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# Peatland afforestation in the UK and consequences for carbon storage

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## SUMMARY

Peatlands are a globally significant store of carbon. During the second half of the 20<sup>th</sup> century new planting techniques combined with tax incentives encouraged commercial forestry across large areas of peat bog in the UK, particularly in the Flow Country of northern Scotland. Such planting was controversial and was ultimately halted by removal of the tax incentives, and policies to prevent new planting. Here we review the literature on UK peatland afforestation in relation to carbon and climate implications, and identify key issues for future research. The effects of conifer planting on peat bog carbon storage in the UK are poorly understood. A large body of research on peatland forestry exists, particularly from naturally forested fen peatlands in Fennoscandia and Russia, but the different conditions in the UK mean that results are not directly transferable. Data on the responses of UK peat bogs to afforestation are required to address this shortfall. Studies are required that quantify the loss of carbon from the peat and evaluate it against the accumulation of carbon above and below ground in trees, considering the likely residence time of carbon in wood products.

**KEY WORDS:** Flow Country, forestry, GHG, greenhouse gases, peat

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## INTRODUCTION

### Carbon storage in peat

Peatlands are globally important stores of carbon. Covering about 3 % of the surface of the Earth (Dise 2009), they are believed to store over 600 Gt of carbon (Yu 2011, Loisel et al. 2017). This is of a similar order of magnitude to the 800+ Gt of carbon in the atmosphere (Batjes 1996, IPCC 2014).

Northern peatlands are globally the most important stores of carbon, and are distributed primarily across Russia, North America, Fennoscandia, Eastern Europe and the British Isles (Mitsch & Gosselink 2015). These northern peatlands are estimated to contain more than 90 % (547 Gt) of the total peatland carbon pool (Yu et al. 2010).

Accumulation of this peat has provided a small negative feedback to climate over the last 1000 years (Charman et al. 2013), with an estimated net sink of carbon of 44 Gt ka<sup>-1</sup> (Yu 2011). Peatlands also influence the climate system as a significant source

of methane (CH<sub>4</sub>), carbon dioxide (CO<sub>2</sub>), aquatic carbon and to a less significant extent other greenhouse gases (N<sub>2</sub>O, VOCs), and have a direct effect on radiation balance through albedo.

### Development of the forest industry in peatlands

Peatlands have historically been viewed as barren and unproductive places, but in reality support many economic activities and provide often unnoticed ecosystem services. Economically, peatlands have important roles for agriculture (in particular grazing), water management, and leisure activities such as shooting and tourism (Whitfield et al. 2011). Ecosystem services include water and carbon storage (Joosten et al. 2012), and maintenance of biodiversity including specialised peatland species (Stroud et al. 1987, Lindsay et al. 1988, Littlewood et al. 2010).

It is estimated that around 20 % of European peatlands are currently drained for forestry (Drosler et al. 2008). Many peatlands, especially those in tropical and boreal regions, have natural tree cover and may be categorised as 'forest'. Other peatlands,

for example many within the Arctic and temperate zone, are naturally treeless. In these landscapes, mixed wet scrub and low wet woodland are restricted to peat bog margins and along the courses of streams (Lindsay 2010). One such area is the United Kingdom, where an estimated 2300 Mt of carbon is stored in peatlands (Billett *et al.* 2010), of which blanket bog is the predominant type. There are a few sites where trees naturally occur on ombrotrophic bog peat and these may have been more widespread in the past, but today almost all UK bogs are open. This changed between the 1950s and the 1980s, when approximately 9 % (190,000 ha) of the UK's deep peats were drained for forestry (Hargreaves *et al.* 2003), although this figure may be an underestimate and may be as high as 17 % in Scotland (Vanguelova *et al.* 2018).

### Early trials in the UK

Numerous attempts at peatland afforestation have been made in the UK since the 18<sup>th</sup> century. For instance, in his history of the county of Peeblesshire, William Chambers (1864) records a drainage initiative by the Duke of Argyll in 1730, in which he made an “attempt to make a quagmire not only into a dry and arable land, but fitted by its amenity for the residence of a man of taste”. This included an early and largely ineffective attempt at drain cutting, with trees being planted on any sufficiently dry areas. After poor results and the death of the Duke in 1761, the plan was abandoned (W. Chambers 1864). Such schemes, driven generally by individuals or individual estates, are typical of the small-scale and uncoordinated efforts common at the time.

Foresters in Britain were slow to take note of developments in continental Europe. By 1836, foresters in Belgium had developed a system of turf planting in which some of the peat was removed, upturned and laid over the remaining surface to give a deeper, drier substrate on which to plant. This was combined with intensive drainage to yield the first significant successes in planting forests on peat, a system not widely adopted in Britain until around 1907 (Zehetmayer 1954).

### UK peatland afforestation in the 20<sup>th</sup> century

A critical moment in the history of British peatland forestry was the establishment of the Forestry Commission, a government body with responsibility for managing forestry. The Forestry Commission was founded under the Forestry Act of 1919, with a remit to increase forest coverage and timber production. As well as aiming to develop an economic resource, this was in part a response to concerns about depleted woodland stocks following the First World War, as a

domestic supply of wooden pit props to support the mining industry was strategically important (Marren 2002). The establishment of the Forestry Commission led to a more coordinated and efficient approach to forestry.

