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1 **Abstract**

2 Turbidity of water due to the presence suspended sediment is measured and interpreted in a variety of
3 ways, which can lead to the misinterpretation of data. This paper re-examines the physics of light-
4 scattering in water, and exposes the extent to which the reporting of turbidity data is inconsistent. It is
5 proposed that the cause of this inconsistency is the fact that the accepted turbidity standards USEPA
6 Method 180.1, ISO 7027 and GLI Method 2 are mutually inconsistent, as these standards give rise to a
7 large number of measurement units that are not based on the optical properties of light absorption and
8 scattering by suspensions in water, but by the arbitrary definition of the degree of turbidity being due
9 to a concentration of formazin or other similar polymer-based calibration standard. It is then proposed
10 that all turbidity-measuring devices should be calibrated with precise optical attenuators such as ND
11 filters. Such calibration would allow for the definition of a beam attenuation coefficient (BAC) for every
12 turbidity measuring instrument which would be cross-comparable with any other instrument calibrated
13 in the same way. The units for turbidity measurements should be based on attenuation and reported as
14 dB m^{-1} . It is also proposed that a new standard should be drafted according to this attenuation-based
15 method, and this new standard should also define the nomenclature for reporting data collected at any
16 specific scattering angle in terms of an attenuation in dB m^{-1} . The importance of multi-parameter
17 turbidity measurements for the improvement of the quality of turbidity data, and the application of
18 parameter-rich data sets to new methods of sediment characterization are discussed. It is suggested that
19 more research into multi-parameter turbidity measurements is needed, as these new methods will
20 facilitate an increase in parity between turbidity and suspended sediment concentration (SSC), a
21 relationship that is subjective.

22 **Key words**

23 Turbidity; Suspended Sediment; Sediment Transport

24 Introduction

25 The term “turbidity” is used widely throughout the physical sciences, and is interpreted in different
26 ways in different contexts. It is commonly used to describe the optical clarity of a fluid (for example,
27 the atmosphere), but for the purposes of this paper it refers to another common usage of the term
28 which is the optical clarity of water. The presence of suspended particulates, dissolved inorganic
29 chemical species, organic matter content and temperature can all affect the turbidity of a body of
30 water. Investigators from different fields (waste water treatment; drinking water quality; forestry; civil
31 engineering, aquaculture and ecology), and from the sub-disciplines within physical geography (fluvial;
32 marine; glacial; coastal and estuarial) use turbidity measurement as a surrogate relative indicator of
33 some other physical property, typically suspended sediment concentration (SSC) or total suspended
34 solids (TSS). The amount of literature available on the subject of water turbidity is large, and a number
35 of reviews have already been undertaken by investigators from some of the sub-disciplinary groups
36 (Bilotta & Brazier 2008; Davies-Colley & Smith 2001; Kerr 1995; Ziegler 2003). There is however, some
37 disagreement about what turbidity actually *means*, partly due to the different sub-disciplinary
38 contexts in which the term is used, and partly because of the way in which the various measurement
39 standards are assumed to be based on a correct *a priori* understanding of the physical processes of
40 light-scattering and absorption.

41 Why is turbidity measurement important? The answer to this question depends on the
42 perspective of the investigator. Some researchers are purely interested in the effect that the
43 attenuation of light has on, for example, aquatic ecosystems, so that knowledge of the mass
44 concentration of the suspended particles is not always the primary concern. In this case other
45 parameters of interest include the reduction of visual range in water (affecting the ability of predators
46 to hunt), and the amount of light available for photosynthesis (Bilotta & Brazier 2008). Other
47 investigators are concerned directly with the study of sediment-transport processes, in which case
48 knowledge of the mass concentration of the suspended particles and other parameters such as the

49 particle-size distribution (PSD) is highly desirable for a number of reasons. Turbidity measurement is
50 important in this context, as although the turbidity measurement itself is heavily biased by the PSD
51 (Gippel 1989), it is not specifically designed to provide detailed information about the PSD. For
52 example, knowledge of particle size is important as the transport of fine sediment derived from
53 different land uses through catchments will impact directly on ecosystem services, such as the
54 provision of drinking water. Fine sediment delivery into river systems is also known to cause problems
55 such as irritation to fish gills whilst it is in suspension (Davies-Colley & Smith 2001). Bilotta & Brazier
56 (2008) summarize the effects of what they refer to as suspended solids (SS) on periphyton and
57 macrophytes, invertebrates and salmonid fish species. The displacement of many fish species can
58 often be due to an increase in turbidity caused by the cumulative effects of fine sediment introduced
59 into the riparian environment as a direct result of human activities such as deforestation (Kerr 1995),
60 or by natural events such as sediment-transport by stormwater runoff. The use of turbidity
61 measurement as a surrogate indicator for parameters such as suspended sediment concentration has
62 been explored by many researchers, as reviewed by Ziegler (2003). It has been shown that the particle
63 size distribution (PSD) of a homogenous sediment can vary temporally from its source (e.g. hillslope
64 runoff) as it is transported through a catchment into a stream, due to a variation in the relative
65 proportion of aggregates (flocs) present in the measured flux (Slattery & Burt 1997). Therefore
66 knowledge of how the PSD varies dynamically in this fluvial context due to a variability in the degree
67 of flocculation (DOF) is important for the study of the transport processes of both sediment and
68 organic species in flocs (Williams et al. 2007). There is clearly some variation in the importance given
69 to the parameters of turbidity by the different sub-disciplinary groups, and so the aim of this paper is
70 to evaluate how relevant turbidity measurement is to the study of sediment-transport processes
71 specifically, and to propose methods for the improvement of the measurement and reporting of
72 turbidity in a general context. The steps required to achieve this evaluation are given by the following
73 list of objectives:

- 74 1. To analyse critically the measurement methodologies described in the literature
75 including any inconsistencies in nomenclature of measurement principles.
- 76 2. To review briefly the physics of light absorption and scattering processes in water in
77 order to provide an underpinning for the discussion of the definition of terms
78 according to various investigators from different sub-disciplinary groups.
- 79 3. To present a critique of the measurement units, calibration methods and standards
80 applicable to the measurement of turbidity, SSC and TSS, and to examine of the origins
81 of the relationship between turbidity measurements and the implied properties of
82 suspended sediment. This step is vital because the cross-comparability of turbidity
83 data obtained in the field is often invalid due to a widespread reliance on the assumed
84 integrity of Formazin calibration methods.
- 85 4. To propose, based on objective 3, that a new turbidity instrumentation standard is
86 required, and to describe its fundamental content.

