presented where among the independent variables multidecade oscillation signals over various oceanic areas are included. These indices are defined in analogy with the Atlantic Multidecadal Oscillation (AMO) index. Unexpectedly we find that a strong multidecade oscillation signal echoing the AMO is also present in the Western and Northwestern Pacific region. The results indicate that naturally induced climate variations seem to be dominated by two internal variability modes of the ocean–atmosphere system: AMO and El Niño Southern Oscillation, with a marked geographical dichotomy in their respective areas of dominance. As the AMO index is directly derived from SST data the finding that the AMO signal is present on a large fraction of the global oceanic surface casts doubt on its use as an independent explanatory

variable in regression analyses of the global surface temperature anomalies.

Keywords Climate change · Linear multiregression · AMO

1 Introduction

While there is very strong evidence that climate change in the past decades has been dominated by human effects on the environment, natural factors still exert a significant influence on climate variations (IPCC 2013). The respective roles of various natural influences, their interplay and their expected variation in the future is, however, still poorly understood.

Global climate simulations have by now reached a degree of development where they are able to model many of these effects; however, there is still a very significant diversity between the outcomes of various simulations depending on model assumptions and the effects considered. The high complexity of these simulations often also makes their results hard to interpret from a physical point of view. (For a detailed discussion see IPCC 2013, Chapter 9).

As a result, at the opposite end of the spectrum of models, taking the simplest possible approach of a linear multiregression analysis of the empirical data has become increasingly popular in recent years. In such studies, the time series of a global indicator of climate change y , most commonly the mean global surface temperature (GST) or the mean sea surface temperature (SST) is fitted by a linear combination of of number of parameters x_i :

$$y(t) = x_0 + c_1x_1(t - \Delta t_1) + c_2x_2(t - \Delta t_2) + \dots + c_Nx_N(t - \Delta t_N) \quad (1)$$

Electronic supplementary material The online version of this article (doi:10.1007/s00382-016-3179-3) contains supplementary material, which is available to authorized users.

✉ Melinda Nagy
m.nagy@astro.elte.hu
Kristóf Petrovay
k.petrovay@astro.elte.hu
Robert Erdélyi
robertus@sheffield.ac.uk

¹ Department of Astronomy, Institute for Geography and Earth Sciences, Eötvös University, Budapest, Hungary

² Solar Physics and Space Plasma Research Center, University of Sheffield, Sheffield, UK

³ Debrecen Heliophysical Observatory (DHO), Research Centre for Astronomy and Earth Sciences, Hungarian Academy of Sciences, P.O. Box 30, Debrecen 4010, Hungary

Here, t is time and the c_i 's and Δt_i 's are fitting parameters. The set of x_i considered are usually indices of human and natural influences on the climate system, such as greenhouse gas (GHG) emission, tropospheric and stratospheric aerosol input (these are predominantly of anthropogenic and natural origin, respectively), solar radiative forcing as well as indices of the most important internal oscillations of the atmosphere–ocean system such as the El Niño Southern Oscillation (ENSO).

Such multiregression analyses rely on two assumptions that inherently limit the validity of their conclusions. The first is the assumption of linearity which can be reasonably assumed to hold as long as the input parameters vary within a certain limited range. How wide that range is and what will be the response of the climate system beyond the range is an issue that can be only be examined in physical models of the climate systems such as global simulations. The second, although often implicit, critical assumption of multiregression analyses is that the set of x_i 's is complete and non-redundant, i.e. that the input parameters are independent and include all physically important factors influencing climate change. Neglecting an important factor will result in assigning an unduly high explanatory power to the factors considered, while a correlation between two x_i time series (whether it is a chance correlation or the result of a physical link between the effects) will result in *aliasing*, i.e. an indecisive situation with a strong cross-talk between the corresponding c_i coefficients.

Perhaps the most influential of these studies was the work of (Lean and Rind 2008, 2009). The input parameters considered were monthly values of the multivariate ENSO index, total solar irradiance, stratospheric aerosol concentration, and a net anthropogenic forcing parameter. The time lags Δt_i were chosen to maximize the explained variance: this generally resulted in lags of a few months except in the case of anthropogenic forcing where an optimal time lag of 10 years was found. The regression was found to explain 76 % of the variance in the CRU¹ monthly GST data from 1889 to 2006.