Expansion of forestry into the uplands during the inter-war years occurred mainly across organo-mineral soils. Deeper peat was considered too challenging for silviculture and unsuitable for the machinery then in use. It was not until after the Second World War that development and modification of the double mouldboard plough combined with efficient tractors with wide tracks allowed the Forestry Commission to commence more widespread trials on deeper peats (Wood 1974, Anderson 1997). The double mouldboard plough pushed cut peat into a ridge on either side of a drainage furrow, creating raised dry ridges typically two metres apart on which trees could be planted (Figure 1). This closely-spaced furrow ploughing was combined with collector drainage ditches at intervals to provide a sufficiently dry environment for tree growth (Harrison *et al.* 1994) and was supplemented by fertiliser application to overcome the paucity of nutrients, particularly phosphorous but also potassium, nitrogen and trace elements (Taylor 1991).

Norway spruce (*Picea abies*), Scots pine (*Pinus sylvestris*), mountain pine (*Pinus mugo*) and species of larch (*Larix decidua*, *Larix kaempferi*) had been trialled for peatland afforestation in the UK by the early 20<sup>th</sup> century, but with limited success (Zehetmayer 1954). Ultimately, forestry in UK peat bogs became feasible with the adoption into European silviculture of trees native to North America, particularly some varieties of lodgepole pine (*Pinus contorta*) and Sitka spruce (*Picea sitchensis*). ‘South Coastal’ varieties of lodgepole pine were initially seen as good candidates for afforesting large areas of peat (Pyatt 1990). This species tolerates high water tables by creating gas pockets within the pericycle of the roots that allow continued oxygenation in waterlogged conditions by diffusion from the air (King *et al.* 1986). Consequently, it roots deeply, drying the peat. However, problems with curvature of the base of the trunk (‘basal sweep’), low wood quality and occasional devastating outbreaks of Pine Beauty Moth (*Panolis flammea*) meant that lodgepole pine was ultimately abandoned as a commercial crop.

Sitka spruce was introduced into the UK as an ornamental species in the late 1820s. Due to its rapid growth and excellent quality wood, it was adopted as a commercial crop in the early 20<sup>th</sup> century (Oosthoek 2013). Sitka spruce is a valuable timber-producing

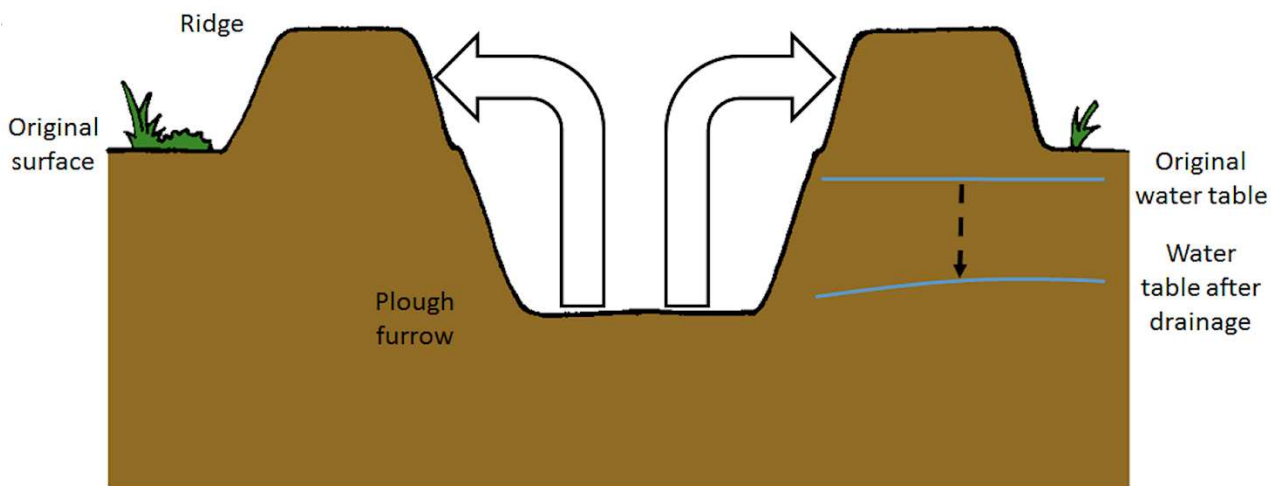


Figure 1. A double mould board plough creates a furrow in the peat and pushes excavated peat into ridges. These ridges are sufficiently raised above the drained water table that the survival chance of planted trees is substantially increased. Unlike the low tillage on Fennoscandian sites, such ploughing on UK forestry sites disturbs the peat and removes much of the bog vegetation.

species and is the most widely grown conifer in the UK, covering 682,100 ha (50.3 % of the total conifer stock) (Forestry Commission 2011). This species alone accounts for 33 % of the total woodland coverage in the Highlands of Scotland (Smith & Gilbert 2003). Sitka spruce grows poorly in waterlogged conditions, so in peat bogs in many parts of the UK it was mostly planted in mixed stands with lodgepole pine, which acted as a ‘nurse species’ (Pyatt 1993). It was hoped that the relative vigour of lodgepole pine in wet conditions, and consequent water interception due to canopy closure, would in turn increase yields of other species during the first rotation (King *et al.* 1986). Ultimately, while lodgepole pine was shown to have a drying effect on the peat, this did not always translate into an improvement in growth of the Sitka spruce (Ray & Schweizer 1994). For this reason the economic benefits of mixed planting were questioned and Sitka spruce monocultures became increasingly common as more stands were planted (Oosthoek 2013).

Sites across the UK were drained and planted by the Forestry Commission in the second half of the 20<sup>th</sup> century. At this time, forest planting was an industrial-scale operation involving extensive landscape change beyond simply planting trees including construction of roads, bridges and fences, and quarrying for building materials.