87 **Turbidity measurement principles and nomenclature**

88 The measurement of turbidity is split into two basic methodologies: *turbidimetry*, in which the degree
89 of *transmission* of light is determined, and *nephelometry*, in which the degree of light-scattering is
90 evaluated (see reviews by Ziegler, 2003 and Lawler, 2005). This division has its roots in the
91 mathematical descriptions employed to model the various phenomena. In the case of turbidimetry,
92 the appropriate theories are due to Beer (1852) and Lambert (1760) ; as for nephelometry, many
93 theories and models have been developed to describe a range of scattering processes, and these
94 models are mostly derived from Mie theory (Mie 1908). Nephelometry itself is sub-divided into three
95 further categories which are *forward-scattering*, *side-scattering* and *back-scattering*. Side-scattering
96 is generally accepted to be a measurement angle of 90° to the incident beam, although the existing
97 standards impose different upper and lower bounds on that value (Table 3). Forward-scattering ($0^\circ <$
98 $\vartheta < 90^\circ$) and Back-scattering ($90^\circ < \vartheta < 180^\circ$, often referred to as *optical back-scattering* or *OBS*)

99 however, do not have a well-defined relative measurement angle. Different instruments employ
100 different measurement angles, and these values are not always reported.

101 **[Insert Figure 1.]**

102 Before continuing with the discussion another ambiguity in terminology must be addressed. The
103 definition of the scattering angle in terms of where the 0° position is located spatially also varies
104 throughout the literature ([Table 1](#)). For example in some cases a forward-scattering angle is stated,
105 which *implies* that the transmitted (direct) beam is located at 0° (Agrawal et al., 2008 and Jansson,
106 1992). Contradictory to this position, Bilro et al. (2010) define the transmitted beam as being located
107 at the 180° position. In one instance two contradictory diagrams are presented in the same paper
108 (Sadar 2004, pp.8-9), and in many other cases the scattering-regime nomenclature is not associated
109 with a specific scattering angle (e.g. Fugate & Friedrichs, 2002).

110 The interpretation that is adopted throughout this paper is that the scattering-angle is
111 specified in terms of a detector placed at a position with respect to the incident beam *after* a physical
112 interaction has occurred in the sample, i.e. the direct beam detector is placed at the 0° position
113 (denoting “pure” attenuation measurement), forward-scattering detectors are placed anywhere from
114 $0^\circ < \vartheta < 90^\circ$, a side-scattering detector is placed at exactly 90° , and back-scattering detectors are
115 placed at $90^\circ < \vartheta \leq 180^\circ$.

116 **[Insert Table 1]**

117 **The physics of light absorption and scattering through turbid water**

118 **A brief review of optical theories**

119 To understand the physics of light scattering by particles suspended in water, it is necessary to have
120 some knowledge of the mathematical models employed to describe the various absorption and

121 scattering processes. Fundamental theory and mathematical model development are continually
122 progressing in this area, but the basic points of interest pertinent to the understanding of turbidity in
123 water for the practical investigator are summarised in this section. Three main theories are discussed:
124 Rayleigh theory, Mie theory and geometric optics. Also discussed are two theories that can be
125 considered as approximations to Mie theory for specific conditions. These are the *Fraunhofer*
126 *diffraction* theory (FDT) and the *Anomalous diffraction* theory (ADT) of Van De Hulst (1957). The
127 reason that these two theories are considered here is that they both yield computationally fast
128 algorithms that are utilised by laser-based particle-sizing instruments. These instruments are used
129 widely in suspended particle analysis (organic and inorganic) both *in situ* and off-line in laboratories,
130 and are extensively employed for suspended sediment characterization.

131

132 Rayleigh and Mie scattering

133 The third Baron Rayleigh formulated his scattering theory to account for the blue colour of the sky
134 (Strutt 1871). Rayleigh scattering involves particles that are much smaller than the wavelength of the
135 incident light, and are also defined as being *optically soft* – meaning that the particles are limited to
136 having a refractive index very close to 1 (air molecules in the case of Rayleigh’s model). Rayleigh
137 demonstrated that scattering from small particles is strongly wavelength dependent in favour of the
138 shorter wavelengths and is spatially isometric (i.e. scattered equally in all directions), hence the blue
139 colour of the sky. He determined that this blue colour is predominant because the scattered light
140 intensity is inversely proportional to the fourth power of the incident light wavelength, i.e. the shorter
141 wavelengths of light (e.g. blue end of the visible spectrum) are scattered more readily than the longer
142 wavelengths of light (e.g. red end of the visible spectrum).

143 Gustav Mie originally developed his theory to explain the colouration of metals in the colloidal
144 state (Mie 1908). Mie theory successfully explains the dominance of forward scattering where

145 particles are of a similar size to or larger than the incident wavelength of light, unlike the case of
146 isotropic scattering of light by much smaller particles as in Rayleigh scattering.

147 In order to get some sense of the particle size ranges that are applicable to the different
148 scattering regimes it is first necessary to define the *dimensionless size parameter* x ,

$$149 \quad x = \frac{2\pi r}{\lambda} \quad (1)$$

150 where r is the spherical particle radius [m] and λ is the wavelength of the incident light [m]. Figure 2
151 shows how the forward-lobed nature of a set of light intensity distribution functions develops as x
152 increases from 0.1 to 10. These spatial intensity distribution functions are also known as *scattering*
153 *phase functions*, which are calculated using Mie theory.

154 **[insert Figure 2.]**

155 Geometric optics

156 Geometric optics, otherwise known as ray optics, describes the light traversing a medium in terms of
157 a straight path (hence “ray”). It explains *refraction*, in which there is a change in direction of a light ray
158 at the interface between two regions with differing refractive indices. It also accounts for *reflection*
159 and *absorption*, and is best applied in situations where the wavelength of light is much less than the
160 size of the scattering particle. Figure 3 depicts a simplified diagram of scattering and absorption
161 processes of a particle suspended in water as viewed from the perspective of ray optics.

