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**Article:**

Perez, JPH, Tobler, DJ, Thomas, AN et al. (4 more authors) (2019) Adsorption and Reduction of Arsenate during the Fe<sup>2+</sup>- Induced Transformation of Ferrihydrite. ACS Earth and Space Chemistry, 3 (6). pp. 884-894. ISSN 2472-3452

<https://doi.org/10.1021/acsearthspacechem.9b00031>

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# Adsorption and reduction of arsenate during the Fe<sup>2+</sup>- induced transformation of ferrihydrite

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**KEYWORDS:** arsenic, ferrihydrite, goethite, green rust, mineral transformation, XAS, XPS

## ABSTRACT

Iron (oxyhydr)oxides play an important role in controlling the mobility and toxicity of arsenic (As) in contaminated soils and groundwaters. Dynamic changes in subsurface geochemical conditions can impact As sequestration and remobilization since the fate of As is highly dependent on the dominant iron mineral phases present and, specifically, the pathways through which these form or transform. To assess the fate of arsenate [As(V)] in subsurface settings, we have investigated the Fe<sup>2+</sup>-induced transformation of As(V)-bearing ferrihydrite (As(V)-FH) to more crystalline phases under environmentally relevant anoxic subsurface conditions. Specifically, we examined the influence of varying Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratios (0.5, 1, 2) on the behavior and speciation of mineral-bound As species during the transformation of As(V)-FH to crystalline iron-bearing phases at circum-neutral pH conditions. At all Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratios, goethite (GT), green rust sulfate (GR<sub>SO4</sub>) and lepidocrocite (LP) formed within the first 2 h of reaction. At low ratios (0.5 to 1), initially formed GR<sub>SO4</sub> and/or LP dissolved as the reaction progressed, and only GT and some unreacted FH remained after 24 h. At Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratio of 2, GR<sub>SO4</sub> remained stable throughout the 24 h of reaction, alongside GT and unreacted As(V)-FH. Despite the fact that majority of the starting As(V)-FH transformed to other phases, the initially adsorbed As was not released into solution during the transformation reactions and ~99.9% of it remained mineral-bound. Nevertheless, the initial As(V) became partially reduced to As(III), most likely because of the surface-associated Fe<sup>2+</sup>-GT redox couple. The extent of As(V) reduction increased from ~34% to ~40%, as the Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratio increased from 0.5 to 2. Overall, our results provide important insights into transformation pathways of iron (oxyhydr)oxide minerals in As contaminated, anoxic soils and sediments, and demonstrate the impact that such transformations can have on As mobility and also importantly oxidation state and, hence, toxicity in these environments.

## INTRODUCTION

Ferrihydrite (FH) is a nanoparticulate ferric oxyhydroxide mineral commonly found in natural and engineered environments (e.g., soils, groundwater, acid mine drainage and acid sulfate soils).<sup>1,2</sup> FH

can sequester considerable amounts of trace or toxic elements via adsorption or co-precipitation due to its high specific surface area (from 120 to 850 m<sup>2</sup> g<sup>-1</sup>) and reactivity.<sup>3-8</sup> However, FH is thermodynamically metastable and usually transforms to more crystalline iron (oxyhydr)oxides (e.g., goethite, hematite, lepidocrocite, green rust or magnetite),<sup>1</sup> whereby any adsorbed or incorporated compounds can be remobilized and re-distributed. FH transformation in oxic, ambient conditions and at circum-neutral pH is very slow (months to years),<sup>9</sup> and the rates, mechanisms and pathways of transformation strongly depend on physico-chemical factors including pH,<sup>10-12</sup> temperature,<sup>11,12</sup> and the presence of inorganic ions<sup>4,12-14</sup> and organic ligands.<sup>15,16</sup>

In anoxic and non-sulfidic environments, FH transformations can occur more rapidly (within hours or days) due to the presence of aqueous ferrous iron (Fe<sup>2+</sup><sub>(aq)</sub>),<sup>17-20</sup> generated by dissimilatory iron-reducing bacteria.<sup>21,22</sup> FH transformation usually starts by an initial adsorption of aqueous Fe<sup>2+</sup> onto FH surface sites and the oxidation of this surface-bound Fe(II) to surface Fe(III) species by loss of an electron to the FH solid. This electron is then conducted through the FH and eventually leads to a release of Fe<sup>2+</sup><sub>(aq)</sub>.<sup>23-26</sup> This electron conduction process creates “reactive” surface sites, which in turn initiates the dissolution of FH and recrystallization to goethite (GT) and/or lepidocrocite (LP).<sup>17,27</sup> If the aqueous Fe<sup>2+</sup> is in excess compared to the solid Fe(III) (oxyhydr)oxide (e.g., FH, GT and LP), they can transform to mixed-valent Fe minerals such as green rust (GR) and magnetite (MGT).<sup>28-30</sup> Hence, FH transformations can lead to a variety of Fe mineral phases, and each of these phases has different sorption and redox properties. In turn, this will affect biogeochemical cycling of iron and nutrients,<sup>21</sup> and importantly also the sequestration of FH-bound toxic elements.

Arsenic is a persistent contaminant affecting groundwater resources worldwide due to its widespread occurrence and distribution.<sup>31,32</sup> Its mobility in the environment can be greatly influenced by its interaction with mineral phases such as iron (oxyhydr)oxides, which have been shown to be highly effective substrates for the sequestration of As in contaminated groundwater. However, the adsorption capacity of iron (oxyhydr)oxides vary dramatically and is also strongly affected by the As oxidation state, which can quickly change during Fe redox transformations. Among the various iron (oxyhydr)oxides, FH,

which is often the first Fe phase forming in subsurface near-neutral environments, exhibits one of the highest adsorption affinity for both As(III) and As(V) while most crystalline Fe phases have far lower As adsorption affinities.<sup>33-38</sup> Under anoxic conditions and in the presence of  $\text{Fe}^{2+}_{(\text{aq})}$ , FH readily transforms to crystalline Fe phases and this can be accompanied by the release and remobilization of As back into the aqueous phase or the As can become associated with the newly-formed Fe phases. However, the mechanisms and pathways of these processes during the inter-transformation of the various iron (oxyhydr)oxides is, however, so far poorly understood or quantified.

To the best of our knowledge, only a few studies examined the  $\text{Fe}^{2+}$ -induced transformation of As-bearing FH under anoxic conditions. Pedersen et al.<sup>39</sup> used  $^{55}\text{Fe}$  and  $^{73}\text{As}$  radiotracers to monitor the transformation of As(V)-co-precipitated FH at pH 6.5 and at varying  $\text{Fe}^{2+}_{(\text{aq})}$  concentrations (0 to 1 mM), an  $\text{Fe(III)}_{\text{FH}}$  loading of 0.5 mM, and As/ $\text{Fe}_{\text{solid}}$  ratios between 0.001 to 0.005. They showed that after 5 days, LP and GT formed at low  $[\text{Fe}^{2+}_{(\text{aq})}]$ , while GT and MGT formed at higher  $[\text{Fe}^{2+}_{(\text{aq})}]$ . They also inferred that the co-precipitated As had little to no effect on the FH transformation rates and that most of the As remained associated with the solids. More recently, Masue-Slowey et al.<sup>40</sup> investigated the  $\text{Fe}^{2+}$ -induced transformation of As(V)-adsorbed FH. They used higher As/ $\text{Fe}_{\text{solid}}$  ratios (0.013 to 0.05), higher  $\text{Fe(III)}_{\text{FH}}$  loadings (20 mM), and also up to 2 mM of  $\text{Fe}^{2+}_{(\text{aq})}$  concentration. They showed that LP and MGT formed instead of GT, and that the pre-adsorbed As retarded FH transformation. These studies have provided insights into the mineralogical changes that occur when As-bearing FH is reacted with varying  $[\text{Fe}^{2+}_{(\text{aq})}]$  and revealed how the transformation rates can be affected by the presence of As. However, the fate, bonding environment, or redox state of the co-precipitated or adsorbed As during the crystallizations remains elusive. The question whether transformation reactions in systems where higher amounts of As are associated with the initial FH will cause As release, and what happens if As is only adsorbed to FH rather than co-precipitated are still open. Moreover, As oxidation state could be affected by these redox reactions, and this would affect the toxicity of As in the subsurface. Lastly, the previously tested conditions do not favor GR formation; however, GR phases may be a key substrate for As sequestration in Fe-rich and oxygen-poor subsurface environments (e.g., gley soils or contaminated aquifers),

particularly as they can adsorb large amounts of As.<sup>33,41</sup> Thus, GR formation, stability and behavior with respect to As has to be evaluated.

Herein, we aim to fill a part of this knowledge gap by describing a study in which we performed batch experiments under anoxic conditions and examined the Fe<sup>2+</sup>-induced transformation of As(V)-bearing FH. Experiments were carried out at pH 6.5 with FH onto which As(V) was adsorbed, and was subsequently reacted at varying Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratios for up to 24 h. In particular, we tested Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> concentrations and ratios that were higher than in the above mentioned studies but that have been shown to favor the formation of GR.<sup>29,41</sup> The mineralogical transformations of As(V)-FH and the fate of As in these processes were assessed using conventional laboratory and synchrotron-based X-ray scattering and spectroscopic techniques and the resulting products were imaged using electron microscopy. Our results provide new insights on the influence of iron (oxyhydr)oxide mineral transformations on the speciation and hence mobility and toxicity of As in contaminated subsurface environments.

## EXPERIMENTAL SECTION

### General methods

All glass- and plastic-wares were cleaned in 5 M HCl for 24 h, followed by thorough rinsing with Milli-Q water (~18.2 MΩ·cm). All chemicals were ACS reagent grade from Sigma-Aldrich and Acros Organics and were used as received. Stock solutions were prepared inside the anaerobic chamber (97% N<sub>2</sub>, 3% H<sub>2</sub>, Coy Laboratory Products, Inc.) using O<sub>2</sub>-free water, which was obtained by purging Milli-Q water with O<sub>2</sub>-free nitrogen for at least 4 h.