The technological developments which permitted peat bog afforestation coincided with a tax and grant regime favourable to forest development in unsuitable areas. Government incentives proved popular as a mechanism for reducing tax liability (Mather 1986, Mather & Murray 1988). All expenses

related to forestry were tax deductible, with loans available which could also be written off against tax. Companies such as Fountain Forestry managed large areas of land for wealthy individuals. Through the 1970s private planting overtook planting by the Forestry Commission (Figure 2), much of it concentrated in Scotland (Mather & Murray 1988). Tree growth was frequently poor and a large proportion of the forests planted during this period would not have been economically viable without tax relief.

#### Public and scientific reaction to afforestation

From the late 1960s the issue of peatland afforestation grew in prominence, with concerns raised over the loss of biodiversity and the risk of eutrophication of freshwaters and damage to fisheries (Moore & Bellamy 1974, Thompson 1987). Public awareness of the large-scale planting of the uplands and the economic factors underpinning it was raised with the revelation that well-known figures such as TV presenter Terry Wogan, singer Cliff Richard and snooker player Alex ‘Hurricane’ Higgins were using forestry-based tax avoidance schemes (Rosie 1986, Anon. 1995).

Between 1987 and 1988, the Nature Conservancy Council - the UK government statutory advisor on wildlife conservation matters at the time - published ‘Birds, Bogs and Forestry’ and ‘The Flow Country - the Peatlands of Caithness and Sutherland’, a linked pair of reports on the biodiversity of the Flow Country and the scale of forestry expansion (Stroud *et al.* 1987, Lindsay *et al.* 1988). The Flow Country is the UK’s most extensive peatland region with over

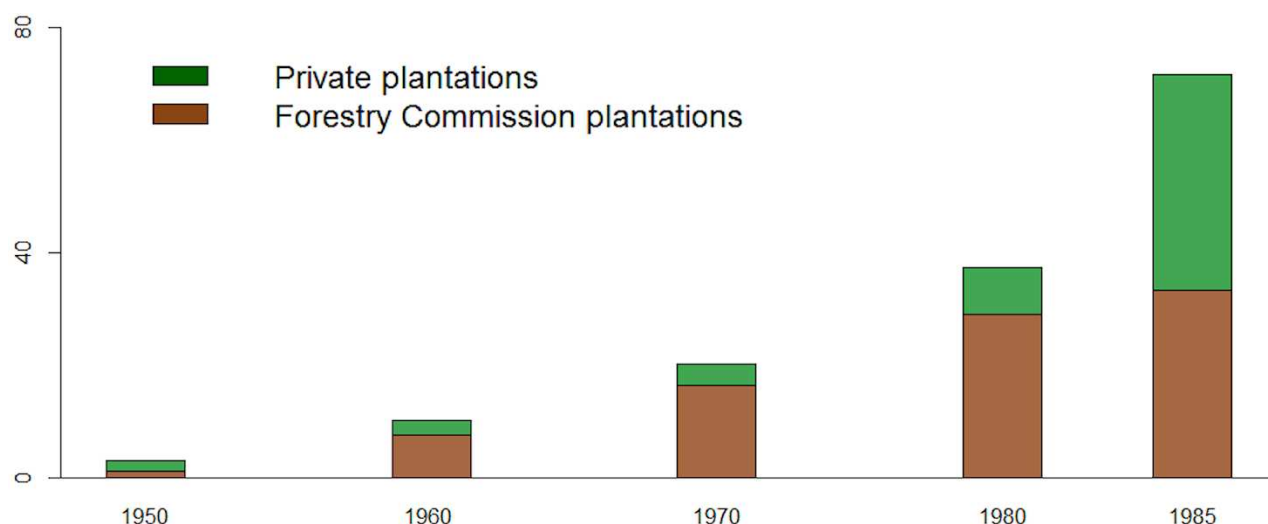


Figure 2. Extent and ownership of forest plantations in Caithness and Sutherland between 1950 and 1985. Adapted from Stroud *et al.* (2015).

400,000 ha of peat and wetland, of which around 67,000 ha (approximately 17 %) had by then been afforested. The reports highlighted the potential disruption that could be caused by forestry and, while the first report generated extensive political controversy, the detailed figures provided in the second report led the Secretary of State for Scotland to afford statutory protection to almost 200,000 ha of un-afforested peatland in the Flow Country as a composite Site of Special Scientific Interest (SSSI), the largest such site in the UK. The fallout from this controversial action is widely believed to have contributed to the subsequent decision of the government of the time to break up the Nature Conservancy Council (Warren 2000). Later, the SSSI was designated as the UK's largest terrestrial Special Area for Conservation (SAC) and Special Protection Area (SPA) within the European Commission's 'Natura 2000' nature protection network.

Controversy over tax avoidance in general, but particularly the schemes set up for forestry, led to legislative changes. With public outcry increasing, the then Chancellor of the Exchequer Nigel Lawson ended the tax breaks in his budget of 1988 (Oosthoek 2013). With the main financial incentive removed, new peat bog forestry planting has been limited since 1990 (Stroud *et al.* 2015) and was effectively halted by later Forestry Commission policy guidance (Patterson & Anderson 2000).

### Restocking or restoration?

Following the intensive afforestation of the twentieth century around 9 % of the UK's deep peats, amounting to a total of approximately 190,000 ha, were drained for forestry (Hargreaves *et al.* 2003).

This forestry is distributed across the UK but is particularly extensive in Scotland. Many plantations are approaching harvesting age and decisions must soon be taken on whether to restock the forests or restore drained bogs as far as possible to their previous state. While there is some debate as to whether restoration should aim to recreate a pre-drainage or pre-afforestation state, the process typically involves the removal of trees and blocking of drainage to raise and stabilise water tables and restore active peatland habitats.

