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**Published paper**

Glacial geomorphology of Marguerite Bay Palaeo-Ice stream, western Antarctic Peninsula

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Glacial geomorphology of Marguerite Bay Palaeo-Ice stream, western Antarctic Peninsula

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This paper presents a glacial geomorphological map of over 17,000 landforms on the bed of a major palaeo-ice stream in Marguerite Bay, western Antarctic Peninsula. The map was compiled using various geophysical datasets from multiple marine research cruises. Eight glacial landform types are identified: mega-scale glacial lineations, crag-and-tails, whalebacks, gouged, grooved and streamlined bedrock, grounding-zone wedges, subglacial meltwater channels, gullies and channels, and iceberg scours. The map represents one of the most complete marine ice-stream signatures available for scrutiny, and these data hold much potential for reconstructing former ice sheet dynamics, testing numerical ice sheet models, and understanding the formation of subglacial bedforms beneath ice streams. In particular, they record a complex bedform signature of palaeo-ice stream flow and retreat since the last glacial maximum, characterised by considerable spatial variability and strongly influenced by the underlying geology. The map is presented at a scale of 1:750,000, designed to be printed at A2 size, and encompasses an area of 128,420 km².

Keywords: marine; palaeo-ice stream; glacial geomorphology; Antarctica; bedform variability; geology

1. Introduction

Rapidly flowing ice streams and outlet glaciers account for dynamic mass loss from the Greenland and Antarctic ice sheets, and recent observations of their thinning and acceleration suggest that their contribution to sea-level rise is increasing (Pritchard, Arthern, Vaughan, & Edwards, 2009; Shepherd et al., 2012). Their pattern of retreat is, however, characterised by a variable, often asynchronous behaviour at short (decadal) time-scales (e.g. Moon, Joughin, Smith, & Howat, 2012; Pritchard et al., 2009). This short-term variability introduces considerable uncertainty in understanding recent ice-sheet changes and predicting near future sea-level rise, as highlighted by the Intergovernmental Panel on Climate Change (IPCC, 2007). To provide a long-term context for recent changes, therefore, numerous workers have recognised the potential of

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reconstructing palaeo-ice streams. There are now well-established criteria for recognising former
ice streams in the geological record (Dyke & Morris, 1988; Stokes & Clark, 1999), which
represent important archives for reconstructing ice-stream dynamics and investigating basal pro-
cesses (e.g. Graham et al. 2009; Ó Cofaigh et al., 2005).

In Antarctica, the advent of multi-beam swath bathymetry systems in the mid 1990s was funda-
mental in facilitating the identification of marine-terminating palaeo-ice streams from the glacial
gemorphic imprint preserved on the sea-floor. Over 30 palaeo-ice streams have now been docu-
mented in deep cross-shelf bathymetric troughs surrounding Antarctica, detailing the expansion
of ice out towards the continental shelf edge prior to deglaciation during the last glacial
maximum (cf. Livingstone et al., 2012 for a review) and several of these have been mapped in
detail (e.g. Camerlenghi et al., 2001; Evans, Pudsey, Ó Cofaigh, Morris, & Domack, 2005;
Graham et al., 2009; Jakobsson et al., 2012; Nitsche et al. 2013; Reinardy et al., 2011).

One of the best-studied palaeo-ice streams, with a wealth of geophysical data, was located in
the area now known as Marguerite Bay on the west side of the Antarctic Peninsula (Figure 1). It
constitutes a 50–80 km wide bathymetric trough, which extends for about 370 km from the inner
bay at the mouth of George VI Sound to the continental shelf edge. The trough is characterised by
a reverse-sloping bed that deepens from 500 m at the continental shelf edge to 1600 m on the inner
shelf (Figure 1) (Jamieson et al., 2012). The inner shelf of Marguerite Bay comprises rugged
bedrock characterised by large, isolated basins (up to 900 m deep), leading westwards into a
network of smaller connected basins 800–1000 m deep (Anderson & Oakes-Fretwell, 2008).