### Synthesis of 2-line FH

Two-line FH was synthesized using the method described by Schwertmann and Cornell<sup>42</sup> by slowly titrating 0.1 M Fe<sub>2</sub>(SO<sub>4</sub>)<sub>3</sub>·5H<sub>2</sub>O with 1 M NaOH to pH ~7. The resulting suspension was washed using 6 cycles of centrifugation (9,000 rpm, 5 minutes) and re-dispersion in Milli-Q water to remove

excess solutes. Afterwards, the FH slurry was purged with O<sub>2</sub>-free N<sub>2</sub> for at least 4 h to remove O<sub>2</sub> and then immediately transferred into the anaerobic chamber. The amount of synthesized FH was determined based on the total iron concentration of an aliquot of the suspension dissolved in 0.3 M HNO<sub>3</sub>. The total Fe concentration was analyzed by flame atomic absorption spectrometry (AAS, Perkin Elmer AAS Analyst 800). Each batch of FH (~88.3 mM Fe(III)<sub>solid</sub>) was prepared fresh and used on the day of synthesis.

### **Batch transformation experiments**

All batch experiments were performed in triplicate at room temperature inside the anaerobic chamber. To prepare As(V)-bearing FH, an aliquot of the washed FH was re-suspended in a 0.1 M NaCl solution buffered at pH 6.5 using 0.05 M morpholinoethanesulfonic acid (MOPS). After pH equilibration, the resulting FH suspension was then spiked with an aliquot from an As(V) stock solution prepared from Na<sub>2</sub>HAsO<sub>4</sub>·7H<sub>2</sub>O. The resulting suspensions [41.6 mM Fe(III)<sub>solid</sub>, 1.33 mM As(V)] were stirred at 350 rpm for 24 h to ensure As(V) adsorption onto FH (Figure S-1). Afterwards, aliquots of 0.5 M FeSO<sub>4</sub> were added to the As(V)-bearing FH suspension to achieve Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratios of 0.5, 1 or 2 (denoted as R0.5, R1 and R2 from here on). A control experiment without FeSO<sub>4</sub> addition (no aqueous Fe<sup>2+</sup>, R0) was also conducted. The resulting mixtures were stirred at 350 rpm for 24 h, with aliquots of the suspension being removed after 0.5, 1, 2, 4, 8 and 24 h. Parts of the collected suspensions were filtered through 0.22- $\mu$ m syringe filters, and the resulting solutions were acidified with HNO<sub>3</sub> and stored at 4°C until the concentrations of aqueous As were analyzed by inductively coupled plasma optical emission spectrometry (ICP-OES, Varian 720ES), following the method described by Perez et al.<sup>33</sup> Further analytical details can be found in the Supporting Information (Text S-1, Table S-1). The remainder of the collected suspensions was used to characterize the solid phase. For this, the suspension was filtered using 0.22- $\mu$ m polycarbonate membrane filters and the obtained solids were dried in a desiccator inside the chamber, ground and stored until use in crimped headspace vials inside the anaerobic chamber.

## Mineral characterization and thermodynamic modelling

The solids were analyzed by a suite of laboratory- and synchrotron-based characterization techniques to determine their structure and composition, particle sizes and morphologies, surface properties, as well as As and Fe redox states. Detailed information on sample preparation to minimize oxidation and on solid characterization can be found in the Supporting Information (Text S-2). Mineralogical changes in the solid phase during the reaction were monitored by X-ray powder diffraction (XRD) using a Bruker D8 powder diffractometer (Cu K $\alpha$  radiation,  $\lambda = 1.5406 \text{ \AA}$ ). The morphology, size, structure and chemical composition of the final solids (collected after 24 h) were characterized by transmission electron microscopy (TEM) and scanning electron microscopy (SEM). TEM micrographs and selected area electron diffraction (SAED) patterns were recorded using a FEI Tecnai G2 F20 X-Twin FEG TEM, operated at 200 keV and equipped with a Gatan Imaging Filter (GIF) Tridiem<sup>TM</sup>. SEM images were acquired using a ZEISS Ultra Plus FE-SEM operated in high vacuum mode at an acceleration voltage of 3 kV with 10  $\mu\text{m}$  aperture size using an InLens secondary electron detector. The local structure was investigated using pair distribution function (PDF) analysis. The high energy X-ray scattering data used for PDF analysis were collected at the 11-ID-B beamline of the Advanced Photon Source (Argonne National Laboratory, USA). X-ray absorption spectroscopic (XAS) analyses were carried out to monitor the changes in As oxidation state and to quantify the Fe phases in the final solids. Fe K-edge extended X-ray absorption fine structure (EXAFS) spectra were collected at the SUL-X beamline of Angströmquelle Karlsruhe (ANKA, Karlsruhe, Germany), and the As K-edge X-ray absorption near-edge structure (XANES) data were collected at the BM23 beamline of the European Synchrotron Radiation Facility (ESRF, Grenoble, France). The Fe K-edge EXAFS spectra of synthetic iron (oxyhydr)oxide mineral samples [i.e., FH,<sup>42</sup> GT,<sup>42,43</sup> LP,<sup>42</sup> GR sulfate (GR<sub>SO4</sub>)<sup>33</sup>] were also collected as reference standards for Fe phase quantification. As K-edge XANES spectra of As(III)- and As(V)-interacted GT samples were also collected and were used as reference standards for the determination of As oxidation state. X-ray photoelectron spectroscopy (XPS) measurements were performed using a KRATOS Axis Ultra DLD to determine the surface chemistry of the solids. To predict Fe and As speciation and Fe phase stability in



the studied system, thermodynamic modelling was carried out using Geochemist's Workbench® (GWB)<sup>44</sup> with the MINTEQ thermodynamic database (see Supporting Information Text S-2 for details). Missing thermodynamic data of mineral phases in the Fe-S-H<sub>2</sub>O system (e.g., GR<sub>SO4</sub>) were manually added to the MINTEQ database.<sup>45,46</sup>

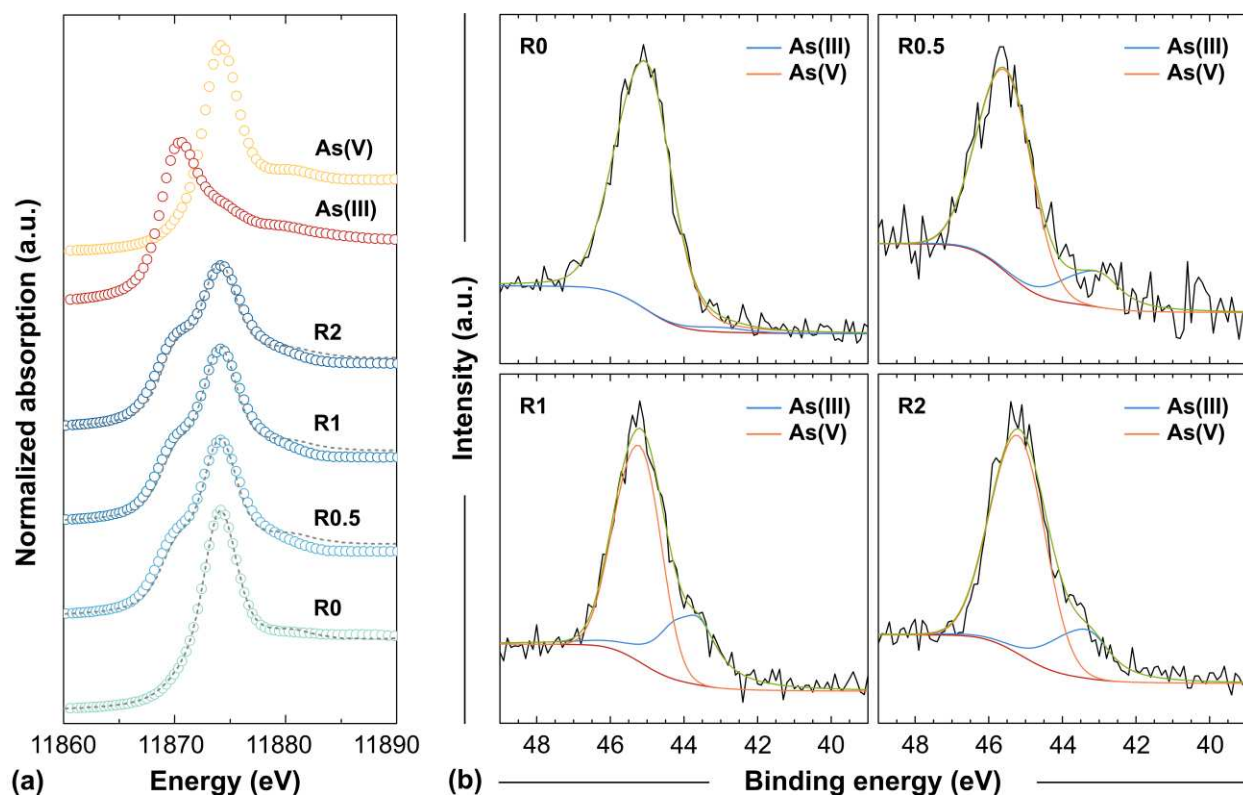
## RESULTS AND DISCUSSION

### Aqueous behavior and speciation of mineral-bound As species

The aqueous concentrations of As revealed that barely any As was released (< 0.15 %, Figure S-1) during the Fe<sup>2+</sup>-induced transformation of As(V)-bearing FH to GT ± GR. This is consistent with the high uptake capacity of synthetic iron (oxyhydr)oxides for As species determined in the adsorption experiments (Figure S-2). Similar minimal As release (<1%) were also reported in previous Fe<sup>2+</sup>-catalyzed transformation experiments of As(V)-bearing ferrhydrite<sup>39,47</sup> and As(V)/Sb(V)-bearing jarosite.<sup>45,46</sup> Moreover, it has also been shown that As removal efficiencies were even higher in experiments wherein As were co-precipitated with iron (oxyhydr)oxides compared to those adsorbed onto pre-synthesized iron (oxyhydr)oxides.<sup>39,48,49</sup>

The oxidation state of As associated with the solids after 24 h of reaction as probed by As K-edge XANES (Figure 1a) showed that the initial FH-bound As(V) was partially reduced to As(III) when the initial As(V)-FH was reacted with aqueous Fe<sup>2+</sup> under anoxic conditions. The degree of As(V) reduction slightly increased from 33.6 ± 1.8% to 42.4 ± 1.8% as the Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratio increased from 0.5 to 2 (see Table 1). This trend was also confirmed by high resolution XPS of the final solids (Figure 1b), that showed the presence of a shoulder at a binding energy of ~44 eV, indicative of As(III) (see Table S-3 for As reference binding energies). Due to the uncertainty of the XPS measurements (see Table S-4), a fully quantitative determination of the As(III) contents was difficult, but XPS confirmed its presence. Furthermore, XPS analyses showed that the initial As (V) was still the primary valence state in the near surface region (top 10 nm of the samples). A possible reduction of As(V) due to X-ray beam damage is negligible as shown by analysis of the control (R0).