In a more recent extension of this approach, Chylek et al. (2014) include the Atlantic Multidecadal Oscillation (AMO) index among the explanatory variables. Further differences in their analysis as compared to Lean and Rind include the use of annual, rather than monthly data and another GST time series (GISS² vs. CRU). Even without the introduction of time lags, the factors considered by Lean and Rind were found to explain 89 % of the variance in the mean annual GST series for the period considered (1900–2012), the higher explanatory power being presumably due to the use of yearly averages. Adding the AMO

index to the natural factors influencing climate was found to increase the explained variance to 94 %. A most surprising finding, however, was that a combination of greenhouse gas (GHG) emission and AMO *alone* explains 93 % of the variance—nearly as much as all the factors combined! (To place this in context we note that GHG alone explains ~81 % of the variance).

The fact that a high explanatory power can be achieved in linear multiregression models employing quite different sets of explanatory variables suggests that a strong aliasing may be present either between the various explanatory variables, or between some of the explanatory variables and the dependent variable (the mean GST). The objective of the present paper is to examine to what extent such aliasing may influence the curious results of Chylek et al.

The structure of the paper is the following. In Sect. 2 we present and discuss the data used. Section 3 focuses on the issue on potential aliasing between the AMO index and the GST time series, while in Sect. 4 we address aliasing between the various natural factors determining climate. Section 5 concludes the paper.

2 Data analysis

In contrast to the previous studies, in most of the present work we restrict the analysis to the time period 1950–2011. This is motivated by the following considerations of homogeneity of the data set.

- The ENSO index is directly available for this period only, earlier values being reconstructions based on temperature data alone.
- The usual definition of the AMO index involves the subtraction of a linear trend purported to represent anthropogenic influence—this assumed linearity of anthropogenic effects is also questionable over a longer base line.
- All standard GST anomaly (GSTA) data sets rely heavily on sea surface temperatures. In the period prior to 1950 major bias corrections are applied to these data the details of which are a source of added uncertainty in the results (Smith and Reynolds 2002; Thompson et al. 2008; Hansen et al. 2010).
- Recently there has been some disagreement regarding a possible temporal inhomogeneity in the calibration of the sunspot number series upon which the solar irradiance reconstruction is based (Clette et al. 2014). This issue also mainly affects data before about 1950. A caveat to be noted in connection with this temporal limitation is that due to a more limited long term variation of solar activity levels during this period, our approach may result in a lower weight attributed to solar irradiation variations in natural climate variations.

¹ CRU = Climate Research Unit (University of East Anglia).

² GISS = NASA Goddard Institute for Space Studies.

Table 1 URLs of data sources

GSTA	http://data.giss.nasa.gov/gistemp/
SST	http://www.esrl.noaa.gov/psd/data/gridded/data.noaa.ersst.html
GHG + aerosols	http://data.giss.nasa.gov/modellforce/Fe.1880-2011.txt
Sunspot numbers (unrevised)	http://www.sidc.be/silso/
ENSO	http://www.esrl.noaa.gov/psd/enso/mei/
NAO	http://www.esrl.noaa.gov/psd/gcos_wgsp/Timeseries/NAO/
AMO	http://www.esrl.noaa.gov/psd/data/timeseries/AMO/

Table 2 Multilingual regressions applied to GST anomaly using subsets of the potential explanatory variables, including natural and anthropogenic effects

Explanatory variables	R^2	R^2_{adj}
Anthropogenic	0.816	0.810
Natural + anthropogenic	0.888	0.878
Natural + anthropogenic + AMO	0.956	0.952
Natural + anthropogenic + NAO	0.889	0.8768
GHG + AMO	0.919	0.917
GHG + NAO	0.815	0.809

R^2_{adj} values in bold letters are those that are also listed in Table 1 of Chylek et al. (2014) (see the results in graphic form in Section S.1, Electronic Supplementary Material)

“Natural” = ENSO + solar and volcanic forcing. AMO index values here are from NOAA

“Anthropogenic” = green house gases (GHG) + tropospheric aerosols (AER), as given by Hansen et al. (2007, 2011)

URLs of our data sources are listed in Table 1.

For GSTA we use the data available at the NASA GISS web site (Hansen et al. 2010) while SST data were downloaded from the home page of NOAA (Smith and Reynolds 2004). Regression analyses of GST and SST data yield very similar results. In this paper, after analysing GST data to confirm some previous results we will focus on the analysis of the SST data set.

Radiative forcing by anthropogenic aerosols, greenhouse gases and aerosols from volcanic eruptions were taken from (Hansen et al. 2007, 2011). Results using greenhouse gas (GHG) data alone or all anthropogenic forcings proved to be virtually indistinguishable; in addition, a strong aliasing exists between these effects, so in what follows anthropogenic forcings will be routinely often represented by GHG alone.