2. Previous work in marguerite bay

Previous marine geophysical and geological studies of the glacial geomorphology of Marguerite Bay
document a number of glacial landforms orientated along the axis of the trough and interpreted to in-
dicate streaming flow to the shelf edge during the last glaciation (Ó Cofaigh, Pudsey, Dowdeswell, &
Morris, 2002). This includes a transition from ice-moulded bedrock, drumlins and subglacial meltwater
channels formed predominantly in crystalline bedrock in the inner bay, to classical drumlins, highly atte-
nuated drumlins and mega-scale glacial lineations (MSGLs) on the mid-shelf. MSGLs up to 20 km in
length are formed in sediment across the outer shelf (e.g. Anderson & Oakes-Fretwell, 2008; Dowdes-
well, Ó Cofaigh, & Pudsey, 2004a, 2004b; Kennedy & Anderson, 1989; Ó Cofaigh et al., 2002, 2005,
Ó Cofaigh, Evans, Dowdeswell, & Larter, 2007). Marine sediment cores recovered from the main
trough in Marguerite Bay suggest a non-linear pattern of ice-stream retreat characterised by rapid de-
glaciation of the outer shelf at ~14 cal. ka BP, followed by a slower phase of retreat through the mid shelf,
thought to be associated with the break-up of an ice-shelf (Kilfeather et al., 2011). Thereafter, retreat
proceeded rapidly to the inner-shelf at ~9 cal. ka BP (Heroy & Anderson, 2007), with George VI
Sound becoming ice free between 6.6 and 9.6 cal. ka BP (Hjort, Bentley, & Ingólfsson, 2001; Smith
et al., 2007; Sugden & Clapperton, 1981). This final phase of deglaciation was accompanied by
rapid ice-sheet thinning in the hinterland (Bentley et al., 2011). However, despite the wealth of geomor-
phological, and sedimentological and chronological data available for the Marguerite Bay Palaeo-Ice
Stream (MBIS), a detailed map of the ice stream bed has not yet been produced. This paper draws
together the available geophysical datasets and compiles a comprehensive map of the glacial landforms
and sediments along MBIS (see main map) in order to develop a baseline dataset for investigating the
basal properties of ice streams and for testing ice-stream models (e.g. Jamieson et al., 2012).

3. Methods

Marine geophysical data were collected on cruises JR59, JR71 and JR157 of the RRS James
Clark Ross (JCR) and NBP0201 of the RV/IB Nathaniel B. Palmer (NBP) (Figure 1). Swath
Figure 1. Overview map of Marguerite Bay Ice Stream. Swath bathymetry tracks are as follows: dark blue line: NBP0201 RV/IB Nathaniel B. Palmer cruise tracks; yellow line: JR59; white line: JR71; and orange line: JR157 RRS James Clark Ross cruise tracks. Ages (in calibrated years before present, BP) derived from marine sediment cores (as per Livingstone et al., 2012) are displayed with 1 sigma error and the dates in bold refer to the most reliable core dates (i.e. those derived from calcareous micro-(fossils) and not affected by iceberg turbation). Note that the dates suggest rapid retreat of the outer-mid shelf at \( \approx 14 \) cal. ka BP, followed by a period of slower retreat through the mid shelf and then another phase of rapid retreat off the inner shelf \( \approx 9 \) cal. ka BP (see Heroy & Anderson, 2007; Kilfeather et al., 2011). APIS: Antarctic Peninsula Ice Sheet; EAIS: East Antarctic Ice Sheet; WAIS: West Antarctic Ice Sheet; LC = Larsen C Ice Shelf; FIS = Filchner Ice Shelf.
bathymetry data were obtained using Kongsberg EM120 (*JCR*) and hull-mounted SeaBeam 2100 (*NBP*) systems. The EM120 system has a $1^\circ \times 1^\circ$ beam configuration and emits 191 beams each with a frequency of 12 kHz and the hull-mounted SeaBeam 2100 system emits 120 beams at 12 kHz. Swath data were processed in MB-System (*Caress & Chayes, 2003*) to remove anomalous data points, corrected to include sound-velocity profiles, and gridded at $15 \times 45$ m cell sizes.