Thermodynamic calculations based on the Eh-pH conditions used in our experiments (Figure S-3) suggest that, at equilibrium, all initially adsorbed As(V) species should have been reduced to As(III) during the transformation. The partial reduction of As(V) to As(III) after 24 h observed in our data is likely a result of kinetic limitations since it might take longer time scales for full reduction.



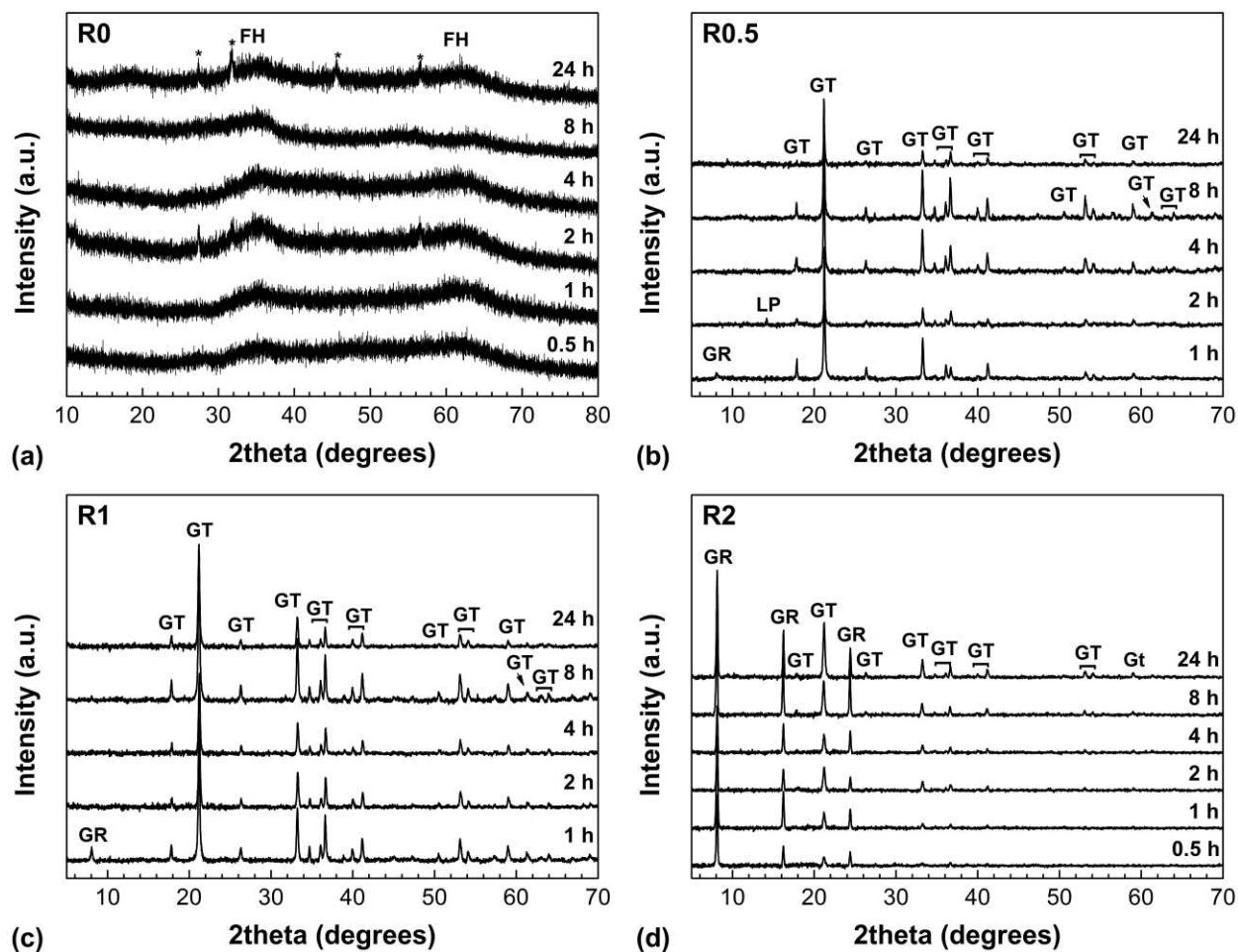
**Figure 1.** (a) Normalized As K-edge XANES spectra of the end-products. Fits (grey dashed lines) are linear combinations of the As reference standards (i.e., As(III) and As(V) adsorbed on GT). (b) Deconvoluted high-resolution As 3d XPS spectra of the end-products (calibrated to yield adventitious C 1s peak at 285.0 eV). Details of the fitting parameters and statistics for the quantification of As speciation based from the As K-edge XANES and XPS data can be found in the Tables S-2 and S-4, respectively.

**Table 1.** Arsenic oxidation state and mineralogical composition of the end-products of As(V)-bearing FH transformation with varying  $\text{Fe}^{2+}_{(\text{aq})}/\text{Fe(III)}_{\text{solid}}$  ratios (R).

Ratio	As oxidation state			Fe phase composition							
	As K-edge XANES			Fe-K edge EXAFS				PDF			Goodness of fit ( $R_w$ )
	As(III)	As(V)	Red. $\chi^2$	FH	GT	GR	Red. $\chi^2$	FH	GT	GR	
0	4.1 ± 0.1	95.9 ± 0.1	0.001	100	-	-	-	100	-	-	
0.5	33.6 ± 1.8	66.4 ± 1.7	0.014	17 ± 4	83 ± 3	-	0.221	70 ± 3	30 ± 1	-	0.205
1	34.3 ± 1.8	65.7 ± 1.8	0.015	15 ± 1	85 ± 1	-	2.663	22 ± 5	78 ± 3	-	0.150
2	42.4 ± 1.8	57.6 ± 1.7	0.013	11 ± 2	84 ± 2	5 ± 1	0.226	-	92 ± 3	8 ± 1	0.175

### Mineralogical transformation of As(V)-bearing FH

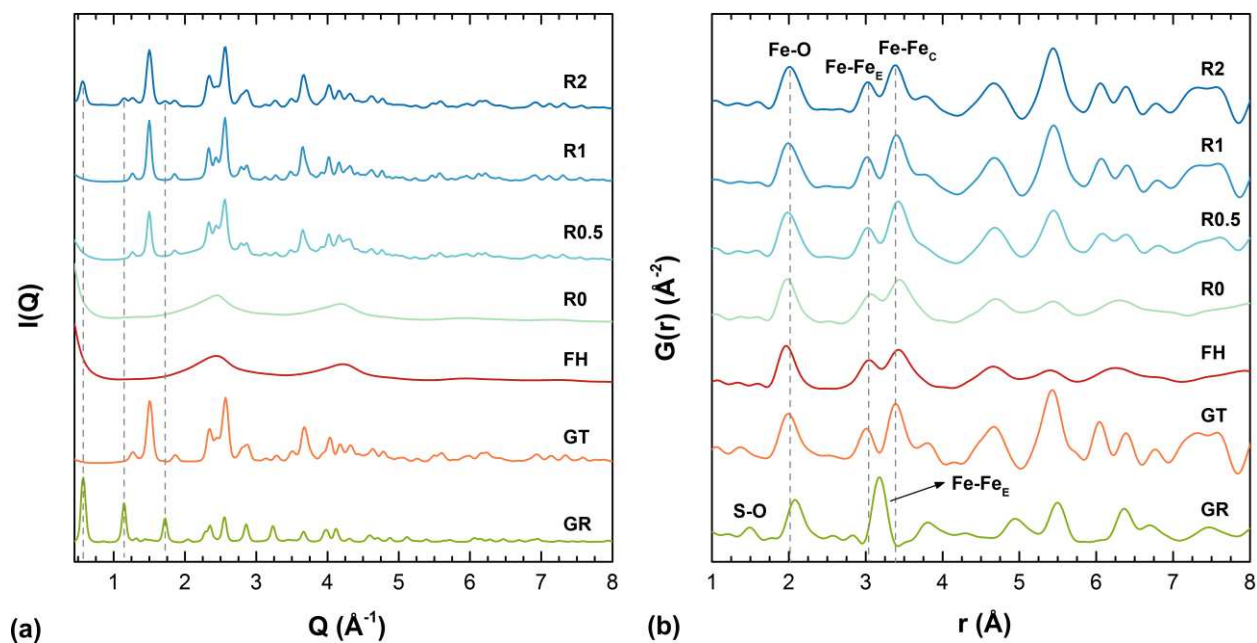
In the absence of aqueous  $\text{Fe}^{2+}$ , the As(V)-bearing FH did not transform to other iron (oxyhydr)oxides (R0, Figure 2a). While barely any As was released during the reactions, exposure of the initial As(V)-bearing FH to varying aqueous  $\text{Fe}^{2+}$  concentrations led to its rapid transformation into more crystalline iron (oxyhydr)oxides. At  $\text{Fe}^{2+}_{(\text{aq})}/\text{Fe(III)}_{\text{solid}}$  ratios of 0.5 and 1 (R0.5 and R1, Figure 2b, c, respectively), goethite (GT) formed within the first hour and dominated the pattern over the remaining 24 h. Small amounts of green rust sulfate ( $\text{GR}_{\text{SO}_4}$ ) and lepidocrocite (LP, only in R0.5) also formed in the R0.5 and R1 experiments. However, both phases dissolved, as supported by aqueous  $\text{Fe}^{2+}$  release (Figure S-1), which then precipitated as goethite after 2 h. At an  $\text{Fe}^{2+}_{(\text{aq})}/\text{Fe(III)}_{\text{solid}}$  ratio of 2 (R2, Figure 2d), both GT and  $\text{GR}_{\text{SO}_4}$  formed rapidly within the first 30 min but both also remained present throughout the 24 h of reaction.



**Figure 2.** XRD patterns showing the change in mineralogical composition in the solid samples during the 24-h transformation of As(V)-FH at varying  $\text{Fe}^{2+}_{(\text{aq})}/\text{Fe}(\text{III})_{\text{solid}}$  ratios (R): (a) control (0), (b) 0.5, (c) 1 and (d) 2. The ‘\*’ denotes peaks for halite from the background electrolyte. XRD patterns for R0.5 and R1 at 0.5 h are not shown because no crystalline mineral phases were detected. Note that the increased peak intensity of  $\text{GR}_{\text{SO}_4}$  (001) comes from preferential orientation of  $\text{GR}_{\text{SO}_4}$  plate-like particles along the [001] zone axis during XRD sample preparation.