162 **[insert Figure 3.]**

163 Fraunhofer diffraction theory (FDT)

164 Fraunhofer diffraction occurs at small angles to the forward-scattered beam, i.e. $<30^\circ$. Under these
165 conditions of wavelength and scattering angle, FDT is a useful approximation to Mie theory, and is
166 popular due to the relative simplicity of its algorithms. Due to the wavelength and particle size

167 restrictions FDT cannot be applied to sub-micron sized particles. For example, the smallest sized
168 sediment particle that could exhibit Fraunhofer diffraction when illuminated by a beam of red light
169 (wavelength 630 nm) would be 6.3 μm , i.e. well above the sub-micron size limit.

170

171 Anomalous diffraction theory (ADT)

172 ADT (Van De Hulst 1957) is a computationally efficient method by which the scattering from small
173 particles can be modelled. The caveat is that the particles must be *optically soft* as in Rayleigh
174 scattering (i.e. they must have a refractive index close to 1), and they must also have a large size
175 parameter $x \gg 1$.

176

177 The single scattering albedo

178 The single scattering albedo, denoted ω , is a useful unitless quantity defined as the ratio of
179 scattering efficiency to total extinction efficiency. If the attenuation observed by a detector placed in
180 the “direct beam” configuration as in Figure 1 was due entirely to absorption, then $\omega = 0$. When the
181 observed attenuation is due to scattering processes alone, then $\omega = 1$. The scattering albedo is useful
182 when describing the particle size range that can be effectively modelled by the various regimes
183 (Rayleigh, Mie etc.). A graph of scattering albedo (ω) versus size parameter (x) is presented by
184 Moosmüller & Arnott (2009, Fig.1, p.1031), which shows the particle size ranges covered by Rayleigh
185 and Mie theory for particles with a refractive index of 1.55 (similar to that of silica). On this graph,
186 the approximate scattering-model regime boundaries are observed, as shown in Figure 4. The *large*
187 *particle limit* of Mie theory is also shown, and the size parameter at which Mie theory converges
188 with this limit is the point at which geometric optics (not shown on the graph) becomes an
189 alternative scattering model (at $x \approx 2000$).

190 **Light absorption and scattering by suspensions in water**

191 In the terminology of physical optics absorption is a *non-parametric* process, i.e. one that is inherently
192 lossy – meaning that energy is dissipated in the absorbing medium. The *parametric* processes that are
193 to be considered do not involve any imparting of energy to the physical system through which the
194 radiation is traversing, i.e. the wavelength of the scattered light is not altered (elastic scattering). The
195 pertinence of these (and other) theories to the study of suspended particles in general, and suspended
196 sediment specifically, must be considered. Rayleigh theory is applicable to small, non-absorbing
197 (dielectric) spherical particles. Mie theory is the most ubiquitous of the models that is applied to the
198 study of light scattering by suspensions in water. It represents a general solution to scattering from
199 absorbing or non-absorbing spherical particles, with no limits on particle size. Rayleigh theory is less
200 complex to apply than Mie theory, but is limited to small particles. The dimensionless size parameter
201 x (Equation 1) for the scattering regimes, and the equivalent approximate particle size ranges are:

202 $x \ll 1$ Rayleigh scattering (2 nm to 75 nm)

203 $x \cong 1$ Mie scattering (20 nm to 765 μm)

204 $x \gg 1$ Geometric optics (>200 μm)

205 The graph of wavelength vs. particle diameter (Figure 4) shows the accepted boundaries
206 between the various scattering regimes, as adapted from Lelli (2014) and confirmed by Moosmüller
207 & Arnott (2009). Also plotted on the graph are the clastic sediment size ranges that are of interest in
208 this paper.

209 **[insert Figure 4.]**

210 Interpretation of this plot must however be considered carefully, as the data it represents are limited
211 to a *single scattering event* from a *purely spherical* particle. The regime boundaries located at $x=0.02$,
212 $x=0.2$ and $x=2000$ (Lelli 2014 and Moosmüller & Arnott (2009) are not strict demarcation lines (i.e. Mie

213 theory includes Rayleigh theory as $x \rightarrow 0$), but are there to suggest the generally accepted view of
214 where the various models are used with respect to particle size parameter x . These boundaries should
215 be considered to be somewhat blurred when applied to multiple-scattering from non-homogenous
216 suspended sediment particles. Considerable model development is needed to account for scattering
217 from large, non-spherical sediment particles. This work will lead to a redefinition of the scattering
218 regime boundaries as depicted in [Figure 4](#), with new models specific to suspended sediment being
219 represented on the graph. There would also be one omission from the graph, namely Rayleigh
220 scattering. As far as light scattering from suspended sediment is concerned, this theory has no
221 application due to the restrictions in particle size (i.e. very small: < 76.4 nm) and refractive index (i.e.
222 $n \approx 1$). Although Mie theory is limited to small, spherical particles only, it has many extensions that
223 describe much more complex scattering regimes (including multiple-scattering and scattering from
224 small non-spherical particles), and also simpler scattering regimes such as FDT (valid for particle
225 diameter $d \geq 10 \lambda$, and scattering angle $\vartheta \leq 30^\circ$). Other theories such as ADT which as with Rayleigh
226 theory was originally designed for optically soft particles (but in this case with a large x value), are also
227 adaptable to cope with higher refractive indices and non-spherical particles (Liu et al. 1998).

228 There is clearly a need to find a light-scattering model framework that is consistent with both
229 small and large particle scattering, and which is also extensible to many-particle analysis. In the case
230 of back-scattering from suspended sediment it has been shown that the reflectivity of the sediment
231 also has a direct effect on the scattered light intensity (Sutherland et al. 2000), suggesting that
232 geometric optics may play a part in future model development. Without a comprehensive
233 understanding of the complex manner by which particle size, shape and concentration affect the
234 absorption and scattering of light, it will not be possible to interpret what a turbidity measurement
235 actually means.

236 **The definition of the beam attenuation coefficient.**

237 The attenuation coefficient Σ is commonly referred to as the *beam attenuation coefficient* (BAC) in the
238 turbidity literature, but these two quantities are defined in different ways by different authors. It is
239 important that the ambiguities in both the definition and application of the BAC as a method for
240 comparing turbidity data obtained by different methods are appreciated, as these ambiguities can
241 lead to the misinterpretation of that data. The following discussion focusses on how the *a priori* Σ is
242 defined, and then leads on to a definition of the BAC as an expression of Σ in terms of observable
243 quantities, i.e. a *measured* attenuation and the *optical path-length* of the measurement instrument.