<table>
<thead>
<tr>
<th>Glacial landform type</th>
<th>Count</th>
<th>Identification characteristics</th>
<th>Mapping style</th>
</tr>
</thead>
<tbody>
<tr>
<td>Mega-scale glacial lineations (MSGLs)</td>
<td>5037</td>
<td>Highly elongate, parallel ridge-groove sets formed in the upper, transparent acoustic unit (soft till).</td>
<td>Crest-line of individual landforms</td>
</tr>
<tr>
<td>Crag-and-tail</td>
<td>452</td>
<td>A ‘tear-drop’ shaped landform with exposed bedrock at the stoss end and with a tail of tapering sediment.</td>
<td>Crest-line of individual landforms</td>
</tr>
<tr>
<td>(Asymmetric) whaleback/ drumlins</td>
<td>2761</td>
<td>Streamlined ‘tear-drop’ or ‘oval’ shaped hill formed in bedrock, often attenuated in the down-flow direction and with a stoss face that is sometimes steeper than the lee face. Often associated with crescentic scours around the stoss head.</td>
<td>Break-of slope outline of individual landforms</td>
</tr>
<tr>
<td>Iceberg scour</td>
<td>4246</td>
<td>Curvilinear furrows formed in sediment and displaying multiple orientations, which often cross-cut one another.</td>
<td>Crest-line of individual landforms</td>
</tr>
<tr>
<td>Gouged, grooved and streamlined bedrock</td>
<td>4000</td>
<td>Elongate, parallel or roughly parallel ridge/groove sets incised into bedrock.</td>
<td>Baseline of individual landforms</td>
</tr>
<tr>
<td>Grounding zone wedge (GZW)</td>
<td>12</td>
<td>Wedge-shaped landform formed in glacigenic sediment, with a steep distal sea-floor ramp and shallow back-slope. The back-slope is often lineated. Typically tens of km’s wide-long and tens of m’s high.</td>
<td>Crest-line of scarp slope</td>
</tr>
<tr>
<td>Subglacial meltwater channels</td>
<td>844</td>
<td>Straight to sinuous channels predominantly cut into the crystalline bedrock (although also rarely observed in sediment); typically form anastomosing networks that connect with basins or as smaller isolated channels that cut across bedrock knolls. Abrupt initiation and termination points and undulating thalwegs.</td>
<td>Break-of slope outline of individual channels or baseline of individual landforms; too narrow to accurately delineate the break of slope.</td>
</tr>
<tr>
<td>Gullies</td>
<td>181</td>
<td>Straight to sinuous channels formed in sediment on the continental slope, often with dendritic patterns that converge down-slope. Typically initiating at the shelf edge, gullies are sometimes eroded back into the shelf, thereby forming ‘amphitheatre’ or ‘cauliflower’ shaped heads.</td>
<td>Baseline of individual landforms</td>
</tr>
</tbody>
</table>
Vertical and horizontal uncertainties are about 1 and 5 m, respectively. Solar relief-shaded visualisations of the compiled grid were used to conduct the mapping. Two orthogonally shaded images were used to avoid azimuth bias (e.g. Smith & Clark, 2005) and the images were vertically exaggerated 20–50 times to aid identification. Glacial bedforms along Marguerite Trough were manually identified and systematically mapped in ArcGIS 9.3. The bedforms were identified according to conventional criteria (e.g. Graham et al., 2009) listed in Table 1 and discussed below.

4. Glacial landforms

4.1 Mega-scale glacial lineations (MSGLs)

MSGLs are highly attenuated parallel sets of grooves and ridges formed in till and commonly observed on the beds of Antarctic marine palaeo-ice streams (e.g. Canals, Urgeles, & Calafat, 2000; Evans et al., 2005; Graham et al., 2009; Shipp, Anderson, & Domack, 1999). Over 5000 MSGLs have been mapped along the length of MBIS, extending all the way to the shelf break, where they are increasingly disrupted by iceberg ploughmarks (Figure 2). MSGLs are confined to regions of weak till (e.g. Ó Cofaigh et al., 2007) and are 100–18,000 m long and 100–600 m wide, with amplitudes and wavelengths typically between 2 and 8 m and 100–700 m, respectively. The longest MSGLs are observed along the central axis of the outer-shelf trough, with the bedrock-dominated inner- and mid-shelf restricted to shorter bedforms. Although
subtle shifts in MSGL orientation are observed along the length of MBIS, often associated with grounding-zone wedge (GZW) positions, they generally conform to the orientation of the trough.

4.2 Crag-and-tails
Crag-and-tails are tear-drop shaped bedforms with exposed bedrock at the stoss end and a tail of tapering sediment (Graham et al., 2009). The tail of sediment relates to the protection or inflow of sediment in the lee of the bedrock crag. These features are restricted to areas of bedrock with an upstream supply of sediment, either as isolated features or broad swaths in the lee of escarpments. Over 450 crag-and-tail bedforms have been mapped along MBIS, predominantly on the mid-shelf where there is a mix of bedrock and unconsolidated sediment (Figure 3). They display similar characteristics to neighbouring MSGLs, including orientation and size.