The high energy XRD pattern [ $I(Q)$ ] of the mineral end-products (Figure 3a) corroborated the laboratory-based XRD data (Figure 2), with the main end-product being GT and with some  $\text{GR}_{\text{SO}_4}$  forming at the highest tested  $\text{Fe}^{2+}_{(\text{aq})}/\text{Fe}(\text{III})_{\text{solid}}$  ratio of 2. The broad humps at  $Q$ -values of  $\sim 2.4$  and  $\sim 4.2$   $\text{\AA}^{-1}$  for R0.5 and, in part, R1 stem from unreacted FH, the presence of which was not unexpected in the

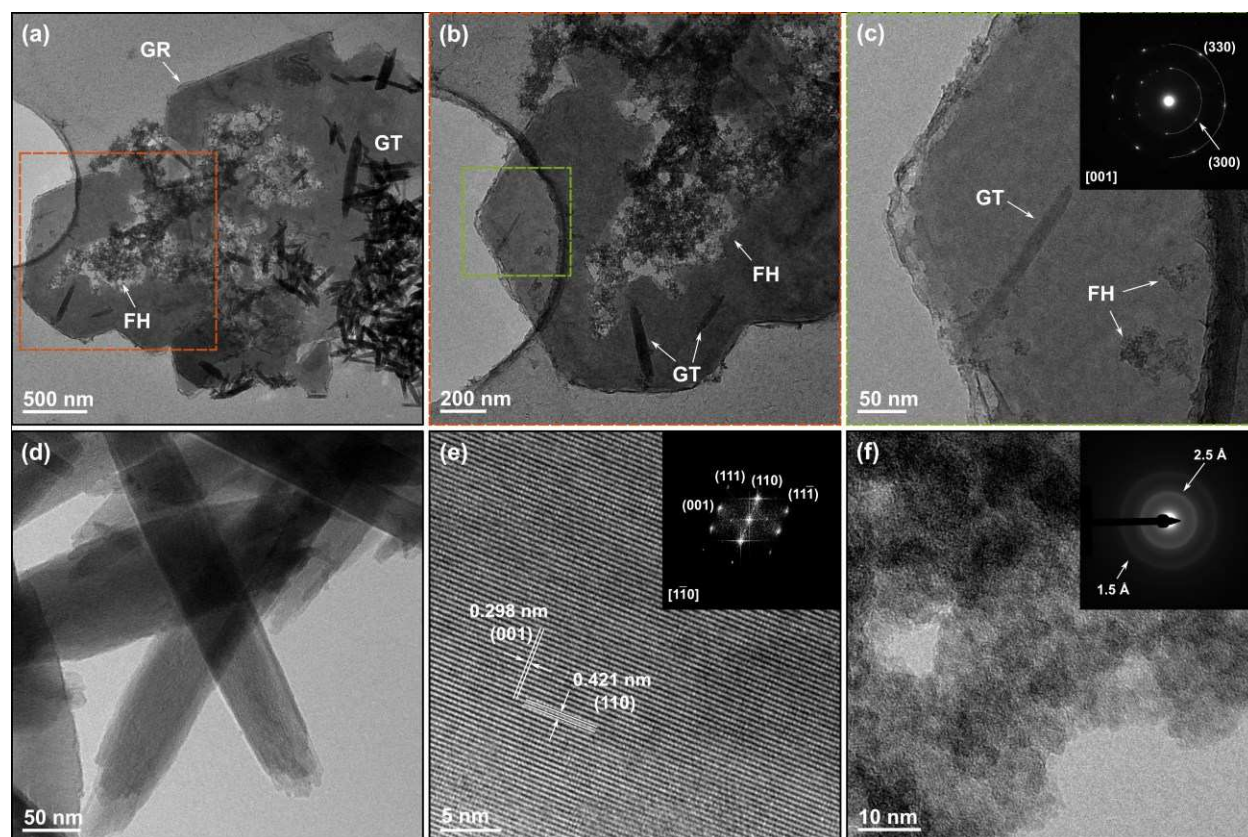
end-product material because previous studies<sup>50-52</sup> have shown that As can slow down the transformation of FH to crystalline iron (oxyhydr)oxides. The PDF analyses (Figure 3b, S-4) were used to derive the characteristic interatomic distances in the mineral-end products. The atomic pair correlations at  $r$ -values  $< 4 \text{ \AA}$  (Figure 3b) correspond to the atomic arrangements in the Fe-O polyhedra in iron (oxyhydr)oxides. The first peak at  $\sim 2.0 \text{ \AA}$  matches first neighbor Fe-O pairs, while peaks at  $\sim 3.0$  and  $\sim 3.4 \text{ \AA}$  represent edge- and corner-sharing Fe-Fe pairs (Fe-Fe<sub>E</sub>, Fe-Fe<sub>C</sub>), respectively. Changes in peak positions and intensities for these Fe-Fe pairs are a consequence of the presence of mixed iron (oxyhydr)oxides (i.e., GT, FH  $\pm$  GR<sub>SO4</sub>) in these solids, when compared with the standard materials (spectra labeled GT, FH, GR in Figure 3a and 3b).



**Figure 3.** High energy X-ray scattering data of the end-products after 24 h  $\text{Fe}^{2+}$ -induced As(V)-FH transformation at varying  $\text{Fe}^{2+}_{(\text{aq})}/\text{Fe(III)}_{\text{solid}}$  ratios (R): (a) high energy XRD patterns [ $I(Q)$ ]. GR (00l) reflections in the R2 end-product are indicated by grey dashed lines while all the other peaks in the transformation end-products can be assigned to GT (except for the R0 end-product, which is naturally still pure As(V)-bearing FH). The patterns of the reference materials (i.e., FH, GT and GR) are shown for comparison; and (b) PDFs [ $G(r)$ ] of the low  $r$ -value region showing the short-range structure of the solids.

The full PDFs are shown in Figure S-4. Fe-Fe<sub>E</sub> and Fe-Fe<sub>C</sub> refer to edge- and corner-sharing pairs, respectively.

TEM and SEM analyses of the transformation end-products confirmed that GT was the main product with FH still present in all experiments after 24 h. As shown before with XRD and PDF, GR<sub>SO4</sub> was only present in reactions with  $\text{Fe}^{2+}_{(\text{aq})}/\text{Fe(III)}_{\text{solid}} = 2$  (Figure 4a, S-5,6). GR<sub>SO4</sub> was identified by its thin hexagonal plate-like particles (Figure 4b),<sup>33,53,54</sup> GT by its distinctive crystalline nanorod (Figure 4d) and FH by its ~3 nm-sized particle aggregates (Figure 4f, S-5). SEM images of the end-products also revealed that particle lengths of the GT nanorods gradually decreased with increasing  $\text{Fe}^{2+}_{(\text{aq})}/\text{Fe(III)}_{\text{solid}}$  ratios (Figure S-6,7). Both TEM and SEM images confirmed that GT was the dominant mineral phase in all experiments (Figure 4a, S-5,6) and that FH was closely associated with GT and GR<sub>SO4</sub> (Figure 4a,b). It is important to note that, often, FH was observed to seemingly “fill” voids in GR<sub>SO4</sub> particles (Figure 4a,b and S-5c). Such features could indicate that the GR<sub>SO4</sub> particles were still forming from the As(V)-bearing FH precursor after 24 h, or that the formed GR<sub>SO4</sub> crystals are dissolving from the center, as previously suggested by Skovbjerg et al.<sup>55</sup> However, dissolution of the GR<sub>SO4</sub> from the exposed crystal edges (Figure 4c) cannot be excluded.

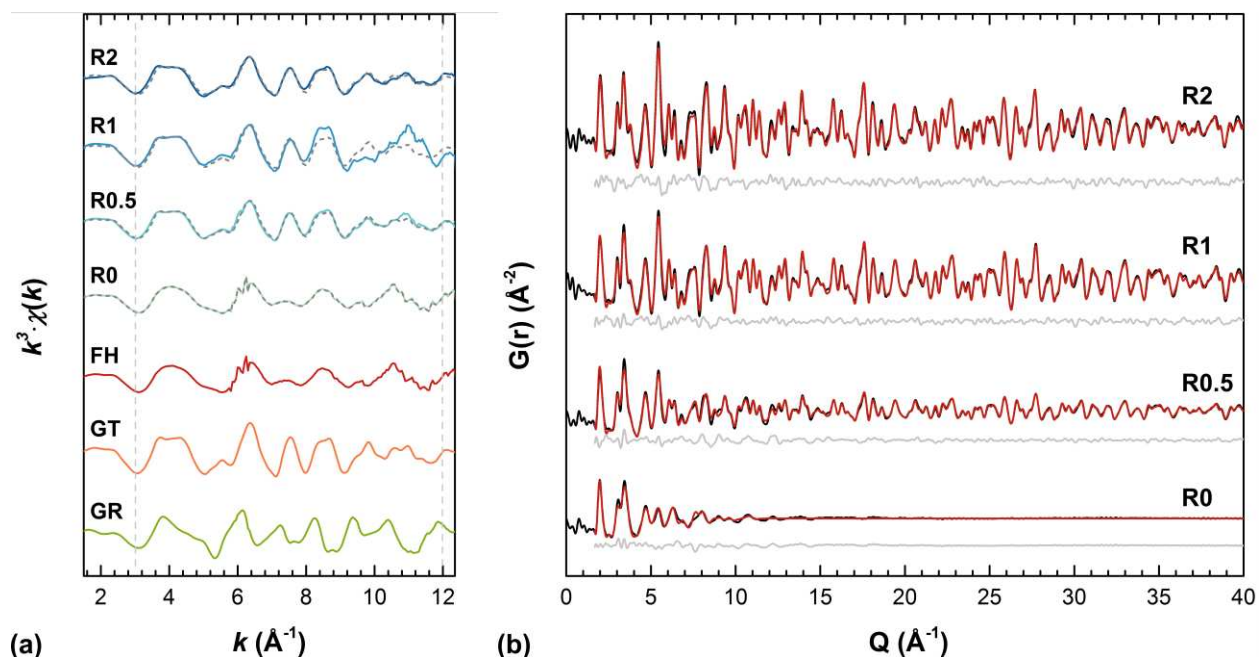


**Figure 4.** TEM images of Fe phases following the 24-h transformation of As(V)-bearing FH at  $\text{Fe}^{2+}_{(\text{aq})}/\text{Fe}(\text{III})_{\text{solid}}$  ratio of 2: (a) overview showing the close association between  $\text{GR}_{\text{SO}_4}$  (dark grey, > 300 nm wide hexagonal platelets), GT (ca. 50 nm wide black rods), and unreacted FH (aggregates of ~3 nm sized particles); (b) blow-up of the orange marked area in (a); (c)  $\text{GR}_{\text{SO}_4}$  particle seen in green marked area in (b) with the SAED pattern in inset; (d) GT nanorods and the corresponding (e) HRTEM image with the fast Fourier transformation (FFT) pattern in the inset showing the lattice fringes for (001) and (110) planes of GT (in *Pbnm* spacegroup); (f) As(V)-bearing FH nanoparticles with the SAED pattern in inset. The SAED pattern of  $\text{GR}_{\text{SO}_4}$  was indexed according to the proposed structure of Christiansen et al.<sup>56</sup>