Instead of using a model dependent reconstruction of solar radiative forcing we characterise solar activity variations simply by the International Sunspot Number series as given by the Royal Observatory of Belgium (unrevised values).

Annual values of the ENSO and NAO indices were obtained from the NOAA data base (Wolter and Timlin 1993, 1998; Jones et al. 1997).

Values of the AMO index are also available from NOAA but in most of our analysis here we calculated this index directly from the SST grids as described in (Enfield et al. 2001) but without applying the smoothing on the data.

3 AMO versus GSTA

In the present paper we extend previous multiregression analyses to address a number of questions raised in earlier work. First, there may be an issue regarding the validity of the use of the AMO index as an independent explanatory variable, given that the definition of this index is based on the mean sea surface temperature over a significant fraction of the globe (i.e., the N-Atlantic basin). It is to some extent natural to expect that, upon linear detrending, these data will reflect naturally induced global temperature variations.

As a first step we made an analysis closely following that of (Chylek et al. 2014), performing a linear regression in order to explain the temporal evolution of the GSTA with different subsets of the potential explanatory variables (anthropogenic, natural, AMO—cf. caption of Fig. 1).

The resulting values of the explained variance R^2 as well as the adjusted variance R^2_{adj} , taking into account the number of variables included to the regression, are summarized in Table 2. Overall, our numbers are in good agreement with the results shown in Table 1 of Chylek et al. (2014). Note that we further test whether the introduction of the North Atlantic Oscillation (NAO) Index as an explanatory variable improves the regressions but, as apparent from the table, the results were negative.

In order to check whether the good explanatory power of AMO is simply a consequence of its definition based on mean temperature anomalies on a non-negligible part of the global ocean surface, we set out to investigate whether the N-Atlantic region is of any special interest from this point of view.

For this purpose, we arbitrarily assign other eight regions of comparable size on the global ocean (see Fig. 1). For each region we determine a “multidecade oscillation” (MO) index using the same procedure as for the AMO (but no smoothing is applied to the annual values). With each of these MO data sets, combined with the GHG data, we

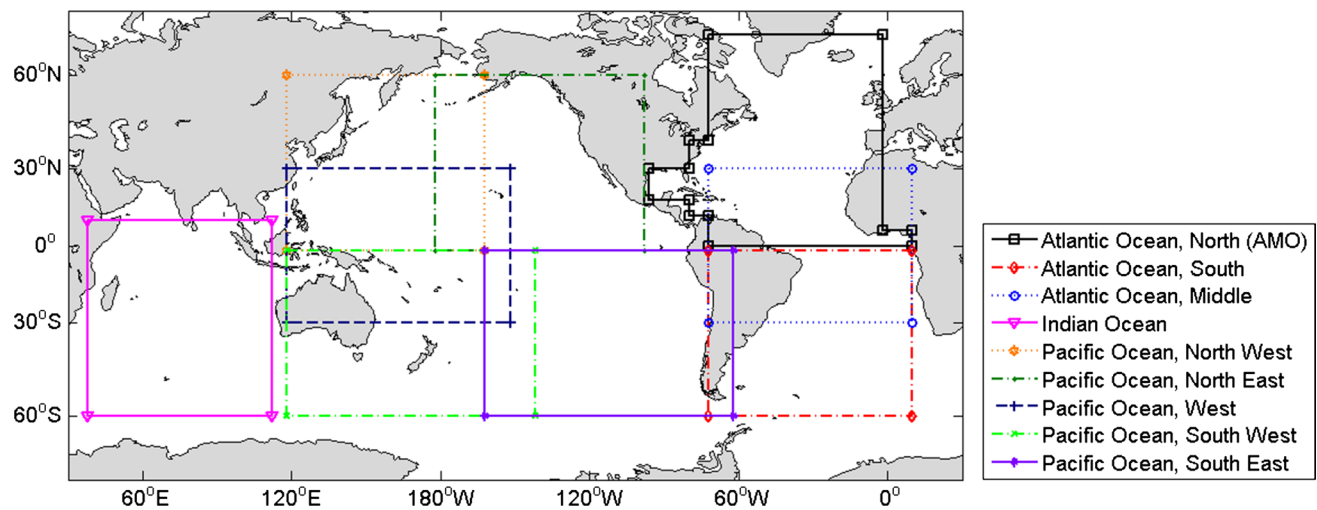


Fig. 1 Sea surface regions defined for comparison with the North Atlantic region, as outlined for the usual definition of AMO index (Enfield et al. 2001)

perform a linear regression to the GSTA record. Results are shown in Table 3. (The first line of the table refers to the AMO index calculated by us from the SST data.)