While not the only factor (biodiversity considerations are especially important; Holden *et al.* 2007), the effect of afforestation on carbon stock and carbon cycling is an important issue in this decision-making given likely consequences for climate change. While peatland restoration was not originally specified by the Kyoto Protocol (beyond a general call for the protection of natural carbon stocks and sinks) or used as a mitigating factor in subsequent calculations of carbon emissions, the Protocol was ultimately amended to allow peatland rewetting to be considered in carbon accounting (Bain *et al.* 2012). Restoration of peatlands is recommended by several international bodies (Joosten *et al.* 2012) including, most recently, the International Union for the Conservation of Nature (Resolution 043; IUCN 2016).

In Scotland, the devolved government aims to restore 40 % of the estimated 600,000 ha of damaged peatlands by 2030 (Scottish Government 2017), which includes restoration of afforested peat bogs (Scottish Natural Heritage 2015). Generally, there is a presumption that any felled woodlands will be restocked, but allowances are made in the Scottish

Government's Policy on the Control of Woodland Removal for not replanting on peatland sites that are a priority for restoration on ecological grounds, and on those peatlands that are not a priority for restoration when there would be a significant greenhouse gas benefit to restoring degraded peat (Forestry Commission Scotland 2009). Guidance (Forestry Commission Scotland 2015, 2016) provides a decision framework for this, but the underpinning evidence is limited in some important areas. Therefore, the question of what effects the drainage and planting have had on peat bogs, and the likely effects of restoration, are issues of critical importance. There are extensive gaps in current knowledge that need to be filled. This article considers the likely effects of forestry on the peatland, and the applicability of currently available research data to the unique circumstances in which UK peatlands were afforested.

## WHAT EFFECTS HAS TREE PLANTING HAD ON RADIATIVE FORCING?

The climatic consequences of afforestation represent the net effect of several interacting processes on the peat bog ecosystem and wider supply-chain considerations. Changes to peatlands encompass physical changes to the peat itself, vegetation changes, changes to carbon sequestration, effluxes of carbon in gaseous and aquatic forms, and other more minor factors which may nevertheless contribute to the overall radiative forcing. This section reviews these processes.

### Drainage and planting effects on carbon accumulation in peat

Undrained peatlands accumulate carbon through primary production, as plants (often non-vascular species such as Sphagnum) photosynthesise. Within an undrained natural bog, carbon sequestered in this way remains within the peat over long timescales (millennia) because dead material will not fully decay within the main body of peat (the catotelm). Drainage and the process of ploughing disrupts the existing vegetation, affecting the amount of carbon sequestered directly to the bog by the living layer (the acrotelm) (Figure 3). Afforestation essentially halts primary production by typical peat-forming bog species, so the ultimate capacity for radiative forcing then largely depends on the fate of carbon sequestered by trees and by the response of the peat stored in the catotelm.

Peat in a natural bog is divided between the aerated acrotelm and the deeper, constantly

waterlogged, catotelm (Ingram 1978, 1983). The boundary between these two layers is the deepest point to which the water table falls under normal conditions. Undrained peat bogs typically have a high water table, commonly within 10–20 cm of the surface of the peat, but this is substantially lowered with afforestation. Lowering the water table through drainage is arguably the most important factor for successful afforestation, providing the aeration that is essential for growth of the roots of most tree species (Braekke 1983), changing the physical and chemical properties of the peat, and affecting hydrology (Braekke 1987, Holden 2004). Planted forests lower the water table further when canopy closure leads to increased interception and evapotranspiration (Sarkkola *et al.* 2010).

Drainage of a peat soil gives rise to three important processes: primary consolidation, secondary compression and oxidative loss (or peat 'wastage', discussed below). Primary consolidation occurs rapidly following drainage and is caused by loss of water from large pore spaces within the peat. Secondary compression occurs because more tightly-bound water is slowly squeezed from the peat matrix by the weight of peat material no longer supported by the bog water. In addition, the peat may be further compacted by the weight of growing trees (Hobbs 1986). These various processes cause subsidence of the ground surface and ultimately cracking of the upper peat, which can lead to deeper aeration (Pyatt & John 1989, Pyatt *et al.* 1992). Furthermore, any clearing or re-grading of the drainage system will stimulate a new round of primary consolidation before the slower, steady processes of secondary compression and oxidative loss resume (Wold 1976).

The horizontal 'zone of impact' associated with forest blocks on adjacent peatlands has yet to be determined for carbon, hydrology and bog vegetation, though effects on peatland birds are well-established (Wilson *et al.* 2014). There has been only limited monitoring of long-term changes in surface morphology, vegetation assemblages, hydrology and peatland microtopography, meaning that current estimates are based largely on relatively short-term studies, often of hydrology. These estimates currently range from 2–3 up to 50–60 metres, but some hydrological models suggest that drainage effects may extend for several hundred metres in some circumstances (Holden 2005).

The net increase in radiative forcing caused by the effect of physical changes in the peat on carbon storage may be added to by the direct radiative effect through changed surface albedo of forest plantations. Trees can affect snow cover and where trees are felled, the surface environment can have a very high