From the evaluation of the Fe K-edge EXAFS and PDF data (Figure 5, Table 1), we determined the relative amounts (% mol Fe) of the reaction transformation end-products. The Fe K-edge EXAFS data (Figure 5a) confirmed GT ( $\geq 84\%$ ) as the main mineral phase in all  $\text{Fe}^{2+}$ -spiked experiments, with GR only accounting for ~5% in the system with  $\text{Fe}^{2+}_{(\text{aq})}/\text{Fe}(\text{III})_{\text{solid}} = 2$ . The EXAFS fitting revealed that the

amount of remnant As(V)-FH after 24 h was inversely proportional to the  $\text{Fe}^{2+}$  concentration added to the As(V)-bearing FH. Interestingly, the amount of As(V)-FH derived from the PDF data (Figure 5b) followed a similar trend to the EXAFS data, but unreacted FH could not be identified in the R2 end-product PDF pattern. This was most likely due to its low relative amount in the sample (from EXAFS  $\sim$  11%). However, the biggest difference in the relative phase amounts between PDF and EXAFS fitting was seen in the R0.5 end-product. PDF indicates  $\sim$ 70% FH compared to  $\sim$ 16% from the EXAFS evaluation, which naturally also impacted the proportion of GT in this sample. Upon closer inspection, PDF of the R0.5 end-product (Figure 3b) seems to lack the characteristic GT features observed in R1 and R2 samples. For example, the small but sharp peak at  $r \approx 3.8 \text{ \AA}$  is missing, and both the peak at  $5.5 \text{ \AA}$  and the double peaks at  $6\text{-}6.5 \text{ \AA}$  are also poorly developed. Thus, the bonding environment at  $\sim 3.8$  to  $\sim 7 \text{ \AA}$  does not exactly resemble GT. From this, we suspect that there is a short-range distortion in the Fe octahedra that is uncharacteristic of GT. This results in the large discrepancy between the PDF and EXAFS Fe phase quantification. Despite the variation between the calculated proportions of Fe phases from EXAFS and PDF data, and considering both experimental, analytical and fitting uncertainties the results show that both the extent of FH transformation to GT and/or GR increases as the  $\text{Fe}^{2+}_{(\text{aq})}/\text{Fe}(\text{III})_{\text{solid}}$  ratios increase.<sup>27</sup>





**Figure 5.** (a)  $k^3$ -weighted  $\chi(k)$  Fe K-edge EXAFS spectra of transformation end-products following the 24 h reaction of As(V)-bearing FH with varying  $\text{Fe}^{2+}_{(\text{aq})}$  concentrations ( $\text{Fe}^{2+}_{(\text{aq})}/\text{Fe(III)}_{\text{solid}}$  ratios from 0 to 2). Fits (grey dashed lines) are least square linear combinations of the reference materials (i.e., lower 3 patterns FH, GT,  $\text{GR}_{\text{SO4}}$ ). Fit boundaries are indicated by the vertical dashed lines ( $k$ -range = 3-12  $\text{\AA}^{-1}$ ). (b) Fits of PDFs of same end-products ( $\text{Fe}^{2+}_{(\text{aq})}/\text{Fe(III)}_{\text{solid}}$  ratio from 0 to 2). The black curves represent the experimental data, whereas red and light grey curves represent the calculated pattern and the residuals. Details of the fitting method for Fe K-edge EXAFS and PDF conducted in Athena<sup>57</sup> and PDFgui<sup>58</sup> can be found in Supplementary Information Text S-9.

Overall, the composition of the mineral end-products as determined with XRD, PDF, TEM, SEM and EXAFS at the end of the 24 h  $\text{Fe}^{2+}$ -induced As(V)-bearing FH transformation are consistent with each other, and also match the predicted phases from thermodynamic calculations for the Fe-S- $\text{H}_2\text{O}$  system (Figure S-8).

## Mechanism of As(V)-ferrihydrite transformation and As redox transformation

Iron redox cycling in subsurface environments highly impacts the mobility and toxicity of As in contaminated sediments and groundwaters. Specifically, mineral transformations involving iron (oxyhydr)oxides are important since specially under reducing conditions such transformation reactions can change the oxidation state of mineral-associated As, which in turn controls As toxicity as well as the extent to which As will be sorbed by minerals. Thus, such reactions may not only release As back into the environment, but these processes could render As to be present in the more toxic form.

Our results demonstrated that the initial As(V)-bearing FH rapidly transforms to GT and to a lesser extent to GR<sub>SO4</sub> and lepidocrocite upon the addition of Fe<sup>2+</sup> (Figure 2). We also showed that the transformation rate of FH increased with increasing Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratios. This is seen, for example, by the appearance of crystalline Fe phases already after 30 min in experiments with an Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratio of 2 (Figure 2d), compared to 1 h at lower ratios, or in the lower relative abundance of FH in the end-products at higher Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratios (Figure 5). Furthermore, the absence of LP at Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratios > 0.5 indicates that the transformation was very fast because LP formation requires low levels of FH-surface-adsorbed Fe<sup>2+</sup>.<sup>17,39,45</sup> Moreover, the smaller GT nanorods obtained at higher Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratios (Figure S-6) indicate faster FH transformation rates, because higher nucleation rates lead to smaller crystals.

GR<sub>SO4</sub> formed under all tested conditions alongside with GT, but disappeared already after 2 h at lower Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratios (< 2) and it transformed into the thermodynamically more stable GT (Figure 2b, 2c). At Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> = 2, GR<sub>SO4</sub> remained throughout the reaction as expected based on previous Fe<sup>2+</sup>-induced FH transformation experiments where a similar Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratio was employed without the addition of As.<sup>28,29,59</sup> However, in contrast to the As-free FH experiments, which only formed GR, the R2 end-products in the current study also contained FH and GT. Arsenic species have been shown to hinder iron (oxyhydr)oxide transformations.<sup>30,45,60</sup> Thus, the incomplete conversion of As(V)-bearing FH into GT and/or GR<sub>SO4</sub> (Figure 4 and 5) is likely a consequence of crystallite poisoning by the surface-bound As species. Specifically, As species have been shown to inhibit Fe-O-Fe

polymerization, thereby inducing distortions in the Fe bonding environment and inhibiting crystal nucleation and growth.<sup>50-52</sup>

The solid-state characterization results and electron microscopy images further suggest that GR<sub>SO4</sub> formed independently of GT during the Fe<sup>2+</sup>-induced transformation of As(V)-bearing FH. The XRD data (Figure 2) document the rapid and simultaneous occurrence of GT and GR<sub>SO4</sub> in the early stages of transformation and, thus, suggest that both Fe phases formed directly from FH. This is also supported by the calculated Gibbs free energies ( $\Delta G_{rxn}^{\circ}$ ), which showed that the formation of GR<sub>SO4</sub> is more thermodynamically favored from a FH precursor (Table 2, Eq. 2) compared to GT (Table 2, Eq. 3). Moreover, the added Fe<sup>2+</sup> rapidly hydrolyzed, as evidenced by the quick decrease in aqueous Fe<sup>2+</sup> concentration (Figure S-1a), and thus the simultaneous formation of GT and GR<sub>SO4</sub> from FH is likely. The formation of GT from FH is well documented,<sup>24,27,61</sup> while the formation pathways and mechanisms of GR phases from other iron (oxyhydr)oxides are far less studied.<sup>28,59</sup> Sumoondur et al.<sup>29</sup> however reported a similar observation wherein GR<sub>SO4</sub> formed directly from pure FH (no As added, Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratios of 0.5 to 2) within the first 10 min of the Fe<sup>2+</sup>-catalyzed transformation reaction as monitored by synchrotron-based *in situ* time-resolved energy dispersive X-ray diffraction.

During the transformation reaction of the As(V)-bearing FH a minor initial release of As (< 0.15%, Figure S-1b) from its surface was observed. The initial As release is a result of the dissolution of FH, which can have surface areas up to 850 m<sup>2</sup> g<sup>-1</sup>,<sup>3</sup> and the formation of GT and GR phases which both have lower surface areas. This released As was quickly adsorbed by the newly-formed GT and/or GR particles (Figure S-1b). However, the possibility of incorporation of As into the structure of GT cannot be ruled out, especially since the ionic radius of As(V) is similar to tetrahedrally-coordinated Fe,<sup>1</sup> although such phenomenon has not been documented yet.<sup>39</sup>

A more relevant finding of this study is that the initial As(V) was partially reduced to As(III) during the Fe<sup>2+</sup>-induced transformation of As(V)-bearing FH, and this reduction (i.e., As(III)/As(V) ratio) increased with increasing Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratio. Based on the calculated  $\Delta G_{rxn}^{\circ}$  values (Table 2, Eq. 4-7), the most thermodynamically feasible reductant in the Fe-As-S-H<sub>2</sub>O system is GR<sub>SO4</sub>, yet no study to date

has been able to document such reduction of As(V) to As(III) by GR.<sup>41,62,63</sup> Moreover, the formation and stability of the GR in the experiments R0.5 and R1 were substantially lower compared to the R2 experiment (Figure 2). This suggests that another redox couple may have induced As(V) reduction. The most likely candidate is the surface-associated Fe<sup>2+</sup> and GT redox couple (Table 2, Eq. 8-9), which has been shown to reduce other groundwater contaminants such as carbon tetrachloride,<sup>64</sup> nitrobenzene<sup>65,66</sup> and chromate.<sup>67</sup> The surface-associated Fe<sup>2+</sup>-GT redox couple might also explain why As(V) reduction was only observed at high Fe<sup>2+</sup> concentrations during the Fe<sup>2+</sup>-catalyzed transformation of As(V)/Sb(V)-jarosite ([Fe(III)]<sub>jarosite</sub> = 21.8 mM, [Fe<sup>2+</sup><sub>(aq)</sub>] = 0 to 20 mM, As/Fe<sub>solid</sub> = 0.003).<sup>45</sup> These authors noted that in their experiments, LP was the dominant mineral phase at low Fe<sup>2+</sup> concentrations while GT was the primary end-product (with minor GR<sub>SO4</sub>, <10%) at higher Fe<sup>2+</sup> concentrations.