It is clear that significant differences are present between the explanatory power of the MO indices of different regions with respect to the GSTA. For some regions the value of R^2_{adj} hardly differs from its counterpart obtained with GHG alone, while for others the inclusion of the MO index in the analysis leads to a spectacular improvement. The N-Atlantic region (and the mid-Atlantic region overlapping with it) yields the highest explained variance of all; yet, interestingly, it is almost equalled by the Western Pacific region (and, to some extent, the NW-Pacific region overlapping with it).

Indeed, the overall parallelism in the MO indices calculated for these regions is borne out in Fig. 2. The characteristic 50–60 year cyclicity of AMO is very clearly reflected in the Western Pacific region, albeit with a higher added “noise” in the form of short-term fluctuations, and also in a smoother but slightly more distorted form in the NW-Pacific. In the first part of the period there is a hint of some phase lag between the Atlantic and Pacific signals but in view of the smoothing applied the only certain statement is that any phase lag is well below the width of the smoothing window (i.e. no more than a few years).

Other areas of the global ocean such as the Indian Ocean, the Eastern Pacific or the Southern Atlantic, on the other hand, do not show such a clear reflection of the AMO signal, nor is their explanatory power with regard to the GSTA even nearly as high.

In view of the relative shortness of the time period considered (1950–2011) the robustness of our findings may

be in question. In order to address this, ignoring the issues related to the use of pre-1950 data listed in Sect. 2, in Fig. 3 we plot smoothed MO indices for the relevant regions starting from 1920. The plot confirms that the parallel variation of smoothed detrended temperature anomalies between the N-Atlantic and W-Pacific regions was present also in the first half of the 20th century, lending support to the robustness of the result.

The question naturally arises: what is the origin of this surprising parallelism between two distant regions of the globe? One possibility may be that the AMO is a global climate phenomenon that is for some reason more strongly imprinted in the regional signal in some areas than in others (e.g. due to being suppressed in ENSO-affected areas). Alternatively, it is also possible that we are dealing with a real teleconnection effect. The physical mechanism of this teleconnection is an open question. One possibility is an atmospheric teleconnection across the Eurasian land mass: the prevailing westerlies may carry the AMO-induced atmospheric anomalies (SAT and moisture) to the Pacific, where the effect is attenuated from West to East. The other possibility of a link via the global thermohalocirculation is even more tentative as it would imply a connection, possibly by advection of deep/bottom water, across the Southern hemisphere. In either case, given that the other main manifestation of this oscillatory phenomenon is in the Western Pacific region, it may be more aptly called the “Atlanto-Pacific Multidimensional Oscillation”.

The mechanism driving the multidecade variation is still open to question; here we only note that internal variations in ocean basins on a time scale of decades naturally arise from aperiodic oscillations [(cf. Huang (2010), Ch.5.).