244

245 **The attenuation coefficient Σ**

246 Light is absorbed by water and this absorption is a function of the wavelength of the incident light
247 (Figure 5). The strongest absorption occurs at a wavelength of $\lambda = 417.5\text{nm}$ (Pope & Fry 1997) which
248 gives a maximum reduction in transmitted light intensity of 0.05% over a distance of 0.1 m, which is
249 the typical limit to the optical path length of existing turbidity instruments. As this is the worst-case
250 scenario, the absorption of light by water is considered to be negligible in the context of turbidity
251 measurement.

252 **[insert Figure 5.]**

253 Light is also absorbed by any other material that may be suspended in the water. In order to determine
254 practically a value for absorption it is necessary to measure the amount of light transmitted through
255 a given sample of water. This is termed the transmittance, T , which is defined as the ratio of the
256 transmitted light intensity I to the light source intensity I_0 , and has units of Wm^{-2} . The transmittance
257 is also related to the optical depth (Equation 2), τ (effectively the opacity of the medium), and the
258 *absorbance*, A :

259
$$T = \frac{I}{I_0} = e^{-\tau} = 10^{-A} \quad (2)$$

260 A quantitative measure of the optical depth τ can be expressed in terms of the natural logarithm of
261 the transmittance or in terms of the absorbance (Equation 3). This in turn leads to a definition of
262 absorbance with units of the Neper (Equation 4), or in terms of the base-ten logarithm (Equation 5)
263 yielding a decibel quantity.

264
$$\tau = -\ln(T) = A \ln(10) \quad (3)$$

265
$$A = \frac{\ln(T)}{\ln(10)} = -\log_{10}(T) \quad (4)$$

266
$$A = -10 \log_{10}(T) \quad (5)$$

267 This definition of absorbance as a logarithmic function of transmittance is useful as it facilitates a linear
268 relationship with the optical path-length. When a linear relationship between *transmittance* and *path-*
269 *length* is established it then becomes theoretically easier to relate the *absorbance* to the
270 *concentration* of a suspension, which will consequently itself be a linear function.

271 The *a posteriori* description of the attenuation of light through a homogeneous medium is credited
272 to Bouguer (1729) and is also associated with Lambert. It has been called Bouguer's law, Lambert's
273 law (Lambert 1760) and the Bouguer-Lambert law. It states that the attenuation is proportional to the
274 distance travelled through the absorbing medium. The extension to this law which includes a term for
275 the concentration of absorbers is known as Beer's law, or more ubiquitously as the Beer-Lambert law
276 (Equation 6 and Equation 7), which states that the attenuation is proportional to the concentration of
277 the absorbers (Beer 1852).

278 The Beer-Lambert law allows the absorbance to be stated under *ideal conditions*, including the
279 assumption that there are *no scattering processes* occurring in the sample, and that the attenuation
280 is linear along the light path. This law enables the absorbance to be directly related to the

281 concentration of absorbers, c , and the path length l (Equation 6). Equation 7 expresses the same
282 quantity as a transmittance:

$$283 \quad A = \varepsilon c l \quad (6) \quad T = e^{-\varepsilon c l} \quad (7)$$

284 where ε is the absorptivity [m^2 , or $\text{m}^2 \text{kg}^{-1}$] of the absorbers in suspension, and is a constant dependent
285 on the physical properties of the absorbers (i.e. dielectric properties). When defined in these terms,
286 the *attenuation coefficient* Σ can be stated as the product of the absorptivity and the concentration of
287 the absorbers:

$$288 \quad \Sigma = \varepsilon c \quad (8)$$

289 Substituting Equation 8 into Equation 6 gives the absorbance in terms of the attenuation coefficient:

$$290 \quad A = \Sigma l \quad (9)$$

291 The attenuation coefficient can be expressed in Napierian terms or as a decadic quantity (i.e. in
292 decibels). The measured *luminance* (Cd m^{-2}) represents the *power* delivered by the transmitted light
293 beam per unit area. In electronic design it is more common to use decadic terminology to specify
294 measurement instrument parameters such as those used for the determination of light attenuation.
295 If Equation 7 is substituted into Equation 5, then the absorbance can alternatively be stated in decibels
296 (Equation 10 and Equation 11).

$$297 \quad A = 10 \Sigma l \quad (10) \quad A = 10 \varepsilon c l \quad (11)$$

298 It is worth noting that the absorbance A is a dimensionless parameter, and the attenuation coefficient
299 Σ has units of reciprocal length (m^{-1}). However, the absorptivity ε may have different units depending
300 on the context in which the concentration c is expressed (Equation 11). For example, in the case where
301 the concentration is simply the number of absorbers N per unit volume, then the units of
302 concentration are reciprocal volume, i.e. m^{-3} or l^{-1} . Therefore, absorptivity ε in this instance has units
303 of m^2 . In the case of suspended sediment, the absorptivity ε would have units of $\text{m}^2 \text{kg}^{-1}$. It is important

304 to recognise the units stated for absorptivity, as other nomenclature could potentially refer to the
305 same physical quantity. For example, the *mass attenuation coefficient* used in chemistry also has units
306 of $\text{m}^2 \text{kg}^{-1}$. Hence it is prudent to examine the mathematical definition being used within a given text
307 to determine what physical quantity is actually being discussed, and not to rely on the accuracy of the
308 nomenclature at all. Another example of ambiguous nomenclature is highlighted by Figure 5, which
309 shows the graph of the light absorption spectrum of water. The range of this function is referred to as
310 the *absorption coefficient*, and as it has units of reciprocal length (m^{-1}) it is equivalent to the Σ of this
311 discussion (i.e. the *attenuation coefficient*). This multiplicity of measurement units has the potential
312 to cause confusion, since the absorption coefficient has *the same units* as the attenuation coefficient
313 Σ . This is an important point as absorption is *not the same* as attenuation. Attenuation is the *end result*
314 of the effects of the *physical properties* of the medium on the *propagation* of the light waves, and
315 represents a loss of measureable light intensity. Any measured attenuation cannot be presumed to be
316 due to absorption alone ([Figure 3](#)). Scattering of light can occur in all directions, and reflection and
317 refraction of light can also distort any attenuation measurement. For example, Gumprecht &
318 Sliepcevich (1953) suggested that *forward scattering* can distort a true attenuation measurement by
319 *adding* to the transmitted light intensity observed by a detector. This forward-scattering component
320 is referred to as the *extinction coefficient* by Clifford et al. (1995, p.774), who describe it as “*the re-*
321 *formation of light after scattering behind the particle*”, and attribute this effect to the presence of
322 suspended particles of diameter less than approximately $4 \mu\text{m}$.