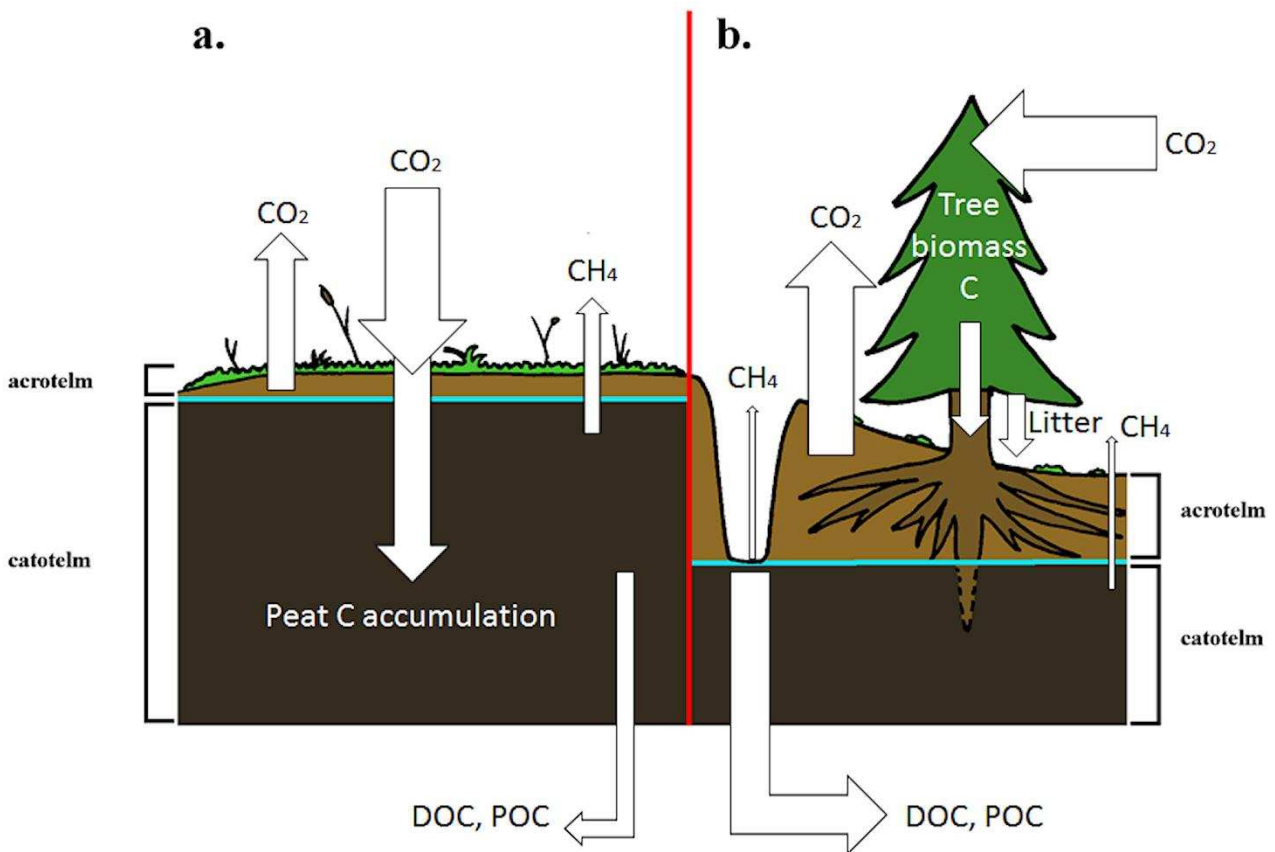


Figure 3. A lowered water table (blue line) gives rise to different rates of carbon loss and accumulation in peatland systems (a) before drainage and (b) after drainage and afforestation, which also cause subsidence. Carbon dioxide ( $\text{CO}_2$ ) production will increase with aeration of the upper layer of peat, with a reduction of methane ( $\text{CH}_4$ ) production from waterlogged peat. Loss of aquatic dissolved and particulate carbon (DOC and POC) may be increased through drainage. Carbon is taken up by vegetation in both scenarios. In an undrained bog some of this will go on to be stored in peat over long timescales, whereas in the drained system it will form tree biomass, eventually reaching the soil as litter and roots or being removed from the site as harvested timber. Unlike Sitka spruce, lodgepole pine is tolerant of waterlogging, and its roots can extend below the water table. The peat beneath the tree crop will have increased dry bulk density compared with the non-afforested peat bog. Flux magnitudes indicated by arrow widths are indicative and open to varying and different degrees of uncertainty and to variation with site conditions.

albedo leading to a cooling effect (Lohila *et al.* 2010). Such effects are rarely considered but may be significant.

### Greenhouse gases

The depth of the water table below the ground surface is a key driver of greenhouse gas (GHG) balance, as this determines the volume of peat exposed to aeration and consequently microbial production of both  $\text{CO}_2$  and  $\text{CH}_4$  (Drosler *et al.* 2008). Lowering the water table during afforestation has the potential to significantly affect the fluxes of both of these GHGs from peat to the atmosphere (Figure 3).

In the permanently waterlogged catotelm, bacterial decomposition is inhibited by low temperature, pH, and oxygen availability (Freeman *et*

*al.* 2001b). In these anoxic conditions  $\text{CH}_4$  is an end product of anaerobic decomposition through several pathways (Lai 2009). As it moves up through the acrotelm a large proportion of this  $\text{CH}_4$  is oxidised by methanotrophic bacteria. Lowering the water table in peatland afforestation increases the depth of air penetration into the normally-waterlogged catotelm peat and thereby the space in which  $\text{CH}_4$  can be oxidised, and thus typically leads to a linear decrease in  $\text{CH}_4$  efflux (Moore and Knowles 1989).

Simultaneous with the reduction in  $\text{CH}_4$  efflux, lowering the water table with peatland afforestation leads to increased efflux of  $\text{CO}_2$  through oxidative loss, or peat 'wastage'. Drainage enables oxygen to penetrate into the catotelm peat, exposing the long-term carbon store to oxidative decomposition by

bacteria and fungi, leading to increased production of CO<sub>2</sub> (Eggelsmann 1975, Hobbs 1986). The loss of the unique structure and function of the aerated acrotelm may lead to the bog becoming a single-layered haplotelm bog (Ingram & Bragg 1984).