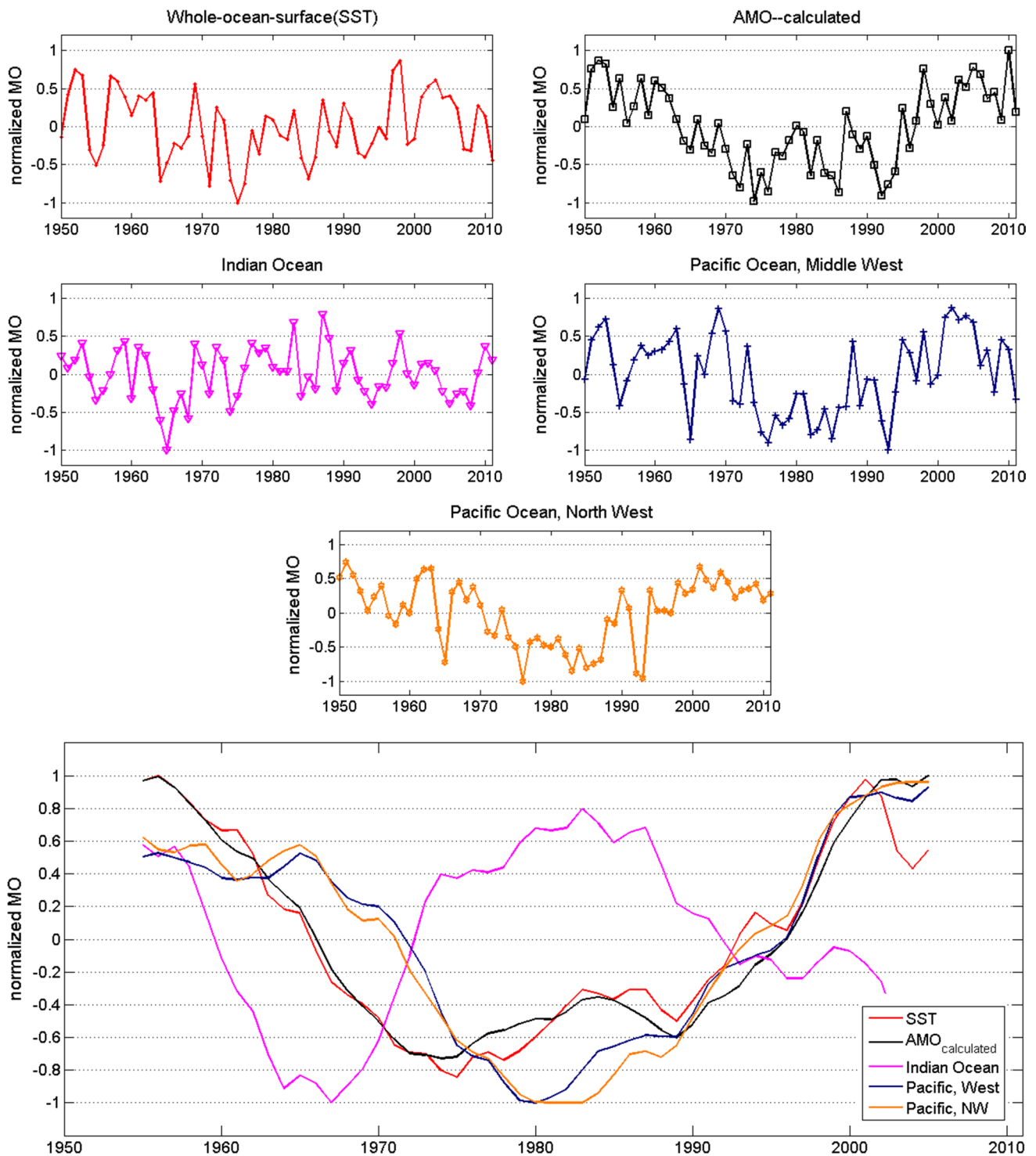


Fig. 2 MO indices as a function of time for selected sea surface regions individually (*top*) and overplotted (*bottom*), after smoothing by an 11-year sliding window

4 AMO versus other natural factors

The results in Table 1 confirm the surprising finding by Chylek et al. (2014) that anthropogenic forcing combined

with the AMO has the potential to explain a very high fraction of the observed variance in GSTA, while adding all the other natural forcings only leads to a slight further improvement in the fit. What makes this result rather puzzling is

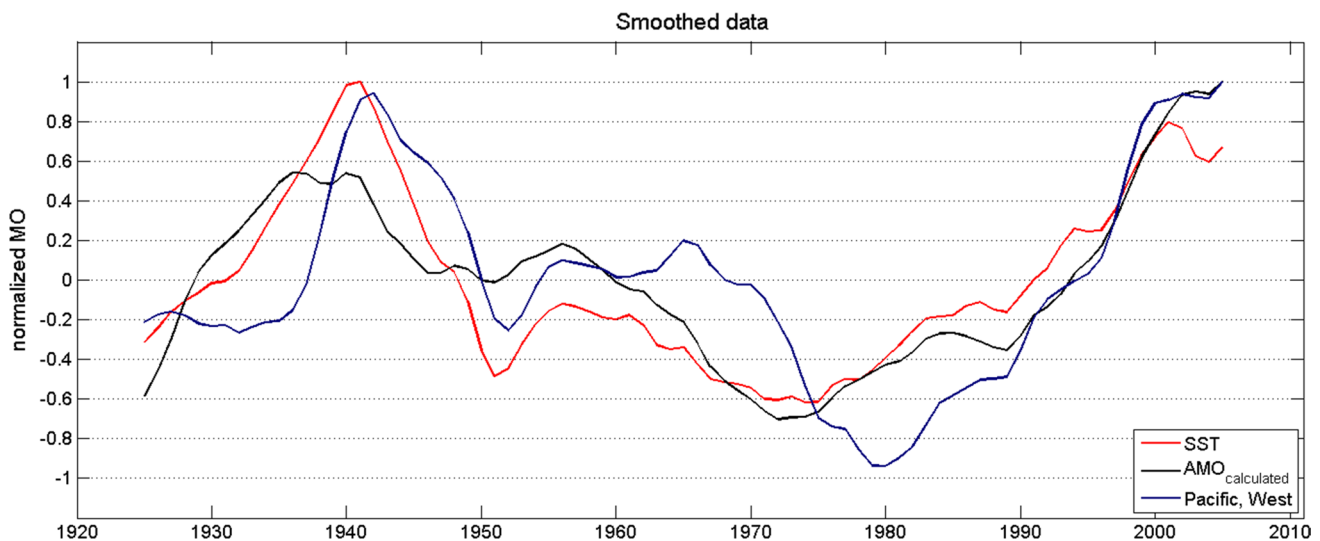


Fig. 3 Same as Fig. 2 (*bottom*) for the base period 1920–2011

Table 3 Results of linear regression of Global Surface Temperature Anomalies when the explanatory variables employed were the GHG and MO indices