323

324 BAC – the beam attenuation coefficient

325 The attenuation coefficient Σ is defined for *ideal* conditions, i.e. situations in which the attenuation of
326 light obeys the Beer-Lambert law and is thus concerned with absorption only, although some
327 definitions of BAC include a term for light-scattering (Kirk 1985). However, light-absorption cannot be
328 measured directly; only the *attenuation* of a light source can be determined by direct measurement

329 of light transmitted through a sample. As this attenuation could be affected by other processes besides
330 absorption (e.g. scattering), the absorption itself is not directly observable. The absorption and
331 scattering processes that occur within the sample do not have any bearing on how a transmitted light
332 intensity is measured at a given angle with respect to the incident beam, as the only available
333 parameters are the measurement angle ϑ , and I/I_0 for each ϑ . It is crucial that the BAC is accepted
334 only as a measurement of light *attenuation*, and it cannot by itself be used to infer any *a priori*
335 mechanism of absorption or scattering. It is however conceptually convenient to consider the
336 definition of the BAC as being based purely on the effects of absorption alone (i.e. the ideal conditions
337 of the Beer-Lambert law). The measurement of transmissivity and hence the attenuation of light due
338 to the turbidity of water is referred to in the literature as *turbidimetry* or *transmissometry*. The class
339 of device for performing this measurement is consequently termed a *turbidimeter* or a
340 *transmissometer*.

341

342 A practical definition of the BAC

343 Many devices exist for the measurement of optical transmissivity in water, and in this sense the word
344 “transmissivity” is synonymous with *attenuation* and refers to the measurement of I/I_0 at an angle ϑ
345 of 0° with respect to I_0 , i.e. the “direct beam” ([Figure 1](#)). This measurement leads to the derivation of
346 the BAC by application of [Equation 4](#), such that the BAC in decibels per metre (dB m^{-1}) can be stated
347 as

$$348 \quad \text{BAC} = \frac{-10 \log T}{l} \quad (12)$$

349 where l is the optical path length (m) as determined by the particular instrument used for the
350 measurement.

351

352 Turbidity measurement units, calibration methods and standards

353 A summary of the major turbidity standards

354 The following three standards are in common use throughout the sub-disciplines of water quality
355 assessment. Although other standards do exist, these three are the most commonly cited by
356 researchers into the properties of natural waters. The summaries of these standards are presented in
357 order to highlight some of the technical imprecision inherent in their measurement methodologies.

358 US EPA Method 180.1

359 This standard has been in use in various revisions since the early 1970s. The most recent revision being
360 2.0 (US EPA 1993), which states that it is applicable to the measurement of turbidity in "*drinking,*
361 *ground, surface, and saline waters, domestic and industrial wastes*" (US EPA 1993, p.1). The standard
362 employs the comparison between the light scattered by the test sample to the light scattered by a
363 "*standard reference suspension*" (US EPA 1993, p.1). This reference suspension consists of a defined
364 mixture of two chemicals, hydrazine sulphate and hexamethylenetetramine, to produce a "*stock*
365 *standard suspension*" known as Formazin (US EPA 1993, p.3). A *primary standard suspension* is then
366 created by diluting 10mL of stock standard in 100mL of reagent water. This concentration is defined
367 as having a turbidity of 40 Nephelometric Turbidity Units (NTU). Another acceptable commercially
368 available primary standard based on styrene divinylbenzene polymer is also stated.

369 The instrumentation parameters for the measurement of scattered light by this standard are the use
370 of a tungsten light source with a colour temperature from 2200-3000K, and a beam path-length of not
371 greater than 0.1 m. The detector response should peak at 400-600 nm, and the measurement angle
372 should be $90^\circ \pm 30^\circ$. Note that this is a very broad range of light wavelengths and scattering angles
373 which encompass forward-, side- and back-scattering geometries.

374

375 **ISO 7027**

376 This standard has been in effect in Europe since 1994. It relies in part on the use of light scattering and
377 attenuation by standard suspensions for comparison with the same measurements in a test sample,
378 as with EPA Method 180.1. A notable difference between the two standards is that ISO 7027 dictates
379 the use of near infrared light ($\lambda = 860$ nm) for all measurements. The standard suggests that at
380 wavelengths greater than 800nm the interferences caused by natural colouration of the water (e.g.
381 by dissolved humic substances) can be significantly reduced, an effect which has been observed by
382 Hongve & Akesson (1998).

383 In addition to the measurement of diffuse radiation (i.e. nephelometry) expressed in Formazin
384 Nephelometric Units (FNU – in the range 0-40), the standard also defines a method for the
385 *“measurement of the attenuation of a radiant flux, more applicable to highly turbid waters (for*
386 *example waste or polluted waters)”* (ISO 1999). This measurement is expressed in Formazin
387 Attenuation Units (FAU), in the range 40-4000 FAU.

388

389 **GLI Method 2**

390 This method is explicitly for the determination of turbidity in drinking water. It is a nephelometric *and*
391 attenuation-based *ratio-metric* method based on infrared light of 860 nm wavelength, in common
392 with ISO 7027. The use of dual-beam instruments that have two light sources and two detectors is
393 specified. Each light source is pulsed sequentially, and for each measurement phase a 90° *active*
394 intensity and a 0° *reference* intensity measurement is acquired (Figure 6). A ratio-based algorithm is
395 then used to calculate an NTU value based on the four data points (i.e. two 0° and two 90°
396 measurements). The accepted reason for employing this method is that it improves instrument
397 stability due to interferences caused by the degradation of the light source, the fouling of sensor
398 windows, and the effects of water colouration. It must be noted that the ratio algorithm is not defined

399 in the standard, which implies that the implementation is left to the instrument designer (the topic of
400 ratio methods is considered in greater detail later). As in the previously discussed standards, formazin
401 suspensions are used for calibration. This is an example of a *multiple parameter* measurement
402 method.

403 **[insert Figure 6.]**

404 [A summary of turbidity measurement units](#)

405 The U.S. Geological Survey has summarized currently used turbidity units and their associated
406 standards as reproduced in [Table 2](#) (USGS 2013), with amendments for the scattering angle
407 convention in use throughout this paper .