4.3 Grooved, gouged and streamlined bedrock
The crystalline bedrock of the inner and mid-shelf has been preferentially eroded parallel to palaeo-ice flow, creating a grooved and streamlined appearance (Figure 4), similar to other regions around Antarctica (cf. Livingstone et al. 2012; Wellner, Heroy, & Anderson, 2006). Grooves, gouges and streamlined ridges tend to be concentrated along the axis of the trough, and across highs separating

Figure 3. Example of crag-and-tails in a large isolated basin on the inner shelf. The left-hand panel is the relief-shaded image and the right-hand panel shows the mapped landforms (colours are the same as in the main map). The relief-shaded image is x20 exaggeration and is shaded from the NE.
and surrounding basins (Figure 4), while the trough flanks and the rough interior of Marguerite Bay are crudely streamlined. The features are typically 10 s of metres deep/high, although streamlined hills up to 200 m high are also observed, albeit rarely. In contrast to MSGLs, the grooves and gouges are often curvilinear, probably related to the structure of the underlying bedrock.

4.4 Whalebacks

Whalebacks are ‘tear-drop’ or ‘oval’ shaped hills formed in bedrock and can be described as symmetric or classically asymmetric (Table 1, e.g. Roberts & Long, 2005). They are commonly found on the inner- and mid-shelf of MBIS in crystalline bedrock, and range from crudely sculpted symmetrical features <3 km long, with elongation ratios of 2:1–4:1, to large asymmetric whalebacks up to 10 km long and with elongation ratios up to 18:1 (Figure 5). In places, whalebacks are associated with crescentic scours around their stoss ends. A clear evolution is apparent between the stubbier, crudely streamlined forms on the inner shelf, especially outside of the main trough, and the more elongate, increasingly asymmetric whalebacks observed towards the mid-shelf and on the floors of the large isolated and connected basins on the rugged inner shelf. Nearly 3000 whalebacks have been mapped along MBIS, whereas drumlins formed in unconsolidated sediment are entirely absent.
4.5 **Meltwater channels**

An extensive network of subglacial meltwater channels incised into crystalline bedrock has been identified on the inner- and mid-shelf sections of MBIS (Figure 6, e.g. Anderson & Oakes-Fretwell, 2008). The channels range from straight to sinuous; have U- to V-shaped cross-profiles; abrupt initiation and termination points; undulating thalwegs; and orientations that vary from parallel to oblique to the palaeo-ice flow direction. Meltwater channels in the rugged eastern interior of Marguerite Bay are small isolated features that typically extend into large basins from bedrock highs and dissect bedrock knolls. These channels evolve westward into a coherent, anastomosing network of large meltwater channels, 300–550 m wide and 30–200 m deep, which connect deep basins and converge towards the main trough. The trough is also characterised by an anastomosing network of well-connected channels that are up to 1.5 km wide and 100 m deep (Figure 6). The meltwater channels become more disparate seawards of the transition from bedrock to sedimentary substrate, and are not observed on the sediment-floored outer-shelf.

4.6 **Grounding zone wedges (GZWs)**

GZWs are common features of Antarctic palaeo-ice streams (e.g. Graham et al., 2009; Larter & Vanneste, 1995; Ó Cofaigh et al., 2005; Shipp et al., 1999), and are characterised by a steep distal
sea-floor ramp and shallow, commonly lineated back-slope (Figure 7). They are composed of diamicton tens of metres thick and extend for tens of kilometres (Livingstone et al., 2012). They are thought to form by the advection and deposition of sediment at the grounding-line during temporary still-stands or minor readvances (e.g. Alley, Blankenship, Rooney, & Bentley, 1989), although the exact process for sediment advection is debated (see Livingstone et al., 2012). Nine GZWs have been identified on the mid and outer shelf of MBIS, occurring both along the central axis (e.g. Figure 7) and flanks of the trough (see also Jamieson et al., 2012). Three additional GZWs have been mapped outside the confines of the main trough close to the shelf edge. These GZWs are associated with the partial preservation of MSGLs in front of their scarps. The GZWs are typically 10–40 m thick (maximum 80 m), 5–14 km wide and 3–14 km long.

4.7 Gullies and channels

Gullies are preserved on the uppermost continental slope in front of and to either side of the trough mouth (Figure 8, also see Dowdeswell et al., 2004a; Noormets, Dowdeswell, Larter, Ó Cofaigh, & Evans, 2009; Ó Cofaigh, Taylor, Dowdeswell, & Pudsey, 2003). They are characteristically straight to sinuous channels formed in sediment, often with a dendritic pattern converging down-slope. The gullies rarely extend more than 10 km down-slope and display a V-shaped
profile. In front of the trough, the gullies display wavelengths between 300 and 1500 m and are typically 120 m deep. Gullies on either side of the trough mouth are typically initiated at the shelf edge, but those offshore of the trough mouth are frequently seeded further down-slope. The gullied upper slope gives way to a smooth lower slope and then into large channels and sediment drifts on the upper continental rise (Figure 8).