Peatland drainage is, therefore, likely to have opposing effects on these two GHGs, increasing CO<sub>2</sub> and reducing CH<sub>4</sub> effluxes. While more carbon is lost to the environment as CO<sub>2</sub>, CH<sub>4</sub> has a global warming potential over 100 yr (GWP<sub>100</sub>) 34 times greater than CO<sub>2</sub> when climate-carbon feedbacks are considered (IPCC 2013). In terms of fluxes from peat, it is likely that the CO<sub>2</sub> increase outweighs the CH<sub>4</sub> decrease and the net effect is to promote climate warming (Martikainen *et al.* 1995, Alm *et al.* 1999).

CO<sub>2</sub> and CH<sub>4</sub> are the most important GHGs arising from peatlands, but nitrous oxide (N<sub>2</sub>O) may also be significant in some situations. The GWP<sub>100</sub> of N<sub>2</sub>O is 298 times that of CO<sub>2</sub> when climate-carbon feedbacks are considered (IPCC 2013). Fluxes of N<sub>2</sub>O in peatlands are typically small but can become substantial in fens or when peatlands are exposed to N in fertiliser, as in some peatland afforestation. However there are few studies which directly consider the effect of afforestation on N<sub>2</sub>O flux (Maljanen *et al.* 2010).

Beyond the direct effect of afforestation on the carbon balance of peat there are other factors which may also result in GHG production. Emissions from vehicles and machinery, as well as road and steel fence construction, also have significant GHG implications for the initial ploughing, planting, interim management and final harvesting of any forestry site (Morison *et al.* 2012).

### **Aquatic carbon**

Aquatic carbon is exported from peatlands via watercourses, principally as dissolved organic carbon (DOC) and particulate organic carbon (POC). Both of these fluxes may be affected by afforestation. DOC concentration in streams correlates positively with the presence of organic soils and peats in a catchment (Hope *et al.* 1997, Aitkenhead *et al.* 1999). Higher outflow of water either drained from the system or lost through consolidation and compression carries with it more aquatic carbon. This process will continue slowly but indefinitely in a drained system. DOC may enter the atmosphere downstream through other degradative pathways, usually through rapid emission as CO<sub>2</sub>, and may be a significant GHG source in upland areas (Freeman *et al.* 2001a). The pathways of POC to the atmosphere are less certain (Rowson *et al.* 2010).

Disruption caused by on-site activity such as ploughing, tree planting and the continuing

maintenance of drains is also associated with increased concentrations of DOC and POC in streams draining the forest stand. Later, disruption to the peat surface caused by tree thinning or felling can lead to further aquatic carbon loss for several years after the trees are removed (Cummins & Farrell 2003). This loss of carbon through aquatic pathways may depend on variables including nutrient content of the peat (Nieminen *et al.* 2015), catchment properties (Holden 2005) and weather patterns (Koehler *et al.* 2009).

### **Carbon accumulation in tree biomass**

Any loss of carbon from peat soils may be offset by gains of carbon stored in tree biomass, litter and new soil organic matter. The true carbon balance then depends partly on the fate of the wood produced (Minkinen *et al.* 2002). The quality and longevity of the wood products that arise from forestry will determine whether or not the harvested portion of the carbon captured by the trees is sequestered over long timescales (Laine *et al.* 1992, Ojanen *et al.* 2013). In areas with high yield and high-quality wood this timber may be used for long-lifespan uses such as construction, effectively storing the carbon for many decades or even centuries. However, forestry crops on bogs in the UK are often of such poor quality that much of the wood goes for pulp, fuel and other low-grade uses, returning carbon to the atmosphere much more quickly (Thompson & Matthews 1989, Artz *et al.* 2013). The portion of the carbon captured by the trees that is left below ground when they are felled consists of roots, litter (root, needle, branch, etc.) and soil organic matter derived from these. In addition, the stumps, branches and top parts of the stems are normally left on the ground after harvesting. The fate of these below-ground and surface components containing tree-derived carbon also influences the true carbon balance (Vanguelova *et al.* 2017). The true climate consequences of peatland forestry are further complicated by the role of the wood produced in the supply chain and the potential for timber to replace alternative materials with high carbon footprints such as plastics and concrete.

The wetness of naturally treeless British bogs may contribute to an increase in trees lost to wind-throw (Figure 4). Many peat bogs used for forestry remain wet even after drainage, leading to the development of shallow and often uni-directional root plates confined by cracks beneath the ploughing furrows (Lindsay & Bragg 2004). This, combined with the very windy climate of many UK peatland forest regions, makes trees more prone to toppling (Ray & Nicoll 1998). Wind-throw will reduce timber yields, and may force earlier harvesting (Gardiner & Quine 2000), reducing the quantity and quality of wood





Figure 4. A wind-thrown lodgepole pine with exposed root plate at Bad a'Cheo, Rumster Forest, Caithness.

products and so reducing the residence time of carbon in the tree biomass. Lodgepole pine is especially prone to wind-throw (Nicoll *et al.* 2006).

#### **Approaches to measuring carbon loss from peatlands**

From the above discussion, it will be clear that afforestation can affect the peatland carbon budget and radiative forcing more generally through many mechanisms. Studies have taken several approaches to quantifying these effects (Table 1). Many studies attempt to assess peat carbon balance by directly measuring the key fluxes of GHGs and aquatic carbon (although aqueous carbon is considered less often in the literature). Methods such as cover boxes ('chambers') or eddy covariance towers use infra-red gas analysis (IRGA) to measure GHG fluxes in real time in the field, replacing older methods using gas sampling for chromatography or quadrupole mass spectrometry (QMS), or recording weight change in soda lime. Using these methods to understand the way in which carbon is imported to or exported from

peatlands can help to understand processes over short timescales. Typically, forestry on bogs requires a programme of site drainage followed by over forty years of tree growth. As a result, short-term studies of carbon fluxes in the system may not accurately describe the carbon change in the system over longer timescales. This is important as GHG emissions can be highly variable over time (Klemedtsson *et al.* 2008).