408 **[Insert Table 2]**

409 Most of the material reviewed for this paper pertains to measurements taken by turbidity instruments
410 that comply with either USEPA Method 180.1 or ISO 7027, and hence the measurement units that are
411 most commonly encountered in the literature are NTU, FNU (specifically for drinking-water
412 assessment) and FAU (specifically for waste-water assessment). The USGS considers these units to be
413 the ones that are most commonly applied to submersible turbidimeters. The other units listed in [Table](#)
414 [2](#) are rarely encountered in the turbidity literature. In addition to the USGS website, another useful
415 summary containing greater detail regarding the applications of the different turbidimeter designs is
416 presented by Sadar (2004). A more concise summary of the standards discussed in this paper is
417 presented by (Ziegler 2003), and this summary is reproduced here ([Table 3](#)) as it provides pertinent
418 and useful aid to the context of this discussion.

419 **[Insert Table 3]**

420 The problem with formazin

421 Formazin is useful as a turbidity standard as it can be reproducibly prepared from raw materials to
422 within $\pm 1\%$, and comprises a wide range of particle shapes and sizes ranging from 0.1 μm to 10 μm
423 (Buzoianu 2000). However, it also has a number of drawbacks as highlighted by Buzoianu (2000):

- 424 • The preparation temperature affects the resulting PSD.
- 425 • Formazin is carcinogenic.
- 426 • Formazin primary standards do not usually state the concentration uncertainty.
- 427 • The stability of formazin standards decreases as the concentration decreases ([Table 4](#)). The
428 dilution ratio can be very high which leads to high uncertainty at low concentrations. This
429 necessitates the use of secondary standards with longer shelf lives, and these standards can
430 have poor repeatability of preparation, they are not formazin (eg latex), and they have
431 different (narrow) PSDs. Hence, the use of secondary standards produces more variation in
432 the response of different measurement instruments to the same nominal turbidity level.

433 **[Insert Table 4]**

434 It is a key fact that all of the units described in the previous section ([Table 2](#) and [Table 3](#)) are derived
435 from a *chemical concentration level* of formazin or a secondary polymer-based standard. By this
436 methodology an increase in concentration is defined as an increase in turbidity. *There is no defined*
437 *relationship between the stated turbidity and the measured light intensity*. The word “concentration”
438 has effectively been replaced by “turbidity” in the definition of these measurement units. For example
439 section 7.3 of US EPA Method 180.1 states “Primary calibration standards: Mix and dilute 10.00 mL of
440 stock standard suspension (Section 7.2) to 100 mL with reagent water. **The turbidity of this**
441 **suspension is defined as 40 NTU**. For other values, mix and dilute portions of this suspension as
442 required.”

443 This definition is a serious issue as “turbidity” in these standard techniques no longer refers to an
444 optical property of water, but rather a chemical concentration of what is in terms of particle

445 classification an unknown distribution of both particle sizes and particle shapes. As the particle-size
446 distribution (PSD) is not known, it is therefore not repeatable between measurements due to factors
447 such as chemical degradation and flocculation during storage of the “stock standards”. Also, the fact
448 that it is deemed acceptable to use secondary standards that will not have the exact same optical
449 response as formazin (Sethi et al. 1997, p.110) suggests a flaw in the methodology at its root, as these
450 “stock standards” are clearly not consistent nor are they traceable.

451 The sphericity of the suspended formazin particles is also not quantified. Sadar (1999) states when
452 describing formazin “*the polymer in solution consists of random shapes and sizes.*” Both PSD (Baker &
453 Lavelle 1984, Ziegler 2003) and sphericity (Gibbs 1978) have been shown to have a significant effect
454 on the light-scattering characteristics of a suspension. Referring back to [Figure 2](#), the dimensionless
455 size parameter x has a large effect on the scattering phase function. For example, nephelometric
456 instruments are most sensitive to particles of $<1 \mu\text{m}$ diameter as in this size-range there is a significant
457 amount of side-scattering, yet the standards do not state the PSD limits required for reference
458 solutions.

459 It has been demonstrated that different instruments measure different turbidity values when
460 calibrated with the same primary standard, due to the differences in instrument design (Buzoianu
461 2000). This is a situation that can occur even when the different instruments are made to comply with
462 the same measurement standard (e.g. EPA Method 180.1), due to the wide design tolerances (e.g. a
463 measurement angle of $90^\circ \pm 30^\circ$). In view of the large uncertainties in the concentrations (and PSDs)
464 of the calibration standards, augmented by the variation in measurement instrument response, there
465 is then a scenario in which *one* stock standard and *two* different measurement instruments (made to
466 the same or different standards) could potentially give rise to not *two*, but *multiple different* initial
467 calibration results ([Figure 7](#)). An inaccurate surrogate model of turbidity has now effectively become
468 synonymous with turbidity itself *by definition* in these standards. This calibration problem has
469 implications for the measurement of turbidity in the field. The cross-comparability of measurements

470 made by different researchers at different sites using different instrumentation is now questionable,
471 even if each researcher has a self-consistent set of repeatable calibration data for their own particular
472 measurement instrument. It is therefore necessary to take a step back and to re-define the chain of
473 measurement at its first and weakest link, which is the Formazin standard, and to establish a new
474 methodology based purely on the calibration of measurement instruments to well-defined light
475 intensities at well-defined wavelengths.

476 **[insert Figure 7.]**

477 **Towards a new turbidity instrumentation standard**

478 In order to move towards a new standard for the design of turbidity instrumentation it is first
479 necessary to take a step back from the accepted suspension-based calibration methods as prescribed
480 by the existing standards. The following discussion attempts to clarify the misconceptions associated
481 with the relationship between SSC, TSS and turbidity, and leads on to a proposed calibration
482 methodology based on the measurement of light-attenuation due to the presence of optical neutral
483 density (ND) filters in the optical beam path. To complete the new standard, a new nomenclature
484 based on the BAC is proposed for the reporting of turbidity at multiple scattering angles and
485 wavelengths of light. To conclude the discussion, some suggestions for the contents of potential
486 secondary standards (based on the newly proposed instrumentation standard) for surrogate SSC
487 determination are then outlined briefly.