Sea surface area	R^2	R^2_{adj}	r_{AMO}
Atlantic Ocean, North (AMO)	0.923	0.920	0.970
Atlantic Ocean, Middle	0.914	0.911	0.604
Atlantic Ocean, South	0.812	0.805	-0.295
Indian Ocean	0.9226	0.846	0.202
Pacific Ocean, Middle West	0.912	0.909	0.664
Pacific Ocean, North East	0.840	0.833	0.0838
Pacific Ocean, North West	0.880	0.875	0.657
Pacific Ocean, South East	0.839	0.834	0.089
Pacific Ocean, South West	0.862	0.857	0.428
Whole Ocean Surface (SST)	0.971	0.970	0.646

Here, the MO index was defined for different sea surface areas shown in Fig. 1. In this table AMO refers to the index that we calculated from the SST data. The last column shows the correlation coefficient between the downloaded AMO data series and MO indices. In the first line the correlation coefficient of downloaded and calculated AMO index is shown (see the results in graphic form in Section S.2, Electronic Supplementary Material)

that these other natural factors, combined with the anthropogenic forcing could also explain a rather high proportion (88 %) of the variance even without the inclusion of AMO. Therefore, if the AMO were indeed the predominant natural factor influencing climate, this high explanatory power of other natural influences would remain unexplained or should be attributed to a chance coincidence.

Note, however, that from Table 1 it is also seen that anthropogenic effects alone can explain 81 % of the variance of the temperature anomaly. External forcing and the ENSO only add 7 % to this while the AMO adds approximately 10 %. AMO and other natural effects combined add 14 % instead of the 17 % that would be expected if they were uncorrelated. The correlation between AMO and other natural factors that is needed to explain this discrepancy is then not so excessive as it might appear at first sight.

To independently check this inference, we have performed a regression analysis with a detrended MO index as dependent variable and other natural factors as explanatory variables. The results are shown in Table 4. It is apparent that in the regions displaying a marked AMO signal natural factors explain typically about 20 % of the variance in the MO index, in agreement with our arguments above.

The four regions found to display a strong AMO signal in the previous section together cover a significant fraction of the total sea surface area inside the $-60 < \phi < +60$ latitude zone. This prompts the question whether the strong imprint of AMO on the global mean SST signal (and, by inference, on the GST signal) is simply due to its predominance in a significant fraction of the base area or else it influences climate on a more global scale.

In order to clarify this issue we consider the mean sea surface temperature anomaly in the part of the $-60 < \phi < +60$ latitude zone *not* covered by the four regions displaying a strong AMO signal (see Fig. 4). The results of a multiregression analysis of this mean anomaly series are shown in Table 5, and an example of regression

Table 4 Results of linear regression when the dependent variable is one of the different MO indices and explanatory variables are the natural factors alone or combined with anthropogenic forcing