Another approach to measuring changes in soil carbon is to use a whole-column inventory of the carbon stock in the peat (Pitkanen *et al.* 2013). This typically involves coring a column of peat, then determining the carbon content through the measurement of dry bulk density followed by either direct elemental analysis or deriving a value from the amount of organic material lost on ignition and an assumption of about 50 % as the proportion of carbon in the organic matter (F.M. Chambers *et al.* 2011). Such carbon analysis allows an assessment of the net exchange of carbon with the environment over long timescales, although this does not identify the

Table 1. Methods used for determining carbon budgets in peatlands.

Assessment type	Methods	Timescale	Advantages	Disadvantages
carbon flux	cover box (GHG)	usually between days and months	continuous, precise data	high cost equipment, flux from trees not measured precisely
	eddy covariance tower (GHG)		continuous, precise data	high cost equipment, ground level processes missed
	gas sampling, gas chromatography (GHG)		precise data	data not continuous, analysis can be expensive
	soda lime measurement (GHG)		low cost	only measures CO <sub>2</sub> , imprecise, prone to underestimates
	water sampling, elemental analysis (DOC/POC)		precise	data not continuous, high cost
carbon stock	coring, bulk density, carbon analysis	the whole age of the peat	provides complete picture of carbon loss or gain, no long-term monitoring	no information about fine-scale processes, only total carbon, reliable stratigraphic markers required
	optical or satellite surveys of subsidence	decades, depending on age of original records	low cost, quick	subsidence an unreliable proxy for carbon loss, original surveys may be of poor quality

specific gas and aqueous components. The use of stratigraphic markers in the peat allows the age of a sample to be identified (Pitkanen *et al.* 2013), meaning that direct comparisons can be made between peat of the same age. This analysis can be paired with analysis of carbon in the trees to determine net balance. Laiho and Pearson (2016) highlight a number of issues which must, nevertheless, be considered when using such an approach.

A less exact method of determining loss of carbon stock on sites that have historical ground level surveys is to use subsidence as a proxy for loss of carbon. While this method is relatively low-cost where historical records of ground levels exist, subsidence is an unreliable indicator of carbon loss as it is often based on initial surveys which can be decades old and of poor quality, with estimates produced in this way “roughly determined” at best (Hommeltenberg *et al.* 2014). In addition, it ignores the compaction and compression that usually occurs.

## AVAILABLE RESEARCH

### **Applicability of previous research to UK peatlands**

Much of the work on the effects of peatland afforestation has been carried out in Fennoscandia and Russia. Forestry is particularly widespread on drained peat in Finland, with up to 25 % of the nation’s exploited forests growing on peat (Laiho & Laine 1997) and an area of 4.8 Mha of peatlands drained for forestry (Ojanen *et al.* 2014). Even so, there is little evidence from such sites of the effect of clear cutting on GHG balance, nor of the long-term balance over a full stand rotation and subsequent rotations. The majority of available data have been obtained from minerotrophic fen sites or naturally wooded bog sites, both of which tend to have greater timber production than do ombrotrophic bog sites (Minkkinen *et al.* 2002, Drosler *et al.* 2008, Maljanen *et al.* 2010). These issues, combined with the differences in climate between Fennoscandia and the UK, mean that any comparison between forestry on peatlands in Nordic countries and afforestation of peat bogs in the UK and Ireland must be made with considerable care and may sometimes be inappropriate.

UK blanket bogs are naturally treeless, at least in their broad central expanses, requiring cultivation and more fertiliser than would be used elsewhere (Laine *et al.* 1995). The natural or pre-existing conditions on many of the Finnish peatlands are very different from the UK, and typically may include dwarf trees and scrub (Laiho & Laine 1997) or even

a significant pre-existing tree cover (Minkkinen *et al.* 2002). Fennoscandian bogs typically have peat with inherently very low hydraulic conductivity (Päivänen 1973), so that water table drawdown in response to drainage is probably more limited in depth and extent, with resulting aeration of the peat more limited. In addition, many peatland sites drained for forestry in Fennoscandia are minerotrophic fens and thus the required site treatments, planting methods and suitability for silviculture differ significantly from blanket bog peatlands of the UK and Ireland (Minkkinen *et al.* 2002, Maljanen *et al.* 2010). In Fennoscandia trees are often not actively planted. Drains are instead dug across peatland systems in order to encourage growth of existing trees which grow sparsely or in a variety of growth forms prior to drainage. While work on site is often required to cope with forest regeneration, forestry activities are generally restricted to deep ditching and fertiliser application, along with appropriate thinning as the forest develops (Päivänen & Hånell 2012). In consequence, such peatlands suffer less direct disruption during site preparation than in the UK, and this allows much of the original vegetation to remain and leaves the peat relatively undisturbed (Laine *et al.* 2009).

The use of closely-spaced plough furrows between the deeper drainage systems is, thus, almost unique to the UK and Ireland and this may explain many of the observed differences between peatland forestry responses here compared with those reported from the rest of northern Europe. These differences are worth considering in detail before Fennoscandian evidence is used to inform UK policy, as emissions from UK peatlands are likely to be much greater.

### **Evidence from Ireland and the UK**

There has been limited work on afforested peatlands in the UK and only a few studies have considered the consequences of peat bog forestry for carbon storage (Table 2). Reviews and carbon accounting studies have often integrated Fennoscandian data to argue that planting on peat would produce a net carbon accumulation in UK peatlands over the first 100 years (Cannell 1999, Worrall *et al.* 2010).