488 **Suspended sediment concentration (SSC) and total suspended solids (TSS): their** 489 **relationship with turbidity and the importance of the PSD**

490 The surrogacy of physical properties for intrinsic optical properties as is the case regarding *chemical*
491 *concentration* becoming a surrogate for *optical turbidity* has raised the possibility of further
492 misinterpretation, due to the undefined PSD of the calibration standards and the inconsistent

493 response of different measurement instruments to the same PSD (Buzoianu 2000). In this section it is
494 necessary to take a step back from turbidity to examine the meanings of the pre-existing terminology
495 for suspensions (of sediment or otherwise) in water. It is important to understand this terminology as
496 the descriptive acronyms actually refer to documented test methods for the determination of
497 sediment concentration and suspended solids concentration. An understanding of these methods will
498 then facilitate a deeper appreciation of the reasons for the conceptual conflation of sediment
499 concentration with turbidity.

500 The US convention regarding the attribution of documented test methods to the acronyms “SSC” and
501 “TSS” has been adopted in this paper. Regarding this terminology, as with that of turbidity, the
502 differences in use in different disciplinary areas arises again. For example Holliday et al. (2003) suggest
503 TSS to mean “total suspended sediment concentration”, rather than “total suspended solids”, i.e. the
504 acronym SSC may have been a better choice.

505 The field techniques and laboratory methods for the measurement of SSC and TSS were reviewed by
506 Gray et al. (2000), who cite Method D 3977-97 (ASTM 1998) for SSC and Method 2540 D (APHA 1971)
507 for TSS. They describe the two different analytical methods as follows:

- 508 • SSC data are produced by measuring the dry weight of all the sediment from a known
509 volume of a water-sediment mixture.
- 510 • TSS data are produced by several methods, most of which entail measuring the dry weight
511 of sediment from a known volume of a subsample of the original.

512 After an analysis of 3235 paired SSC and TSS measurements was performed, it was concluded that SSC
513 was the more reliable methodology (Gray et al. 2000), especially when the amount of sand in a sample
514 exceeds approximately one quarter of the dry sediment mass. The main reason given for this disparity
515 of results is that the SSC analytical method utilises the entire sample (including all sediment present),
516 whereas the TSS methods typically involve the analysis of only a sub-sampled aliquot of the total
517 sample. The decanting and pipetting techniques employed to obtain this aliquot do not capture a

518 complete representation of the sediment population of the original sample. The resulting sub-sample
519 is therefore sediment deficient, particularly of the larger sand-sized sediment fraction. Gray et al.
520 (2000) go on to suggest that the reason for this loss of sediment during TSS analysis arises from the
521 fact that TSS methods were originally designed for analysis of waste-water samples that were to be
522 collected *after* an initial settling phase, hence larger sediment particles were never intended to be
523 part of the analysis. They finally conclude that SSC and TSS analysis of natural water samples are not
524 comparable, and that SSC is the only viable method for the determination of the sediment
525 concentration of natural waters.

526 In order to relate a subjective turbidity reading to a real physical property such as SSC, a calibration
527 procedure is typically performed. This relationship between the optical properties of suspended
528 sediment and its mass concentration must therefore be understood, requiring the characterisation of
529 its lithology. The size of the sediment particle is frequently measured either directly (e.g. filtering and
530 sieving), or analytically (by LASER diffraction) in the case of smaller size fractions. LASER-based particle
531 size measurements give a *volume* concentration value, which then requires further knowledge of the
532 specific density and mineralogy of the sample in order for an estimate of the *mass* concentration to
533 be obtained. This process is known as *end-member calibration*.

534 The problem now arises that the detector response has been pre-calibrated to a primary standard,
535 with arbitrary units for turbidity based on unstable calibration methods. It has already been suggested
536 ([Figure 7](#)) that these units (NTU etc.) are not comparable between calibrations made on instruments
537 constructed to the same standard. It is therefore highly unlikely that calibrations made by *different*
538 instruments (constructed to the same or different standards) can ever be accurately compared due to
539 the invalidity of these extrinsic turbidity units. It is therefore necessary to determine the true
540 instrument response by a different method entirely. Only then can an end-member calibration have
541 any chance of being meaningful.

542 Optical neutral density filters (ND filters) are regularly employed for the calibration of transmission-
543 based optical instruments, but are seldom employed in turbidimetry or nephelometry. These filters
544 provide a consistent optical density (OD) which in turn will attenuate a well-defined percentage of the
545 transmitted light. One such example of an attempt to calibrate a turbidimeter against a known light
546 attenuator is Finlayson (1985). By not only calibrating a turbidimeter against Formazin suspension, but
547 also against ND filters, Finlayson has devised a method by which direct comparison between
548 attenuation measurements made on the same sample by different devices could potentially be
549 developed. It can be seen that Formazin concentration does not in fact have a linear relationship to
550 measured light attenuation ([Figure 8](#)). Although the calibration data are sparse in the upper range of
551 the instrument in this case (Finlayson 1985), there is a good fit of the data to a power law ($R^2 = 0.9954$).
552 The only two useful axes on this graph are “meter reading” and “neutral density filters”, as these two
553 alone are all that is required to accurately establish the response of the instrument to attenuation
554 ([Figure 9](#)). Only when this detector attenuation curve has been established can further selective end-
555 member calibrations be performed to determine the effect the PSD has on the response of a particular
556 instrument to a given sediment. Each ND filter represents an optical density, d , which is directly
557 equivalent to the absorbance A , as in [Equation 4](#). So in order to calculate the BAC in dB m^{-1} for an
558 instrument with path-length l , the following equation can be applied ([Equation 13](#)):

559
$$BAC = \frac{10 d}{l} = \frac{10 ND}{l} \quad (13)$$

560 **[insert Figure 8.]**

561 **[insert Figure 9.]**

562 Instrumentation parameters and calibration methods

563 To arrive at a *consistent* methodology for the measurement of turbidity it is necessary to accept that
564 the only quantity that can be readily measured optically in this context is the transmitted light
565 intensity, and hence *attenuation* with respect to the light source (i.e. I/I_0). It is the methodology for

566 taking this measurement that should be rigorously specified, regardless of the measurement angle ϑ
567 with respect to I_0 . The implementation section of the standard should address this methodology, and
568 focus purely on the desired response of the instrument to light at defined intensities and wavelengths.
569 This aspect of work would involve the definition of parameters such as sensor type, variable intensity
570 light source specification (including coherence and polarization), detector amplifier gains and ranges,
571 ND filter calibration procedure involving multiple beam paths, beam path-length and collimation
572 arrangements. It is then necessary to decide which instrument parameters (e.g. ϑ , λ and I) should be
573 specified as mandatory for all turbidity measuring instruments, and which ones should be considered
574 as being application-specific.