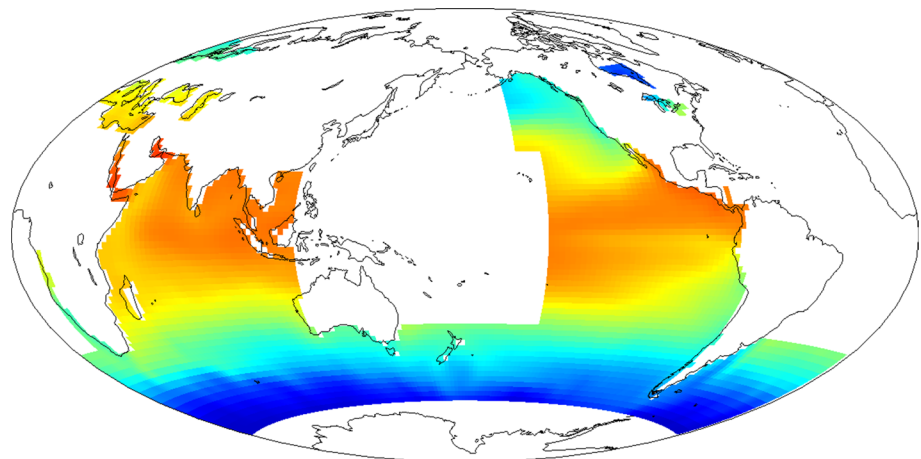
Sea surface area	Natural effects only		Natural and anthropogenic effects	
	R^2	R^2_{adj}	R^2	R^2_{adj}
Atlantic Ocean, North (AMO)	0.229	0.189	0.500	0.455
Atlantic Ocean, Middle	0.200	0.158	0.742	0.719
Atlantic Ocean, South	0.076	0.028	0.793	0.775
Indian Ocean	0.488	0.198	0.933	0.859
Pacific Ocean, Middle West	0.127	0.082	0.660	0.629
Pacific Ocean, North East	0.591	0.570	0.648	0.616
Pacific Ocean, North West	0.223	0.183	0.555	0.516
Pacific Ocean, South East	0.766	0.754	0.909	0.901
Pacific Ocean, South West	0.249	0.210	0.792	0.773
Whole Ocean Surface (SST)	0.497	0.471	0.911	0.903

See the results of regression in the case of both natural and anthropogenic effects in graphic form in Section S.3, Electronic Supplementary Material

Table 5 Results from multilinear regressions to SST anomalies calculated over regions unaffected by the AMO

Explanatory variables	R^2	R^2_{adj}
Anthropogenic	0.819	0.810
Natural + anthropogenic	0.965	0.961
Natural + GHG + AMO	0.961	0.957
GHG + AMO	0.771	0.763
GHG + ENSO	0.942	0.941

Fig. 4 A snapshot of the distribution of the MO index defined over the sea surface area not affected by the AMO signal. (Blue-to-red scale with arbitrary normalization)



on Fig. 5. It is clearly seen that the explanatory power of AMO with respect to the mean SST anomaly series is drastically reduced if the main AMO regions are omitted.

5 Conclusion

In this paper we have presented results from a series of linear multiregression analyses of global and regional SST.

A most surprising and unexpected find was the very strong appearance of the AMO signal in the detrended SST anomaly time series over the Western tropical Pacific and over the NW Pacific. To the best of our knowledge, such a clear indication of a strong AMO imprint in this oceanic region (or indeed *any* oceanic region far from the North Atlantic) has not been demonstrated before. Hints of an in-phase variation of temperature variations in the NW-Pacific and N-Atlantic can be seen in both proxy reconstructions and coupled atmosphere-ocean models (Delworth and Mann 2000). The definitions of the so-called Pacific Decade Oscillation (PDO) and Interdepartmental Pacific Oscillation (POI) involve much larger areas over which the AMO signal becomes inter tangled with the long-term residual of ENSO and other effects. Principal component analysis (PA) and other similar methods applied to the spatio-temporal distribution of SST data have in fact already demonstrated the presence of a component related to AMO with a rather high amplitude in the NW Pacific (McCabe and Palecki 2006; d’Orgeville and Peltier 2007), while the other dominant component is a decade-scale ENSO residual. Our results are in agreement with these earlier findings