4.8 Iceberg scours

Iceberg scours are present on the sediment-floored mid- and outer-shelf forming cross-cutting curvilinear furrows that are superimposed onto all the other mapped glacial features (Figure 9). They occur in water depths of 300–750 m, with a mean of 480 m, and become more prevalent towards the outer shelf, where the bed shallows. Although the iceberg scours display multiple orientations, a general SW-NE trend is apparent.

5. Discussion

The map depicts over 17,000 glacial landforms in Marguerite Bay, which record the extent of a significant (>370 km long) palaeo-ice stream along the trough that was fed by ice from George
VI Sound and surrounding tributary regions on the inner shelf (cf. Ó Cofaigh et al., 2002, 2005). The distribution of MSGLs indicates that ice flow along the central axis of the over-deepened trough was rapid and that ice reached the shelf edge at some point during the last glacial. The identification of twelve GZWs on the landward-sloped outer- and mid-shelf documents the advection and build-up of glacigenic sediment during temporary pauses in the ice stream’s recession. These GZWs are within a region thought to have undergone rapid (millennial-scale) retreat from the outer-shelf (Kilfeather et al., 2011) and must, therefore, have formed rapidly.

The spatial distribution and evolution of glacial bedforms along the MBIS is typical of Antarctic palaeo-ice streams (e.g. Camerlenghi et al., 2001; Canals et al., 2000; Evans et al., 2005; Graham et al., 2009; Jakobsson et al., 2012; Wellner et al., 2006). The rough bedrock-dominated inner-shelf region is characterised by a wide variety and spatial arrangement of bedforms, including patches of lineated sediment and a deeply incised anastomosing network of meltwater channels, which converge on the main trough. The highly variable size, elongation and depth of bedrock incisions imply a complex ice-flow history conditioned by the bedrock properties. The bedforms associated with bedrock (channels, whalebacks, and gouged, grooved and streamlined bedrock) were probably eroded over several glacial cycles (and different stages of ice-stream activity) and therefore reflect a long history of seascape evolution heavily influenced by the bedrock structure and geology (e.g. Anderson & Oakes-Fretwell, 2008; Canals et al., 2000;
Graham et al., 2009; Nitsche et al., 2013; Wellner et al., 2006). Conversely, the bedforms in glacial sediments (crag-and-tail, GZWs and MSGLs) are likely to have formed during the most recent phase of ice streaming (e.g. Graham et al., 2009; O’Cofaigh et al., 2005).

The outer shelf is dominated by highly elongate MSGLs formed in till (e.g. O’Cofaigh et al., 2007), which probably relate to the final imprint of palaeo-ice stream activity. Although there is considerable variation in MSGL elongation, it generally increases downstream and towards the centre of the trough. Unlike the bedrock-floored inner- and mid-shelf there is no evidence of meltwater channels or related fluvial sediments on the sedimentary substrate of the outer-shelf (e.g. Kilfeather et al., 2011; O’Cofaigh et al., 2007). Bedrock channels on the inner shelf are interpreted to have formed over many glacial cycles (e.g. Anderson & Oakes-Fretwell, 2008) and thus reflect an entirely separate temporal signature of subglacial water production and flow. It is therefore not surprising that similar features are not documented on the sediment-floored outer shelf. However, it is unknown how subglacial meltwater drained through this sector of the palaeo ice-stream. Possible mechanisms include: (i) water flow through meltwater channels or canals below the resolution of the instruments; (ii), flow as part of a dynamic network, where channels were constantly evolving due to the ‘competition’ between water flux and sediment creep; or (iii), via Darcian flow through the sediment (see Noormets et al., 2009).

Figure 9. Example of iceberg scours on the outer-shelf. Note that the lineations are increasingly disrupted by iceberg scours and eventually disappear in a westward direction. The left-hand panel is the relief-shaded image and the right-hand panel shows the mapped landforms (colours are the same as in the main map). The relief-shaded image is x20 exaggeration and is shaded from the NE.
6. Conclusions and implications

This paper presents a detailed geomorphological map of the glacial landforms along the MBIS. It shows a complex landform arrangement characterised by a progression from bedrock moulded wha-""