**Table 2.** Calculated Gibbs free energies ( $\Delta G_r^\circ$ ) at 25 °C.

	Chemical reaction	$\Delta G_{rxn}^\circ$ (kJ mol <sup>-1</sup> ) <sup>a</sup>
<i>Mineral formation</i>		
1	Fe <sup>III</sup> (OH) <sub>3</sub> → $\alpha$ -Fe <sup>III</sup> OOH + H <sub>2</sub> O	-20.4
2	4Fe <sup>2+</sup> + 2Fe <sup>III</sup> (OH) <sub>3</sub> + SO <sub>4</sub> <sup>2-</sup> + 6H <sub>2</sub> O → Fe <sup>II</sup> <sub>4</sub> Fe <sup>III</sup> <sub>2</sub> (OH) <sub>12</sub> SO <sub>4</sub> + 6H <sup>+</sup>	124.2
3	4Fe <sup>2+</sup> + 2 $\alpha$ -Fe <sup>III</sup> OOH + SO <sub>4</sub> <sup>2-</sup> + 8H <sub>2</sub> O → Fe <sup>II</sup> <sub>4</sub> Fe <sup>III</sup> <sub>2</sub> (OH) <sub>12</sub> SO <sub>4</sub> + 6H <sup>+</sup>	598.4
<i>Redox reactions</i>		
4	Fe <sup>II</sup> <sub>4</sub> Fe <sup>III</sup> <sub>2</sub> (OH) <sub>12</sub> SO <sub>4</sub> + 2H <sub>2</sub> As <sup>V</sup> O <sub>4</sub> <sup>-</sup> ⇌ 6 $\alpha$ -Fe <sup>III</sup> OOH + 2As <sup>III</sup> (OH) <sub>3</sub> + SO <sub>4</sub> <sup>2-</sup> + 2H <sub>2</sub> O	-122.2
5	Fe <sup>II</sup> <sub>4</sub> Fe <sup>III</sup> <sub>2</sub> (OH) <sub>12</sub> SO <sub>4</sub> + 2HAs <sup>V</sup> O <sub>4</sub> <sup>2-</sup> + 2H <sup>+</sup> ⇌ 6 $\alpha$ -Fe <sup>III</sup> OOH + 2As <sup>III</sup> (OH) <sub>3</sub> + SO <sub>4</sub> <sup>2-</sup> + 2H <sub>2</sub> O	-202.0
6	Fe <sup>II</sup> <sub>4</sub> Fe <sup>III</sup> <sub>2</sub> (OH) <sub>12</sub> SO <sub>4</sub> + 2H <sub>2</sub> As <sup>V</sup> O <sub>4</sub> <sup>-</sup> + 4H <sub>2</sub> O ⇌ 6Fe <sup>III</sup> (OH) <sub>3</sub> + 2As <sup>III</sup> (OH) <sub>3</sub> + SO <sub>4</sub> <sup>2-</sup>	0.2
7	Fe <sup>II</sup> <sub>4</sub> Fe <sup>III</sup> <sub>2</sub> (OH) <sub>12</sub> SO <sub>4</sub> + 2HAs <sup>V</sup> O <sub>4</sub> <sup>2-</sup> + 4H <sub>2</sub> O + 2H <sup>+</sup> ⇌ 6Fe <sup>III</sup> (OH) <sub>3</sub> + 2As <sup>III</sup> (OH) <sub>3</sub> + SO <sub>4</sub> <sup>2-</sup>	-79.6
8	2Fe <sup>2+</sup> + H <sub>2</sub> As <sup>V</sup> O <sub>4</sub> <sup>-</sup> + 3H <sub>2</sub> O ⇌ 2 $\alpha$ -Fe <sup>III</sup> OOH + As <sup>III</sup> (OH) <sub>3</sub> + 3H <sup>+</sup>	21.4
9	2Fe <sup>2+</sup> + HAs <sup>V</sup> O <sub>4</sub> <sup>2-</sup> + 3H <sub>2</sub> O ⇌ 2 $\alpha$ -Fe <sup>III</sup> OOH + As <sup>III</sup> (OH) <sub>3</sub> + 2H <sup>+</sup>	-18.5
10	2Fe <sup>2+</sup> + H <sub>2</sub> As <sup>V</sup> O <sub>4</sub> <sup>-</sup> + 5H <sub>2</sub> O ⇌ 2Fe <sup>III</sup> (OH) <sub>3</sub> + As <sup>III</sup> (OH) <sub>3</sub> + 3H <sup>+</sup>	62.2
11	2Fe <sup>2+</sup> + HAs <sup>V</sup> O <sub>4</sub> <sup>2-</sup> + 3H <sub>2</sub> O ⇌ 2Fe <sup>III</sup> (OH) <sub>3</sub> + As <sup>III</sup> (OH) <sub>3</sub> + 2H <sup>+</sup>	22.3
12	2Fe <sup>2+</sup> + H <sub>2</sub> As <sup>V</sup> O <sub>4</sub> <sup>-</sup> + 3H <sup>+</sup> ⇌ 2Fe <sup>3+</sup> + As <sup>III</sup> (OH) <sub>3</sub> + H <sub>2</sub> O	23.2
13	2Fe <sup>2+</sup> + HAs <sup>V</sup> O <sub>4</sub> <sup>2-</sup> + 4H <sup>+</sup> ⇌ 2Fe <sup>3+</sup> + As <sup>III</sup> (OH) <sub>3</sub> + H <sub>2</sub> O	-16.7

<sup>a</sup> Values calculated from the standard Gibbs free energies ( $\Delta G_f^\circ$ ) of minerals and aqueous species (Table S-6).

It must be noted, however, that As(V) reduction has not been observed previously upon interaction with Fe<sup>2+</sup>-activated synthetic GT (e.g., Amstaetter et al.<sup>68</sup>), who examined the interactions at a Fe<sup>2+</sup><sub>(aq)</sub>/Fe(III)<sub>solid</sub> ratio of 0.03, which is approximately 15 to 55 times lower than the ratios used in this study. Since the reduction reaction is driven by Fe<sup>2+</sup> concentration, the low Fe<sup>2+</sup> concentration used in

their study could explain why they did not observe any As(V) reduction to As(III) in their system. However, a question arises whether As(V) could be reduced to As(III) at lower  $\text{Fe}^{2+}_{(\text{aq})}/\text{Fe(III)}_{\text{solid}}$  ratios and  $\text{Fe(III)}_{\text{FH}}$  loadings similar to those reported by Pedersen et al.<sup>39</sup> and Masue-Slowey et al.<sup>40</sup>, especially since the mineralogical composition of the end-products is different from what we observed in our study.

Overall, these redox transformations have important implications for the mobility and toxicity of As. The partial reduction of As(V) to As(III), as documented in this study, is an unexpected and also detrimental consequence as such reduction results in the generation of far more toxic and mobile As species.<sup>69</sup> On the positive side, the sorption capacities of these Fe mineral phases towards As species is very high, and therefore we observed no significant As release. Noteworthy, however, is the fact that invariably real subsurface environments are significantly more complex. The presence of many different mineral substrates and the variation in mineral sorption capacities will be affected by Eh/pH conditions,<sup>70</sup> the presence of other inorganic ions<sup>33,69</sup> (e.g., silicate and phosphate anions) or organic ligands<sup>71-74</sup> all competing with As for active surface sites and influencing the mechanisms and pathways Fe (oxyhydr)oxide transformation.

## CONCLUSION

In subsurface environments, iron-bearing mineral transformations can massively impact the mobility and toxicity of contaminants, since these mineral phases serve as toxic element sinks that can control and even prevent release and further transport of contaminants in soils and groundwaters. In this study, we followed the transformation of As(V)-bearing ferrihydrite, catalyzed by aqueous  $\text{Fe}^{2+}$ , under anoxic conditions as it converts to more crystalline iron (oxyhydr)oxides. Higher  $\text{Fe}^{2+}$  concentrations resulted in the formation of both GT and GR phases, while lower  $\text{Fe}^{2+}$  concentrations led to a GT end-product. However, at all the tested conditions, the conversion of ferrihydrite was incomplete, and our data indicate that this was a consequence of As surface complexation. Analyses of the mineral-bound As species also revealed partial reduction of initial As(V) to As(III), although no significant release of As was observed during the transformation. Overall, our results highlight the need to understand such inter-

transformations among iron (oxyhydr)oxide in subsurface environments where aqueous  $\text{Fe}^{2+}$  is present as it will impact As sequestration, mobilization and transport.

## **ASSOCIATED CONTENT**

### **Supporting Information**

Detailed information on mineral characterization and analytical techniques can be found in the Supporting Information. Additional mineral characterization data (TEM, SEM, XAS, PDF, thermodynamic calculations), XPS reference data and thermodynamic calculations can also be seen in this section. This material is available free of charge on the ACS Publications website.

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### **Notes**

The authors declare no competing financial interest.

### **Funding**

This project has received funding from the European Union's Horizon 2020 Marie Skłodowska-Curie Innovative Training Network Grant No. 675219. LGB and HMF acknowledges the financial support from the Helmholtz Recruiting Initiative (award number I-044-16-01). Use of the Advanced Photon Source was supported by the U. S. Department of Energy, Office of Science, Office of Basic Energy Sciences, under Contract No. DE-AC02-06CH11357. DJT and KD acknowledge financial support from the Danish Council for Independent Research (via DANSCATT) for travel to APS.