In County Galway, Ireland, Byrne & Farrell (2005) examined CO<sub>2</sub> fluxes from afforested blanket peat. They found that CO<sub>2</sub> loss from drained and planted peat was similar to estimated uptake of carbon by the tree stands, suggesting that there would be no net loss of carbon (Byrne & Farrell 2005). However, DOC and POC export from the site were not considered, meaning that total carbon loss was likely to be greater than uptake by the trees. Furthermore, the ‘soda lime’ method was used to

Table 2. Published empirical studies on carbon in afforested peat bogs in the UK and Ireland.

Authors	Year	Location	Type of peatland	Timescale of study	Measurement	Method
Byrne & Farrell	2005	Cloosh Forest, County Galway, Ireland	ombrotrophic blanket bog	two 24-hour measurements, repeated 13 times	CO <sub>2</sub>	soda lime
Hargreaves et al.	2003	Auchencorth Moss, Midlothian, Scotland	extensively drained ombrotrophic blanket bog	22 months, continuous	CO <sub>2</sub>	eddy covariance
		Bealach Burn, Sutherland, Scotland	ombrotrophic blanket bog	month-long continuous measurements, repeated at different aged forest stands		
		Channain Forest, Sutherland, Scotland	peat of 1m depth			
		Mindork Moss, Newton Stewart, Scotland	peat of 2m depth			
Yamulki et al.	2013	Flanders Moss Forest, Scotland	ombrotrophic raised bog	two years, 2–4 week intervals	CO <sub>2</sub> , CH <sub>4</sub> , N <sub>2</sub> O	chamber flux



measure the CO<sub>2</sub> flux from the peat; an approach that the authors acknowledge underestimates CO<sub>2</sub> relative to direct instrumental measurements. And the sampling was for relatively short time periods only.

Hargreaves *et al.* (2003) conducted a study of carbon flux from three afforested bogs in Scotland representing different tree maturities, and one unplanted control site. This study had continuous eddy covariance assessment for over a year in the control site, but only extrapolated from shorter periods of measurements in the afforested sites. The article concluded that afforested peatlands will accumulate more carbon due to forestry than would be lost because of planting and drainage. This was believed to hold true for 90–190 years, after which restoration should take place because the amount of carbon in the tree biomass and peat would fall below that which would have been sequestered by undrained peat (Hargreaves *et al.* 2003). However, these conclusions are questionable. In fact, the site used to provide the baseline control had previously been drained extensively, and so provided unusually low carbon accumulation values. Also, in the ‘mature’ afforested site, canopy closure was not complete and the stand was up to 30 yr from a full rotation, meaning that the carbon loss from a large proportion of the life of the forest stand was not properly accounted for, nor were DOC and POC losses considered (Lindsay 2010).

Another important UK-based study is that of Yamulki *et al.* (2013), who studied gas fluxes and DOC loss from sites at West Flanders Moss in Scotland. While suggesting that drainage increases GHG emissions by 33 %, they concluded that increased CH<sub>4</sub> emissions from rewetted bogs would outweigh the reduced CO<sub>2</sub> emissions, meaning that restoring forest to bog is likely to increase potential warming effects on climate (Yamulki *et al.* 2013). This article has been criticised by Artz *et al.* (2013), who pointed out that there were problems with the control being unrepresentative of undrained bog, with higher than expected CH<sub>4</sub> fluxes, and that there were calculation errors. Further investigation revealed that the control was in an area that had been dug out as a reservoir for flushing cut peat into a nearby river around 100 years previously. The flux work also ignored above-ground tree respiration, comparing below-ground CO<sub>2</sub> flux under forest stands to total above- and below-ground flux in the control. In addition it was noted that carbon budgets of restored sites change over time, and as restored sites mature the vegetation cover becomes less ‘patchy’, producing a stronger CO<sub>2</sub> uptake which would make restoration seem more beneficial (Artz *et al.* 2013).

## DISCUSSION

The evidence base for the effects of afforestation on UK peat bog carbon is weak, and research is often underpinned by data taken from other regions, particularly Fennoscandia. Such studies rely on assumptions that may not hold for conditions in the UK. There is also a bias within the research towards measurement of gas flux without considering other pathways of carbon loss from the system. At present it cannot be reliably determined whether afforestation of open UK peatlands exacerbates or ameliorates climate change.

As existing forests on peat come to harvesting age, decisions must be taken to either restock trees or, where possible, to restore bog habitats. The benefits of restoration on biodiversity are well understood. As the effects on carbon are more uncertain, work is urgently required to plug gaps in current knowledge (IUCN 2014).

Better data on the yields, quality and ultimate use of peat bog forests in the UK are needed. There must also be a proper quantification of other aspects of climate effects including fossil fuel use in ploughing, planting, fencing, fertilising, drain maintenance, road building and the effects on albedo, emissions from transport, and the fate of the wood products.

Further use of whole-column inventories should be made to provide peat carbon budgets over the life of a plantation, particularly if the ground is to undergo restocking. Such carbon stock research must be integrated with flux studies to provide a complete long-term picture of total changes in carbon storage and the processes by which these changes occur, which will determine the loss of carbon to the atmosphere relative to accumulation in tree biomass and quantify any resulting global warming potential.

A wide range of organisations (government, academic, charity and non-government) are now addressing the effects of peatland forestry. A coordinated effort is required to plan and share peatland forestry research, to provide a sound body of evidence for approaching policy decisions. Work on the carbon effects of forestry needs to be understood in relation to research on the economic and ecosystem services provided by peatlands. This is a particular priority in the Flow Country, the UK’s most extensively afforested peatland region and focus of this special issue of *Mires and Peat*.

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## AUTHOR CONTRIBUTIONS

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