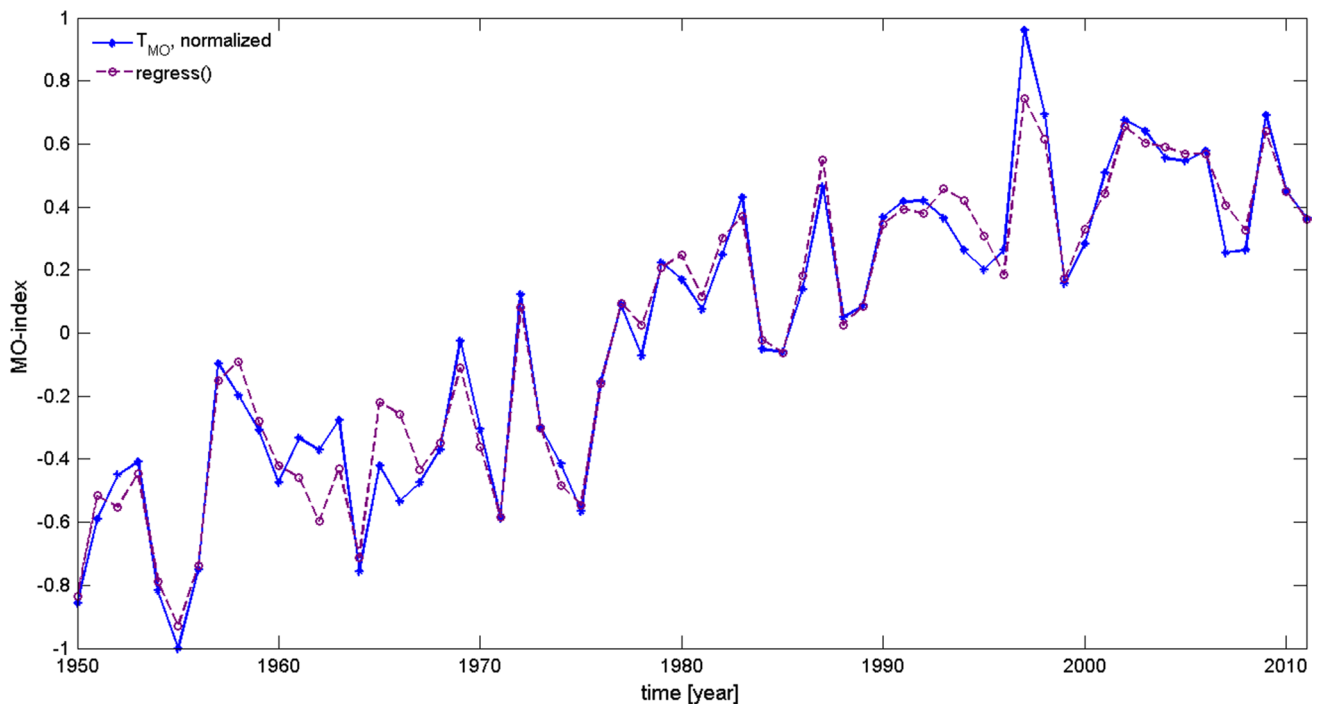


Fig. 5 Time series of MO index, according to Fig. 4—blue line. Purple dashed line indicates the result of linear regression, when the explanatory variables are the natural and anthropogenic effects

but also display a surprisingly sharp geographical dichotomy between oceanic areas affected by one or the other principal component. These areas predominantly (though not exclusively) lie on the Northern and Southern hemispheres for the AMO and ENSO components, respectively, supporting also the use of the term “Northern Multidimensional Oscillation” as recently suggested by Steinman et al. (2015).

Aliasing between the AMO signal and other natural factors is found to be relatively weak ($\sim 20\%$). On the other hand, the fact that a strong and evident AMO signal is displayed in detrended SST anomalies over a significant fraction of the global oceanic surface implies a strong aliasing link between the AMO index and the global mean SST anomaly.

Indeed, if the AMO signal, defined on the basis of SST values over a larger area is present over a significant fraction of the ocean surface, this casts doubt on the use of AMO as an explanatory factor in multiregression analysis. “Explaining” global mean temperature variations by the mean temperature variations on a large fraction of the globe is a rather tautological exercise. It may seem that AMO is part of the problem rather than part of the solution. One may even argue that this is also true for ENSO to some extent, even though SST data are only one of six inputs that are combined to form this particular index.

Without delving too deeply into the philosophical issue of the meaning of “explanatory variable” we may suggest that a sensible approach is to consider a statistically significant correlation “explanatory” when the physical link is at least not implausible while the possible alternatives (an inverse cause-effect relationship or the two time series being due to a common cause) are a priori much less plausible; at the same time the two data sets are truly independent. These criteria are satisfied for external climate forcings such as TS, stratospheric aerosols or anthropogenic forcings (note that aliasing between these effects may still be an issue). For internal variabilities independence is difficult to ascertain: even when an index (e.g. the SOI) is based on data that do not include SST it can be argued that it is just another manifestation of the same weather pattern that is also reflected in [a subset of] the SST data analysed. In short of a definitive answer to this conundrum it is still important to stress the importance of choosing an internal variability index that has as little direct dependence on the data set being analysed as possible.

To summarize our findings, naturally induced climate variations seem to be dominated by two dominant internal variability modes of the ocean–atmosphere system, with a marked geographical dichotomy in their respective areas of dominance. The influence of external forcings is non-negligible but secondary to these internal variabilities. However,

for a proper assessment of the influence of forcings a multi-regression analysis where internal variabilities, esp. AMO, are treated as explanatory variables on an equal footing with the forcings seems to be inappropriate.

Acknowledgments This research was supported by STOCK Grant ST/J001430/1 and The Royal Society Grant (ref nr.: R/142495). The possibility that an atmospheric teleconnection mechanism may link the N-Atlantic and W-Pacific regions, mentioned in Sect. 3, was suggested to us by an anonymous referee.

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