### **Acknowledgement**

ICP-OES analyses were carried out at the Helmholtz Laboratory for the Geochemistry of the Earth Surface (HELGES) at GFZ Potsdam. The authors would like to thank David Uhlig of HELGES for his help during ICP-OES analysis of some of the samples and Sathish Mayana of Interface Geochemistry group at GFZ Potsdam for his help during SEM imaging. JPHP and ANT thank ANKA for access to SUL-X beamline, and beamline staff Jörg Göttlicher and Ralf Steininger for their competent support and advice during collection of Fe K-edge EXAFS data. The As K-edge XANES data were collected at the BM23 beamline at ESRF (experiment no. EV-338), and JPHP, LGB, DJT and KD thank Sakura Pascarelli for assistance during beamtime. DJT and KD thank Olaf Borkiewicz and Kevin A. Beyer for support with X-ray total scattering measurements at APS beamline 11 ID-B, Argonne, USA. JPHP acknowledges the help of Case Van Genuchten and Hongyan Wang during the XAS beamtime experiments. JPHP would also like to thank Leonard Daniël Samson for his help with the statistical analysis of the particle size distribution data.

## References

1. Cornell, R. M.; Schwertmann, U., *The Iron Oxides: Structure, Properties, Reactions, Occurrences and Uses*. 2nd ed.; Wiley-VCH Verlag GmbH & Co. KGaA: Weinheim, FRG, 2003.
2. Karimian, N.; Johnston, S. G.; Burton, E. D., Iron and sulfur cycling in acid sulfate soil wetlands under dynamic redox conditions: A review. *Chemosphere* **2018**, *197*, 803-816.
3. Jambor, J. L.; Dutrizac, J. E., Occurrence and constitution of natural and synthetic ferrihydrite, a widespread iron oxyhydroxide. *Chemical Reviews* **1998**, *98* (7), 2549-2586.
4. Cornell, R. M., The influence of some divalent cations on the transformation of ferrihydrite to more crystalline products. *Clay Minerals* **1988**, *23* (3), 329-332.
5. Vu, H. P.; Shaw, S.; Brinza, L.; Benning, L. G., Crystallization of hematite ( $\alpha$ -Fe<sub>2</sub>O<sub>3</sub>) under alkaline condition: The effects of Pb. *Crystal Growth & Design* **2010**, *10* (4), 1544-1551.
6. Vu, H. P.; Shaw, S.; Brinza, L.; Benning, L. G., Partitioning of Pb(II) during goethite and hematite crystallization: Implications for Pb transport in natural systems. *Applied Geochemistry* **2013**, *39*, 119-128.
7. Brinza, L.; Vu, H. P.; Shaw, S.; Mosselmans, J. F. W.; Benning, L. G., Effect of Mo and V on the hydrothermal crystallization of hematite from ferrihydrite: An *in situ* energy dispersive X-ray diffraction and X-ray absorption spectroscopy study. *Crystal Growth & Design* **2015**, *15* (10), 4768-4780.
8. Brinza, L.; Vu, H. P.; Neamtu, M.; Benning, L. G., Experimental and simulation results of the adsorption of Mo and V onto ferrihydrite. *Scientific Reports* **2019**, *9* (1), 1365.
9. Schwertmann, U.; Stanjek, H.; Becher, H. H., Long-term *in vitro* transformation of 2-line ferrihydrite to goethite/hematite at 4, 10, 15 and 25 °C. *Clay Minerals* **2004**, *39* (4), 433-438.
10. Schwertmann, U.; Murad, E., Effect of pH on the formation of goethite and hematite from ferrihydrite. *Clays and Clay Minerals* **1983**, *31* (4), 277-284.

11. Das, S.; Hendry, M. J.; Essilfie-Dughan, J., Transformation of two-line ferrihydrite to goethite and hematite as a function of pH and temperature. *Environmental Science & Technology* **2011**, *45* (1), 268-275.
12. Shaw, S.; Pepper, S. E.; Bryan, N. D.; Livens, F. R., The kinetics and mechanisms of goethite and hematite crystallization under alkaline conditions, and in the presence of phosphate. *American Mineralogist* **2005**, *90* (11-12), 1852–1860.
13. Jang, J.-H.; Dempsey, B. A.; Catchen, G. L.; Burgos, W. D., Effects of Zn(II), Cu(II), Mn(II), Fe(II), NO<sub>3</sub><sup>-</sup>, or SO<sub>4</sub><sup>2-</sup> at pH 6.5 and 8.5 on transformations of hydrous ferric oxide (HFO) as evidenced by Mössbauer spectroscopy. *Colloids and Surfaces A: Physicochemical and Engineering Aspects* **2003**, *221* (1), 55-68.
14. Vempati, R.; Loeppert, R. H., Influence of structural and adsorbed Si on the transformation of synthetic ferrihydrite. *Clays and Clay Minerals* **1989**, *37* (3), 273-279.
15. Cornell, R. M.; Schneider, W., Formation of goethite from ferrihydrite at physiological pH under the influence of cysteine. *Polyhedron* **1989**, *8* (2), 149-155.
16. Cornell, R. M., Comparison and classification of the effects of simple ions and molecules upon the transformation of ferrihydrite into more crystalline products. *Zeitschrift für Pflanzenernährung und Bodenkunde* **1987**, *150* (5), 304-307.
17. Hansel, C. M.; Benner, S. G.; Fendorf, S., Competing Fe(II)-induced mineralization pathways of ferrihydrite. *Environmental Science & Technology* **2005**, *39* (18), 7147-7153.
18. Tronc, E.; Belleville, P.; Jolivet, J. P.; Livage, J., Transformation of ferric hydroxide into spinel by iron(II) adsorption. *Langmuir* **1992**, *8* (1), 313-319.
19. Liu, H.; Li, P.; Zhu, M.; Wei, Y.; Sun, Y., Fe(II)-induced transformation from ferrihydrite to lepidocrocite and goethite. *Journal of Solid State Chemistry* **2007**, *180* (7), 2121-2128.
20. Yang, L.; Steefel, C. I.; Marcus, M. A.; Bargar, J. R., Kinetics of Fe(II)-catalyzed transformation of 6-line ferrihydrite under anaerobic flow conditions. *Environmental Science & Technology* **2010**, *44* (14), 5469-5475.
21. Kapler, A.; Straub, K. L., Geomicrobiological cycling of iron. *Reviews in Mineralogy and Geochemistry* **2005**, *59* (1), 85-108.
22. Fortin, D.; Langley, S., Formation and occurrence of biogenic iron-rich minerals. *Earth-Science Reviews* **2005**, *72* (1), 1-19.
23. Hiemstra, T.; van Riemsdijk, W. H., Adsorption and surface oxidation of Fe(II) on metal (hydr)oxides. *Geochimica et Cosmochimica Acta* **2007**, *71* (24), 5913-5933.
24. Gorski, C. A.; Scherer, M. M., Fe<sup>2+</sup> sorption at the Fe oxide-water interface: A revised conceptual framework. In *Aquatic Redox Chemistry*, Tratnyek, P. G.; Grundl, T. J.; Haderlein, S. B., Eds. American Chemical Society: 2011; Vol. 1071, pp 315-343.
25. Katz, J. E.; Zhang, X.; Attenkofer, K.; Chapman, K. W.; Frandsen, C.; Zarzycki, P.; Rosso, K. M.; Falcone, R. W.; Waychunas, G. A.; Gilbert, B., Electron small polarons and their mobility in iron (oxyhydr)oxide nanoparticles. *Science* **2012**, *337* (6099), 1200-1203.
26. Williams, A. G. B.; Scherer, M. M., Spectroscopic evidence for Fe(II)–Fe(III) electron transfer at the iron oxide–water interface. *Environmental Science & Technology* **2004**, *38* (18), 4782-4790.
27. Boland, D. D.; Collins, R. N.; Miller, C. J.; Glover, C. J.; Waite, T. D., Effect of solution and solid-phase conditions on the Fe(II)-accelerated transformation of ferrihydrite to lepidocrocite and goethite. *Environmental Science & Technology* **2014**, *48* (10), 5477-5485.
28. Ahmed, I. A. M.; Benning, L. G.; Kakonyi, G.; Sumoondur, A. D.; Terrill, N. J.; Shaw, S., Formation of green rust sulfate: A combined *in situ* time-resolved X-ray scattering and electrochemical study. *Langmuir* **2010**, *26* (9), 6593-6603.
29. Sumoondur, A.; Shaw, S.; Ahmed, I.; Benning, L. G., Green rust as a precursor for magnetite: An *in situ* synchrotron based study. *Mineralogical Magazine* **2008**, *72* (1), 201-204.
30. Wang, Y.; Morin, G.; Ona-Nguema, G.; Brown, G. E., Arsenic(III) and arsenic(V) speciation during transformation of lepidocrocite to magnetite. *Environmental Science & Technology* **2014**, *48* (24), 14282-14290.