575

576 The reporting of turbidity measurement data

577 The standardization of the reporting of turbidity as attenuation data (Ziegler 2003) and the use of a
578 more descriptive nomenclature is proposed, which will allow for the easy identification of application-
579 specific data such that incompatible measurements will not be inadvertently compared to each other.
580 It is suggested that significant progress could be made if the measurement concepts for turbidimetry
581 and nephelometry were unified, i.e. by treating them *both* as an attenuation process. The only
582 difference being that for scattered light measurement the effective concentration of scatterers is
583 *inversely proportional* to the BAC measured at a specific angle to the incident beam. However, for that
584 to be achieved formulations of the BAC at specific angles must then be defined, for example BAC_0 for
585 a standard transmissivity measurement and BAC_{90} for the nephelometric counterpart at 90° . For the
586 nephelometric case the relationship between the scattered light intensity and the concentration could
587 be viewed as an *inverse attenuation*, since a higher concentration of particles will produce stronger
588 scattering (until the concentration is too high, at which point multiple-scattering and grain-shielding
589 will dominate and interfere with the measurement of the side-scattered light). Measurement-
590 instrument calibration now becomes somewhat critical, as any drift in the incident light intensity or

591 the sensor response will affect the sensitivity of the system to the low light intensities that need to be
592 detected due to side- or back-scattering. This nephelometric BAC_{90} measurement results in potentially
593 larger percentage errors than those that are likely for measurements based on BAC_0 , as greater
594 electronic amplification is required to detect the weaker scattered-light signal which can be inherently
595 noisy. In order to formulate a generic equation for the BAC as a function of measurement angle it is
596 necessary to include two terms: one for attenuation and one for scattering. The use of these terms is
597 in no way a new idea (e.g. Kirk 1985), however the interpretation of scattered light intensity as an
598 *inverse absorbance* has not been previously considered. In this new method the same measurement
599 units could be employed for practical comparison between data obtained under different conditions
600 using different instruments, so long as those instruments complied with the same instrumentation
601 standard, and the reporting of said data is consistent (Ziegler 2003). For example Kirk (1985) suggested
602 using the correct description of the measurement method, such as “side-scattering”, when stating
603 results – or preferably BAC_{90} in this case.

604

605 Standards for surrogate SSC determination

606 Further standards for the determination of surrogate properties such as SSC should refer to
607 instruments that are specified according to the new instrumentation standard. In order to estimate
608 SSC accurately, optical instruments must be capable of producing data rich enough to facilitate
609 suspended sediment characterization. Methods for the determination of the PSD (and other
610 properties) of a suspended sediment by multi-parameter measurements need to be developed, which
611 could include the use of LASER diffraction techniques. Other potential methods of sediment
612 characterization should also be explored more thoroughly.

613

614 **Suspended sediment characterization**

615 For a deeper understanding of sediment transport to be realized, it is essential to know how the
616 different size-classes of sediment respond to different flow conditions, especially the larger sand-sized
617 particles that can be transiently in suspension long enough to affect turbidity measurements. A
618 knowledge of sediment particle shape in terms of *sphericity* and *roundness* can also provide an insight
619 into the distance travelled by sediment particles that have previously been entrained in a flow of
620 water. There is a clear need therefore to characterize the suspended sediment to determine the
621 particle sizes present. This characterization can be achieved by traditional gravimetric sampling
622 methods, but there is an increasing need to gather data for research purposes *in-situ* and quickly. In
623 some cases, these measurements could be made “off-line” by optical means, which would still be
624 much faster than can be achieved by gravimetric methods. LASER-based optical measurements are
625 the most commonly employed for this purpose, although there have been attempts to derive particle-
626 size information from multi-parameter turbidity measurements. The effect that particle shape has on
627 such measurements could also be exploited as a characterization technique.

628 **Measurement ratios and multi-parameter method development**

629 The designers of some turbidity meters (i.e. any commercially available instrument that claims
630 compliance with GLI Method 2) have adopted the use of multi-parameter measurements in order to
631 improve instrument performance. This innovation has included the measurement of light intensities
632 at multiple scattering angles, and the use of the ratios of those intensities to infer some of the physical
633 properties of the scattering suspension, e.g. sphericity (Gibbs 1978), or to negate the effect of water
634 colour as an interference to the turbidity measurement (Lawler 2005, Lambrou et al. 2009). An
635 example of another multi-parameter approach to turbidity measurement is presented by Yang & Hogg
636 (1979), wherein two different wavelengths of light are used to predict the PSD of the scattering

637 suspension. These and other multi-parameter approaches to turbidity measurement should be the
638 focus of further research, and will aid the development of new turbidity standards.

639 Conclusions

- 640 1. The use of turbidity purely as an indicator of water clarity is entirely acceptable assuming the
641 development of more consistent standards. The problem is that the existing standards have
642 introduced a set of measurement units that actually represent a surrogate for turbidity and
643 therefore cannot be used to describe water clarity.
- 644 2. Simple turbidity measurements when used as a surrogate for suspended sediment
645 concentration are only viable under highly constrained conditions. Bias toward the fine
646 sediment fraction is usually considered unimportant, but this is not always the case.
- 647 3. Sand-sized sediment fractions are not consistently accounted for by existing turbidity
648 measurements, due to their high settling velocities. The SSC method is also required in order
649 to quantify the sand fraction fully.
- 650 4. The development of new light-scattering models will permit more sophisticated approaches
651 to turbidity measurement, in particular by the use of parameter-rich data sets obtainable from
652 multi-parameter methods. This approach will facilitate the improvement of turbidity
653 standards, and could increase the accuracy of large sediment particle detection.
- 654 5. A new turbidity instrumentation standard needs to be drafted, based purely on the principle
655 of attenuation for calibration and reporting purposes. It should specify the reporting of the
656 BAC in dB m^{-1} (or derived units) for a range of measurement angles and wavelengths of light.
657 This standard should be a root standard from which other secondary standards are derived,
658 e.g. standards for suspended sediment characterisation or total suspended solids assessment
659 by optical turbidity measurement.

660 6. A further standard for suspended sediment determination by simple multi-parameter
661 turbidity measurements needs to be devised (leading on from point 4 above). This standard
662 should include basic sediment characterisation as an outcome of optical turbidity
663 measurements (e.g. PSD and sphericity).

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