31. Vaughan, D. J., Arsenic. *Elements* **2006**, 2 (2), 71-75.
32. Smedley, P. L.; Kinniburgh, D. G., A review of the source, behaviour and distribution of arsenic in natural waters. *Applied Geochemistry* **2002**, 17 (5), 517-568.
33. Perez, J. P. H.; Freeman, H. M.; Schuessler, J. A.; Benning, L. G., The interfacial reactivity of arsenic species with green rust sulfate (GR<sub>SO4</sub>). *Science of The Total Environment* **2019**, 648, 1161-1170.
34. Raven, K. P.; Jain, A.; Loeppert, R. H., Arsenite and arsenate adsorption on ferrihydrite: Kinetics, equilibrium, and adsorption envelopes. *Environmental Science & Technology* **1998**, 32 (3), 344-349.
35. Mamindy-Pajany, Y.; Hurel, C.; Marmier, N.; Roméo, M., Arsenic adsorption onto hematite and goethite. *Comptes Rendus Chimie* **2009**, 12 (8), 876-881.
36. Tang, W.; Li, Q.; Gao, S.; Shang, J. K., Arsenic (III,V) removal from aqueous solution by ultrafine  $\alpha$ -Fe<sub>2</sub>O<sub>3</sub> nanoparticles synthesized from solvent thermal method. *Journal of Hazardous Materials* **2011**, 192 (1), 131-138.
37. Lin, S.; Lu, D.; Liu, Z., Removal of arsenic contaminants with magnetic  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub> nanoparticles. *Chemical Engineering Journal* **2012**, 211-212, 46-52.
38. Feng, L.; Cao, M.; Ma, X.; Zhu, Y.; Hu, C., Superparamagnetic high-surface-area Fe<sub>3</sub>O<sub>4</sub> nanoparticles as adsorbents for arsenic removal. *Journal of Hazardous Materials* **2012**, 217-218, 439-446.
39. Pedersen, H. D.; Postma, D.; Jakobsen, R., Release of arsenic associated with the reduction and transformation of iron oxides. *Geochimica et Cosmochimica Acta* **2006**, 70 (16), 4116-4129.
40. Masue-Slowey, Y.; Loeppert, R. H.; Fendorf, S., Alteration of ferrihydrite reductive dissolution and transformation by adsorbed As and structural Al: Implications for As retention. *Geochimica et Cosmochimica Acta* **2011**, 75 (3), 870-886.
41. Jönsson, J.; Sherman, D. M., Sorption of As(III) and As(V) to siderite, green rust (fougerite) and magnetite: Implications for arsenic release in anoxic groundwaters. *Chemical Geology* **2008**, 255 (1-2), 173-181.
42. Schwertmann, U.; Cornell, R. M., *Iron Oxides in the Laboratory: Preparation and Characterization*. 2nd ed.; Wiley-VCH Verlag GmbH & Co. KGaA: Weinheim, FRG, 2000; p 188.
43. Atkinson, R. J.; Posner, A. M.; Quirk, J. P., Adsorption of potential-determining ions at the ferric oxide-aqueous electrolyte interface. *The Journal of Physical Chemistry* **1967**, 71 (3), 550-558.
44. Bethke, C. M., *Geochemical and Biogeochemical Reaction Modeling*. Cambridge University Press: 2010.
45. Karimian, N.; Johnston, S. G.; Burton, E. D., Antimony and arsenic behavior during Fe(II)-induced transformation of jarosite. *Environmental Science & Technology* **2017**, 51 (8), 4259-4268.
46. Karimian, N.; Johnston, S. G.; Burton, E. D., Antimony and arsenic partitioning during Fe<sup>2+</sup>-induced transformation of jarosite under acidic conditions. *Chemosphere* **2018**, 195, 515-523.
47. Gomez, M. A.; Jim Hendry, M.; Hossain, A.; Das, S.; Elouatik, S., Abiotic reduction of 2-line ferrihydrite: Effects on adsorbed arsenate, molybdate, and nickel. *RSC Advances* **2013**, 3 (48), 25812-25822.
48. Asta, M. P.; Cama, J.; Martínez, M.; Giménez, J., Arsenic removal by goethite and jarosite in acidic conditions and its environmental implications. *Journal of Hazardous Materials* **2009**, 171 (1), 965-972.
49. Park, J. H.; Han, Y.-S.; Ahn, J. S., Comparison of arsenic co-precipitation and adsorption by iron minerals and the mechanism of arsenic natural attenuation in a mine stream. *Water Research* **2016**, 106, 295-303.
50. Waychunas, G. A.; Rea, B. A.; Fuller, C. C.; Davis, J. A., Surface chemistry of ferrihydrite: Part 1. EXAFS studies of the geometry of coprecipitated and adsorbed arsenate. *Geochimica et Cosmochimica Acta* **1993**, 57 (10), 2251-2269.
51. Richmond, W. R.; Loan, M.; Morton, J.; Parkinson, G. M., Arsenic removal from aqueous solution via ferrihydrite crystallization control. *Environmental Science & Technology* **2004**, 38 (8), 2368-2372.

52. Rancourt, D. G.; Fortin, D.; Pichler, T.; Thibault, P.-J.; Lamarche, G.; Morris, R. V.; Mercier, P. H. J., Mineralogy of a natural As-rich hydrous ferric oxide coprecipitate formed by mixing of hydrothermal fluid and seawater: Implications regarding surface complexation and color banding in ferrihydrite deposits. *American Mineralogist* **2001**, *86* (7-8), 834-851.
53. Freeman, H. M.; Perez, J. P. H.; Hondow, N.; Benning, L. G.; Brown, A. P., Beam-induced oxidation of mixed-valent Fe (oxyhydr)oxides (green rust) monitored by STEM-EELS. *Micron* **2019**.
54. Perez, J. P. H.; Mangayayam, M. C.; Rubio, S. N.; Freeman, H. M.; Tobler, D. J.; Benning, L. G., Intercalation of aromatic sulfonates in 'green rust' via ion exchange. *Energy Procedia* **2018**, *146*, 179-187.
55. Skovbjerg, L. L.; Stipp, S. L. S.; Utsunomiya, S.; Ewing, R. C., The mechanisms of reduction of hexavalent chromium by green rust sodium sulphate: Formation of Cr-goethite. *Geochimica et Cosmochimica Acta* **2006**, *70* (14), 3582-3592.
56. Christiansen, B. C.; Balic-Zunic, T.; Petit, P. O.; Frandsen, C.; Mørup, S.; Geckeis, H.; Katerinopoulou, A.; Stipp, S. L. S., Composition and structure of an iron-bearing, layered double hydroxide (LDH) – Green rust sodium sulphate. *Geochimica et Cosmochimica Acta* **2009**, *73* (12), 3579-3592.
57. Ravel, B.; Newville, M., ATHENA, ARTEMIS, HEPHAESTUS: Data analysis for X-ray absorption spectroscopy using IFEFFIT. *Journal of Synchrotron Radiation* **2005**, *12* (4), 537-541.
58. Farrow, C. L.; Juhas, P.; Liu, J. W.; Bryndin, D.; Božin, E. S.; Bloch, J.; Th, P.; Billinge, S. J. L., PDFfit2 and PDFgui: Computer programs for studying nanostructure in crystals. *Journal of Physics: Condensed Matter* **2007**, *19* (33), 335219.
59. Usman, M.; Hanna, K.; Abdelmoula, M.; Zegeye, A.; Faure, P.; Ruby, C., Formation of green rust via mineralogical transformation of ferric oxides (ferrihydrite, goethite and hematite). *Applied Clay Science* **2012**, *64*, 38-43.
60. Su, C.; Puls, R. W., Significance of iron(II,III) hydroxycarbonate green rust in arsenic remediation using zerovalent iron in laboratory column tests. *Environmental Science & Technology* **2004**, *38* (19), 5224-5231.
61. Yee, N.; Shaw, S.; Benning, L. G.; Nguyen, T. H., The rate of ferrihydrite transformation to goethite via the Fe(II) pathway. *American Mineralogist* **2006**, *91* (1), 92-96.
62. Wang, Y.; Morin, G.; Ona-Nguema, G.; Juillot, F.; Guyot, F.; Calas, G.; Brown, G. E., Evidence for different surface speciation of arsenite and arsenate on green rust: An EXAFS and XANES Study. *Environmental Science & Technology* **2010**, *44* (1), 109-115.
63. Randall, S. R.; Sherman, D. M.; Ragnarsdottir, K. V., Sorption of As(V) on green rust (Fe<sub>4</sub>(II)Fe<sub>2</sub>(III)(OH)<sub>12</sub>SO<sub>4</sub>·3H<sub>2</sub>O) and lepidocrocite (γ-FeOOH): Surface complexes from EXAFS spectroscopy. *Geochimica et Cosmochimica Acta* **2001**, *65* (7), 1015-1023.
64. Amonette, J. E.; Workman, D. J.; Kennedy, D. W.; Fruchter, J. S.; Gorby, Y. A., Dechlorination of carbon tetrachloride by Fe(II) associated with goethite. *Environmental Science & Technology* **2000**, *34* (21), 4606-4613.
65. Stewart, S. M.; Hofstetter, T. B.; Joshi, P.; Gorski, C. A., Linking thermodynamics to pollutant reduction kinetics by Fe<sup>2+</sup> bound to iron oxides. *Environmental Science & Technology* **2018**, *52* (10), 5600-5609.
66. Gorski, C. A.; Edwards, R.; Sander, M.; Hofstetter, T. B.; Stewart, S. M., Thermodynamic characterization of iron oxide–aqueous Fe<sup>2+</sup> redox couples. *Environmental Science & Technology* **2016**, *50* (16), 8538-8547.
67. Buerge, I. J.; Hug, S. J., Influence of mineral surfaces on chromium(VI) reduction by iron(II). *Environmental Science & Technology* **1999**, *33* (23), 4285-4291.
68. Amstaetter, K.; Borch, T.; Larese-Casanova, P.; Kappler, A., Redox transformation of arsenic by Fe(II)-activated goethite (α-FeOOH). *Environmental Science & Technology* **2010**, *44* (1), 102-108.
69. Roberts, L. C.; Hug, S. J.; Ruettimann, T.; Billah, M. M.; Khan, A. W.; Rahman, M. T., Arsenic removal with iron(II) and iron(III) in waters with high silicate and phosphate concentrations. *Environmental Science & Technology* **2004**, *38* (1), 307-315.

70. Root, R. A.; Dixit, S.; Campbell, K. M.; Jew, A. D.; Hering, J. G.; O'Day, P. A., Arsenic sequestration by sorption processes in high-iron sediments. *Geochimica et Cosmochimica Acta* **2007**, *71* (23), 5782-5803.
71. Zhou, Z.; Latta, D. E.; Noor, N.; Thompson, A.; Borch, T.; Scherer, M. M., Fe(II)-catalyzed transformation of organic matter–ferrihydrite coprecipitates: A closer look using  $\epsilon$  isotopes. *Environmental Science & Technology* **2018**, *52* (19), 11142-11150.
72. Chen, C.; Sparks, D. L., Fe(II)-induced mineral transformation of ferrihydrite–organic matter adsorption and co-precipitation complexes in the absence and presence of As(III). *ACS Earth and Space Chemistry* **2018**, *2* (11), 1095-1101.
73. Chen, C.; Kukkadapu, R.; Sparks, D. L., Influence of coprecipitated organic matter on  $\text{Fe}^{2+(\text{aq})}$ -catalyzed transformation of ferrihydrite: Implications for carbon dynamics. *Environmental Science & Technology* **2015**, *49* (18), 10927-10936.
74. Hu, S.; Lu, Y.; Peng, L.; Wang, P.; Zhu, M.; Dohnalkova, A. C.; Chen, H.; Lin, Z.; Dang, Z.; Shi, Z., Coupled kinetics of ferrihydrite transformation and As(V) sequestration under the effect of humic acids: A mechanistic and quantitative study. *Environmental Science & Technology* **2018**, *52* (20), 11632-11641.

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